

Part III Conclusion and Proposal

Chapter 1: Conclusion

This project started in anticipation of whether we can find Noril'sk style copper-nickel-PGE sulfide ore deposits in the Paraná flood basalts. Studies on LIPs (Large Igneous Provinces), including the flood basalt that formed ore deposit of this type, have just made rapid development since 1995. Because discussions are still on the progress on the characteristics and origin of LIP, researchers have naturally different opinions.

Given this situation, this survey began, based on existing literature, by reviewing the characteristics and genesis of flood basalts around the world, chemical stratigraphy and geological structure of the Noril'sk ore deposits and Siberian trap, and considering the exploration criteria of the Noril'sk style ore deposits. According to the exploration criteria studied, the approaches from analyzing both the geological structure and the geochemistry of the basalts were conducted in order to extract an area with similar geological environment to the Noril'sk region in the Paraná Basin. That is, the geological structure analysis of the entire basin was carried out using the existing documents, satellite images, and gravity/geomagnetic data. At the same time, the analysis of the major chemical compositions and trace elements, such as Pt, Pd, Au, and REE, and isotope analysis for rock samples collected during the survey were also carried out. This resulted in defining the geochemical characteristics of the Paraná flood basalts. Based on comprehension of these results, we carried out study for the area that has a potential for existence of ore deposits.

1-1 Genesis of Flood Basalt and PGE Sulfide Ore Deposit

A number of theories exist concerning the genesis of flood basalt, including a theory that supports active involvement of the mantle plume and a theory that emphasizes ascensions of the asthenosphere in continent breaking up. However, there has been no established theory yet concerning the genesis of flood basalt. Since 1995, a new theoretical model has been attracting researchers. According to this model, vast amounts of oceanic crust (slab) that subducted from continental margins deposited at the deep of the mantle, and ascended as heterogeneous mantle plume. The flood basalt magma characteristically erupted huge amounts of lava in a short period generally within 1 Ma or less. The heterogeneous plume that consists of basalt constituent having low melting temperature and peridotite constituent. In addition to such variety of the theory, magma activities of flood basalt vary depending on the location, so

the theory of flood basalt formation does not necessarily lead to direct linkage to the existence of PGE ore deposits.

Most PGE concentrate in the earth's core and are rarely included in the continental crust that is markedly differentiated. Therefore, it is reasonable to think that the original magma of PGE deposit was brought from the deep part of the earth. The source materials of flood basalt that yielded the PGE mineralization may have been brought from mantle plume that ascended from the core and mantle boundary zone. The magma is also expected to be generated with large scale melting of plume and surrounding mantle. During the formation of flood basalt, large magma quantities are generated compared to periods of ordinary igneous activity. At this stage, PGE that ordinarily concentrates in solid phases is considered to have a high possibility to be caught in the melt with relatively large quantity. Since PGE concentration is affected by factors of the supplies of SiO₂ and sulfur by crustal contamination, and the local tectonic setting, such factors that influence to PGE concentration are important for forming PGE ore deposits from the magmas generated by the above process.

1-2 Exploration criteria of the Noril'sk style deposit

Assuming that the Noril'sk ore deposits in Russia belong to a type of orthomagmatic sulfide ore deposit, we considered the geological environment that satisfies the common requirements for the genesis of orthomagmatic sulfide has been formed connected to the flood basalt magmatism in the Noril'sk region. Based on such standpoint, a document survey regarding the world orthomagmatic sulfide deposits and the Noril'sk ore deposits was carried out.

The magmas related to the genesis of the Noril'sk ore deposits were initially sulfur undersaturated and rich in PGE, and were also picritic to basaltic compositions of high temperature. The Noril'sk ore deposits were formed in an area with large quantities of such magmas, that is the center of plume related magmatic activity. Considering the forming environment of the Noril'sk ore deposits from the view point of the condition of orthomagmatic sulfide ore deposit forming, the follows are presented as the exploration criteria for orthomagmatic ore deposit accompanied with flood basalt:

- (1) The existence of sulfur undersaturated PGE rich magmas as lavas or intrusions.
- (2) The existence of magmas showing the signatures of crustal contamination and sulfide segregation associated with above PGE rich magmas as lavas or intrusions.
- (3) It is desirable that the magmas are high temperature picritic to basaltic composition that easily contaminates the crust materials.

- (4) The center of the magmatic activity where the large volume of silicate magmas are supplied through crustal suture zone.

1-3 Geology and Geological Structure of the Paraná Basin

After Paleozoic in age, intracratonic basins such as Parana Basin, etc. were formed in the Gondwana continent. M. S. M. Mantovani et al. (2001) calculated the thickness of lithosphere of South-American plates using the tidal gravity anomalies. According to the results, the thickness of lithosphere of Parana Basin is 10 to 20 km or more thinner than that of surrounding basement rocks. At this kind of portion with thin lithosphere, intracratonic basins are formed by changes of regional stress field and mantle plume (Bott, 1971; Makenzie et al., 1978). The Noril'sk ore deposit was formed in Gydansk-Omsk rift (Zonnenshain et al., 1990) and Duluth ore deposit in Midcontinent Rift (MCR) of North America.

From Paleozoic to Mesozoic in age, the Parana Basin was repeatedly deposited and eroded due to dilation and constriction of mantle plume. The present total thickness of the sedimentary rocks attains a maximum of more than 4,000 meters. In Early Cretaceous in age, tholeiitic basaltic magma was active in rift and subsequently Gondwana continent was ruptured (Rezende, 1972). The sedimentary rocks and volcanic rocks of Parana Basin can be divided into the following six Supersequences (SS) according to geologic ages (Milani et al., 1998; Fig. II-1-4-9). They are the Rio Ivai SS (Ordovician to Silurian), the Parana SS (Devonian), Gondwana I SS (Late Carboniferous to Early Triassic), the Gondwana II SS (Early to Middle Triassic), the Gondwana III SS (Early Cretaceous) which mostly comprises tholeiitic flood basalt (maximum thickness 1,723 m) and the Bauru SS (Late Cretaceous).

As the flood basalt lava is distributed thickly in the NE-SW direction along the Parana River, the area is estimated to have been a rift. Drillings for oil exploration revealed that a large amount of sills and dykes associated with flood basalt magmatism intruded into the rift. There are two zones which accumulated thickness of sills and dykes are larger. The one is an area where extends in the NW-SE direction from the east of Parana River to Ponta Grossa Arch and the other an area where are located in the boundaries between Matto Grosso do Sul Province and Sao Paulo province in the northern area of the Basin and extends in the N-S direction. In the TM false color image (RGB=234), they are expressed as geographic features uplifted in the form of thin strings and sometimes continue over 100 kilometers.

Based on the residual gravity anomaly map (M. C. L. Quintas, 1995), A large and high anomaly is present along Parana River in the NE-SW direction of the central area of Basin. The high gravity anomaly of the rift may be due to the density difference between asthenosphere (3.33 g/cm^3) and lithosphere (2.80 g/cm^3). Other large high anomalies and large low anomalies are considered to indicate basement structures blocked by faults in the N-S

direction, NE-SE direction, and NW-SE direction.

With respect to the basement structure of the central region of the Basin based on aeromagnetic anomalies, large and high pseudogravity anomalies are recognized in the N-S direction in the eastern part and the western part, while a large and low pseudogravity anomaly is recognized in the N-S direction at the central part of the central region. The former are inferred to be regions with basement rocks in comparatively shallow depth and the latter a region where basement rocks collapsed due to flood basalt magmatism. The high pseudogravity anomaly in the western part may be possibly a part of Paraguay-Araguay belt and the high pseudogravity anomaly in the eastern part a part of Sao Francisco craton under the flood basalt lava pile. In the large low anomaly of the central part, high anomalies and low anomalies are recognized in the NE-SW direction, which run nearly in the same direction as Parana River. The low anomaly along the right side of Parana River coincides with high residual gravity anomaly and thick part of flood basalt lava.

Many magnetic lineaments which indicate faults are present on the pseudogravity anomaly map and so on. Principal directions of faults are NW-SE direction, NE-SW direction, and E-W direction. The magnetic lineaments in the NW-SE direction are continuous faults which have sharp boundaries. The magnetic lineaments in the E-W direction are poorly continuous. Most of faults in the E-W direction seem to be secondary shear faults derived from sinistral wrench faults in the NW-SE direction. The unclockwise rotations of them coincide with those mentioned by Turner et al. (1994) and W.M. de Rezende (1972). A large amount of sills and dykes are recognized in the areas where Rio Alonzo fault in the NW-SE direction and rift along Parana River in the NE-SW direction intersect.

1-4 Promising Areas of Emplacements of Ni-Cu-PGE Ore Deposits Estimated from Geological structure

We will present a comprehensive explanation of Precambrian basement rocks, sedimentary rocks and flood basalts in Post- Early Paleozoic, and faults in the Parana Basin and consider the relation among the rift zone and the regional stress field, and promising areas of emplacement of Noril'sk-type Cu-Ni-PGE ore deposit (Fig.II-1-5-25).

A large low anomaly zone in the central part of the pseudogravity anomaly map overlaps with a zone where sedimentary sequence and flood basalt lava are thicker, with a zone where a large quantity of sills and dykes intruded, with a high residual gravity anomaly zone, and with a high attenuation factor zone. Based on these facts, it is concluded that the large low anomaly zone was rift zone associated with flood basalt magmatism.

The direction of the rift zone extends in the N-S directions in the northern Basin; in the NE-SW directions the central Basin, and branches into the NE-SW and NW-SE directions in the

southern Rio Alfonzo fault. The inside of the rift zone has repetitions of high anomaly zones and low anomaly zones in the NE-SW direction. These anomaly zones in the NE-SW directions were formed by stair-step normal faults (vertical direction: maximum compressive principal stress axis σ_1) in the extensional stress in the NW-SE directions field (minimum compressive principal stress axis σ_3).

Many magnetic lineaments which represent faults in the NW-SE direction are recognized on the pseudogravity anomaly map. For example, Rio Alfonzo fault on the geological map are represented by several magnetic lineaments on the pseudogravity anomaly map. The faults in the NW-SE directions are estimated to be sinistral transcurrent faults because they have sharp boundaries and are accompanied with many secondary shear faults in E-W directions. In the Duluth ore deposit, many secondary shear fractures occurred in the form of Cymoid Loop-like tensional fractures in areas where the rift zone and transcurrent faults cross. The Duluth Complex intruded there and formed the ore deposit. The Noril'sk ore deposit was also formed in the volcano-center where magma extruded most intensely in the Gydansk-Omsk rift zone (Tamrazyan, 1962).

As the results mentioned above, we propose the following four promising areas, which rift zone and transcurrent faults cross, of emplacements of Ni-Cu-PGE ore deposits.

- (1) The area where the Tiete structural line in the NW-SE direction crosses through Paran River
- (2) The area where the Guapiara fault in the NW-SE direction crosses through Parana River.
- (3) The area where Rio Alfonzo fault in the NW-SE direction crosses through Parana River.
- (4) The western part of Torres - Posadas fault in the southern region of Parana Basin

All of these areas are estimated to have been the volcano-center of flood basalt magmas. The largest volcano-center is found in the area (3) where the Rio Alfonzo fault crosses through Parana River. Among four areas, it is concluded that areas where the Guapiara fault and the Rio Alfonzo fault in the NW-SE directions cross through Parana River are the most promising area of emplacements of Ni-Cu-PGE ore deposits.

1-5 Geochemical Characteristics of Paraná Flood Basalts

(1) Classification of Magma Type

Six types of lava units of Paraná flood basalts are divided into Low-Ti type and High-Ti type by Peate et al. (1992). Since the Ti contents of Paranapanema and Ribeira are intermediate in this survey, the two above were geochemically judged not to be classified into Low-Ti type nor High-Ti type. Therefore, the two above were newly classified as the

Intermediate-Ti type, and the magma types of Paraná flood basalts were classified as follows. The difference of Ti contents among the three types is obvious, and each type shows different trends of fractionation in Mg[#]-TiO₂ diagram.

	Peate et al. (1992)	This survey
Low-Ti type	Gramado	Gramado
	Esmeralda	Esmeralda
	Ribeira	
Intermediate-Ti Type	—	Paranápanema-Ribeira
High-Ti Type	Paranápanema	Pitanga
	Pitanga	Urubici
	Urubici	

Geochemical examinations were carried out for both lava samples and intrusion samples in this survey. Both samples show entirely the same geochemical trend. Therefore, both lava and intrusion samples were classified into the above five magma types. As for distribution, Pitanga is distributed in the northern part of the Paraná Basin, Pitanga and Paranápanema-Ribeira are distributed in the central part, and Gramado, Esmeralda and Urubici are distributed in the southern part. Pitanga and Paranápanema-Ribeira are intermingled in the intrusions of the Ponta Grossa Arch in the central part. The results of the geochemical examinations are summarized in Table II-4-5-1.

(2) Partial Melting and Crustal Contamination

Pitanga, High-Ti type, is distributed from the northern to the central part of the Paraná Basin. This magma type shows the smallest degree of partial melting. It is thought to have been generated in relatively deeper portion in the upper mantle. Paranápanema-Ribeira, the Intermediate type, is distributed in the central area of the Paraná Basin. The degree of partial melting of this type is shown larger than that of Pitanga. This means that it has been generated in a shallower portion than Pitanga was generated. The effect of the crustal contamination is considered limited in these two types of magma because of poor contents of U and Th that concentrate in the crust. Mantovani et al. (2000) described that the two types above were generated from a similar magma source with different degrees of partial melting. The large quantities of the two magma types above erupted and were piled with a thickness of approximately 1,800 meters along the Paraná River in the central part of the Paraná Basin, which is thought to be the center of the magmatic activity.

The Low-Ti type magmas (Gramado and Esmeralda) show relatively large degree of partial melting. They are thought to have been generated in relatively shallower portions compared to the other magma types. The increases of LIL elements such as Th, U and the likes that concentrate in crust, and Sr isotopic ratio are observed in both magma types. It is thought to be the effect of crustal contamination. The effect is particularly marked in Gramado. Esmeralda is considered to be the most primitive magma among all the magma types because of its low contents of light rare earth elements and a relatively high Nd isotopic ratio.

As regards the Low-Ti types of magma, the following three points are thought to have a close relationship. The first is that the magma was generated in a shallow portion. The second is that the magma was highly affected by crustal contamination, and the third is the existence of acidic rocks, which are considered to be the products of lower crust melting, in the southern part of the Paraná Basin. More specifically, the following process is thought to have occurred. That is, the acidic magma was generated from the melt of the lower crust by the mantle plume heating. After generating the acidic magma, this magma and the magma generated by the melting of the mantle plume itself mixed and yielded Gramado type magma. Finally, just after the commencement of expanding the Atlantic, Esmeralda type magma is thought to have been generated that is less influenced by crustal contamination although the origin is the same as Gramado.

(3) PGE Contents of the Paraná Flood Basalts

The original magmas of the Paraná flood basalts are thought to be PGE rich as a whole, particularly Pd rich (Refer to Table II-3-5-1) although there are some differences among the magma types. The most PGE enriched magma is Paranapanema-Ribeira. The average content of Pt and Pd summing is approximately the same as that of the Tuklonsky lava, the most PGE enriched lava in the Noril'sk region. Paranapanema-Ribeira is also rich in Cu. The magma of this type as lavas is distributed along the Paraná River in the central part of the Paraná Basin, and much native copper is observed in the cavities. The geochemical anomaly in PGE, Ni, Cu, and other elements that is observed in the western part of the state of Paraná (by the stream sediment geochemical survey carried out by MINEROPAR) possibly shows the distribution of Paranapanema-Ribeira that has high PGE contents. Paranapanema-Ribeira as intrusions is distributed in the Ponta Grossa Arch intermingled with Pitanga of High-Ti type. The magma type of the higher PGE content next to Pitanga is Esmeralda that has the most primitive chemical composition. The PGE depletion by sulfur saturation is rarely observed in Paranapanema-Ribeira and Esmeralda.

Regarding Gramado and Pitanga, there are two populations of samples, PGE enriched and depleted. These two types are thought to have been initially PGE enriched. The depletion of PGE, however, is thought to have occurred due to sulfur saturation in a part of the magma. As

the cause of sulfur saturation, the crustal contamination in Gramado, and the crystallization differentiation in Pitanga are both thought to have occurred. Pd contents of 133.8 ppb were obtained by analyses of the intrusion sample (KN104) of Pitanga type that was collected in the Ponta Grossa Arch. This value exceeds the Pd content of ordinary basaltic magma, and suggests the existence of sulfide that includes Pd.

According to the facts above, all magma types of Paraná basalts are considered to be initially sulfur undersaturated and rich in PGE. They have the potential to form the ore deposits if satisfy other requirements demonstrated above.

1-6 Mineralization Potential of the Paraná Flood Basalts

In the Noril'sk region, the majority of the lava and intrusion is basaltic, however, the sills which embed ore deposits contain picritic parts and the assimilation of country rock on a large scale in the picritic sill is observed. Picritic magma, however, is rarely observed in the Paraná flood basalts, and the basalts are more differentiated than in the Noril'sk region. The magma temperature of the Paraná flood basalts is thought to be lower than that of the Noril'sk region. It is thought to be possible that this low temperature obstruct the assimilation of country rock and the large scale segregation of immiscible sulfides. As assumed in Pitanga, though, there is a possibility of segregating immiscible sulfides by the progress of the simple crystallization differentiation without the influence of contamination. The saturation mechanism on sulfur mentioned above is the same process considered in the Skaergaard intrusion in East Greenland (Anderson et. al., 1998) and the Sonju Lake intrusion in Minnesota, USA(Refer to Appendix 2).

All the samples of intrusion including the Ponta Grossa Arch sill that could be collected in this survey were distributed in the marginal area of the Basin, the edge area away from the center of magmatic activity. The sills are mostly poor in sulfide, and the maximum thickness is approximately 200 m, which is comparatively thin. The rock compositions show basaltic, and although show high PGE contents, the magma temperature is lower than that of the magma of Noril'sk. The contamination effect of country rock in such intrusions is weak. This is considered to be the reason why only a faint dissemination of sulfides is observed (pyrite/pyrrhotite/chalcopyrite/sphalerite). The segregation of sulfides in a large scale magma chamber like the Skaergaard intrusion under the high R factor is necessary to generate sulfide ore deposit from basaltic magma with high PGE contents.

The location with the highest possibility of generating sulfide ore deposit is thought to be the eruption center of Paranápanema-Ribeira and Pitanga types, which have the largest quantities of magmas among the Patana flood basalts. The eruption center is thought to be located in the central area of the Paraná Basin along the Paraná River where the lava piles are the thickest. Based on the existing data analysis, the strong shrinkage of lithosphere by the

ascension of the asthenosphere is observed in this area, and magma intrusion is estimated to be the most active in this area.

Chapter 2: Proposal

From analyses of existing geological papers, geological survey on the ground surface, geophysical anomaly maps, drilling data of coal exploration and so on resulted in identifying geotectonic characteristics of the Parana Basin and geochemical characteristics of flood basalt magmas and possibility of emplacements of Ni-Cu-PGE ore deposits such as the Noril'sk, Duluth and Skaergaard ore deposits.

According to the dykes and sills in the basement rocks, although the drilling cores of coal exploration and the outcrops were investigated, no Ni-Cu-PGE mineralization was almost identified.

Hereafter, diamond drillings are proposed for identifying mineralization of dykes and sills in the basement rocks in the promising area where have been located at volcano-centers of flood basalt magmas. But, as most of volcano-centers are located at more than 1,000 meters below the ground surface, it is difficult for ore deposits to be economic.

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APPENDIX

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APPENDIX 1

Report of
the 9th
International Platinum Symposium

- Summary of the symposium -

9th International Platinum Symposium

Summary

The 9th International platinum symposium was held at Billings City, State of Montana, USA from July 21 to 25, 2002. Such a symposium has been held every four years to report results of platinum ore deposit studys around the world. This symposium had 70 presentations of lectures and 61 exhibitions of posters.

The participants totaled 204, including 83 from the United States, 49 from Canada, 13 from South Africa, 10 from Finland, 9 from the United Kingdom, 6 each from Brazil and Japan, 4 from China, 3 each from Germany, France, Albania, and Australia, 2 each from Russia and India, and 1 each from Denmark, Austria, Italia, and Egypt. The titles of abstracts and major presentations of platinum ore deposit studys are stated below.

Platinum-Group Elements in the Upper Zone of the Rustenburg Layered Suite of the Bushveld Complex: Preliminary Results from the Bellevue Borehole

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The Rustenburg Layered Suite (RLS) is generally divided into five zones (Hall, 1932): the basal Marginal Zone (0-800 m of norites), overlain by the Lower Zone (800-1300 m of cyclic units of harzburgite and pyroxenites), the Critical Zone (1300 to 1800 m of cyclic units of chromitite, pyroxenites, and norites), the Main Zone (3000 to 3400 m of norites, gabbronorites and anorthosites), and the Upper Zone (2000 to 2800 m of cyclic units of magnetites, gabbronorites anorthosites, and diorites). The Lower Zone and Critical Zone are enriched in Platinum-group elements (PGE) relative to the marginal rocks (Fig. 1). This enrichment culminates in the formation of the world famous Merensky and UG-2 reefs. In contrast the Main Zone contains little PGE. The PGE content of the Upper Zone is poorly documented. Page et al. (1982) published an initial survey, of Pt, Pd and Rh concentrations of the Upper Zone. For most of the rocks the PGE were present at less than the detection limits. However, Page et al. (1982) do report one location with high values in the lower portions of Upper Zone. Harney et al. (1990) investigated 17 samples from the 300 meters of stratigraphy around the Main Magnetite layer. They found that the samples above the Main Magnetite layer have low levels of PGE (Pt+Pd = 5-10 ppb). Samples below the Main Magnetite layer were somewhat richer (Pt+Pd=40-100 ppb). Subsequently, Harney and von Gruenewaldt (1995) quoted company reports that suggest the presence of up to 6 ppm Pt + Pd in the upper part of the Main Zone. They attribute this enrichment and the enrichment reported by Page et al. (1982) to local hydrothermal action.

As part of a larger project "The Bushveld Complex: its origin, crystallization and interaction with the Kaapvaal lithosphere" headed by Prof. Lew Ashwal we have determined PGE, Ni, Cu, S and full major and trace element analyses in 56 samples of Upper Zone from the Bellevue borehole. This borehole provides a complete section through the Upper Zone and the upper part of the Main Zone. The aims of our project are:

a) To characterize the distribution of the PGE in Upper Zone;

b) To investigate whether there is any enrichment of Pd and Au as is found in the upper parts of the Skaergaard intrusion (Andersen et al., 1998). It is intriguing to note that in the Bushveld the reefs have higher Pt/Pd ratios than the silicate liquids from which they formed (2.3 vs. 1.5) requiring that Pd be concentrated elsewhere (Barnes and Maier, 2002). In contrast the Platinova reef of Skaergaard contains little Pt implying that it has been extracted from the magma before the formation of the Platinova reef. Further the Bushveld reefs form from a relative unfractionated magma with chromite and orthopyroxene on the liquidus. The Skaergaard Platinova reef forms from a highly fractionated magma. Could there be a Platinova type reef in the Upper Zone of the Bushveld, or a Merensky type reef in the lower parts of Skaergaard?

c) To investigate whether there is any PGE enrichment of the type reported by Harney and von Gruenewaldt (1995) and Page et al. (1982).

d) To investigate whether magnetite will concentrate Os, Ir, and Ru. Capobianco et al. (1994) have shown in experiments run at high oxygen fugacity that magnetite can concentrate Ru by analogy one would expect Os and Ir to also be concentrated in the magnetite.

Our preliminary results show that for most of the Upper Zone the PGE concentrations are extremely low. In most rocks Pt and Pd do not exceed 1 ppb. Osmium and Ru are present at less than detection limits of 0.2 and 1 ppb respectively. Iridium concentrations range from 0.08 to 0.3 ppb, Au from 0.1 to 7 ppb, Rh from <0.1 to 0.5 ppb. There does not appear to be an enrichment of Pd over Pt so the Pd "missing" from the reef rocks remains to be found. There is an enrichment in all of the PGE, with Pt+Pd = 3 ppm in the cyclic unit of the lower most magnetite layer of the Upper Zone. While it is difficult to make a direct stratigraphic correlation with the enrichments reported by Page et al. (1982) and Harney and von Gruenewaldt (1995), the enrichments are in approximately the same stratigraphic position, i.e. close to the change from the Main Zone to the Upper Zone. The magnetite layers do not appear to be enriched in Os, Ir, and Ru and hence magnetite

does not appear to have concentrated these elements. Iridium, Pd, Re, Au, Se and Cu all show a positive correlation with S suggesting that sulfides control these elements.

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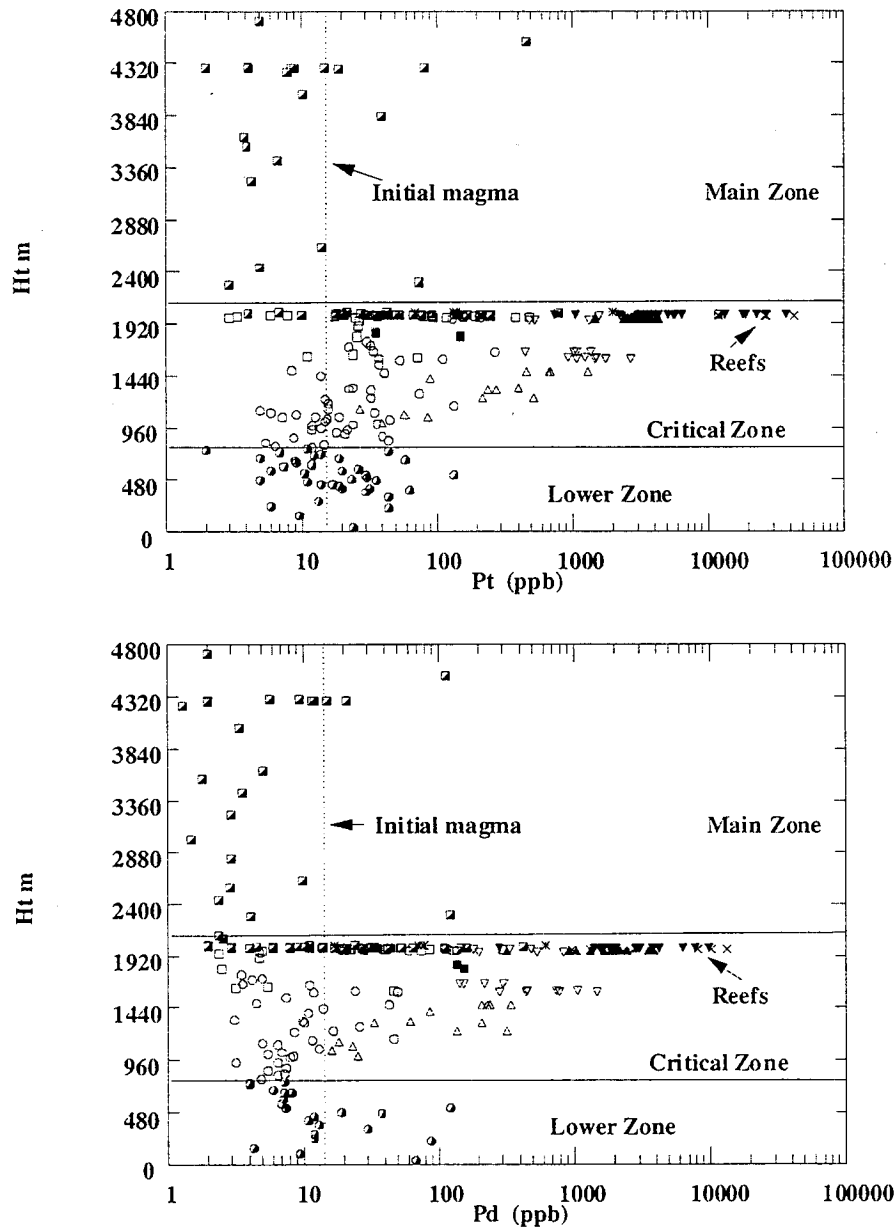


Figure 1. Variations in Pt and Pd content through the Bushveld Complex. Data from Barnes and Maier (in press) and references therein. Circles = peridotites and pyroxenites, triangles = chromitites, x = reefs; squares = mafic rocks.

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The Stillwater and Bushveld Magmas Were Wet!

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Basic magmas that crystallized layered igneous intrusions such as the Stillwater complex have traditionally been considered to have been relatively "dry", largely because of the paucity of hydrous minerals. However, investigations of halogen variations suggest that the high-Cl, high-Mg parent liquids of the Stillwater and Bushveld complexes, like their modern-day boninitic counterparts, contained a much higher volatile content than previously recognized - possibly in excess of 1.0 wt. % H₂O (Boudreau *et al.*, 1997). Although volatile components can have a marked effect on crystallization behavior of these magmas, most existing models for the crystallization of these intrusions and the formation of their economically important platinum-group element deposits either ignore or downplay their potential importance.

Evidence for high volatile concentrations comes from several lines of evidence: 1) Presence of hydrous phases as common late (interstitial) minerals, 2) high Cl/F ratio of halogen-bearing minerals, 3) Cl and H isotopes consistent with a primary origin of the volatile components and 4) boninite-like nature of Stillwater and Bushveld magmas.

The range of apatite compositions in the

Stillwater and Bushveld Complex and a few other intrusions are shown in figure 1, as well as chilled margin and associated contemporaneous sills dikes. In most intrusions, apatite is predominantly fluorian, with the mole fraction chlorapatite component typically less than 40 mole % and more commonly less than about 20 mole %. Furthermore, the apatite compositions generally become poorer in Cl (and richer in F) in the more evolved cumulates. These Cl-poor compositions are reflected in apatite in associated sills, dikes and chilled margins (Boudreau 1995). In contrast, those from the Stillwater and Bushveld complexes at and below the major PGE reefs are unusually Cl-rich.

This is consistent with both Cl and H isotopic evidence as well (fig. 2). While Stillwater Cl isotope values are equivocal (but see below), most Bushveld samples overlap the range of values for pristine MORB and altered seafloor samples ($\delta^{37}\text{Cl} \sim +4 \text{‰}$ relative to sea water). This contrasts with the sedimentary floor rocks ($\delta^{37}\text{Cl} \sim +0 \text{‰}$), with little evidence of contamination away from the marginal rocks. The Cl isotope composition and high Cl/F ratio is therefore a fundamental characteristic of the first magmas and not derived from marginal samples.

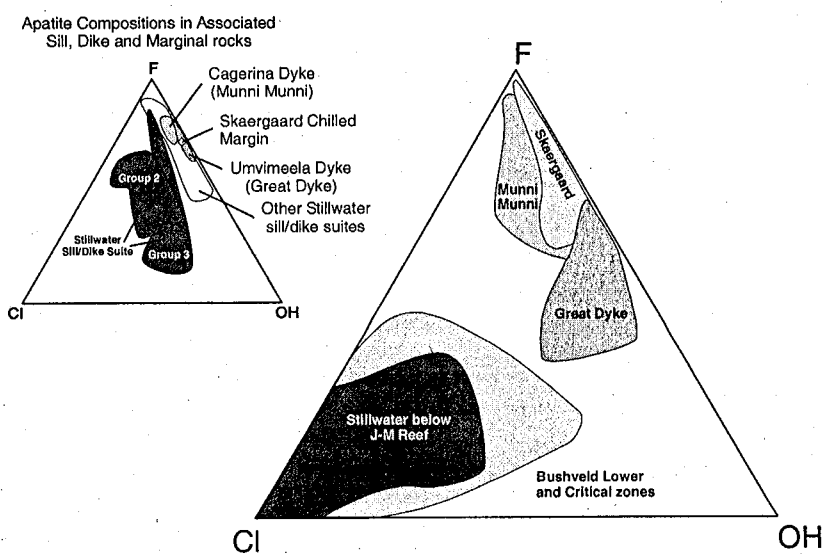


Figure 1. Composition apatite from some layered intrusions. Inset at left: Composition of apatite from coeval sills and dikes as well as those from chilled margins.

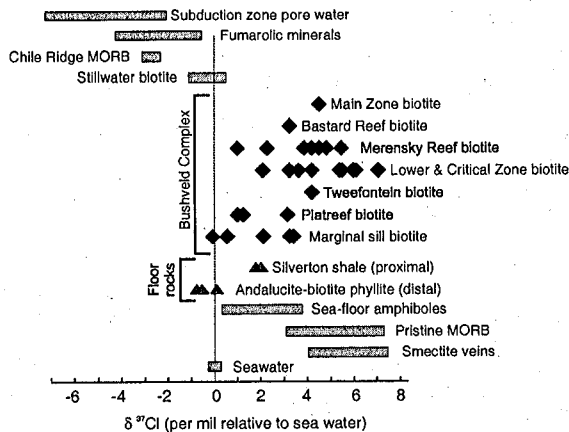


Figure 2. Comparison of Cl isotopic variations of the Stillwater and Bushveld complexes with those of other geologic environments. Data from Boudreau et al (1997), Willmore et al. (2002) and reference therein.

Similarly, a number of studies have shown that the O and H isotopes are largely constant with the volatile components being magmatic and not introduced. Of interest for this discussion is the overlap of Bushveld H isotopes with those of boninites (fig. 3).

The Boninite Analog

The Stillwater group 2 and 3 coeval sills and dikes, as well as the B1 sills of the Bushveld Complex, are high-Mg basalt with boninitic affinities, i.e. generally low TiO₂, but high K₂O and light REE (LREE) contents (fig.2). As has been proposed for boninites (e.g., Hickey and Frey, 1982), these compositional characteristics suggest that these magmas were generated from mantle that had undergone a previous basalt extraction event—but was subsequently enriched by a metasomatic (fluid) agent (Hatton and Sharpe, 1989; Wooden et al., 1991). An early melting event would have resulted in loss of halogen-bearing minerals, leaving a source depleted in F as well as Cl. Later infiltration by a Cl-rich metasomatic agent (fluid or volatile-rich silicate liquid) would have produced a metasomatized mantle with a high Cl/F ratio.

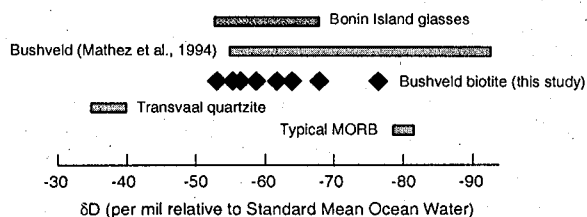


Figure 3. Comparison of H isotopes of the Bushveld complex with those of other geologic environments.

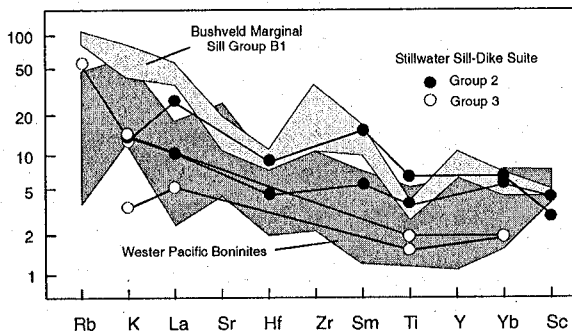


Figure 4. Spider diagram comparing Stillwater, Bushveld and boninite compositional characteristics. Data from Helz, 1985; Hatton and Sharpe, 1989.

In their discussion of potential isotopic sources for Stillwater magmas, Wooden et al. (1991) suggested that the metasomatic component was derived from subducted crustal rocks. As summarized by McCallum (1996), this model satisfies most of the known isotopic constraints for a Stillwater parent liquid, as it would elevate Pb and Sr but lower Nd isotopic ratios while having little effect on Os ratios as the high Os content of the mantle is unaffected by crustal contamination.

Incorporation of subducted material also provides a straightforward mechanism to recycle Cl back into the mantle (e.g., Anderson, 1974). In addition, as discussed by Keppler (1996) and You et al. (1996), a Cl-bearing hydrous fluid will preferentially incorporate alkalis, Pb, Sr and the light REE over Ti and the heavy REE as compared with a silicate liquid at mantle pressures. Thus a slab-derived fluid could give rise to observed elevated Cl/F, K and the LREE in the Stillwater source regions.

Subduction also appears able to fractionate the Cl isotopes, with light Cl characteristic of subduction zone pore water (Fig. 2). The light Cl of the Stillwater biotite and Chile Ridge MORB Cl zone as compared with more typical mantle values (Fig. 3) is consistent with the influenced of shallow subducted crustal fluids. This suggests more incorporation of light (shallow) crustal Cl (and water?) in the case of the Stillwater than occurred in the Bushveld.

A fluid involvement in the Stillwater and Bushveld source regions is also consistent with the high Cl/F weight ratio of Stillwater assemblages. Cl/F weight ratios in modern magmas are typically higher in arc-related magmas (melts generally considered to be the result of fluxing by volatiles derived from the subducted slab) as compared with MORB magmas (melts generated by decompression of upwelling mantle) (Table 1). MORB magmas

typically have Cl/F ratios less than about 1.0. Where unusually high Cl concentrations are reported in MORB, they are associated with higher water contents. In contrast are data from subduction-related magmas, particularly boninites, for which Cl/F ratios are typically an order of magnitude higher. As for arc-related magmas, the high Cl/F ratio of Stillwater assemblages suggest that melting in the Stillwater source region was initiated by an infiltrating fluid.

In addition, as noted by Schilling et al. (1980) for oceanic magmas and illustrated by the data of Table 1 for basic magmas in general, the Cl/F weight ratio is a useful indicator of overall volatile contents. In this regard, the high Cl/F ratio of the Stillwater and Bushveld mineral assemblages is indicative of an overall higher volatile content in the primitive magmas than is generally assumed. If the compositional comparison with modern boninites noted previously holds true for volatile elements as well, these parent liquids could have had in excess of 1 wt % H₂O.

Finally, the higher water content of Bushveld and Stillwater magmas needs to be incorporated in estimates of initial magma compositions as well as in crystallization models. The added water allows for the stabilization of olivine and has a strong effect on liquidus temperatures. In addition, the eventual separation of a volatile-rich fluid phase from interstitial liquids can transport the PGE and cause extensive recrystallization and even melting within the crystal pile.

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Table 1. Some representative Cl, F and H₂O contents of basic magmas from different geologic environments.

Environment	Cl (ppm)	F (ppm)	Cl/F	H ₂ O (wt. %)
<u>Divergent margins (decompression melts)</u>				
Mid-Atlantic Ridge(average)	100	200	0.5	0.21
Galapagos Spreading Center	400 - 3200	100-700	4 - 12	0.13 - 0.96
Flanking seamounts, East Pacific Rise	30 - 2260	40-240	1.9 - 15	0.1- 1.0
<u>Convergent margins (flux melts)</u>				
Continental arcs				
Columbia	830 ± 520	430 ± 200	1.1 - 9.1	<~2.0
Etna	1700	470	3.6	1.42
Crater Lake	1880 ± 70	400 ± 40	4.7	3.8-5.3
Island Arc Basalt				1.0 - 2.9
Boninites				1.2 - 2.5
Bonin Island		800 - 1740	67-116	8.6 - 15
Stillwater sill and dike group 2 and 3	Estimate from apatite composition		10-100	~1.0
Bushveld Lower and Critical zones	Estimate from apatite composition		10-100	~1.0

References: Boudreau et al (1997) and reference therein, Willmore et al. (2002)

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PGM and Complex Ni-Fe-Cu-Co Arsenide-Sulfide Paragenesis Associated with Fe-Ti-V Oxides of the Gulçari Magnetite Pod, Rio Jacaré Sill, Bahia, Brazil

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Introduction

Mafic and ultramafic layered intrusions are well known repositories of important magmatic metallic ore deposits. The best example of such intrusions is the Bushveld Complex in South Africa which exhibits mineralization of chromite, platinum-group metals, and Fe-Ti-V magnetic iron ores. Orthomagmatic processes driven by differentiation of large magma chambers are commonly used to explain such mineralization. Special regard has been historically given to Cu-Ni sulfide-related PGE mineralization and also to PGE-enriched sulfide-bearing chromitites. As the magnetite ore bodies in layered intrusions have been mined solely for the Fe-Ti and V their role as potential PGE carrier have been seldomly studied. In this paper we intend to contribute to this particular issue by investigating one of these magnetite ore deposits as an alternative source of PGE. We present petrographic and mineral chemistry studies of the magnetic iron ores of the layered gabbroic Rio Jacaré Sill (RJS), Bahia (Brazil) and the associated PGE mineralization.

Magnetite pod-like bodies occur in the layered units of the Rio Jacaré Sill. The Gulçari pod is the main RJS magnetite deposit; it is located in the Lower Transition Zone of this intrusion. The pod consists a sequence of magnetitite, pyroxenite and gabbro layers. The known reserves are $2,0 \times 10^6$ tons of vanadiferous iron ore, bearing a mean grade of 2% V₂O₅. The magnetic iron ore also exhibits PGE values up to 4 ppm Pt, 1ppm Pd, and average grade of 400 ppb total PGE. Complex platinum-group minerals, PGM, were identified in the anomalous PGE intervals which were studied by Sá et al (1993). They form an oxide-PGM-arsenide/sulfide paragenesis which shows three distinct textural features. The first, considered as a primary feature, refers to exsolved Pt-Cu alloys within complex sulfide grains poikilitically enclosed in cumulus magnetite and ilmenite. The second one corresponds exsolution fringes of PGM on arsenides that occupy intercumulus space. The third textural type is represented by fracture-filling tiny PGE alloy particles in both silicate and Fe-Ti-V-oxide grains.

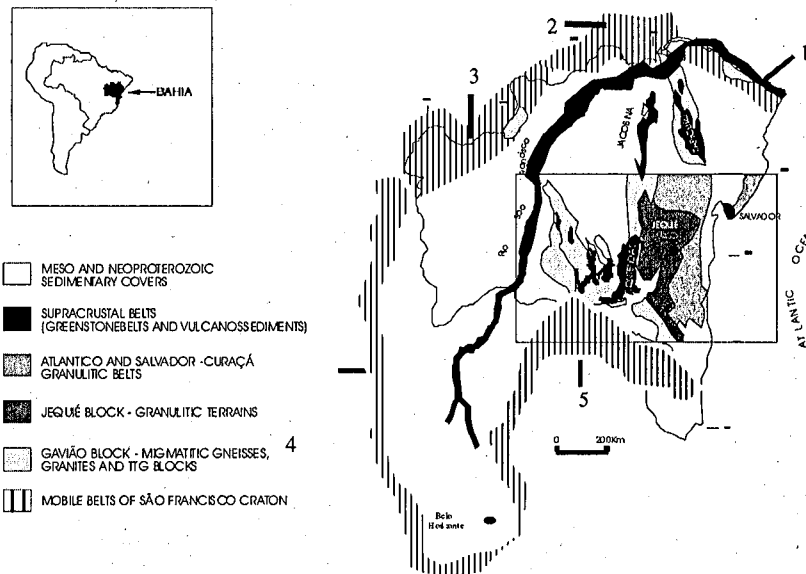


Figure 1. Geological outline of the S. Francisco Craton and the localization of the Rio Jacaré Sill. After Mascarenhas, (1976) and Almeida and Hasui, (1981).

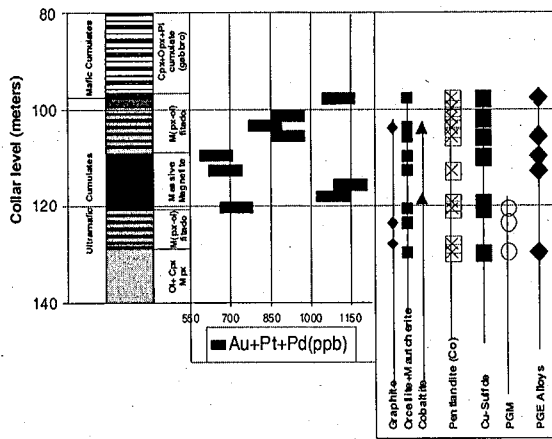


Figure 2. PGE, PGM and associated sulfide-arsenide paragenesis of the Fe-Ti-V oxide mineralization of the FGA40 drill core of the Gulçari Farm prospect. (After Brito 2000).

Geology

The Rio Jacaré Sill is a 70km long, 1km wide synvolcanic Archean ($2,841 \pm 68$ Ma Sm-Nd age, Brito et al. 1999) layered intrusion. It is intruded into folded and metamorphosed basaltic and andesitic rocks interbedded with continental sediments of the Archean Mirante Complex (Brito 2000). This Complex consists of the Mirante Formation (MFm.) and RJS; it is equivalent to the intermediate unit of the Contendas-Mirante Group of Marinho et al. (1994). The complex lies along the southern part of the Contendas-Jacobina Lineament, (fig1), which defines the limit between the Jequié and Gavião Blocks in the central-northern part of the São Francisco Craton.

The Rio Jacaré Sill

R.J.S. is divided up into three zones: a Lower Zone (LZ) which is 300m thick, a Transition Zone (TZ), 200-100 thick and an Upper Zone (UZ) which is 600 to 1000m thick. LZ consists of medium-grained gabbros that exhibit increase in clinopyroxene modal proportion and plagioclase grain size decrease towards TZ. TZ is made up of ultramafic cumulates consisting of cumulus olivine, clinopyroxene, magnetite and ilmenite that grade from magnetite peridotite, magnetite with pyroxene and magnetite pyroxenite and mafic cumulates with cumulus plagioclase and clinopyroxene and minor hypersthene. Magnetite is usually a cumulus phase but sometimes it may be an intercumulus mineral. Monomineralic and bimodal cumulates define a microrhythmically layered sequence of pyroxenite and gabbros with variable amounts of magnetite. Varied-textured fine to medium-grained gabbroic rocks occur in TZ as

enclaves of decimetric to block size of fine-grained gabbro, sometimes pillow-shaped; they also may be present as thin sill-like sheets of diabasic rocks. These varied-textured rocks are interpreted as the result of mixing between different magmas. The Upper Zone consists of two subzones: UZ1 is gabbroic to pyroxenitic and UZ2 is gabbroic and leucogabbroic to anorthositic. UZ1 is a rhythmically banded sequence of micro-layered gabbro-pyroxenite-magnetite, pyroxenite-ferrogabbro and magnetite-bearing anorthosite. UZ2 consists of modally layered rocks divided in three units, a, b, and c. Unit a is made of medium to coarse-grained leucogabbro, unit b is composed of coarse-grained leucogabbro and unit c consists of medium to coarse-grained leucogabbro and anorthosite.

Mineralization

RJS hosts massive magnetite pod-like bodies confined to a layered sequence of mafic and ultramafic cumulates. They are named lower and upper magnetite seams.

The lower magnetite seam is represented by the Gulçari pod which occurs within the lower Transition Zone. This pod is a 300m long, 150m thick sequence of magnetite, pyroxenite and gabbro layers carrying 2 million tons of vanadiferous iron ore with mean grade of 2% V₂O₅, that displays PGE values up to 4 ppm Pt, 1ppm Pd and average grade of 400ppb total PGE. Magnetite layers are interbedded with ultramafic and mafic cumulates (fig 2). The ultramafic cumulates are transgressive towards the contact with the Lower Zone gabbros and consist of olivine-magnetite cumulates, and clinopyroxene-magnetite heteradcumulates.

The massive magnetite layers are made up of ilmenite-magnetite heteradcumulates that form 2 cm to 3 meters thick layers that contain variable amounts of clinopyroxene (En 32-38). Associated mafic cumulates are rhythmically microlayered gabbro, magnetite, and magnetite-pyroxenite bands. The outer contacts of the magnetite pod exhibit hornblende-rich rocks and dunite in places. These features are suggestive of a zoned pattern that Brito (1984) interpreted as similar to the magnetite pipe-like bodies of the Bushveld Complex (Willemse 1979). The upper Transition Zone pod-like magnetite bodies are groupings of magnetite seams and pyroxenites that form 150m long, 20m thick masses of 150,000 tons of vanadiferous iron ore with mean grade of 0.5% V₂O₅ and maximum total PGE contents of 1.3 ppm and mean grade of 380 ppb. The upper Transition

Zone also contains a low grade Cu sulfide mineralization which is confined to the lowermost layers of the upper magnetite seams.

Ore Mineralogy and Microtextural Features

Fe-Ti-V Oxides. The main ore minerals are the Fe-Ti-V oxides, magnetite and ilmenite, which occur together with olivine (Fo 46), clinopyroxene, and metamorphic minerals such as amphiboles (cummingtonite, and Fe-tschermackite), and serpentine. Clinopyroxene is altered to chlorite, especially at the borders of grains. Magnetite and ilmenite show cumulus texture; ilmenite is the early crystallizing phase.

Individual grains of magnetite and ilmenite show round shapes suggesting accumulative

growth. Magnetite and ilmenite also show polygonized shapes, curved borders, and polycrystalline mosaics with 120 degrees dihedral angles at triple junctions. This suggests that annealing processes took place during subsolidus cooling. Magnetite shows rare broad ilmenite lamellar intergrowths of the so-called sandwich microtexture. It also shows trellis microtextures, which is a network of fine ilmenite lamellae exsolved at acute angles. Ilmenite also forms round to irregular shaped exsolved grains at the trellis intersections.

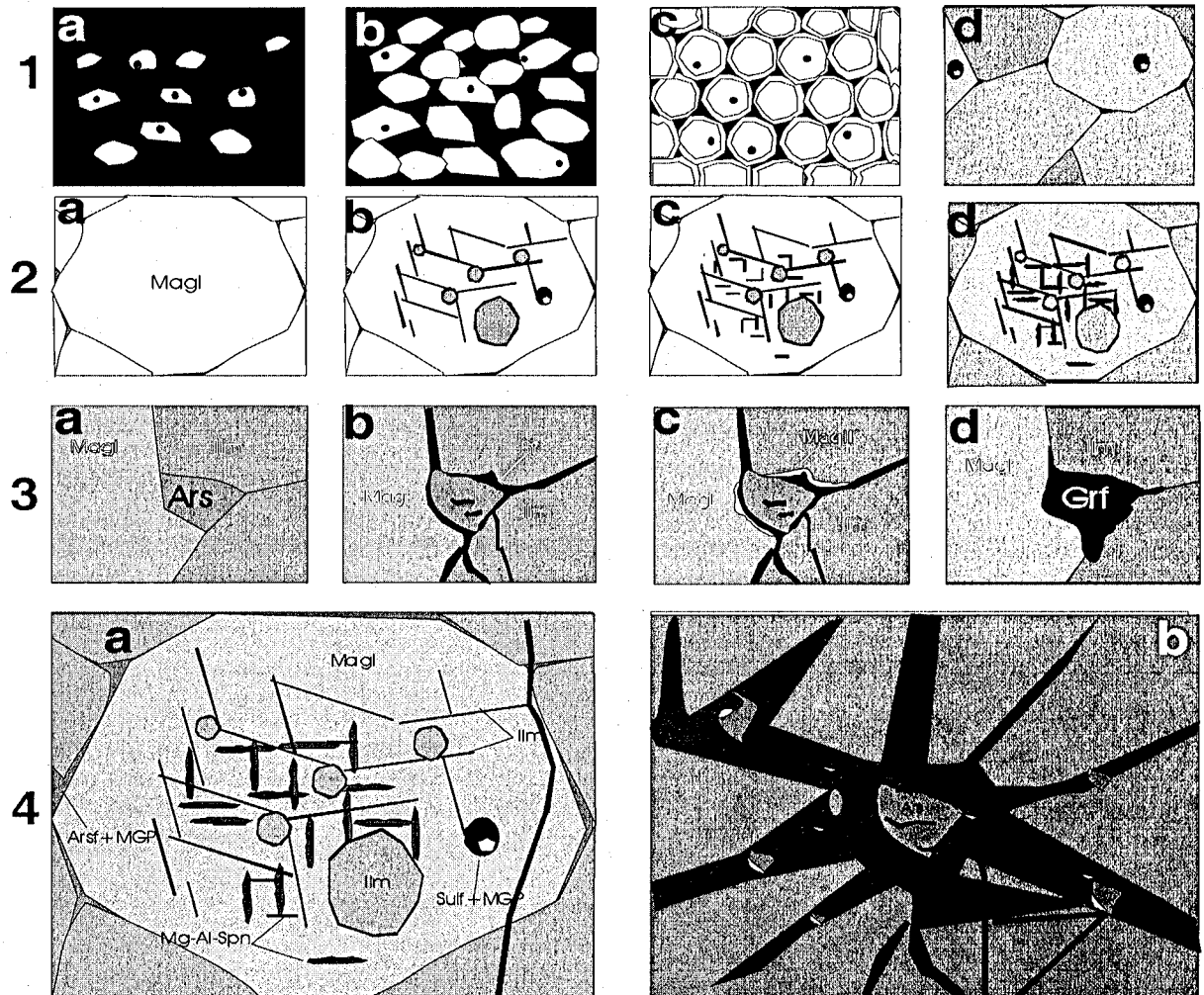


Figure 3. Microtextural evolution of the mineralization: 1-Cumulus stage: 1a-nucleation of ilmenite and magnetite around EGP-rich sulfides. 1b and 1c, accumulative overgrowth. 1d- Annealing and Al-rich spinel formation external to magnetite. 2- Exsolution: 2a.-Arsenic-bearing fluid injection at triple junctions. 2b- Formation of tellis-type oxiexsolution. 2c - Separation of ulvospinel via cloth texture exsolution. 2d- Exsolution of Mg-Al spinel from ulvospinel. 3- Hydrothermal mineralization: 3a. Increasing fluid pressure promotes separation of grains by dilation of contacts. 3b-Sulphur-enriched fluid arrival at arsenide and oxide interfaces. 3d. - Oxidation of sulfides. 4- Current aspect of the mineralization: 4a.- Syngenetic Fe-Ti-V oxide-hosted PGE mineralization. 4b- Hydrothermal PGE mineralization associated with arsenides, sulfides and fracture-bound PGE alloys. Abbreviations : Mag1=primary magnetite; Mag2= Secondary (hydrothermal) magnetite; Ilm = Ilmenite; Sf = Sulfide; Ars = Arsenide; Grf = Grafite; Spn = spinel.

Magnetite also shows the typical cloth microtexture which the exsolution of ulvospinel lamellae in the 100 plane of the magnetite crystals. Aluminous magnesian spinel also occurs as fine exsolution lamellae within magnetite grains. Geothermo-barometric studies indicate that these ores underwent interoxide re-equilibration under reducing conditions defined by ilmenite. Ilmenite and olivine show equilibrium temperatures from 950 to 550°C. Ilmenite and magnetite coexisting grains yielded temperatures as low as 450-550°C at low f_{O_2} .

Platinum-Group Minerals and associated sulfides and Arsenides

Eleven PGM-bearing phases have been identified. These minerals are associated with sulfides and arsenides within the iron-oxide grains and sometimes they are silicate-hosted. Sperrylite occurs as tiny (10µm) round grains within magnetite; sometimes it is silicate-hosted. It also shows external exsolution of gversite which is associated to cobaltite. Gversite also is associated with cobaltite in grains where it forms coronas around maucherite. Cabriite occurs as very tiny grains (1µm) external to orcellite, in contact with silicate (cummingtonite). Cabriite also forms fracture-filling tiny grains (1µm) associated with oxidized chalcopyrite, orcellite and secondary magnetite (pure Fe₃O₄). Cu-Pt alloys are usually enclosed in magnetite and form complex grains, that exhibit intergrowths with pyrrhotite, chalcopyrite and bornite. These Pt-Cu alloys have been identified as isoferroplatinum. Platinum also occurs as platinian awaruite (5% Pt) forming 5µm round euhedral grains. Pt-free awaruite usually forms larger grains (10µm).

Discussion

The PGM, arsenides, sulfides, Fe-Ti-V oxides and silicate paragenesis exhibit textural features indicative of a sequence of magmatic and hydrothermal events. Aiming to explain the exotic relationship between Fe-Ti-V spinels and PGE mineralization we present a simple model of textural evolution via a three stage mineralizing process (figure 3). We suggest a first stage represented by the classic PGE collection by sulfides which were later captured within Fe-Ti-V oxides during cumulus processes. The second stage

consisted of PGE-bearing arsenic post-cumulus fluid injections which were deposited as PGE-bearing arsenides in intercumulus space. The third stage refers to the redistribution of these syngenetic PGE concentrations by late post-magmatic/metamorphic fluids.

The first stage represents orthomagmatic PGE mineralization which could to a certain extent be compared with PGM particles enclosed in chromite grains as shown in other layered intrusions. The second stage could be related to hydrothermal late arsenic-rich synmagmatic fluids possibly related to crystallization compaction. The last phase may be ascribed to post-magmatic redistribution of PGE via a metamorphic fluid.

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Review of Brazilian Chromite Deposits Associated with Layered Complexes: Constraints for the Postulated Genetic Models

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Introduction

Chromitites represent a special case of cumulate rock where chromite is the only cumulus phase. Formation of chromitites thus requires that phase relations of the appropriate system be somehow changed to allow the system to fall into the chromite stability field. Several mechanisms for the formation of chromitite seams have been postulated. They include the contamination of the basic parental magma by sialic rocks, the change in total pressure, the influx of fresh magma into the magma chamber, and changes in oxygen fugacity. In order to test the postulated models, chromitites and hosting mafic and ultramafic rocks from three layered complexes in Brazil were investigated (Fig. 1). The main purpose of these studies was to provide an integrated understanding of the primary igneous evolution of chromitite seams and their host rocks, focusing on magma chamber processes leading to the formation of chromitites. In this paper we summarize the previous results and use them to review the genetic models.

The layered intrusions investigated: Bacuri Complex (Amapá), Niquelândia Complex (Goiás), and Ipeira-Medrado Sill (Bahia), have distinct geological settings and igneous stratigraphy, thus providing an opportunity to look at chromitites formed in different geological-petrological environments.

The Chromite Deposits of the Bacuri Complex

The tectonic setting, geology and chromite deposits of the Bacuri mafic-ultramafic layered complex (Fig. 1) were recently investigated (Spier and Ferreira Filho, 2001; Pimentel et al., 2002; Prichard et al., 2001). The chromite deposits of the Bacuri Mafic-Ultramafic Complex (BMUC) consist of 8.8 Mt. of ore grading 34% Cr₂O₃, representing the second largest reserves of chromite in Brazil. The BMUC is a major Paleoproterozoic (ca. 2.2 Ga) layered intrusion overprinted by ductile deformation and associated regional amphibolite facies metamorphism of the Transamazonian Cycle (ca. 2.0 Ga). The BMUC is intrusive into gneiss-migmatite terranes of the Guyana Shield.

The BMUC consists of a Lower Mafic

Zone (> 500 meters-thick), an Ultramafic Zone (30-120 meters-thick) and an Upper Mafic Zone (> 300 meters-thick). The Ultramafic Zone consists of interlayered serpentinite (olivine cumulate) and chromitite (chromite cumulate). Most of the chromite is concentrated in a few meters-thick chromitite layer located at the base of the Ultramafic Zone in direct contact with the underlying Lower Mafic Zone (Fig. 2). Cryptic variation data (Fig. 2) are consistent with extensive fractionation within the UZ. The stratigraphic position of the main chromitite and cryptic variations in the UZ strongly supports a model for its origin associated with a major new influx of primitive magma, and mixing with more fractionated resident magma. The composition of chromite in massive chromitite of the BMUC is compared with other deposits in Table 1. The Mg ratio of chromite from the BMUC extends over a wide range (0.15-0.55), while the Cr ratios are very high (0.66 - 0.89) and vary through a smaller range.

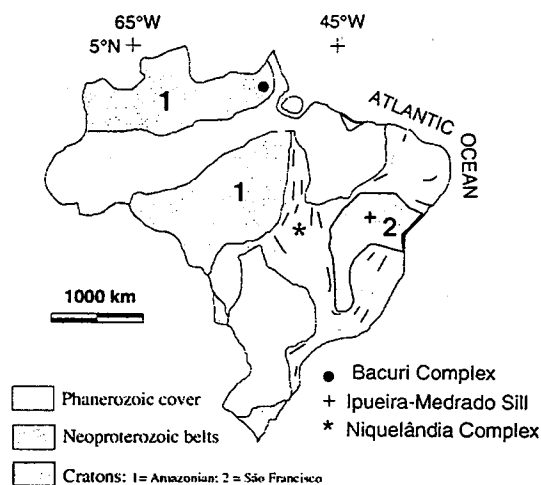


Figure 1. Schematic location of the chromite mineralized layered complexes in Brazil.

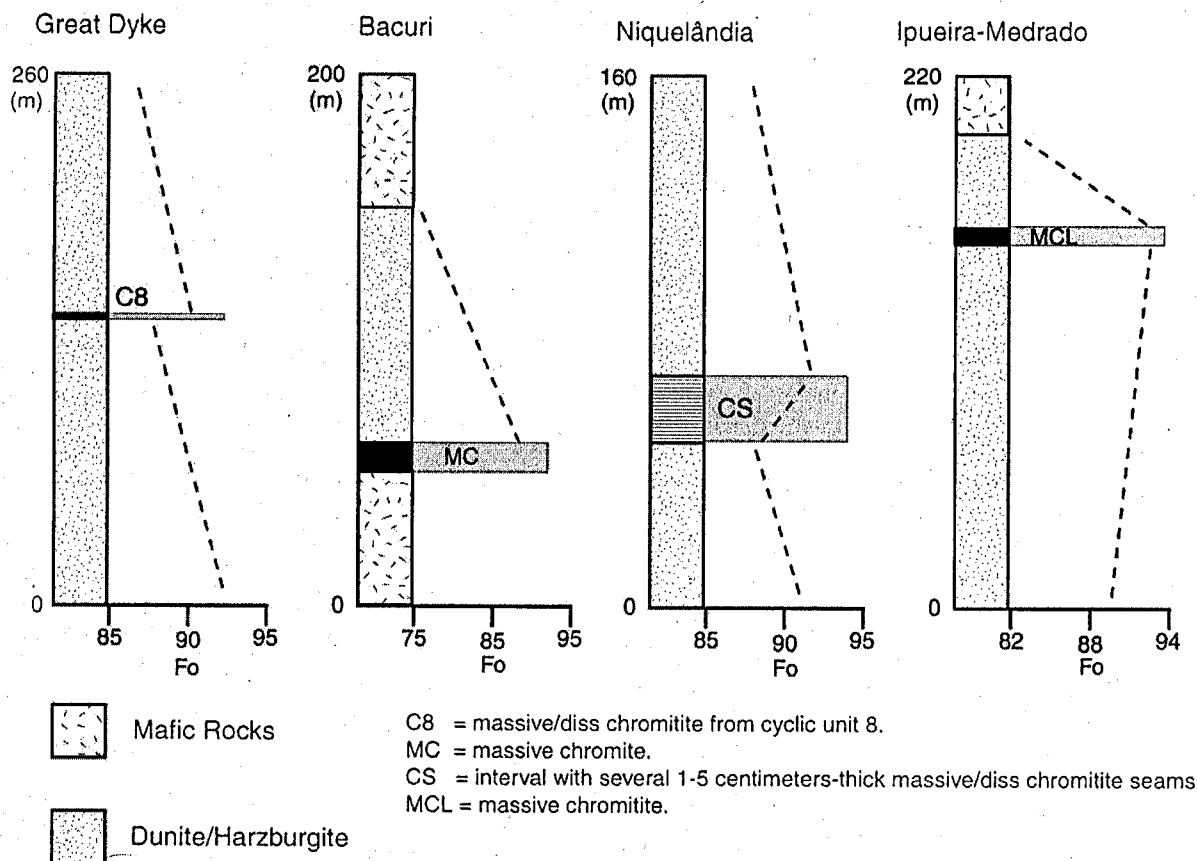


Figure 2. Schematic sections of chromitites from different layered complexes. Olivine cryptic variations indicate the general fractionation trend closely associated to chromitites. Data from the great Dyke are from Wilson et al. (1982).

The Chromitites of the Niquelândia Complex

The tectonic setting, geology and petrology of the Niquelândia Complex were considered in several papers and recently reviewed by Ferreira Filho et al. (1998; 2000). Previous investigations of the non-economic chromitite seams of the Niquelândia Complex were mainly focused on their anomalous PGE contents (Ferreira Filho et al., 1995). The Niquelândia Complex is a major layered intrusion that was affected by tectonic-metamorphic events during the Neoproterozoic. The estimated age for the Lower Layered Series (LS) of the Niquelândia Complex is Paleoproterozoic (ca. 2.0 Ga). The LS has an estimated thickness of several km, consisting of a Lower Mafic Zone (mainly gabbro-norite), an Ultramafic Zone (interlayered dunite/pyroxenite) and an Upper Mafic Zone (mainly gabbro-norite and pyroxenite).

The chromitite seams consist of several few centimeters-thick layers restricted to a 20

meters-thick interval of the estimated 3 km-thick Ultramafic Zone of the LS. Detailed study of the mineralized interval (Ferreira Filho et al., in prep) indicates that chromitites are interlayered with homogeneous dunite (olivine ± chromite adcumulate) of the Ultramafic Zone (Fig. 2). Cryptic variations indicate that the 20 meters-thick interval marks a slight reversal of the fractionation path (Fig. 2). The data support a model for its origin associated with new influx of primitive magma, and mixing with slightly more fractionated resident magma. The composition of chromite in massive chromitite of the Niquelândia Complex is compared with other deposits in Table 1. The Mg ratio of chromite from the Niquelândia Complex extends over a narrow range (0.48-0.57), while the Cr ratios extend over a wider range (0.58-0.71). The small range of Mg ratio is compatible with the very restricted variation of olivine (Fo 87-93) from interlayered dunites.

Table 1. Compositional features of chromite and olivine from different chromite-mineralized layered complexes.

Layered Complex	Chromite in massive chromitite			Olivine		Age (Ga)
	Mg ratio	Cr ratio	Ti ⁺⁴	Fo	NiO	
Bushveld	0.24-0.58	0.60-0.75	0.13-1.06	77-88	na	2.05
Great Dyke	0.36-0.67	0.70-0.80	0.02-0.13	84-92	0.18-0.44	2.74
Stillwater	0.39-0.57	0.60-0.66	0.02-0.28	79-90	0.11-0.33	2.46
Bacuri	0.15-0.55	0.66-0.89	0.03-0.70	76-91	0.18-0.66	ca. 2.2
Ipueira-Medrado	0.45-0.72	0.58-0.68	0.09-0.44	84-93	0.18-0.47	ca. 2.0
Niquelândia	0.48-0.57	0.58-0.71	0.04-0.07	87-93	0.20-0.45	ca. 2.0

Adapted from Spier & Ferreira Filho (2001), Marques & Ferreira Filho (2002), and Stowe (1994)
na = systematic data not available.

Mg ratio = $Mg/(Mg+Fe^{+2})$; Cr ratio = $Cr/(Cr+Al)$.

The Chromite Deposit of the Ipueira-Medrado Sill

The geology, petrology and chromite deposits of the Ipueira-Medrado sill, a Paleoproterozoic (ca. 2.0 Ga) mafic-ultramafic intrusion, were recently investigated (Marques and Ferreira Filho, 2002; Marques et al., 2002). The chromite deposit of the Ipueira-Medrado sill consists of 4.5 Mt. of ore grading 30-40 wt. % Cr₂O₃, representing one of several chromite-mineralized intrusions of the Jacurici Complex, the largest reserve of chromite in Brazil with over 30 Mt of ore. The Jacurici Complex consists of a NS trending (70 km-long to 20 km-wide) swarm of mafic-ultramafic chromite mineralized bodies hosted by granulite-gneiss terranes of the São Francisco Craton (Fig. 1).

At the Ipueira-Medrado Sill, a 5-8 meters-thick massive chromitite layer (MCL) is hosted by a 200-300 meters-thick layered intrusion. The sill is subdivided in a Marginal Zone, consisting of sheared gabbro and harzburgite; an Ultramafic Zone consisting of dunite, harzburgite, chromitite and minor orthopyroxenite; and an upper Mafic Zone consisting of leuco- to melanorites. The MCL is located at the upper part of the Ultramafic Zone. Cryptic variations of olivine and orthopyroxene throughout the stratigraphy indicate that the sill consists of two intervals with distinct magmatic evolution. The interval located below the MCL is characterized by slow evolution of mineral compositions toward more primitive compositions. The interval located above the MCL is characterized by fast upward evolution toward more fractionated compositions (Fig. 2). The data suggest that the interval below the MCL crystallized in a dynamic magma chamber undergoing frequent replenishment with primitive

magma. The magmatic evolution above the MCL suggests that crystallization occurred in a mainly closed magma chamber. The MCL is located at the transition from dynamic open system to mainly closed system magma chamber. The most primitive compositions are observed just below or at the MCL. Cryptic variations are the opposite to what is expected as the result of mixing new influxes of primitive liquid with residual fractionated liquid in the magma chamber. Re-Os and Sm-Nd data indicate strong contamination of the magma at or above the MCL, suggesting that chromite crystallization was triggered by changes of physical-chemical conditions associated to crustal contamination. The composition of chromite in massive chromitite of the Ipueira-Medrado sill is compared with other deposits in Table 1. The Mg ratio of chromite from the Ipueira-Medrado sill extends over a wide range (0.45-0.72), while the Cr ratios (0.58 - 0.68) vary through a smaller range.

Conclusions

Chromitites occur at the base of cyclic units in several layered complex suggesting that chromitites are formed by mixing new influxes of primitive liquid with residual fractionated liquid in the magma chamber (Irvine 1977). Evidence for such model is suggested by changes in the cryptic variation and/or the sequence of cumulus phases illustrated at the Great Dyke, Bacuri and Niquelândia complexes (Fig. 2). However, cryptic variations at the Ipueira-Medrado chromite deposit do not support such model for the origin of the main chromitite layer. Isotope data suggest that chromite crystallization was triggered by changes of physical-chemical conditions associated to crustal contamination at the Ipueira-Medrado sill. The data indicate that chromitite layers are always

associated with major changes in the magma chamber. This study also indicates that the specific mechanism responsible for appropriate phase changes leading to chromite crystallization may be different for each deposit. The compositions of chromite from massive chromitites investigated in our studies occupy a broader compositional field than previously stated for stratiform chromites (Stowe, 1994). Compositional fields currently considered to be typical of stratiform chromite should be reviewed.

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Multimillion-Ounce PGE Deposits of the Portimo Layered Igneous Complex, Finland

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Structure and Megacyclic Units

The ca. 2440 Ma Portimo Layered Igneous Complex (Portimo Complex) is located in the central part of an intrusion belt that crosses Finland almost along the Arctic Circle. Other intrusions in this belt include the Kemi and Penikat intrusions containing a chromitite deposit and PGE reefs, respectively. The Portimo Complex is composed of three principal structural units (Fig. 1): 1) the Portimo Dikes, 2) the Narkaus intrusion, and 3) the Suhanko-Konttijärvi intrusion with the last one having two separate igneous bodies, Suhanko and Konttijärvi.

The Narkaus and Suhanko-Konttijärvi intrusions contain a marginal series and an overlying layered series. The two marginal series differ in thickness and their prevailing rock types. The Narkaus Marginal Series generally varies from 10 to 20 m in thickness, while the Suhanko-Konttijärvi Marginal Series reaches several tens of meters. The Narkaus Marginal Series is mainly composed of pyroxenites, whereas olivine cumulates commonly constitute the upper half of the Suhanko-Konttijärvi Marginal Series.

A striking difference between the layered series of the two intrusions is the presence of marked reversals in the Narkaus intrusion, as shown by thick ultramafic olivine-rich cumulate layers, whereas crystallization in the Suhanko-

Konttijärvi intrusion proceeded without any notable reversals (Fig. 1). The major reversals in the Narkaus Layered Series resemble those of the Penikat intrusion and enable its layered series to be divided into three megacyclic units (MCU). The Megacyclic unit II is, however, found only in the central part of the Narkaus intrusion and fades away eastwards.

Mafic and ultramafic dikes called the Portimo Dikes are encountered in the basement below the Suhanko-Konttijärvi intrusion in the Konttijärvi and Ahmavaara areas (Fig. 1). The dikes have not been dated, and their inclusion in the same magmatic event as the intrusions proper is based on geochemistry. The dikes are subparallel to the basal contact of the intrusion and merge with it locally, so that a dike can actually form the basement of the intrusion.

The MCUs of the intrusions and the dikes were generated from two kinds of parental magmas differing slightly in composition: Cr-richer and Cr-poorer, in this order. The former type resembles SHMB or the Bushveld B1 magma. The Portimo Dikes and MCU I-II have been interpreted to have crystallized from the Cr-richer magma, while the Cr-poorer magma was parental for the Suhanko-Konttijärvi intrusion and MCU III of the Narkaus intrusion.

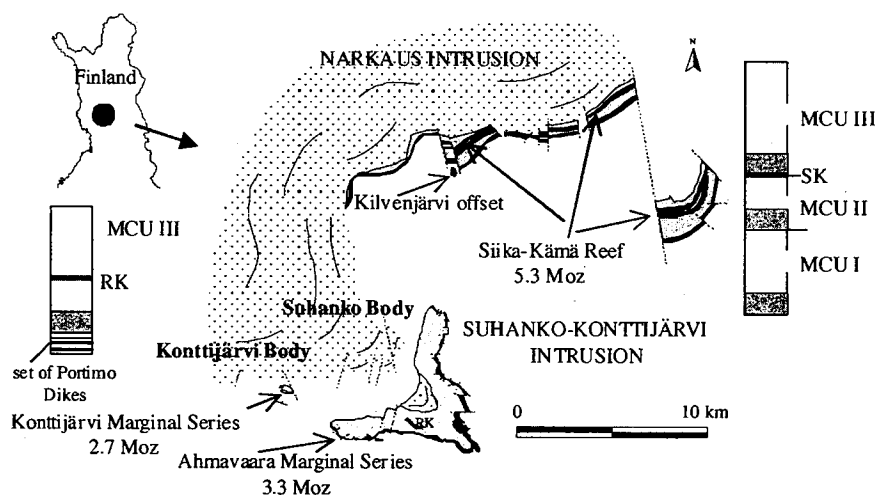


Figure 1. General geology, stratigraphic columns of the Suhanko-Konttijärvi (left) and Narkaus intrusions (ultramafic cumulates in gray). Archean rocks in white, layered intrusions in yellow (gabbros) and green (ultramafic) and younger supracrustal rocks stippled. SK = Siika-Kämä Reef and RK = Rytikangas Reef. Pt+Pd+Au resources, Outokumpu oyj press release.

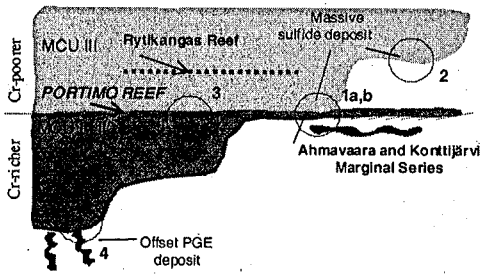


Figure 2. Structural interpretation of the Portimo Complex and the location of the PGE mineralization. Circles 1 - 4 refer to drill hole intersections depicted in Figures 3 - 6, respectively.

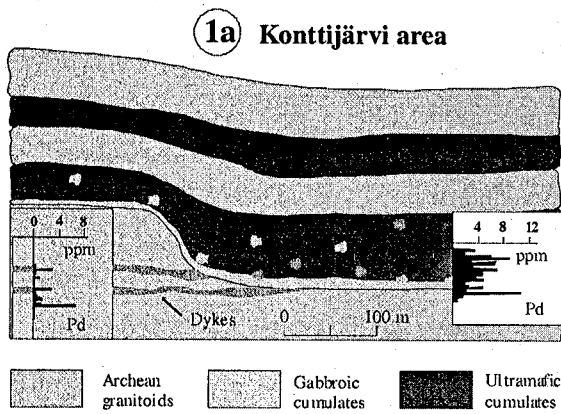


Figure 3. A: Incline cross-section through the Konttijärvi Marginal Series.

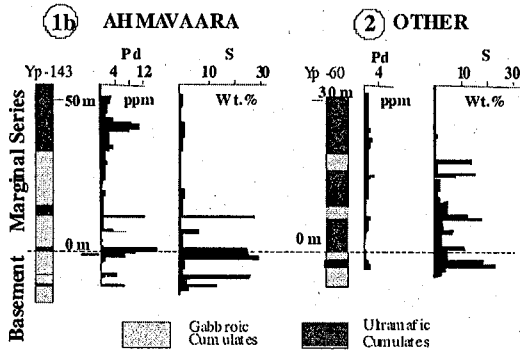


Figure 4. Comparison of Ahmavaara and another marginal series section of the Suhanko Body.

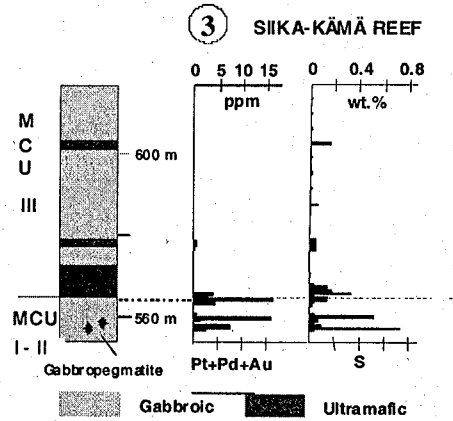


Figure 5. Drill hole section through the Siika-Kämä Reef.

PGE-Cu-Ni Mineralization

Among the layered intrusions in Fennoscandia, the Portimo Complex is exceptional in hosting a variety of PGE mineral deposits. Published resources are indicated in Fig. 1. The mineralization types include (Figs 2-6): a) PGE-bearing Cu-Ni sulfide disseminations in the marginal series of the Suhanko-Konttijärvi intrusion and Portimo Dikes; b) predominantly massive pyrrhotite deposits located close to the basal contact of the Suhanko-Konttijärvi intrusion; c) the Siika-Kämä PGE Reef in the Narkaus Layered Series; d) the Rytikangas PGE Reef in the layered series of the Suhanko Body; e) the offset Cu-PGE mineralization below the Narkaus intrusion.

Disseminated PGE-bearing base metal sulfide concentrations (10-30 m thick) are encountered throughout the marginal series of the Suhanko-Konttijärvi intrusion (Alapieti et al., 1989; Iljina, 1994). Their distribution is erratic, and they generally extend from the lower peridotitic layer downwards for some 30 m into the basement. The PGE content varies from only weakly anomalous values to 2 ppm in most places in the marginal series of the Suhanko Body, but rises to over 10 ppm in several places in the Konttijärvi and Ahmavaara areas (Ni and Cu contents are generally less than 1 wt.% combined). In the same areas, the Portimo Dikes locally carry PGE up to several ppm (Circle 1a, Figs. 2 and 3). Co-precipitation of thorium and uranium minerals with Pd-Se minerals is also found within the Portimo Dikes.

Recent exploration has revealed that the small Konttijärvi Body has deep embayment structure in its center (Fig. 7). Highly PGE-enriched marginal series as encountered in the Konttijärvi-Ahmavaara areas are rare in layered intrusions, and the only other well-known occurrence is the Plat-

reef in the northern Bushveld Complex.

Massive sulfide concentrations are characteristic of the marginal series of the Suhanko Body. The massive sulfide deposits are slab-like concentrations generally varying in thickness from 20 cm to 20 m, and the individual slabs within deposit are separated from each other by more silicate-rich cumulate layers or granitoids. The slabs vary in location from 30 m below the basal contact of the intrusion to a position 20 m above it, within the marginal series. The massive sulfide concentrations vary in size from less than one million tonnes to more than ten million tonnes. The sulfide paragenesis is composed almost exclusively of pyrrhotite, except in the Ahmavaara deposit, which also contains chalcopyrite and pentlandite. Accordingly, the sulfide fraction of Ahmavaara contains 2.7 wt. % Ni and 2.4 wt. % Cu, while the other massive deposits hardly reach a half of these values.

The massive pyrrhotite deposits show low PGE values, with the maximum Pt+Pd normally reaching a few ppm (exemplified by Circle 2, Figs 2 and 4). Similarly to the sulfide disseminated marginal series, the PGE concentrations are much higher in the Ahmavaara deposit, attaining a level of 20 ppm (Circle 1b, Figs 2 and 4).

The Siika-Kämä PGE Reef of the Narkaus intrusion is located most typically at the base of MCU III (Figs 1, 2 and 5), but it can lie even tens of meters below that or in the middle of the olivine cumulate layer of MCU III (Huhtelin et al., 1989; Iljina, 1994). The reef is continuous over the whole strike length of ca. 15 km. The thickness varies from less than one meter to several meters, and many drill cores contain a number of mineralized layers separated by PGE-poor layers, which can be some meters thick. The PGE concentration varies from anomalous values of several hundred ppb to some tens of ppm. The Siika-Kämä mineralization

is one of the sulfide-deficient PGE mineralizations in the Portimo Complex, sometimes containing no visible sulfides at all and hardly ever having a whole-rock sulfur content in excess of 1 wt. %.

The main PGE occurrence in the layered series of the Suhanko-Konttijärvi intrusion is the Rytikangas PGE Reef (Figs 1-2). Its position is known over a distance of 1.5 km. The Rytikangas Reef is hosted by poikilitic plagioclase, plagioclase-bronzite, and plagioclase-bronzite-augite orthocumulates, all containing augite oikocrysts. The orthocumulate layer varies in thickness from 30 cm to 10 m. The thickness of the Reef itself is 30–50 cm, and it is usually encountered on top of the poikilitic orthocumulate layer. PGE + Au concentration are up to 20 ppm.

The offset mineralization is sporadically distributed in the basement gneisses and granites below the Narkaus intrusion (Figs. 1, 2 and 6). The largest deposit, and also the best-known one, is that situated below the Kilvenjärvi block. The offset mineralization represents the richest PGE deposit type within the Portimo area with its Pt + Pd content reaching occasionally 100 ppm. The offset is a Pd-dominated deposit with much higher Pd/Pt than in the other Portimo deposits and extremely low concentrations of Os, Ir, Ru, and Rh. Furthermore, it is irregular in form, being composed of disseminated sulfide-PGM 'clouds', massive sulfide veins or bodies, and breccias, in which sulfide veins brecciate granitoids.

The proportions of base metal sulfides and PGM are variable, but the massive sulfide bodies are always rich in PGE, while some samples containing hardly any visible sulfides can have several tens of ppm of Pd as well. In general terms, the more sulfide-rich occurrences are situated closer to the basal contact of the intrusion and those poorer in sulfides are encountered in a wider zone below the intrusion (Fig. 6).

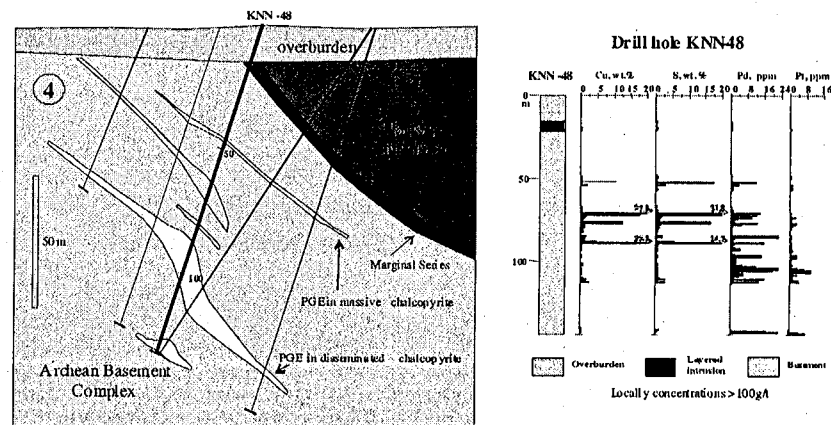


Figure 6. An example of an offset deposit, Kilvenjärvi.

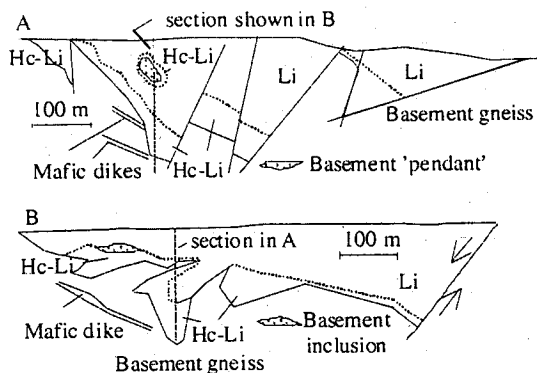


Figure 7. Cross-section (A) and longitudinal section (B) of the Konttijärvi Body. Li, layered intrusion, Hc-Li, heavily contaminated layered intrusion. Redrawn from Arctic Platinum Partnership Oy material.

Figure 2 shows a structural model for the Portimo Complex and the position of the above described mineralization sites and deposits. Taking the boundary region of the two parental magmas as a reference level, it can be seen that the Siika-Kämä Reef, the highly mineralized Ahmavaara and Konttijärvi Marginal Series, and the mineralized Portimo Dikes are located in the same positions in terms of the magma stratigraphy. Accordingly, they were referred to as a group name the 'Portimo Reef' by Iljina (1994). In the Konttijärvi and Ahmavaara areas, the pulses of the earlier parental magma are represented by the Portimo Dikes lying immediately below the marginal series. It is crucial that a marked decrease in PGE values and Ni and Cu contents of the sulfide fraction in the Suhanko-Konttijärvi Marginal Series takes place as soon as the Portimo Dikes disappear below it. In addition, regardless of the quantity of base metal sulfides present, the PGE ratios and chondrite-normalized PGE patterns are practically constant throughout the Portimo Reef. This is the case in the Ahmavaara Marginal Series, for instance, where the variation from disseminated to massive sulfides has only a marginal effect on the PGE patterns.

Discussion

The 'Portimo Reef' seems to lie in an ideal position in the light of the magma mixing model for PGE ore genesis. The two parental magmas inevitably mixed, at least in the Narkaus chamber, while the Portimo Dikes may have had thickenings in the Konttijärvi and Ahmavaara areas, which were not totally solidified at the time of subsequent magma injection, allowing the magmas to mix. The magma mixing model entails, however, certain

problems. One is that the Reef can sometimes be located in cumulates or gabbropegmatites far below the first signs of a new magma. Another problem is found in the Konttijärvi and Ahmavaara areas, where higher tenor of PGE is sporadically found in the footwall granitoids and Portimo Dikes. These PGE enrichments are disseminated in character and therefore not indicative of a stream of immiscible sulfide liquid. Local co-precipitation of U- and Pd-Se minerals refers to co-transportation and mineralization through a fluid phase. The above-mentioned features rule out magma mixing as the sole cause of ore formation.

The location of the Portimo Reef excludes the possibility that its genesis could be exclusively related to the processes that took place in the present intrusions, but suggests that the ore-forming processes started much earlier in some intermediate magma chamber in the crust below the Narkaus and Suhanko-Konttijärvi intrusions and the last batch of the Cr-richer magma was already highly enriched in PGE at the time of emplacement. The PGE was possibly carried by a fluid phase and/or small amounts of immiscible sulfides. The multi-million ounce PGE deposit in the Konttijärvi Body demonstrates that the small size of a mafic body does not make it non-prospective when it is related to large volumes of magma of the whole intrusion belt.

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Exploration for Platinum-Group Element (PGE) Deposits in Mafic and Ultramafic Rocks

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Most economic PGE deposits belong to a continuum of Ni-Cu-PGE sulfide mineralization styles from those that tend to be sulfide-poor and mined mainly for their PGE contents, through to those that are sulfide-rich and are mined mainly for their Ni and Cu contents. These sulfide deposits occur in association with mafic and ultramafic rocks in a variety of settings:

- Narrow "reefs" with low sulfide content within layered mafic/ultramafic intrusions (e.g., the Merensky Reef, South Africa; the J-M Reef of the Stillwater Complex, USA; and the Great Dyke, Zimbabwe).
- Wider zones of PGE sulfide enrichment within mafic/ultramafic intrusions (e.g., Munni Munni, Australia; River Valley and East Bull Lake Intrusions, Canada).
- Stratiform lenses of massive to disseminated Ni-Cu (PGE) sulfides at the base of komatiitic flow units (e.g., Kambalda, Australia; Cape Smith Fold Belt, Canada).
- In association with the outer contact phase of the Sudbury Igneous Complex, Canada, and concentrated in early-formed Offset dykes and the proximal country rocks.
- Stratabound zones of disseminated Ni-(Cu-PGE) sulfides within komatiitic rocks (e.g., Mount Keith, Australia).
- Within feeder zones of magmatic systems (e.g., Voisey's Bay, Canada).
- Within subvolcanic sills which acted as feeders to continental mafic volcanic sequences (e.g., Noril'sk, Russia; and possibly the Thompson Nickel Belt, Canada).

Many Ni-Cu sulfide deposits contain appreciable amounts of the PGEs, although the PGE tenor (calculated in 100% sulfide) varies enormously. Most of the mineralization was formed from magmas that were initially S-under-saturated but became S-saturated either by assimilation of S-bearing crustal rocks (e.g., komatiite-associated Ni sulfide deposits) or by magma mixing (e.g., Bushveld). During the fractionation of S-under-saturated mafic magmas,

incompatible elements such as Cu, S, Se, Pd, and Pt accumulate in the residual silicate melt, whereas compatible elements such as Ni, Ir, Ru and Os are removed with the early silicate and/or oxide phases (Keays, 1995). Once a magma achieves S-saturation and fractionates even small quantities of immiscible magmatic sulfides, it will become strongly depleted in the PGE because of their very high Nernst partition coefficients (i.e. concentration element in sulfide melt/concentration element in silicate melt). The depletion in the PGE will be much more pronounced than the depletion in the other chalcophile elements such as Cu, S, and Se because the PGE have partition coefficients 1-3 orders of magnitude higher than these siderophile and chalcophile elements. A good example of this is provided by the Ni-Cu-PGE sulfide ores of the Sudbury camp; the noritic rocks of the Main Mass of the Sudbury Igneous Complex overlying the Sudbury ores are depleted in Cu and Ni, and, they are extremely depleted in PGE (Keays and Lightfoot, 1999).

In the case of layered intrusions, the ratios of the PGE to elements such as S, Cu and Se are much more useful exploration tools than absolute PGE concentrations. We show that data sets from the Munni Munni Complex, Great Dyke, Stillwater Complex, and the Bushveld Complex exhibit cross-overs in PGE/chalcophile metal ratios proximal to mineralized horizons. This is illustrated for the Merensky Reef of the Bushveld Complex (Fig. 1). The Pd/Cu and Pt/Cu ratios of rocks occurring above the mineralized horizons are all significantly lower than for rocks occurring below the mineralized horizons in these intrusions. The reason for this is that prior to ore formation, the magmas below the mineralized horizons were S-under-saturated whereas those from above it were S-saturated. The advantage of this approach to exploration for PGE deposits is that stratigraphically controlled analyses of ordinary, non-mineralized rocks will provide valuable clues as to the probable presence and location of PGE mineralization within an intrusion.

A feature of all of these mineralized layered intrusions is that the S-under-saturated magmas, which formed the rocks underlying the mineralized horizons, were high Pd- and high Pt-

bearing Siliceous High Magnesian Basaltic Magmas (SHMB). In all but the Munni Munni Complex, the ultramafic rocks are dominated by orthopyroxene, and so they crystallized from Si-rich melts. It is unclear whether these SHMB magmas were boninites or crustally contaminated komatiites. However, it appears that these magmas were initially S-under-saturated, a feature that we suggest is critical to the formation of important magmatic PGE deposits.

The Bushveld magmas appear to have contained more than an order of magnitude higher concentrations of the PGE than "normal" S-under-saturated magmas, such as komatiites that contain about 10 ppb Pd. The UG-2 and the Merensky Reef lie at the bases of cyclic units of rocks, termed the UG-2 Cycle and the Merensky Cycle that are both located in the Upper Critical Zone (Figure 1). They

formed from discrete units of magma and possibly within double diffusion convection cells. All of the PGE decline in abundance from the base to the top of each cycle (Figures 1 and 2). In the UG-2 cycle, Pd drops off from 3100 ppb Pd at the base to 0.7 ppb Pd at the top. These variations indicate that the magma that formed the UG-2 was initially S-under-saturated but became S-saturated as it cooled; the first sulfides to segregate from the magma were Pd-rich reflecting the high partition coefficient. Sulfides that formed at the top of the UG-2 cycle have very low Pd contents; by the time they formed, the bulk of the Pd had been removed from the magma. Mass balance considerations suggest that the magma that formed the UG-2 contained 57 ppb Pd and 73 ppb Pt, while the Merensky Reef magma contained 106 Pd and 266 Pt.

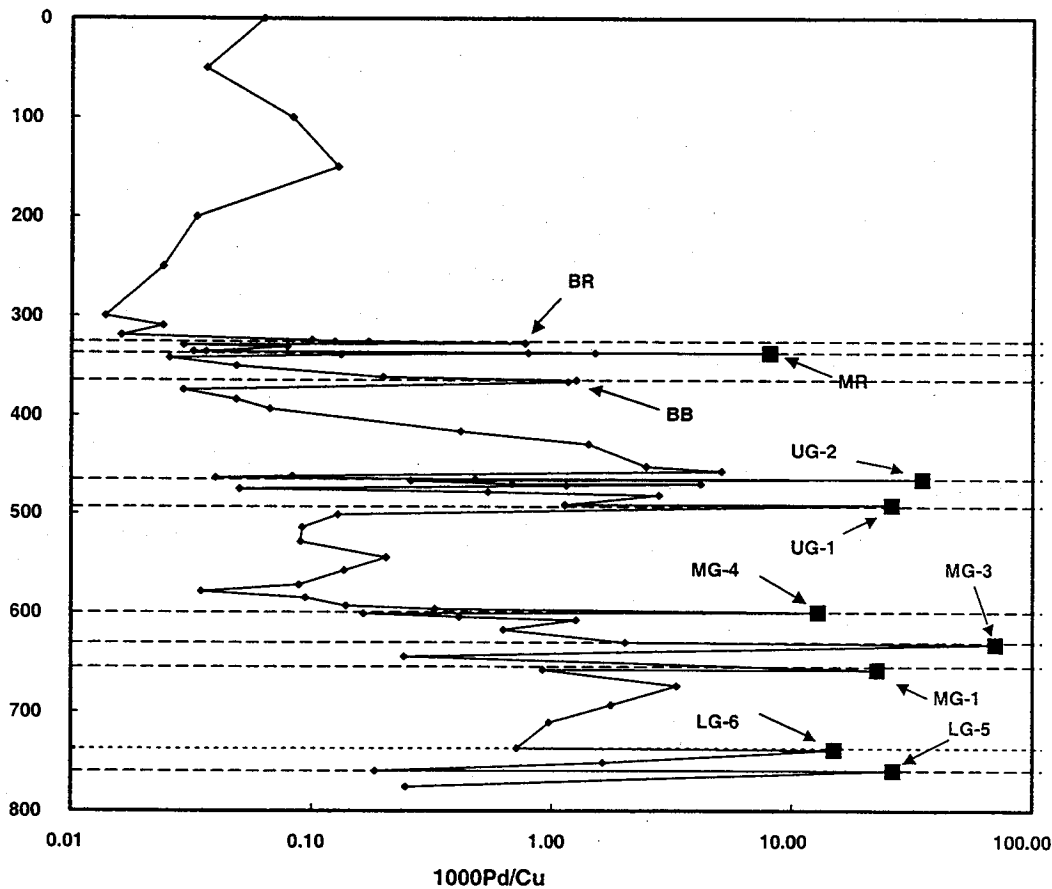


Figure 1. Variation in Pd/Cu ratios of rocks from a drill hole traverse across the Rustenberg Mine (Bushveld Complex). Note that the Pd/Cu ratios decrease upwards from the chromitite seams (solid black squares) located at the base of each package of rocks. Except for the low Pd/Cu values at the top of these packages, Pd/Cu ratios in rocks below the Merensky Reef are higher than those of rocks above. Copper and Pd data for the chromite samples are from Hiemstra (1986) and Lee & Parry (1988). Symbols: BR-Bastard Reef, MR-Merensky Reef, BB-Boulder Bed, UG-1 & UG-2 are Upper Group chromitite seams, MG-1, -3 and -4 are Middle Group chromitite seams and LG-5 and -6 are Lower Group Chromitites.

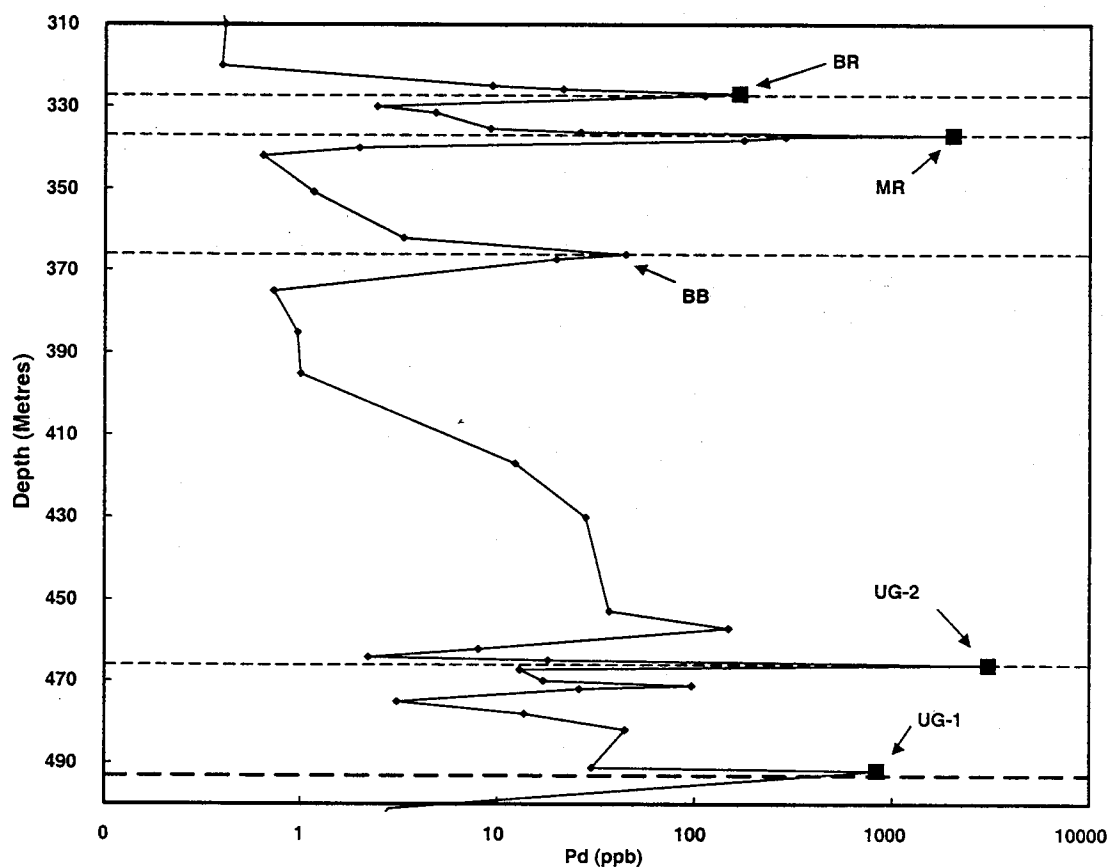


Figure 2. Variation in Pd contents of samples from the base of the UG-1 to the top of the Bastard Reef in samples from a traverse across the Upper Critical Zone of the Bushveld Complex in the Rustenberg Mine. Note that Pd contents in the UG-2 cycle decrease from 3100 ppb Pd at the base of the cycle (in the UG-2) to 0.7 ppb Pd at the top of the cycle (immediately below the Boulder Bed). See Figure 1 for symbols.

A mechanism is required to explain the very high PGE content of the Bushveld magmas. It is suggested that PGE-rich sulfides were initially deposited in a magma chamber at a deeper level than that of the present Bushveld Complex, as a result of interaction between a mantle-derived magma and the crust. Evidence for such interaction is provided by the strongly radiogenic Os-isotope signature of samples from the UG-2 and the Merensky Reef (Schoenberg et al., 1999). At some later stage, these PGE-rich sulfides were dissolved by a second, strongly S-under-saturated magma, such as a SHMB, and introduced into the Bushveld magma chamber where they eventually became S-saturated due to cooling and consequently precipitated the PGE-rich sulfides in the UG-2 and the Merensky Reef.

Many other Ni-Cu-(PGE) sulfide deposits (e.g. Noril'sk, Voisey's Bay) are probably products of two-stage processes initially involving depos-

ition of Ni-Cu-PGE sulfides at depth due to interaction of mantle-derived magmas with crustal materials and subsequent transport of these sulfides to higher levels in the crust. However, unlike the Bushveld, instead of dissolving the PGE-rich sulfides, the second "carrier" magma simply transported these as immiscible sulfide droplets into the higher-level magma chambers.

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Chromite is the Key to PGE

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Introduction

Chromitite layers in the Bushveld Complex have long been known to host anomalous concentrations of PGE. Indeed, the first recorded occurrence of PGE is from a chromitite in the area around Mineral Range in 1908. Initially these discoveries did not flourish due to metallurgical constraints, and the discovery in 1924 of the pipe-like (Onverwacht, Mooihoek and Driekop) bodies with payable Pt-Fe alloys, and then the chromite bearing (but S-enriched) Merensky Reef (both metallurgically amenable) suppressed interest in chromitite as a source of PGE (see Cawthorn, 1999 for a historical overview of the discovery of the Merensky Reef), until in the late 1970's and 1980's when metallurgical breakthroughs and increased demand led to mining of the UG2.

In the 1980's the model of Naldrett and co-workers. (see e.g. Naldrett, 1989), in which a very high ratio of magma to an immiscible sulfide liquid (R-factor) associated with magma influx and mixing is invoked, gained favour due to the presence of sulfide in the Merensky Reef. However, this model has recently been discredited for PGE in the Bushveld Complex (Cawthorn, in press).

Mineralogy, Geochemistry and Sr-Isotope Data

The chromitite layers of the Bushveld Complex host platinum group elements as platinum group minerals (PGM) which principally include PGE-sulfide minerals, (laurite, braggite and cooperite), but antimonides (geversite, stibio-paladinite), arsenides (sperrylite), bismuthides (insizwaite), tellurides (merenskyite, moncheite) and alloys (isoferro-platinum) are major contributors (see e.g. Lee, 1996). Base metal (Ni-Cu-Fe) sulfides also occur, but are not sufficiently abundant to be the residue of a PGE enriched immiscible sulfide liquid. Furthermore, many of these minerals occur as euhedral crystals enclosed in chromite grains. The R-factor modelling depends on large sulfide liquid/silicate liquid partition coefficients which obey Henry's Law. The presence of liquidus PGM negates this assumption – which can only be sustained by special pleading (e.g. assumed post-crystallization sulfur loss).

Chromitite is highly enriched in PGM, and chromite compared to most silicate host rocks and potential primary magmas (see e.g. Scoon &

Teigler (1994)). The enrichment of chromium in chromitite (Cr $\pm 20\%$ in UG2) compared to a tholeiitic liquid (± 250 ppm) is about 800, and the enrichment in PGE in chromitite (8 ppm for the UG2) relative to a potential tholeiitic liquid (± 40 ppb (Davies & Tredoux, 1985)) is about 200. These ball-park figures are within a factor of 4 and suggest that the *process* that led to the crystallization of chromite simultaneously led to the crystallization *and* co-concentration of PGM. This requires the processing of a large volume of liquid (some 200-800 times the mass of the precipitated chromitite).

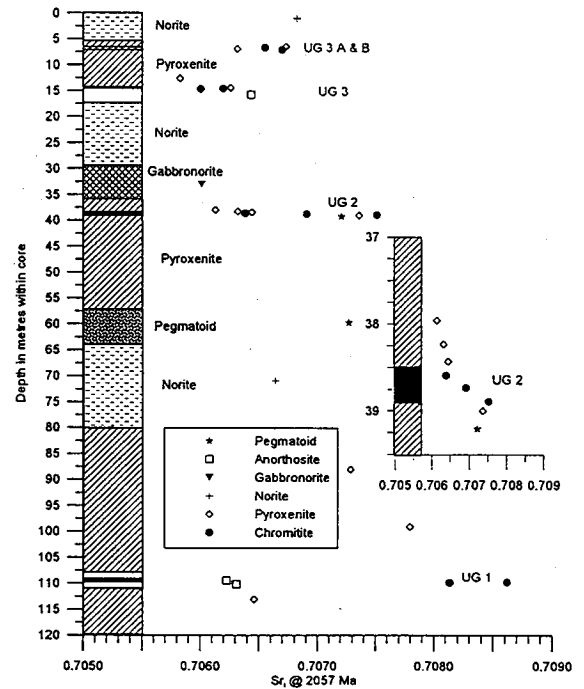


Figure 1. Strontium isotope profile in upper Critical Zone silicates and chromitite-hosted plagioclase from 10m below UG1 to 5m above UG3 in KF17 core, north of the Steelpoort lineament. The expanded inset for isotopic ratios through the UG2 chromitite illustrates the upward decrease in initial ratios from base to top of the chromitite. Source of data, Richardson (2002).

Seminal papers by Irvine (1975 & 1977) showed that chromite crystallization can be induced by felsic (silica) contamination (1975) or by mixing of "parental" magma and more evolved residua (1977); the process of mixing dissimilar magmas forcing the mixture into the chromite phase field.

Recently, Schoenberg *et al.* (1999) showed that chromitite layers have exotic enriched Sr-isotope systematics which are completely different from those of the host silicate rocks, and which imply sudden and major contamination by a 'granitic' crustal component. Furthermore, this contamination was only active while the chromitite was forming and ceased once the chromitite stopped forming. These observations have been confirmed and extended by Kinnaird *et al.* (in press) an example of which is shown in Figure 1.

These mineralogical, geochemical and isotopic data suggest a mechanism for chromitite formation that revives Irvine's (1975) model in combination with his 1977 parental magma mixing model, but negates the later (1981) double diffusive convection model and the Irvine & Sharpe (1986) U- and A-type magma mixing models for these rocks.

A Model for Accumulation of Chromite and PGM

In this model we envisage that influxes of new magma into the Bushveld magma chamber rose up

as fountains that interacted with and entrained granophyric roof-rock melt. This process contaminated both the new magma and the resident entrained melt with a SiO_2 rich component that induced crystallization of copious chromite and the concomitant crystallization of PGM. The small PGM crystals are collected and incorporated by the more abundant chromite grains as suggested by Hiemstra (1979). The chromite with adherent PGM and magma was carried to the floor of the chamber by the collapsing fountain to form chromitite layers enriched in PGE. Chromite (and PGM) only crystallize while there is roof-melt interaction or there is sufficient compositional contrast to force the mixture into the chromite phase field. The contamination process ceases when the head of the fountain ceases to impinge on the floating roof-rock melt. Continued magma influx may occur which is uncontaminated by the roof-rock melt. This process is illustrated in Figure 2

This process involves mixing three component magmas to produce chromitite. These are the resident (residual) magma, the new influx of magma and the roof-rock (granophyric) melt. The resident magma need not be a derivative of the second new magma as required by some models for PGE mineralisation and chromitite formation such as those of Campbell & Naldrett (1983) or Irvine (1977).

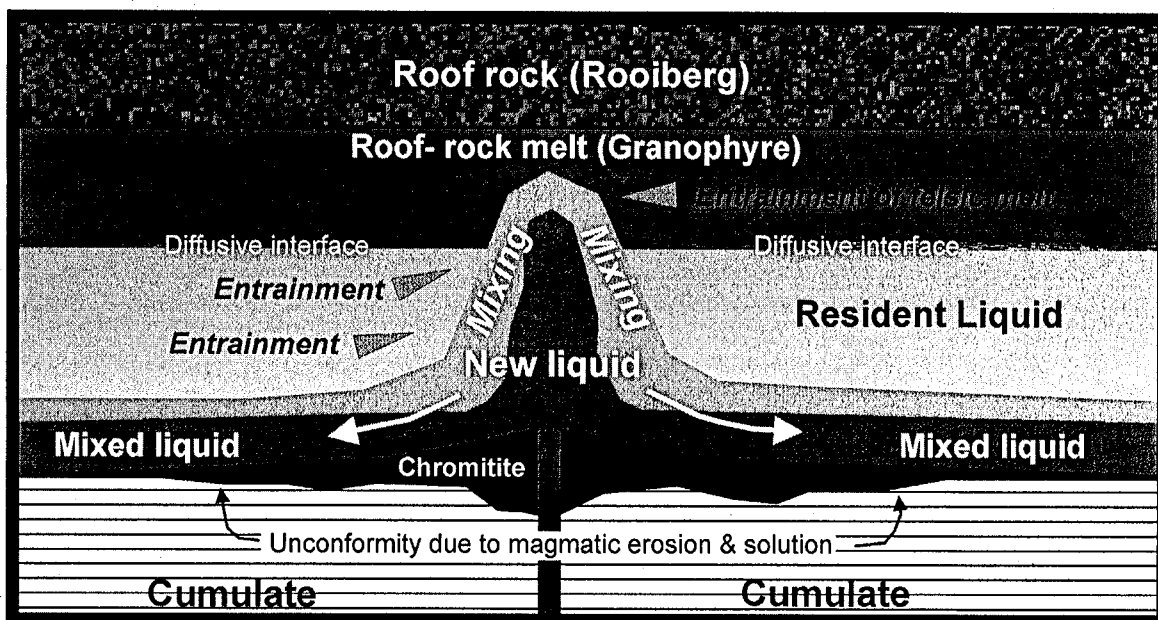


Figure 2. Introduction of a new magma ($Ro = 0.705 - 0.706$) as an active fountain results in entrainment of the resident mafic liquid and if there is sufficient upward momentum roof-rock melt ($Ro > 0.72$) is also entrained. This results in contamination by a silica-rich component with the resulting forced crystallisation of chromite. The mixed liquids are out of equilibrium with the floor cumulates and react and erode to form an unconformity onto which the chromite/PGM ore is deposited.

In the Bushveld Complex, the influxing magmas are not a single primitive "parental" magma, but a diverse series of different magmas as shown by Harmer & Sharpe (1985).

Conclusion

The model outlined here provides an explanation for the following features of chromitites and their PGM mineralization:

1. The apparent random stratigraphic positions of chromitites with respect to the evolving layered silicates, which cannot be predicted in the sequence as they are induced by magma addition events and not fractional crystallization.
2. The apparent unconformable relationships to underlying sequences.
3. The co-enrichment of Cr and PGE.
4. The unusual isotopic, mineralogical and chemical properties.

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Discrimination between Magmatic and Hydrothermal Ni-Cu-PGE and PGE Mineralization

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Introduction

Mafic-ultramafic-hosted Ni-Cu-(PGE) and PGE deposits are believed to have formed by segregation of immiscible sulfide melts and/or alloys from mafic-ultramafic magmas in dynamic magmatic systems such as lava channels, feeder dikes, and magma chambers. However, some of the mineralization in some deposits (e.g., hydrothermal veins at Kambalda: Lesher & Keays, 1984; Stillwater J-M reef: Boudreau et al., 1986; Sudbury footwall ores: Farrow & Watkinson, 1996) is associated with hydrous and/or halogen-bearing phases and is interpreted to have formed from or been modified by hydrothermal fluids.

Different base, precious, and semi-metals have different solubilities and partition differently between silicate magmas, sulfide melts, and supercritical fluids/volatiles (e.g., Keays & Crocket, 1970; Keays et al., 1982; Barnes & Maier, 1999; Mathez, 1999; Wood, in press). For example, the solubilities of Fe, Co, Ni, Cu, PGEs, and Au are different in sulfide melts and hydrothermal fluids; Au, Cu, and platinum-group PGEs (PPGE: Pt, Pd, Rh) behave differently than Ni, Co, and iridium-group PGEs (IPGE: Ru, Ir, Os); and Ru and Ir behave differently than other IPGEs. Therefore, it should be possible to distinguish between those ores deposited by magmatic processes and only modified by hydrothermal fluids and those ores deposited directly from hydrothermal fluids. Unfortunately, very few deposits have been systematically analyzed for complete suites of base, precious, and semi-metals, but it is possible to make some preliminary interpretations.

Discrimination

Although the behavior of the different metals varies with magma/fluid composition, temperature, pressure, fO_2 , and fS_2 , the general order of partitioning into sulfide melts appears to be PGE ~ Au > Cu > Ni > Co > Fe, whereas the general order of solubility in hydrothermal fluids appears to be Fe > Cu ~ Au > Pt ~ Pd > Ni > Co >> Ru ~ Ir. Differential crystallization/precipitation processes can further fractionate different elements, but it is possible, at least in most cases, to distinguish between magmatic and hydrothermal

mineralization on the basis of Au/Cu/Pt/Pd, Ir/Ni, and Cr/Fe ratios. These differences are best illustrated on primitive mantle-normalized variation diagrams (Figs. 1-5)

For example, the majority of the ores at Kambalda (Fig. 1), Langmuir (not shown), Donaldson West (Fig. 2), and Sudbury (Fig. 3) exhibit relatively systematic variations in metal ratios with fractionations consistent with magmatic segregation from komatiitic, basaltic komatiitic, and noritic magmas (see Barnes & Naldrett, 1985), respectively, and local mobilization of Au, Pd, Pt, and/or Cu. However, hydrothermal veins and Ni-enriched metasediments in these localities are strongly depleted in Ir, sometimes depleted in other IPGE (e.g., Rh at Sudbury, Rh-Ru-Os at Donaldson), and also depleted (where data are available) in Cr. In contrast, ores at Lac des Iles, New Rambler, and Rathbun Lake (Fig. 4) are strongly enriched in PPGE relative to IPGE and strongly depleted in Ir relative to Ni and/or other PGE. Their compositions are consistent with hydrothermal modification of magmatic sulfides (Lac des Iles and some Rathbun Lake samples) or with hydrothermal transport and deposition (New Rambler and some Rathbun Lake samples).

The ores in the Bushveld and Stillwater complexes (Fig. 5) are depleted in Ni and Cu relative to other metals, but are not significantly depleted in Ir or other IPGE. Their Ir contents are more consistent with magmatic segregation (e.g., Campbell et al., 1983) and hydrothermal modification than with hydrothermal generation (e.g., Boudreau et al., 1986). However, Stillwater ores appear to have lower Cu and much higher Pd and Pt contents, consistent with stronger hydrothermal modification.

Clearly, more experimental data are required to establish the relative mobilities of PGEs in hydrothermal fluids, but if the observed fractionations in known hydrothermal ores are characteristic, then they suggest that the PGE mineralization in the Bushveld and Stillwater Complexes segregated via magmatic processes and was subsequently modified by hydrothermal processes.

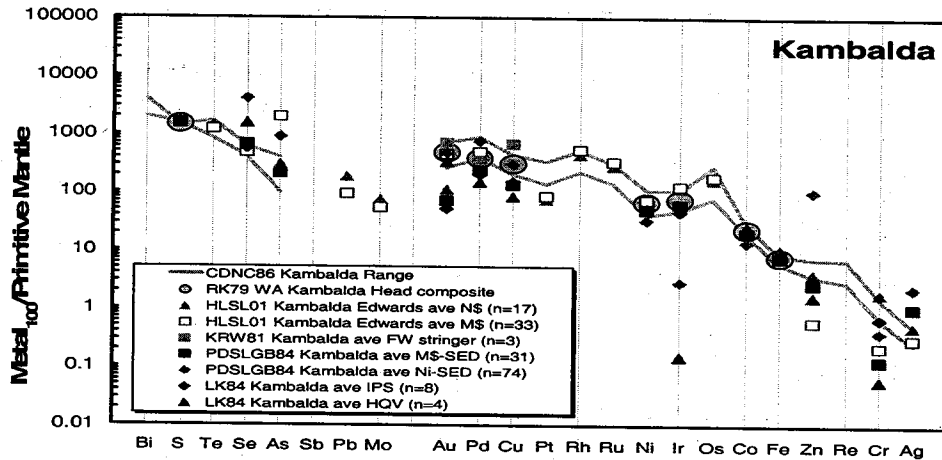


Figure 1. Kambalda ores. For this and subsequent figures, data sources shown as initials of authors and year; ore types: DS = disseminated, N\$ = net-textured, RNS = reverse net-textured, M\$ = massive, IPS = Interpillow, Ni-SED = Ni-enriched metasedimentary, VS = hydrothermal vein.

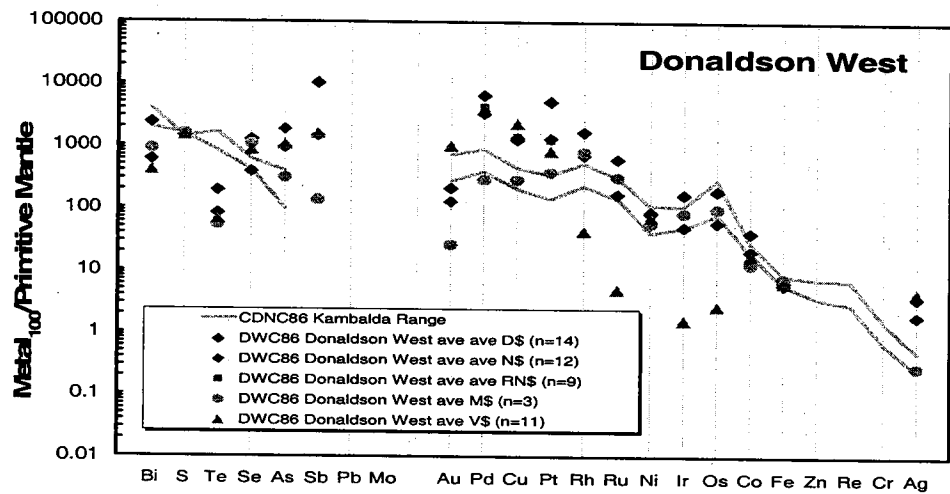


Figure 2. Donaldson West ores.

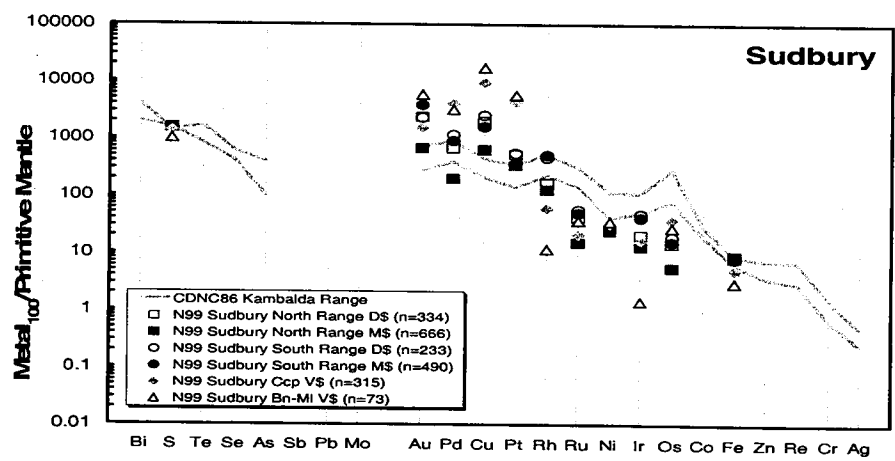


Figure 3. Sudbury ores.

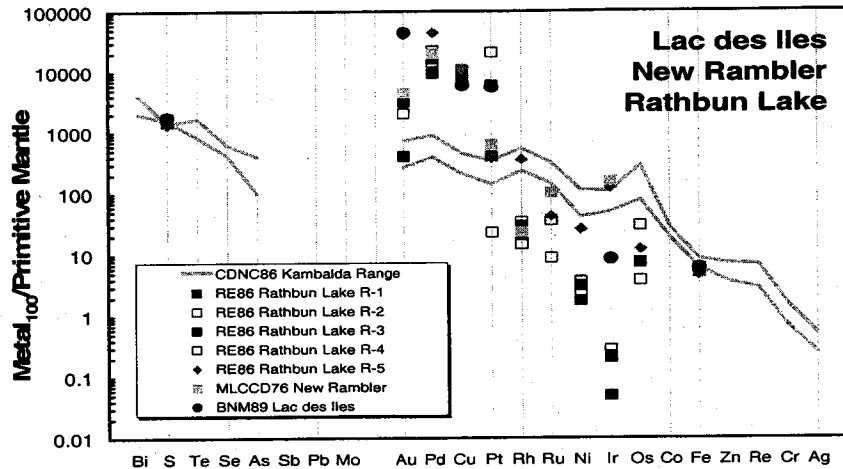


Figure 4. Lac des Iles, New Rambler and Rathbun Lake ores.

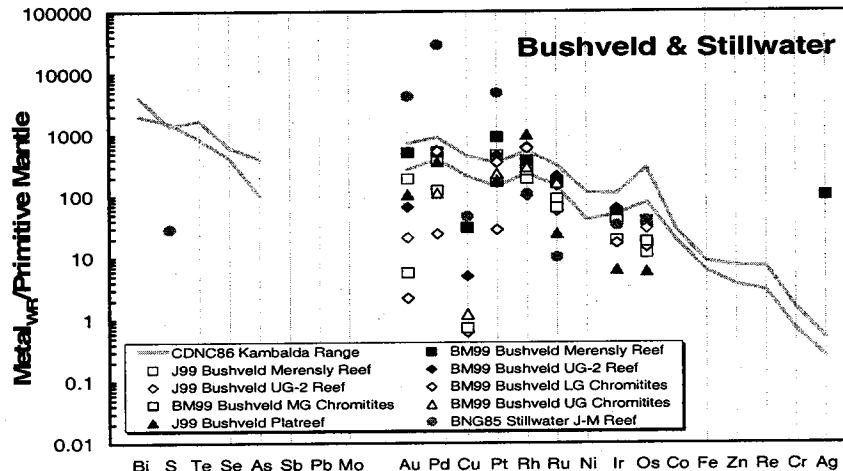


Figure 5. Bushveld and Stillwater ores.

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Olivine and S Isotope Systematics of Ore-Bearing Intrusions at Noril'sk, Siberia

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Introduction

The Ni-Cu-PGE ore-bearing intrusions of the Noril'sk area such as the Noril'sk I and the Main and NW Talnakh intrusions have been recognized to be comagmatic with the voluminous Permo-Triassic Siberian Traps flood basalts (Lightfoot et al., 1992; Naldrett et al., 1992; Czamanske et al., 1994; Fedorenko, 1994). These intrusions contain at least 50 times more sulfides than the magmas represented by their rocks could have dissolved (Naldrett et al., 1992). Although there is a common consensus that much of the sulfides and their metals were possibly derived from the associated lavas as well as the intrusive rocks, opinions as to where the sulfide segregation took place are still divided. Naldrett et al. (1995, 1996), Lightfoot and Hawkesworth (1997) and Naldrett and Lightfoot (1999) have proposed that the ore-bearing intrusions were volcanic feeders in which immiscible sulfide liquids segregated and reacted with continuous surges of magmas *en route* to the surface. Others have suggested that the ore-

bearing intrusions are blind sills formed by sulfide-laden magmas derived from a dynamic staging magma chamber (Naldrett et al., 1992; Brüggmann et al., 1993; Naldrett and Lightfoot, 1993; Lightfoot et al., 1994) or a series of staging magma chambers (Fedorenko, 1994). This study is concerned with the systematic variations of olivine and S isotopic compositions that can be used to discriminate these two competing genetic models.

Types of Olivine and Their Compositional Variations

Olivines from the ore-bearing intrusions occur as (1) inclusions in augite and plagioclase oikocrysts (Fig. 1); (2) discrete grains with granular or subpoikilitic pyroxenes and plagioclase, and interstitial sulfides; and (3) poikilitic crystals enclosing pyroxenes and plagioclase. There is a special type of olivine that occurs as olivine-rich schlieren and masses referred to as "olivinites" by many Russian geologists.

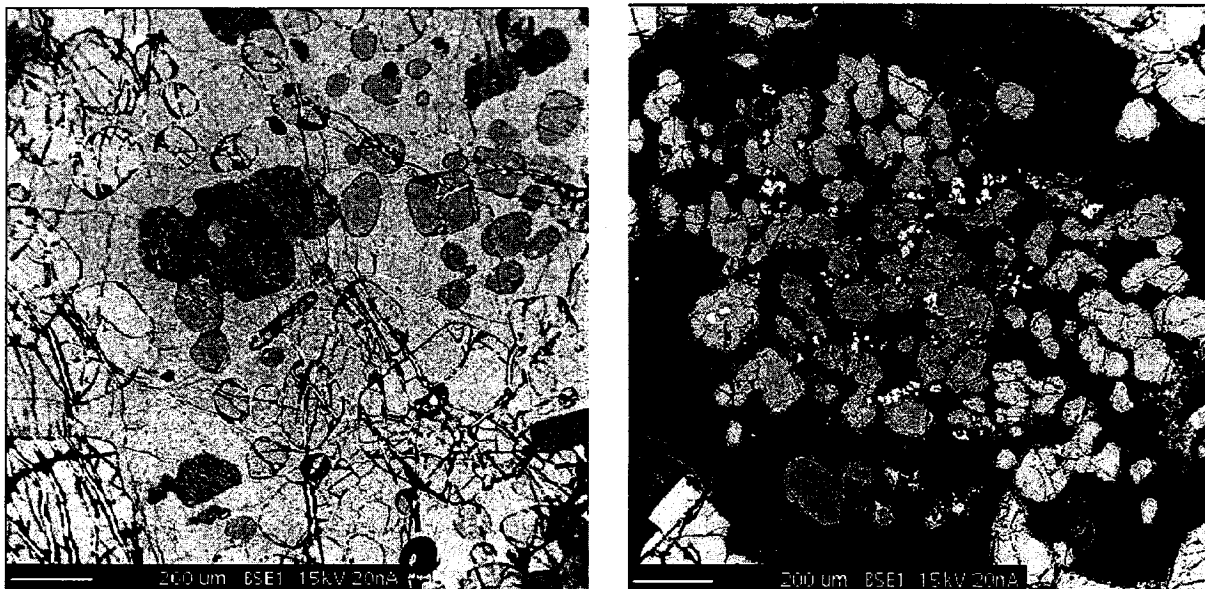


Figure 1. BSE images showing olivine inclusions in augite (left) and plagioclase (right)

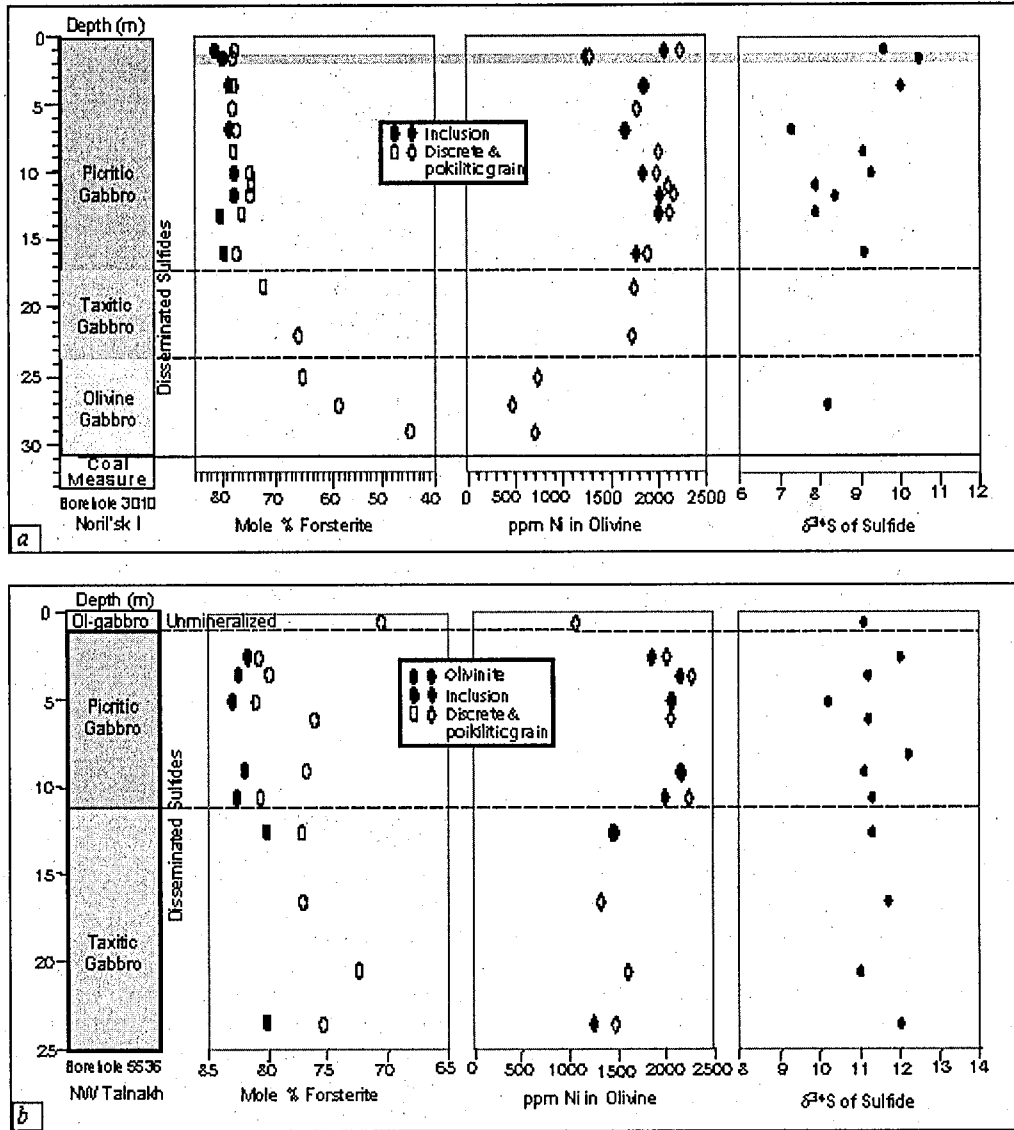


Figure 2. Stratigraphic variations of olivine and S isotopic compositions at Noril'sk (a) and Talnakh (b).

The compositional relationships between olivine inclusions and coexisting discrete olivines are different in sulfide-poor (<0.3 wt.% S) and sulfide-bearing samples. In sulfide-poor samples, olivine inclusions contain higher Fo and Ni than coexisting discrete olivines, consistent with the results of fractional crystallization. In sulfide-bearing samples, discrete olivines also have lower Fo contents than coexisting olivine inclusions. However, the Ni contents of the discrete olivines are higher instead of lower. In addition, the discrete olivines together define an inverse Ni-Fo relationship that apparently has resulted from subsolidus re-equilibration of olivine with trapped sulfide liquid. The difference in olivine composition between an olivinite and its sulfide-

bearing gabbroic host is similar to the difference between olivine inclusions and coexisting discrete olivines in the sulfide-bearing samples.

The main points inferred from the compositional variations of different types of olivine are summarized below: (1) olivine inclusions were early cumulus crystals that have been largely protected by their host oikocrysts from re-equilibration with trapped silicate and/or sulfide liquids; (2) olivinites were either cognate inclusions of early olivine cumulates or they represent *in situ* segregation of cumulus olivine aggregates; and (3) discrete and poikilitic olivines crystallized later and their compositions were variably modified by trapped silicate and/or sulfide liquids.

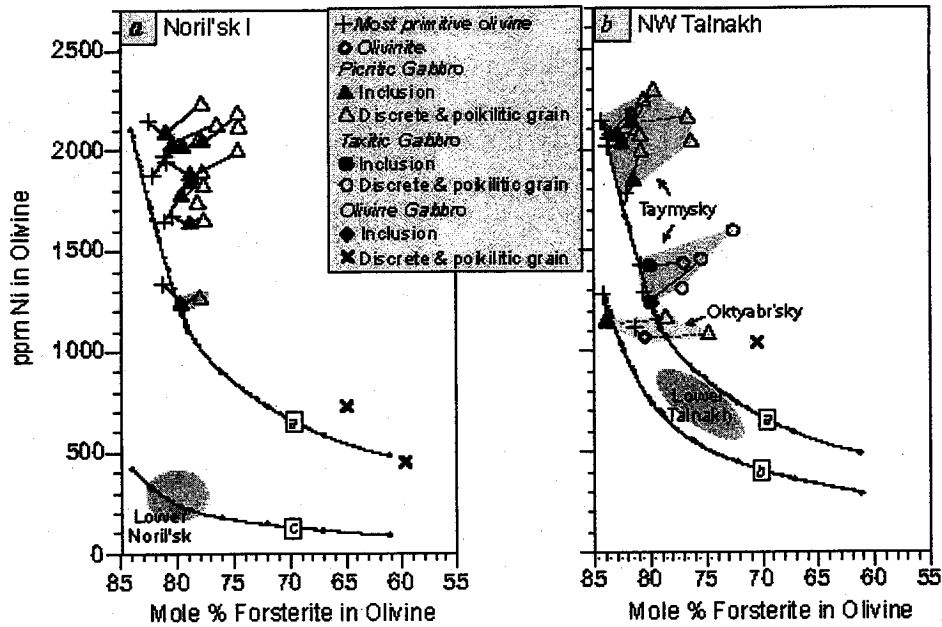


Figure 3. Modeling of olivine fractionation. a. Noril'sk I and Lower Noril'sk intrusions. b. NW Talnakh and Lower Talnakh intrusions. See text for discussions.

Stratigraphic Variations of S Isotopic and Olivine Compositions

Fig. 2a illustrates the stratigraphic variations of olivine and S isotopic compositions in borehole 3010 in the northern part of the Noril'sk I intrusion. The Fo contents of olivines from the olivine gabbro unit decrease downward from 67 mole % to less than 45 mole %. In the overlying taxitic gabbro unit, the Fo contents of olivines are slightly higher and appear to extend the trend of the gabbro unit. But the Ni contents of olivine in the taxitic gabbro unit are much higher, varying between 1700-1900 ppm. In the picritic gabbro, olivine inclusions contain Fo up to 81 mole %. The Fo variations of olivine inclusions from this unit are rather small, mostly less than 3 mole %. The Fo contents of discrete olivines are several mole % less than coexisting olivine inclusions. The contents of Ni in both olivine inclusions and coexisting discrete olivines display an inverse correlation with their Fo contents. Olivines from a sample close to the top of the unit have much lower Ni contents than others from this unit, which may be related to lower Ni content in the sulfide. The content of Ni in the bulk sulfide of this sample is 7 wt%. The contents of Ni in other samples of this unit vary mostly between 9 and 12 wt%. The $\delta^{34}\text{S}$ values of the samples from this borehole vary between 7.9 and 10.5 ‰. Within the picritic gabbro unit, the sample with low Ni content in olivine has a higher $\delta^{34}\text{S}$ value than other samples.

The stratigraphic variations of olivine and S isotopic compositions in borehole 5536 located at the Taymysky Mine in the northern part of the NW Talnakh intrusion is shown in Fig. 2b. The $\delta^{34}\text{S}$ values of sulfides from different rock units in this borehole are similar, varying between 10.2 and 12.2 ‰. In contrast, olivines from different units have distinctly different compositions. Olivines from the picritic gabbro unit are the most primitive and have the highest Ni contents. Olivines from the underlying taxitic gabbro unit have slightly lower Fo contents and significantly lower Ni contents. Olivines from the top olivine gabbro unit are the most evolved and have the lowest Ni contents.

In summary, there are significant discontinuities of olivine compositions, particularly Ni contents, at the contacts of different rock units in the Noril'sk I and Talnakh intrusions. In contrast, the stratigraphic variations in S isotopic compositions are only subtle or minor. This is only explicable if multiple magmas with distinctly different Ni contents and similar S isotopic compositions were involved in the development of the intrusions.

Comparison of Different Intrusions

The compositions of olivine from different intrusions are compared in Fig. 3. The model curves represent the compositional variations of the olivines expected to crystallize from magmas with variable initial Ni contents and the same major

element compositions during fractional crystallization. The model lines labeled *a*, *b* and *c* represent initial Ni contents of 300, 180 and 60 ppm, respectively. The major element compositions are from the average values of the Noril'sk-type sills given by Zen'ko and Czamanske (1994). The simulation of crystallization was performed for 1 kb at the QFM buffer using the silicate liquid model MELTS of Ghiorso and Sack (1995). Olivine is predicted to crystallize at 1230 °C. Plagioclase starts to join olivine at 1198 °C, and clinopyroxene starts to join both olivine and plagioclase at 1175 °C. This crystallization sequence is in good agreement with petrographic observation. Nickel fractionation was calculated assuming that D^{Ni} for olivine, clinopyroxene and plagioclase are 7, 1 and 0, respectively.

Olivine inclusions from the Noril'sk I intrusion have Fo contents similar to the same type of olivine from the Lower Noril'sk intrusion, but have much higher Ni contents and thus require a much higher initial Ni content in magma (300 ppm, line *a* versus 60 ppm, line *c*). It is seen that 5% crystallization in accordance with the model of 300 ppm initial Ni in magma accounts well for the most primitive olivines from the Noril'sk I intrusion, except for one sample which has a much lower Ni content and falls closer to the model line at >20% of crystallization. The displacement of the coexisting discrete olivines to the right can be attributed to reaction with trapped silicate liquid and/or sulfide liquid. Olivines from the olivine gabbro unit of the Noril'sk I intrusion are broadly consistent with the modeled compositions after >50% crystallization.

Compared to the Noril'sk I intrusion, the contents of Fo in olivine inclusions from the picritic gabbro unit of the NW Talnakh intrusion at Taymysky are slightly higher (Fig. 5b). Most of the olivine inclusions from this unit fall closer to the model of 300 ppm initial Ni in magma at <5% crystallization. The olivine inclusions from the taxitic gabbro unit of the NW Talnakh intrusion at Taymysky are more fractionated and fall near the model line at 10-12 % crystallization. Olivine inclusions from the picritic gabbro unit of the NW Talnakh intrusion at Oktyabr'sky have significantly lower Ni contents than the same type of olivine from the same type of rock unit at Taymysky, and require a much lower initial Ni content in magma (180ppm, line *b*). Olivines from the olivinites have compositions similar to the olivine inclusions from the same rock units. Similar to the Noril'sk I intrusion, there are notable displacements to the right for the discrete olivines from the NW Talnakh intrusion, possibly due to re-equilibration with trapped silicate liquid and/or sulfide liquid. Olivines from the Lower Talnakh intrusion plott

slightly to the right of the model *b* fractionation trend (180 ppm initial Ni content in magma).

Lower Ni contents in olivine from the Oktyabr'sky mine than from the Taymysky mine are coupled with lower Ni, Cu and PGE contents in their associated sulfide ores. At Oktyabr'sky, the bulk sulfide ores of the picritic gabbro unit contain 2.0-3.6 wt% Ni, 5.2-9.4 wt% Cu and 8-19 ppm Pd (Naldrett et al., 1996). At Taymysky, the bulk sulfide ores of the same rock unit intercepted by borehole 5536 contain 3.6-9.3 wt% Ni, 7.1-11.5 wt% Cu and 15-28 ppm Pd. It seems that a higher R-factor would explain the higher chalcophile element contents in the sulfide ores at Taymysky. But this is not supported by the similar sulfur isotopic compositions at both locations. Similarly, the disseminated sulfides in the different rock units at Taymysky are also not related to each other by different R-factors. The content of Pd in the bulk sulfide of a sample from the olivine gabbro unit is <4ppm, which is >3 times less than the contents of Pd in the bulk sulfides of the underlying picritic gabbro unit. The $\delta^{34}S$ values of the sulfides from these two units however, are similar and all close to ~ 11 ‰.

The main conclusions from the proceeding discussions are summarized below: (1) with respect to major element compositions, the parental magmas of the picritic gabbro units of the Noril'sk and Talnakh intrusions were similar; (2) with respect to Ni contents, however, the parental magma of the picritic gabbro unit at Oktyabr'sky was significantly depleted in Ni as compared to that of the same type of rock unit at Taymysky and Noril'sk (180 versus 300 ppm); and (3) the disseminated sulfides in the same type of rock units in different intrusions, as well as in different rock units of a single borehole, were often not related to each other by different R-factors. Rather, compositional differences appear to be related to emplacement of multiple magmas that contained sulfide liquids with similar S isotopic compositions but significantly different chalcophile element concentrations.

Summary

Our S isotopic and olivine data are more consistent with the interpretation that sulfide segregation occurred in staging chambers. The sulfide liquids were then episodically upgraded in their chalcophile element concentrations by new surges of magmas into the chambers. Some of these sulfide magmas were simultaneously transported by the episodic surges of magmas to higher levels to form the composite sills of the ore-bearing intrusions.

Attributes of Skaergaard-Type PGE Reefs

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Stratiform, platinum group element (PGE) deposits have long been known to occur in ultramafic-mafic intrusive complexes such as Bushveld and Stillwater (Naldrett, 1989b). Commonly known as PGE reefs, such deposits are typically found near the transition from ultramafic to mafic cumulates where they occur as 1-3 meter thick intervals enriched in PGEs (1-20 ppm) and trace to moderate amounts of sulfide (0.5-5 wt %). However, relatively recent discoveries have demonstrated that potentially economic stratiform PGE mineralization may also occur in tholeiitic mafic layered intrusions. Au- and PGE-bearing layers have been found most notably in the Middle Zone of the Skaergaard Intrusion (Bird et al., 1991; Andersen et al., 1998), but also in related Palaeogene intrusions of East Greenland (Arnason et al., 1997; Arnason and Bird, 2000); in the Jurassic Freetown Layered Complex of Sierra Leone (Bowles, 2000); in thick Mesozoic tholeiitic diabase sheets of the eastern United States; in the 990 Ma Rincón del Tigre complex (Prendergast, 2000); in several intrusions associated with the 1.1 Ga Midcontinent Rift of North America (Miller, 1999, this volume; Miller et al., 2002); in the Paleoproterozoic Lake Owen intrusion of Wyoming (Loucks and Glasscock, 1989); and in the 1.27 Ga Muskox Intrusion (Barnes and Francis, 1995). These and other discoveries reinforce previous suggestions that reef-type PGE deposits hosted by tholeiitic intrusions make up a new class of mineral deposits, which is best represented by the Platinova reefs of the Skaergaard Intrusion (Nielsen and Brooks, 1995; Prendergast, 2000). In this presentation, we summarize the distinctive attributes of this type of deposit, which we refer to as Skaergaard-type PGE reefs and discuss how their metallogensis may differ from classic PGE reefs.

Tholeiitic mafic layered intrusions refer to a class of intrusions formed from aluminous, olivine tholeiitic parent magmas that are commonly associated with active rift settings, especially in plume-affected areas. Troctolite, anorthosite, olivine gabbro and oxide-rich olivine gabbro are common rock types in such intrusions whereas peridotite, pyroxenite, and dunite are usually quantitatively minor. When well-differentiated, such intrusions display a simplified cumulus

stratigraphy following the scheme: Ol or Pl only → Ol + Pl → Pl + Cpx + FeOx ± Ol → Pl + Cpx + FeOx + Ol + Ap. They have a strong cryptic layering towards iron-enrichment indicative of Fenner-type differentiation. They differ from classic PGE reef-bearing intrusions by lacking a significant ultramafic component (early olivine-only crystallization is minor to absent in most tholeiitic intrusions) and by being poor in orthopyroxene (inverted pigeonite is rarely a cumulus phase).

Skaergaard-type PGE reefs in well-differentiated tholeiitic intrusions are similar to classic reefs hosted by ultramafic-mafic complexes such as Bushveld and Stillwater complexes, in that they occur as sulfide-poor, PGE-rich intervals that are meters in thickness and are conformable with igneous layering. However, Skaergaard-type PGE reefs differ from classic reefs in many significant ways as summarized in Table 1. Also like many classic reef deposits, Skaergaard-type PGE reefs appear to have formed by predominantly orthomagmatic processes related the saturation, exsolution, and settling of magmatic sulfide melt from silicate magma and the accumulation of sulfide in stratiform cumulate layers. And like classic reefs, some believe that postcumulus hydromagmatic processes may be predominately responsible for the formation of Skaergaard-type reefs as well (Boudreau and Meuer, 1999).

One very significant difference in their metallogensis, however, appears to be the manner by which sulfide becomes saturated in the magma system. Although magma mixing appears to be an important triggering mechanism for sulfide saturation (or oversaturation) in classic PGE reef systems (Naldrett, 1989b), the nearly constant solubility of sulfur during differentiation of tholeiitic magmas implies that magma recharge may not be capable of producing sulfide saturation in these systems.

To demonstrate this characteristic of sulfide solubility in a differentiating tholeiitic, the sulfide solubility curve and the liquid line of descent for FeO, S and temperature were modelled for the Sonju Lake intrusion (SLI) of northeastern Minnesota (Figure 1A). The SLI is a 1-km-thick, unidirectionally differentiated, sheet-like intrusion

that formed as a closed magmatic system during Midcontinent Rift magmatism (Miller and Ripley, 1996). The variation in FeO and temperature during crystallization differentiation of the SLI have been calculated by applying the fractional crystallization model of Nielsen (1990) to an estimated parent magma composition with 14.7% FeO (Miller and Ripley, 1996). The FeO model curve is similar to a FeO liquid line of descent curve (Fig. 1A) that has been calculated from mass balance of whole rock compositions through the SLI. Based on the experimental studies of Haughton et al. (1974), a 14-20% range of FeO contents at a temperature of 1200°C corresponds to sulfide solubility increasing from 0.11 to 0.19%. Over this range of increased FeO, temperature decrease has an opposite effect on sulfide

solubility. The effect of temperature is not completely known in detail, but a number of experimental studies conducted between 1400 and 1000 °C reported by Naldrett (1989a) suggest a decrease in sulfide solubility by a factor of 3 to 8.5 per 100°C drop in temperature (or a 0.04 -0.09 wt. % decrease in sulfide solubility per 100°C), with a greater effect at lower overall temperatures. The estimated variations in sulfide solubility due to the changes in model FeO and temperature during fractional crystallization of the SLI are shown in Figure 1B. The resultant flat solubility curve (Fig. 1B) demonstrates the offsetting effects of increasing FeO and decreasing temperature throughout most of the SLI's crystallization history.

TABLE 1: Comparison of Attributes of Skaergaard-type and Classic PGE Reefs.

	Skaergaard-type PGE Reefs	Classic PGE Reefs
Tectonic Setting	Plume-influenced continental rifts, occur as subvolcanic intrusions	Unclear, occur as large, isolated intrusions within Precambrian cratons
Age	Mesoproterozoic and younger	Neoproterozoic to Paleoproterozoic
Parent Magma	Aluminous, olivine tholeiite (E-MORB)	High-Cr/Mg/Si boninite/komatiite (U-type) with aluminous tholeiite (A-type)
Host Rock	Gabbroic to ferrogabbroic cumulates (Pl+Cpx±Ox±Ol); some associated with layering.	Pyroxenitic to noritic cumulates; locally associated with chromitite layers, early anorthositic cumulates, and pegmatite.
Sulfide Composition of Mineralized Intervals¹	Cu/Ni >100 (closed systems) Cu/Ni >1 (open systems)	Cu/Ni = 0.3-5.0 (open systems)
Precious Metal Ratios of Mineralized Intervals¹	Pt/Pd = 7 - 0.1 Pt+Pd/Au = 1 - 10	Pt/Pd = 2.5 - 0.2 Pt+Pd/Au > 10
Stratigraphic Distribution of Peak Concentrations¹ (⇒ below, ⇔ coincident with, or)	Pt ⇒/⇔ Pd ⇒⇒ Au Pt + Pd ⇒ Cu (±Ni) ⇔ Au	Pd ⇒/⇔ Pt ⇒/⇔ Au Pt + Pd ⇒/⇔ Ni ⇔ Cu ⇔ Au
Examples²	Closed systems - Skaergaard, Sonju Intrusion, New Haven diabase Open systems- Layered Series at Duluth, Kap Edvard Holm Complex, Freetown Complex, Intrusion, Rincón del Tigre Complex	Lake Bushveld Complex, Stillwater Complex, Great Dyke, Munni Munni Complex, Kimberlana Intrusion, Penikat, Portimo and Lake Owen Other Fenno-scandian intrusions, Rum Intrusion, Muskox Intrusion, Rincón del Tigre Complex

¹ Based on data from this study; Arnason and Bird, 2000; Prendergast, 2000; Andersen et al., 1998; Lee, 1996; Barnes and Francis, 1995; Naldrett, 1989b.

² PGE reefs in the Muskox and Rincón del Tigre intrusions have attributes in common with both reef types.

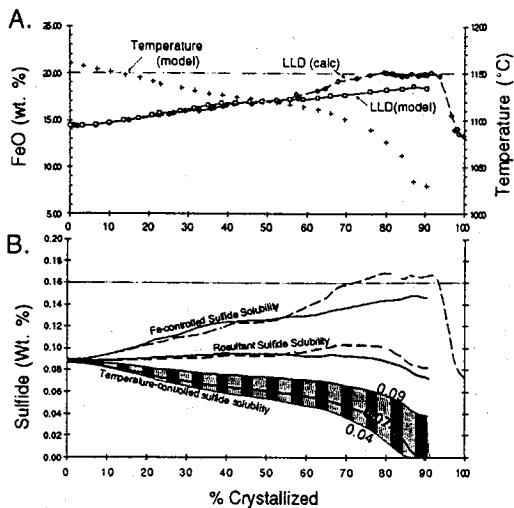


Figure 1. Model curves for FeO, temperature, and sulfide solubility variation during crystallization differentiation of the Sonju Lake intrusion. A) Variation in FeO and temperature based on fractional crystallization model of Nielsen (1990) and variation in FeO based on mass balance calculation (Miller and Ripley, 1996). B) Variation in sulfide solubility due to effects of temperature (factor range of 0.04–0.09 wt% sulfide/100 °C after Naldrett, 1989a) and FeO concentration (based on experimental data from Haughton et al., 1964). Resultant solubility curves are based on a median temperature factor of 0.07 wt% sulfide/100 °C.

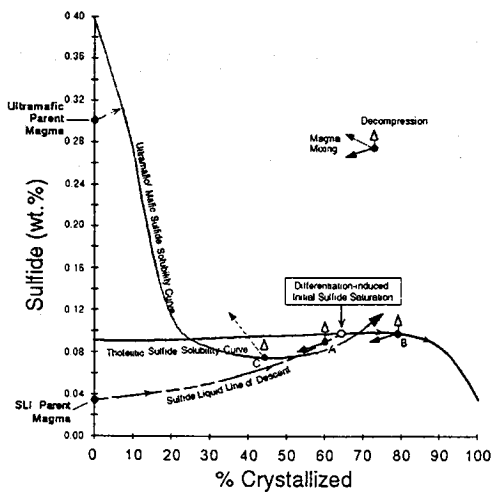


Figure 2. Resultant sulfide solubility curve for Sonju Lake intrusion compared to solubility curve estimated for ultramafic/mafic magma systems (Naldrett, 1989b). Liquid line of descent for sulfide in the SLI magma calculated by Rayleigh distillation and based on sulfide saturation at 65% crystallized. Vector arrows show effects of magma mixing and decompression for three different situations. Point A shows effects on tholeiitic SLI magma just prior to sulfide saturation. Point B shows effects on sulfide-saturated SLI magma. Point C shows effects on evolved ultramafic magma system.

Chemostratigraphic studies show that the Sonju Lake intrusion reached sulfide saturation at 65% crystallized (Miller, 1999; this volume). Treating sulfide as an incompatible component up to the point of saturation, the liquid line of descent of sulfide during fractional crystallization can be calculated by simple Rayleigh distillation. Assuming a sulfide concentration of 1.1 wt.% at saturation, the calculated liquid line of descent curve implies an initial sulfide concentration of the parent magma of 0.35 wt.% (Fig. 2). It is clear that magmatic recharge of this parental composition will not result in saturation or oversaturation of sulfide. Also plotted in Figure 2 is the sulfide solubility curve estimated for ultramafic/mafic magma systems by Naldrett (1989b), which shows how magma mixing can result in sulfide oversaturation in such systems. Although stratiform PGE mineralization in the Kap Edvard Holm intrusion appears to be related to magma mixing (Arnason and Bird, 2000), it is not clear how this has occurred from an uncontaminated tholeiitic parent magma.

Rather than magma recharge, magmatic processes of crystallization differentiation, decompression due to magma venting, and changes in phase equilibrium appear to be important in the formation of most Skaergaard-type reef occurrences. The Sonju Lake intrusion looks to be the most clear cut example of sulfide saturation triggered by straightforward crystallization differentiation, though a detail examination of the reef is underway to confirm this interpretation. Andersen et al. (1998) have similarly interpreted the Pd and Au mineralization of the Skaergaard to have formed by fractional crystallization driving the magma to sulfide saturation. However, the cyclical nature of PGE mineralization and modal layering in the host gabbro may suggest a more dynamic triggering mechanism. Miller (this volume) has interpreted the PGE mineralization and the related macrocyclic layering in the Layered Series at Duluth as resulting from periodic decompression due to magma venting from a shallow (<5km) depth. Perhaps the Platinova reefs of the Skaergaard formed in a similar manner. Sulfide saturation triggered by changes in phase equilibrium, especially the onset of Fe-oxide crystallization, is evident in the Rincón del Tigre complex (Prendergast, 2000).

In conclusion, the suggestion by Nielsen and Brooks (1995) that the stratiform PGE and Au mineralization discovered over a decade ago in the Skaergaard intrusion may constitute an important new type of mineral deposit now seems prophetic in light of similar discoveries in other igneous provinces. Although an economically viable deposit has yet to be found, the search for

Skaergaard-type PGE reefs has only just begun. Like their classic reef cousins, exploration for this subtle style of mineralization requires systematic sampling and a thorough understanding of the crystallization history of the host intrusion.

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On the Formation of Platinum Group Element-Rich Continental Flood Basalt and the Platinova Reefs of the Skaergaard Intrusion

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Platinum Group Element reefs associated with layered mafic intrusions require a parental magma that contains at least several parts per billion (p.p.b.) Pd or Pt. The extremely high distribution coefficient for PGE between sulphide and magma predicts that PGE-rich magmas can only form if all sulphides are melted out of the mantle. For standard peridotite mantle with 250±50 p.p.m. S (McDonough and Sun, 1995), the mantle sulphides are exhausted only if the mean degree of mantle melting (F) exceeds 20%. Komatiite is of course the type example of a Pd-rich magma (8–21 p.p.b., Fig. 1) formed by high F : well above 20% (Keays, 1995). Mid-ocean ridge basalt (MORB) mostly form by moderately high F (8–20%; Klein and Langmuir, 1987), but rarely contains more than 1 p.p.b. Pd (Fig. 1); this is ascribed either to retention in mantle sulphides or to subsequent sulphide saturation during basalt differentiation (Barnes et al., 1987).

For continental flood basalts, the constraints on mantle melting conditions are conflicting. On the one hand, the presence of prominent PGE reefs of continental mafic intrusions such as Noril'sk, Bushveld, Stillwater, and Skaergaard associated with flood basalts implies PGE-rich parental magmas. This would suggest high F . On the other hand, the lithophile trace element composition, in particular the rare-earth elements (REE), indicate that most continental flood basalts formed by moderate F in the range from 5 to 15% (Fram and Lesher, 1993). Although most continental flood basalts are inferred to originate from hot mantle plumes, the reason F remains moderate is that the top portion of the melting regime is truncated by thick non-melting lithosphere – this is referred to as the 'lid'-effect (Ellam, 1992). In this contribution we explore the effects of the lithospheric lid on the PGE and REE composition of continental flood basalt with particular emphasis on the Paleogene East Greenland Igneous Province, following Momme et al. (in press, in review), and the Platinova Pd–Au reef of the Skaergaard intrusion (Andersen et al., 1998).

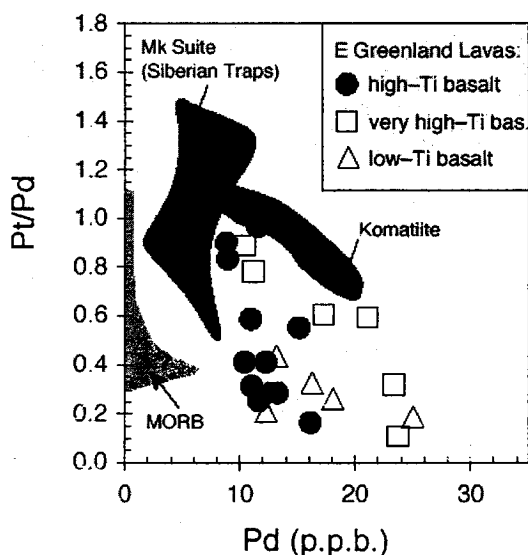


Figure 1. Pd and P: composition of the Paleogene East Greenland flood basalt province (Momme et al., 2002), recent MORB and old komatiite (Rehkämper et al., 1999), and PGE-rich flood basalts of the Permo-Triassic Siberian traps (Mokulavsky Suite; Brüggmann et al., 1993).

PGE of the East Greenland Continental Flood Basalt Province

This Paleogene flood basalt province formed during continental rifting immediately before the opening of the Northeast Atlantic. The major element and REE composition show the 6 km thick lava pile of the Blossville Kyst is composed mainly of high-Ti basalt [mg#, = molar Mg/(Mg+Fe) range from 60 to 43] and with subordinate and intercalated low-Ti and very-high-Ti basalt, all largely unaffected by crustal contamination (Tegner et al., 1998). All three magma suites are PGE-rich with Pd ranging from 3 to 25 p.p.b., but we will focus only on the dominant high-Ti suite that contains 6 to 16 p.p.b. Pd (Fig. 1). The concentration of Pd increase with decreasing mg# (Fig. 2), demonstrating that Pd was excluded during igneous differentiation. This can only be obtained if the magma was undersaturated with respect to S and thus did not reach the point

where sulphide droplets are exsolved (Momme et al., 2002). In summary the East Greenland flood basalts contain moderate amounts of PGE between that of komatiite and MORB (Fig. 1) and did not exsolve sulphide.

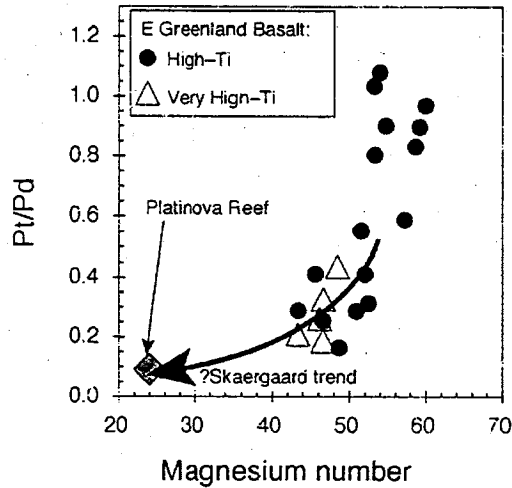


Figure 2. Variation in Pd and Pt with mg# [molar Mg/(Mg+Fe)] for high-Ti and very high-Ti basalt lavas of the East Greenland flood basalt province. Also shown is the typical Pd/Pt of the Platinova Pd-Au reef of the coeval Skaergaard intrusion.

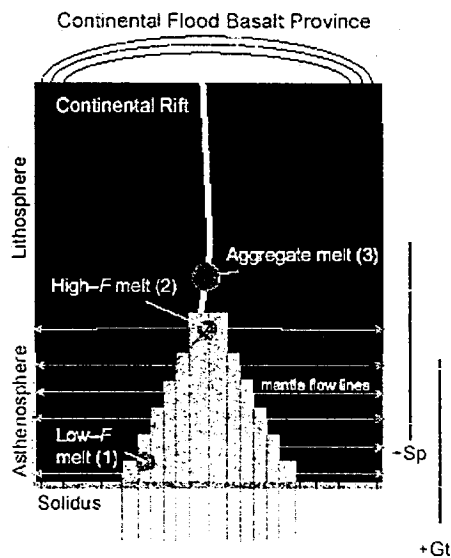


Figure 3. Schematic cross-section showing the melting regime of a corner-flow model (Langmuir et al., 1992) for upwelling mantle plume beneath a continental rift. The mantle solidus is intersected at ~120 km depth and the top of the melting column is at ~60 km depth (Tegner et al., 1998). Predicted melts labelled 1, 2 and 3 are: 1) low-F, high La/Sm_N, low PGE content, and S-saturated; 2) high-F, low La/Sm_N, high PGE content, and S-undersaturated; 3) aggregate melt, intermediate F, PGE content, and La/Sm_N.

Mantle Melting Model

Forward modelling of the REE composition suggests the high-Ti basalts from Greenland formed by mean degrees of melting (F) ranging from 5 to 9%, of mantle depleted slightly (0.5%) relative to Primitive Mantle (Tegner et al., 1998). In this model it was assumed that the melting regime is triangular and governed by upwelling and corner flow as suggested for melting beneath mid-ocean ridges (Langmuir et al., 1992). As depicted in Fig. 3, the mantle starts melting when its solidus is crossed and mantle parcels undergo progressive melting as they ascend. In a forward model, this is approximated by melting mantle parcels at the pressure of the solidus intersection by one percent, extract the magma, recalculate the residual mantle composition and mineralogy, and so forth with pressure of steps decreasing by 1 kbar for each computation. If this is the way the mantle melts, the mantle parcels that reach the top of the melting regime undergo a large degree of melting whereas mantle parcels at the side of the triangle only melts by a few percent before they flow to the side and out of the melting zone (Fig. 3). If melt at any given time is pooled from the entire melt zone it will be a mixture of low and high degree melts. Clearly, the low-degree melts stem from mantle in which sulphides retain the PGE inventory and they will be S-saturated and contain little or no PGE, and will have high concentrations of incompatible lithophile trace elements such as REE. On the other hand, the higher degree melts from the top of the melting column eventually will reach the point at which the sulphides are dissolved and the magmas become PGE-rich. The geometric relations predict that F is half of the maximum degree of melting reached for the topmost mantle parcel of the triangle. Rehkämer et al. (1999) was the first to propose that such a melting model may explain the formation of ocean-island basalts with moderate PGE as the result moderate F .

The results of our forward modelling of the PGE composition of basalt formed during continental breakup is summarised in Fig. 4 and indicate that if the mantle contains 200 S and the magma can dissolve 1000 p.p.m. S, then magmas with 10 p.p.b. Pd may be produced for F larger than 10%. In summary, forward modelling of REE suggests that the high-Ti East Greenland flood basalts may form for F ranging from 5 to 9% and forward modelling of PGE suggests the same basalts may form only for F above 10% for reasonable mantle compositions. The F estimated from REE modelling obviously depends more on assumptions about source composition than that of PGE modelling. Hence, the major element, the REE, and the PGE composition of the high-Ti

flood basalts may all be explained by approximately 10% mantle melting.

The 'Lid'-Effect Drives Magma to S-Undersaturation

The pooled melts of the above model is a mixture of S-undersaturated (high-degree melts) and S-saturated basalts (low-degree melts) and it is not clear whether this mixed product is over- or undersaturated at the depth of the source. In any case, it appears unlikely that the S-undersaturated nature of the East Greenland flood basalts may be so strong without additional effects that drive the magma to S-undersaturation. In fact, there are two mechanisms that may increase the amount of S that may be dissolved in basalt, that is FeO content and pressure. Poulson and Ohmoto (1990) showed that the S-capacity of a magma increases with its FeO content. Because the high-Ti basalts differentiate towards iron enrichment (Brooks et al., 1991) it appears that the S-capacity of the magma may increase solely as a consequence of differentiation. Further, the REE's indicate garnet is present in the source. Hence the lid-effect predicts that the mean pressure of melting, P , is elevated. The FeO content of magma increase with P during melting and thus suggest that continental flood basalts have high S-capacity relative to MORB. Finally, but perhaps most important, Mavrogenes and O'Neill (1999) showed that magmas at high pressure can dissolve less S than the same magma at low pressure. For the reasons outlined above, we believe the high P of melting followed by adiabatic ascent of the magma may be very important in driving continental flood basalts to S-undersaturation.

Formation of the Platinova Reef, Skaergaard intrusion

The Platinova Reef consists of several orthomagmatic palladium- and gold-rich magnetite-gabbro layers at the top of the Middle Zone (Andersen et al., 1998). In terms of igneous differentiation, the host pyroxene and plagioclase composition indicates the reef formed from an evolved magma with mg# close to 23. Compared to PGE reefs of the Bushveld and Stillwater complexes, the Platinova Reef is not only unusual in that the magma was evolved, but also in that it is a Pd reef with Pt/Pd below 0.1. A surprising feature of the Greenland high-Ti basalts is that the Pt content, which ranges from 2 to 11 p.p.b., decreases slightly with fractionation, indicating Pt is incorporated into the fractionating gabbro assemblage (D^{bulk} is 2-3), perhaps as tiny Platinum-Group Minerals hosted by olivine (Momme et al., 2002). Hence Pd is fractionated from Pt and Pt/Pd decreases with differentiation (Fig. 2). In the same way as the Pt/Pd of the high-

Ti basalts decrease with mg# ranging from 60 to 43, we speculate that the low Pt/Pd of Platinova Reef (mg# ~25) may result from prolonged fractionation where Pt is included into and Pd excluded from the cumulates. This is illustrated as the Skaergaard trend in Fig. 2.

Implications for the Siberian Trap Magmatism

The Noril'sk mafic intrusions are associated with the Siberian Traps flood basalts and the parental magmas must have contained considerable amounts of PGE. Almost all the flood basalts of the region are depleted in PGE, saturated in S, and show isotopic evidence for the assimilation of S-rich sediments (Brügmann et al., 1993). However, in the top section of the volcanic pile there are lavas that contain significant amounts of PGE (Brügmann et al., 1993) and may not be S-saturated.

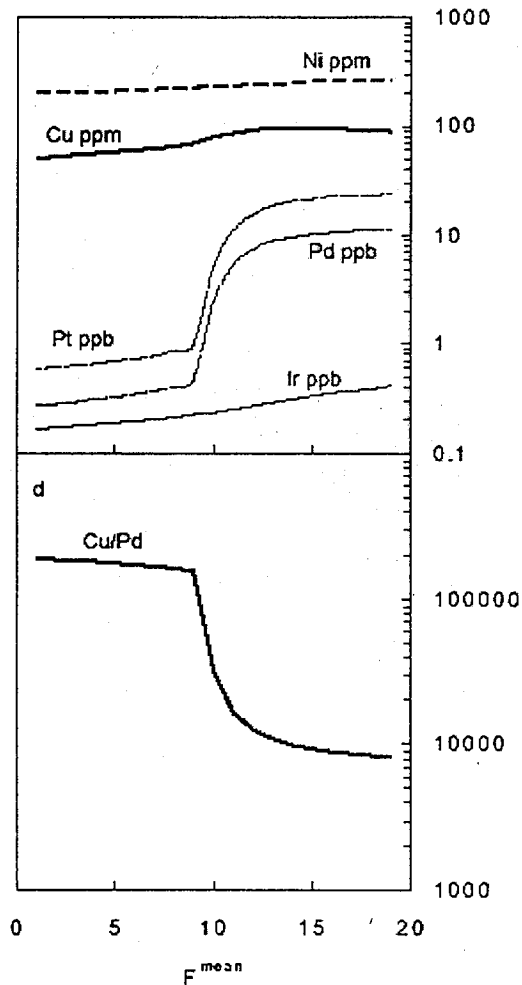


Figure 4. Modelled Ni, Cu, Pt, Pd and Ir composition with mean degree of mantle melting (F). Model calculations assume corner-flow melting regime as shown in Fig. 3, the mantle contains 200 S, and the magma can dissolve 1000 p.p.m. S (Momme et al., in review).

In the Mokulaevsky Suite, the concentration of PGE is similar to the high-Ti basalts of East Greenland (Fig. 1) whereas the REE patterns are steeper. The differences in REE may reflect a more enriched source for the Siberian flood basalts. We speculate that the primary mantle-derived versions of the Siberian Traps flood basalts was very similar to those of the East Greenland high-Ti basalts and that they could have formed at comparable mantle melting conditions. If so, the main difference between the two provinces lie in the crustal processes indicating abundant assimilation in the Siberian case and very little crustal contamination in the East Greenland basalts. However, we speculate whether mantle plume melting beneath a thick lithospheric 'lid' could also account for the composition of these occurrences.

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Local Variations and Regional Trends in PGE Geochemistry and Mineralogy in the Main Sulfide Zone of the Great Dyke, Zimbabwe

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Introduction

The Great Dyke of Zimbabwe constitutes the world's second largest reserve of PGE after the Bushveld Complex in neighbouring South Africa. In the Great Dyke, economic concentrations of PGE are restricted to sulfide disseminations of the Main Sulfide Zone (MSZ). The present work is based on the study of samples from new exposures of economic grade primary ores of the MSZ created in the course of exploration and mining at Hartley, Mhondoro, Ngezi, Unki and Mimosa (Fig. 1). New findings are reported on the geochemical distribution and the mineralogical siting of the PGE, the mineralogy of the PGM, and the regional variations of these parameters. The findings are discussed in view of the contrasting "magmatic sulfide" and "hydrothermal fluid" genetic models.

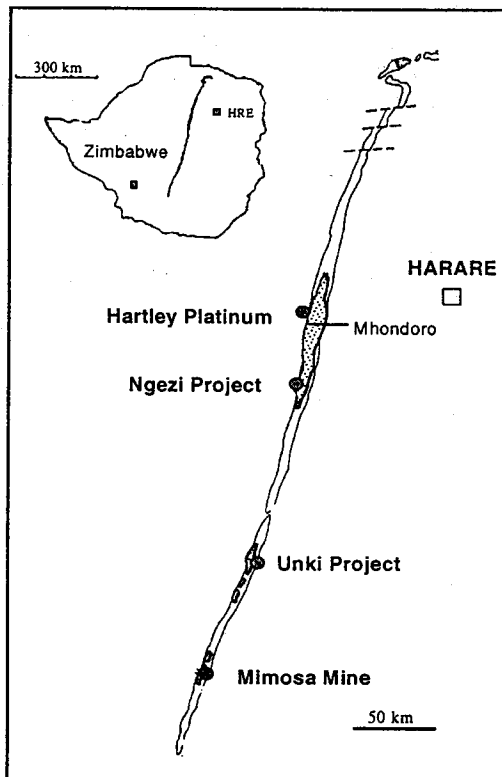


Figure 1. Sketch of the Great Dyke showing the localities of mines and prospects mentioned in the text.

Geologic Setting

The 2575.4 ± 0.7 Ma Great Dyke layered intrusion has a linear shape and strikes over 550 km NNE at a maximum width of about 11 km, cutting Archean granites and greenstone belts of the Zimbabwe craton. Stratigraphically, it is divided into a lower Ultramafic Sequence of dunites, harzburgites, olivine bronzitites and pyroxenites, together with narrow layers of chromitite, and an upper Mafic Sequence mainly consisting of plagioclase-rich rocks (norites, gabbro-norites, olivine gabbros). Economic concentrations of PGE, Ni and Cu in the form of disseminated sulfides occur some meters below the transition from the Ultramafic to the Mafic Sequence, in the several meters thick "Main Sulfide Zone", in pyroxenitic host rocks.

Samples and Methods

Twenty-one vertical profiles across the MSZ, between 1.5 and 5 m long, were taken from drill cores (Hartley, Mhondoro, Ngezi, Unki) or were cut underground (Hartley, Mimosa). The whole rock samples were analysed by XRF and ICP-MS. PGE contents were determined by INAA after Ni-sulfide extraction. Polished sections were investigated by reflected light microscopy, XRD, SEM/EDX, and electron microprobe.

Geochemical Distribution of PGE in the MSZ

Our study of 21 MSZ sections corroborated that, apart from minor variations, the MSZ is characterized by a typical vertical pattern of base metal sulfide and PGE distribution as known from the work of e.g. Prendergast (1988), Prendergast and Wilson (1989), and Wilson (2001). This pattern is characterized by "offsets", i.e. a certain degree of decoupling, and separation of the respective element distribution patterns and peak concentrations, namely of Pd from Pt, and also of all PGE from the base metals. Our data also confirm the subdivision of the MSZ into a lower PGE subzone and an upper BMS subzone, and the further subdivision of the PGE subzone into a lower (Pd>Pt) and an upper part (Pd<Pt).

The vertical distributions of the base metals and sulfur (BMS; Cu, Ni, S), Pt and Pd are shown in Fig. 2, which is an idealized profile across the MSZ, about 3 metres wide, mainly based on data from Hartley mine. Notable are three variably pronounced peaks in the PGE subzone. (i) PGE-1 close to the base is a broader area characterized by slightly increasing PGE contents at low levels of base metals and sulfur, Pd > Pt, and low contents of Σ PGE relative to Σ IPGE (Σ PGE/ Σ IPGE ratios are ca. 9 at PGE-1 compared to ca. 20 at positions PGE-2 and PGE-3 at Hartley mine). Further upward, Σ IPGE remain at nearly constant levels, Pd values exceed those of Pt, and the base metal and sulfur contents continue to increase slightly. (ii) At position PGE-2, the peak of Pd is reached, which coincides with a subsidiary peak of Pt and slightly elevated BMS contents. (iii) PGE-3 is the peak of Pt, distinctly offset from the peak of Pd. At this position, shortly before or overlapping with the strongest BMS peak, Pd shows a subsidiary peak and Rh, Ru, Ir and Os reach maxima. This justifies a further refinement of the structure of the PGE subzone (basal, centre and top part; indicated by the peaks PGE-1, PGE-2 and PGE-3).

Chondrite-normalized patterns of the averages of the PGE subzone of the MSZ show steep upward trending curves from Os to Pd (Fig. 3). This type of pattern is regarded typical for PGE mineralization produced by magmatic scavenging of PGE through segregating sulfides. The tight array displayed by the various MSZ localities indicates consistent contents and proportions of PGE within the MSZ over wide areas.

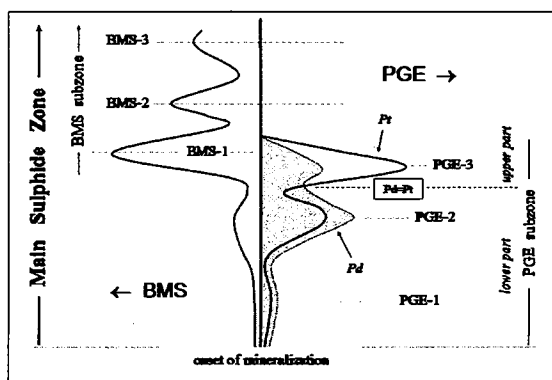


Figure 2. Sketch of key features of element distributions and subdivision of the MSZ. BMS = base metal sulfides.

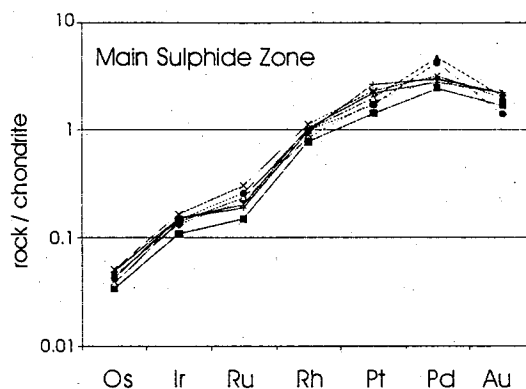


Figure 3. C-1 normalized average PGE contents of the MSZ (PGE subzone) from various Great Dyke localities.

Platinum-Group Minerals (PGM)

The pyroxenites of the MSZ contain between 0.5 and 10 vol% of sulfides – pyrrhotite, pentlandite, chalcopyrite and subordinate pyrite – which mainly occur interstitially to cumulus orthopyroxene. The PGM are usually included in pyrrhotite or chalcopyrite, rarely in pentlandite, or they occur at sulfide/sulfide or sulfide/silicate contacts or within silicates. PGM intergrown with pyrite are extremely rare. PGM grain sizes (apparent maximum diameters) range from < 5 to 50 μ m in general, but may reach up to 373 μ m in the longest dimension.

Altogether, 801 grains of PGM were analyzed by electron microprobe. The suite of PGM (proportions by number) comprises (Pt,Pd)-bismuthotellurides, mainly moncheite, maslovite, merenskyite, and michenerite (50.1%), followed by sperrylite (19%), the PGE-sulfarsenides hollingworthite, platarsite, irarsite and ruarsite (11.9%), the (Pt,Pd)-sulfides cooperite and braggite (8.5%), laurite (5.0%), and some rarer phases.

The various PGM are heterogeneously distributed within the MSZ sequence. At Hartley for example, sperrylite occurs throughout the PGE subzone of the MSZ, whereas cooperite/braggite only occur in samples from and below the Pt peak. The (Pt,Pd)-bismuthotellurides concentrate in the upper portion of the PGE subzone of the MSZ only, mainly in the regions around PGE-3 and BMS-1.

Trace PGE Contents in Sulfides

Micro-PIXE studies had shown that pentlandite from the lower PGE subzone of the MSZ at Hartley, Unki and Mimosa had elevated contents of Pd (maximum value 2,236 ppm Pd), and Rh (max. 259 ppm Rh), and SIMS analysis revealed that pyrite was a carrier of Pt (0.4-244, mean 35.5 ppm; n=37). Electron microprobe analysis using the CSIRO-Trace program was performed on pentlandite from Hartley, Unki and

Mimosa. Maximum Pd and Rh contents measured are 2,506 ppm Pd and 562 ppm Rh. In the example below (Fig. 4; core CD-02, Hartley mine), Pd contents in pentlandite appear to attain and stay at a plateau level of ca. 1,200 ppm Pd through most of the PGE subzone of the MSZ (maximum average value = 1,379 ppm Pd), interrupted by a return to lower values (≈ 300 ppm Pd) between PGE-1 and PGE-2. Further up-sequence, Pd contents in pentlandite drop to values below the detection limit of the method (about 40 ppm Pd) just before reaching the Pt peak (PGE-3). The data therefore underline that most of the Pd is generally hosted by pentlandite in the lower and central parts of the PGE subzone.

Figure 5 summarizes the major geochemical and mineralogical variations and trends in the MSZ. Note the elevated contents of Pd in pentlandite in the lower part of the PGE subzone, overlapping distributions of the various PGM groups in the profile, and the preferred concentration of laurite in the basal part of the PGE subzone. Gold and the tellurides of Bi, Ag, Ni and Pb mainly occur in the transition from the PGE to the BMS subzone of the MSZ. Changes in pyrrhotite compositions, and the presence of pyrite in the upper part of the PGE subzone and the BMS subzone indicate increasing fS_2 up sequence.

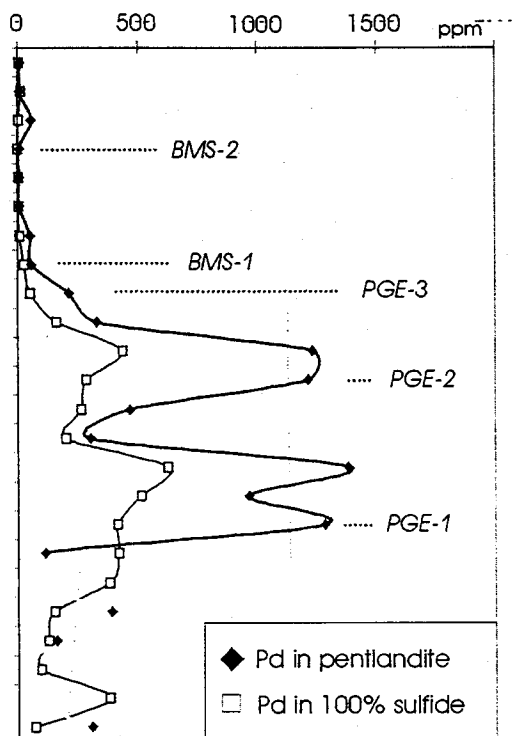


Figure 4. Vertical distribution of Pd in pentlandite (ppm) analyzed by electron microprobe, and Pd in 100% sulfide (ppm) calculated from whole rock data.

Regional Variations in PGM Proportions

Major regional differences were observed in PGM proportions (Table 1) and grain sizes. The MSZ samples of the North Chamber (Hartley + Mhondoro + Ngezi) have higher proportions of (Pt,Pd)-bismuthotellurides (63.8%) and cooperite plus braggite (12.7%), and lower proportions of sperrylite (11.0%), PGE-sulfarsenides (2.4%), and laurite (1.7%) compared to those of the South Chamber. The respective South Chamber (Unki + Mimosa) values are: (Pt,Pd)-bismuthotellurides (35.2%), cooperite/braggite (3.9%), sperrylite (27.6%), PGE-sulfarsenides (22.1%) and laurite (8.6%). The different PGM proportions and accompanying different apparent average PGM grain sizes (39 μm in the North Chamber versus 18.5 μm in the South Chamber) must be considered in the metallurgical treatment of the ores.

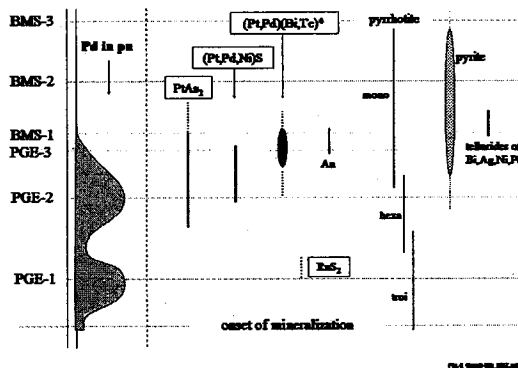


Figure 5. Summary of key geochemical and mineralogical features of MSZ profiles as exemplified at Hartley mine. Pd in pn = Pd contents in pentlandite. (Pt,Pd)(Bi,Te)* = (Pt,Pd)-bismuthotellurides. Pyrrhotite: mono = monoclinic; hexa = hexagonal; troi = troilite.

Table 1. Comparison of PGM proportions by number of various MSZ localities. (Pt,Pd)(Bi,Te)* = (Pt,Pd)-bismuthotellurides. PGE-AsS⁺ = PGE-sulfarsenides.

locality →	Hartley	Mhondoro	Ngezi	Unki	Mimosa
PGM (n) →	181	43	199	250	134
type [%] ↓					
(Pt,Pd)(Bi,Te)*	71	58	60	39	28
PtAs ₂	11	9	11	28	26
(Pt,Pd)S	11	5	16	4	3
PGE-AsS ⁺	2	4	2	16	33
Pt-Fe alloys	1	11	6	—	—
others	4	11	5	12	13

Discussion and Conclusions

The combination of geochemical and mineralogical data suggests the following model for the formation of the PGM in relation to PGE distributions. It is postulated that all PGE were primarily scavenged by sulfide in the magmatic stage due to their high partition coefficients in favour of sulfide. The metal profiles indicate sulfide accumulation in a number of distinct successive layers, each with its own geochemical characteristics. The general trend within the MSZ is one of upward increases in $\Sigma\text{PGE} / \Sigma\text{IPGE}$, Pt/Pd and Cu/Ni ratios, as well as $f\text{S}_2$. Apparently, the various elements were extracted from the magma in consecutive batches which show successive enrichments of the chalcophile elements according to their apparent D values (in case of the MSZ, $D_{\text{Pd}} > D_{\text{Pt}}$, for example). Upon cooling, in the subliquidus stage, the PGE-bearing sulfide droplets recrystallized to form aggregates of pyrrhotite, pentlandite, chalcopyrite and some pyrite. The development down-temperature was accompanied by continuous subsolidus annealing of the sulfides and their internal crystal ordering. Ballhaus and Ryan (1995) demonstrated that Pd and Rh are accommodated within the crystal lattice of pentlandite down to low temperatures, a finding corroborated by our direct measurements of these elements in pentlandite. In contrast, Pt is incompatible with any of the sulfides forming (except for low levels in pyrite) and was thus expelled from their crystal lattices. Pt released from sulfides during the annealing process readily reacted with As, Bi and Te due to the high affinity of these metalloids for Pt, to form sperrylite and bismuthotellurides. Although occasional intergrowths of cooperite/braggite and (Pt,Pd)-bismuthotellurides indicate that these mineral groups do not exclude each other, their distribution patterns within the MSZ are antipathetic.

Furthermore, the high concentrations of Bi and Te combined with the availability of Bi and Te at the position of the main BMS peak (BMS-1) led to a pronounced competition also for Pd. It is postulated that a chemical gradient existed between sulfide-bound Pd in pentlandite and Te and Bi which favoured the formation of Pd-rich bismuthotellurides. It is assumed that these reactions largely took place under isochemical conditions on a sample scale, and interstitial magmatic-hydrothermal fluids probably aided the mass transfer. It can also be speculated that a general upward gradient within the MSZ (peak concentrations of reactive Te and Bi at BMS-1) led to a certain chromatographic control, which may have supported the small-scale redistribution, the separation of certain metals, and the reactions that

formed the PGM in a vapor-refining zone as fluids evolved during solidification of the cumulus pile (Boudreau and Meurer 1999). Furthermore, the elevated $\Sigma\text{IPGE} / \Sigma\text{PGE}$ ratios at the base of the MSZ, closely following the onset of mineralization, are conspicuous. They may be explained by the preferential extraction of IPGE first, probably due to the formation of laurite, as indicated by elevated proportions of laurite in the basal part of the MSZ.

PGE distribution patterns and average contents in the MSZ are near-identical at the various localities studied. In contrast, PGM proportions differ widely between the ores of the North and South Chambers. This finding underlines that the formation of the PGM is mainly controlled by the local availability of reaction partners for the PGE. The elevated proportions of PGE-arsenides and -sulfarsenides in the South Chamber (together 49.7%) probably indicate higher fugacities of arsenic in the South Chamber magmas at the stage of MSZ formation, possibly due to contamination of the magmas by assimilation of country rocks (e.g. Oberthür et al. 2002).

The findings provide new indications towards the mechanisms that led to the distinct vertical element and mineral distribution patterns of the MSZ. Sulfide segregation appears to be the most important factor in the primary magmatic concentration of the PGE. It is concluded here that the orthomagmatic model elegantly explains the observed metal zonation patterns in principle. The degree to which internal fluids and certain metalloids (As, Te, Bi) have influenced the redistribution of the PGE and the formation of PGM still has to be ascertained.

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Evaluation of PGE Potentiality of Layered Gabbros in the Samail Ophiolite, Oman: Results of Geochemical Exploration and Petrological Examination

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Introduction

Major PGE deposits are found in the mafic-ultramafic layered intrusions in the continental environment, and most of them are located in the stable cratons. Discoveries of PGE and gold mineralizations have been reported from the Skaergaard intrusion and several other Tertiary mafic-ultramafic layered intrusions of North Atlantic Igneous Province in Greenland and northwest Scotland (Bird et al., 1991; Bird et al., 1995; Butcher et al., 1999). Magmatism formed those intrusions in the North Atlantic Igneous Province was resulted from mantle plume activity at the opening of the north Atlantic Ocean (Saunders et al., 1997). Therefore, the discoveries of the PGE mineralizations in the gabbros related to the opening of the north Atlantic Ocean encourage to search the PGE mineralization in the oceanic environment. Large gabbro bodies are present in the oceanic crust, though exposures of the gabbro bodies are limited. The layer 3 of oceanic crust is composed of gabbro with typical thickness of 4km. The oceanic crust is exposed in ophiolite bodies representing upper part of oceanic lithosphere. The ophiolite consists of thin sediments, pillow basalts, sheeted dyke complexes, gabbros, and ultramafic mantle unit. Lower part of the gabbros is typically layered with ultramafic cumulate, thus, the oceanic gabbros have lithology similar to the typical layered mafic-ultramafic intrusions which host the major PGE deposits. However, limited evaluations of oceanic gabbros for PGE resources have been conducted. The Samail ophiolite in Oman is one of the best exposed ophiolite complexes in the world, therefore it is best place to evaluate the potentiality of PGE mineralization in the oceanic gabbros. Furthermore, at one locality in the Samail ophiolite, magmatic sulphide ores were found in the gabbro (Lachize, et al., 1991; Lorand and Juteau, 2000). Although, Prichard et al. (1996) considered that the Samail ophiolite is a member of low PGE type ophiolite. Present study evaluates PGE and gold potentiality in the gabbros of the Samail ophiolite, and examines the exploration methodology in the ophiolite area.

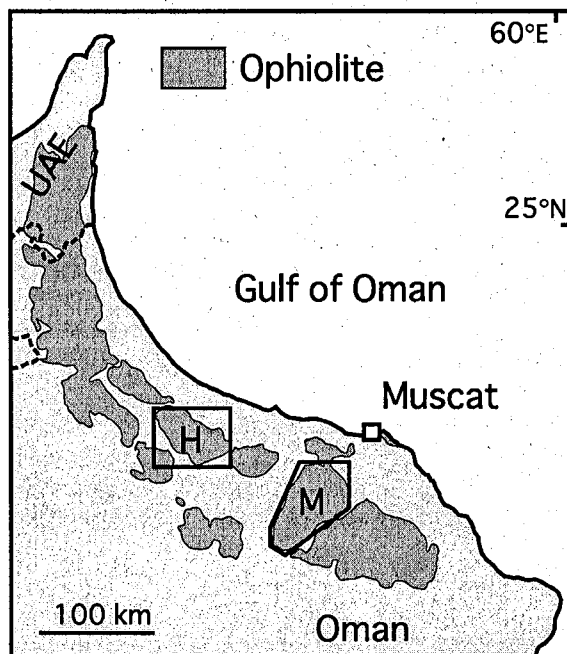


Figure 1. The Samail ophiolite and studied areas. H: Haylayn block. M: Maqsad block.

General Geology

The Samail ophiolite comprises of rocks derived from the Cretaceous oceanic crust and mantle which were obducted on to the Arabian continent during the late Cretaceous, and occurs in northern Oman (Fig. 1). The ophiolite is divided into 13 blocks. Magmatic sulphide ores were found from the Wadi Haymiliyah in the Haylayn block (Lachize, et al., 1991; Lorand and Juteau, 2000). The Haymiliyah outcrop is only known location of gabbro-hosted magmatic sulphide ores in the Samail ophiolite. The Haylayn and Maqsad blocks (Fig. 1) are selected for present study. In the Haylayn block, the ophiolite unit has northwest general trend. Gabbros occur in the central part of the block, and lower mantle unit is present southwest of the gabbros. In the Maqsad block, gabbros are present subhorizontally on the mantle ultramafic rocks, and occur mainly in the north and southeastern parts of the block. Layered gabbros are found in the lower portion of the gabbro unit,

and laminate and massive gabbros are present in the upper portion of the unit.

Geochemical Survey

As the exploration in East Greenland resulted in discovery of PGE and gold mineralization of the Skaergaard intrusion had used the stream sediment geochemical survey at the initial stage (Brooks, 1989), the geochemical survey was conducted in present study. Streams or rivers in the Oman Mountain are generally dry, and are described as wadi. However, occasional rainfalls are sufficient to form thick stream sediments in the dry river. We have collected stream sediments of under 60-mesh grain size by dry sieving at the sampling locations. Samples were collected mainly in the area where gabbros are exposed. Major and trace element data were obtained by XRF. Standard fire assay using PbO as collector with ICP-MS finish was used to obtain concentrations of Pt, Pd, and Au. Selected samples were analyzed with NiS fire assay with ICP-MS finish. Most of samples were also analyzed by ICP-AES with sample dissolution by acid reaching method, and the results were compared with whole rock XRF data. Our survey also checked host lithology and structure of gabbros and possible occurrences of sulfide minerals in the gabbros.

Results

SiO₂ contents of the stream sediments from the study areas range from 42 % to 49 %, indicating that those have similar chemical composition to the ophiolite unit. N-MORB normalized plot (Fig. 2) of average chemical compositions of the stream sediments for two blocks shows that the stream sediments are depleted in incompatible elements and enriched in MgO, Cr, and Ni. Although the stream sediments are collected in the area with exposure of gabbros, the difference of the chemistry from N-MORB may be resulted by selective depletion or enrichment of some elements during erosion, transportation, and deposition to form the stream sediments. High Cr concentrations must be resulted by incorporation of chromite grains in the sediments. Furthermore, as the area contains sub-recent alluvial terraces typically consist of calcrete, high concentrations of CaO and Sr must be derived from those. The average MgO, Ni, and Cr values of the stream sediments from the Haylayn block are higher than those of the Maqsad block, suggesting the stream sediments of the Haylayn block contain more materials from ultramafic part of the ophiolite.

Concentrations of Pd, Pt, and Au in the stream sediments are generally low. As detection limit of Pt is high (10ppb), most of the stream sediments have Pt concentration less than the

detection limit. Pd content ranges from less than 2ppb to 16ppb. Although the range of the Pd concentration is small, Pd values show systematic geochemical relationships with other elements. Pd vs. Cr₂O₃ diagram (Fig. 3) indicates two trends. The first trend indicates small Pd enrichment with increase of Cr₂O₃, suggesting enrichment of Pd associated with chromites. The other trend is Pd enrichment without or with only small increase of Cr₂O₃ value. This trend might suggest the presence of PGE mineralization related with primary magmatic sulfide.

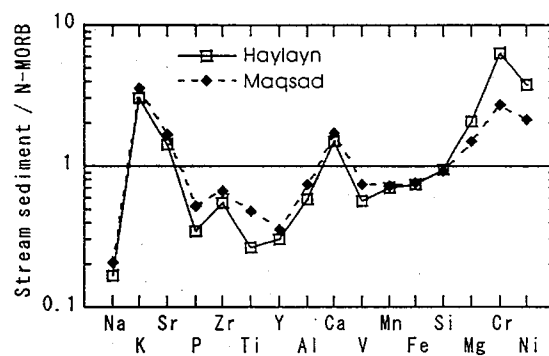


Figure 2. N-MORB normalized diagram for average compositions of stream sediments from the Haylayn and Maqsad blocks in the Samail ophiolite. N-MORB values are calculated from data of Sun et al. (1979). Elements are positioned according to their relative incompatibility for partial melting of mantle rocks.

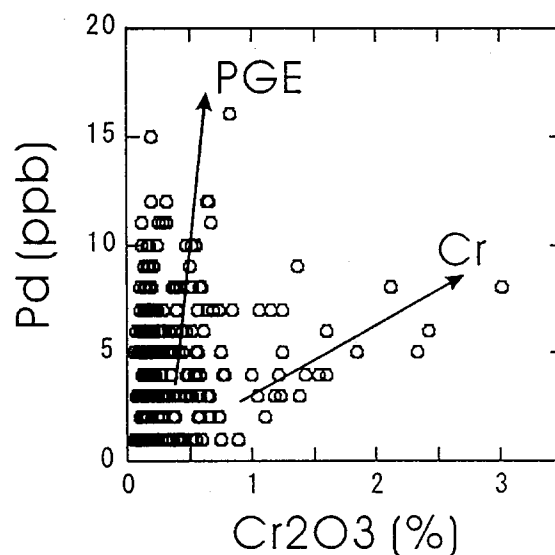


Figure 3. Pd vs. Cr₂O₃ variation diagram for stream sediments from the Haylayn block in the Samail ophiolite. The small Pd enrichment trend associated with chromites is indicated by the arrow (Cr). Another trend (PGE) may be associated with sulphide minerals.

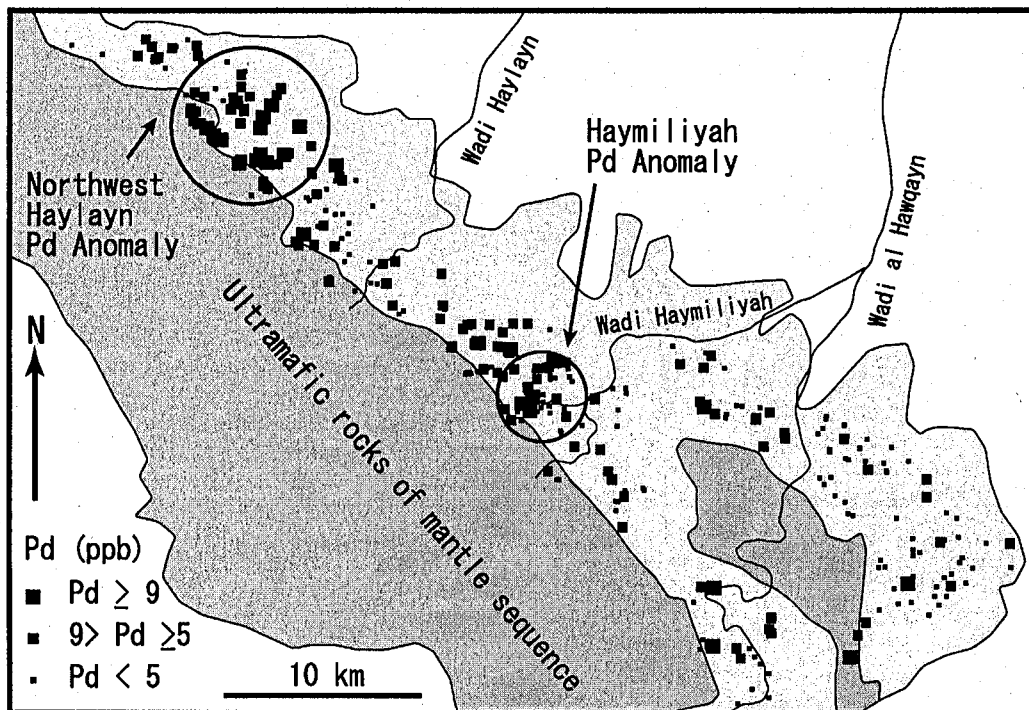


Figure 4. Stream sediment geochemical anomaly map of Pd for the Haylayn block. Light gray area is crustal sequence of the ophiolite unit.

Various geochemical maps have been constructed. Most of the maps reflect main lithology, and are conformable to geologic maps. For example, MgO, Cr, and Ni values increase in the area closed to the mantle sequence. Stream sediments with high Pd values are found in the two areas from the Haylayn block (Fig. 4). Small anomaly in the central part of the block is situated around Wadi Haymiliyah. Occurrence of sulphide ores (Lachize, et al., 1991; Lorand and Juteau, 2000) is located in the center of the anomaly, indicating the stream sediment geochemical survey is an effective method to show presence of such sulfide ores. Another area is located in the northwestern part of the block. The area is considered to be potential area for the PGE mineralization, and further detailed stream sediment sampling and geological survey have been conducted.

During the stream sediment sampling, observation of the gabbros has been made to find any sulphide bearing zone. Copper stains are found at many outcrops, indicating presence of chalcopyrites in the gabbro. Primary sulphide bearing zones have been found at two new locations. However, PGE concentrations of

samples from the sulphide rich zones are low, and are similar level to those from the Wadi Haymiliyah (Lorand and Juteau, 2000). Characteristically yellow alteration zones that were resulted by weathering of primary sulphide minerals, have been found at various location in the gabbro. In the centre of the northwest Pd anomaly of Haylayn block, extensive yellow alteration zones are found. Full PGE analysis of an altered sample indicates slightly high Pd concentration.

Summary

Present study intends to evaluate PGE and gold potentiality in the gabbros of the Samail ophiolite, and to examine the exploration methodology in the ophiolite area. The stream sediment geochemical survey was conducted in the two blocks of the Samail ophiolite. The geochemical characteristics of stream sediments are resulted by two components, ophiolitic rocks and alluvial sediments. Pd in the stream sediments shows two enrichment trends, one with small enrichment with chromites and other may be related with sulphide ores. Two Pd geochemical anomaly areas were found in the Haylayn block. Although absolute level of Pd anomaly is not too high, it is

significantly higher than the back ground level. The one of Pd anomaly area indicates the location of sulphide ores, suggesting the stream sediment geochemical survey is an effective method to find such sulfide ores. Yellow alteration zones that were resulted by weathering of primary sulfide minerals, have been found at various locations in the gabbro. In addition to the alteration zones, primary sulphide mineral bearing zones in the gabbro have been found at two new locations. The findings of additional primary sulphide zones in the gabbro suggest that unusual condition to form primary sulphide zones in the oceanic gabbro may not be necessary.

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APPENDIX 2

Report of
the 9th
International Platinum Symposium

*- Geology, mineral occurrences and PGE
exploration in the Duluth Complex -*

The survey team attended the pre-symposium excursion of the 9th International Platinum Symposium, "PGE Occurrences in Mafic Intrusions around Western Lake Superior, USA and Canada", which are held from July 13 to July 20, 2002. We summarize as follows the geology and mineral deposits of the Duluth Complex based on the pre-symposium excursion guide, Miller and Severson (2002), Severson et al. (2002), and Miller et al. (2002), and so on. The purpose of attending the symposium and the excursion was securing new insight and exploration criteria on Cu-Ni-PGE deposits, which is a mineralization being targeted in the present project of the Parana Basin Area, Brazil.

1 – 1 . Geology of the Duluth Complex

1 – 1 – 1 . Tectonic Setting

The Duluth Complex and associated Keweenaw intrusions in northeastern Minnesota constitute one of the largest mafic intrusive complexes in the world, second only to the Bushveld Complex of South Africa. These rocks cover an arcuate area of over 5,000 square Kilometers (Fig.1, Fig.2, and Fig.3). The intrusive rocks and comagmatic flood basalts underlying most of northeastern Minnesota were emplaced during the development of the Mesoproterozoic Midcontinent rift about 1.1 billion years ago.

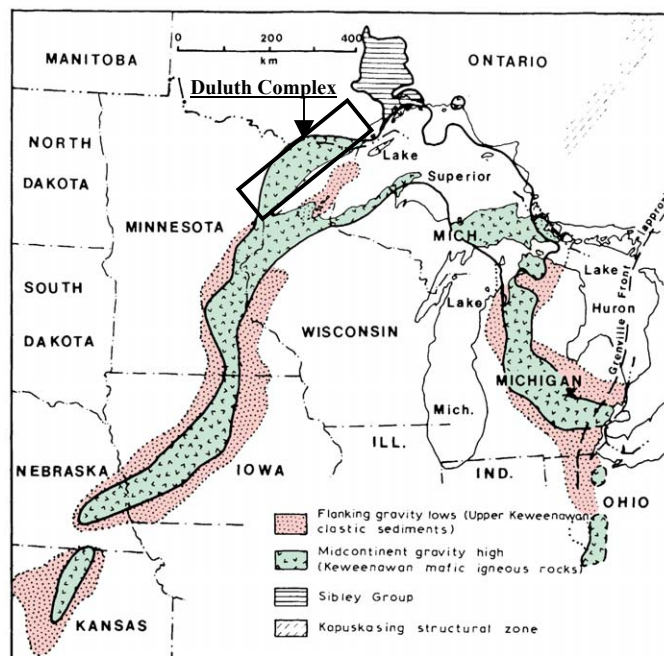


Fig.1 Generalized geological map of the Midcontinent Rift of North America(Weiblen, 1980).

The Midcontinent rift can be traced by its geophysical signature from its exposure in the Lake Superior region along a 2,000- kilometer-long, segmented, arcuate path from Kansas to lower Michigan. The rift is infilled with a lower sequence of volcanic rocks, locally as much as 20- kilometersthick, and an upper sequence of fluvial sedimentary rocks. Conceptually, the evolution of the Duluth Complex and related igneous rocks is interpreted in terms of the interactions between a deep mantle plume and mature continental lithosphere. Geochronologic studies indicate that the magmatic activity occurred in four distinct stages within the span of about 23 million years.

The early magmatic stage (1109 to 1107 Ma) occurred during a time of reversed magnetic polarity and was characterized by the rapid and voluminous eruption of initially primitive magmas derived directly from the plume. This quickly gave way to more evolved and crustally contaminated mafic magmas as well as felsic melts. This stage is thought to represent the impact of the mantle plume head with the initially cool and brittle lithosphere, the rapid heating of the lithosphere by mantle-derived melts, and the eventual contamination of those melts by their staging in the lower crust. This mafic underplating likely resulted in felsic melts being formed by partial melting of the lower crust.

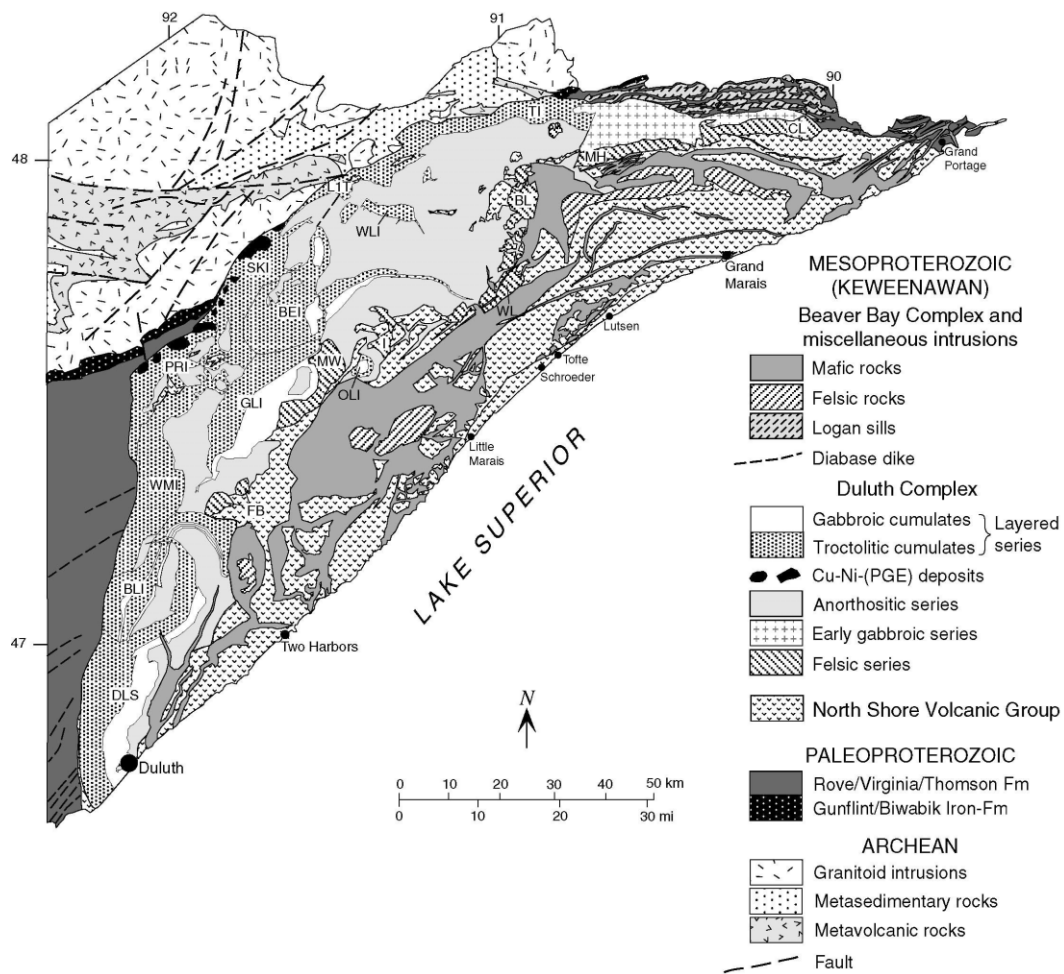
The latent magmatic stage (1107 to 1102 Ma) was a period during which volcanic activity was largely dormant, except for possibly rhyolitic volcanism. This stage is thought to represent a period of continued mantle plume upwelling and melting, extensive crustal underplating, and lower crustal melting. Crustal anatexis created a rheological and density barrier to the passage of mafic magmas to the upper crust.

The main magmatic stage (1102 to 1094 Ma) was a period of normal magnetic polarity conditions during which volcanic activity was renewed. This stage was characterized by moderate rates of eruption of uncontaminated (except for rhyolite) but diverse magma compositions. It is thought to represent the onset of upper crustal separation, the evacuation of lower crustal magma chambers, and continued mantle plume melting.

The late magmatic stage (1094 to 1086 Ma) was a period of waning volcanic activity but continued subsidence of the rift basin. This led to the development of evolved composite volcanoes and the interbedding of lava flows and basin sediments. This period probably represents the loss of the plume heat source by plate drift and the thermal collapse of the rift basin.

U-Pb ages indicate that various intrusions of the Duluth Complex were emplaced during the early and main magmatic stages of the Midcontinent rift. Early stage intrusions include the large granophyre bodies occurring in the roof zone of the Duluth Complex (the Cucumber Lake, Misquah Hills, and Whitefish Lake granophyres) and the gabbroic intrusions occurring at the northeastern end of the Duluth Complex (the Crocodile Lake and Poplar Lake intrusions). Gradational

relationships between the roof zones of these gabbros and the overlying granophyre bodies imply that the felsic rocks preceded emplacement of the gabbros. The overall evolved nature of these mafic and felsic intrusions implies that they were emplaced late in the early magmatic stage, probably contemporaneous with the eruption of the evolved and contaminated Hovland lavas. A precise age of 1106.9 ± 0.8 Ma for the basal gabbro of the Poplar Lake intrusion is consistent with this interpretation. The large volume of anorthositic rocks and the eleven discrete mafic layered intrusions that make up the remainder of the Duluth Complex were emplaced during the main magmatic stage.



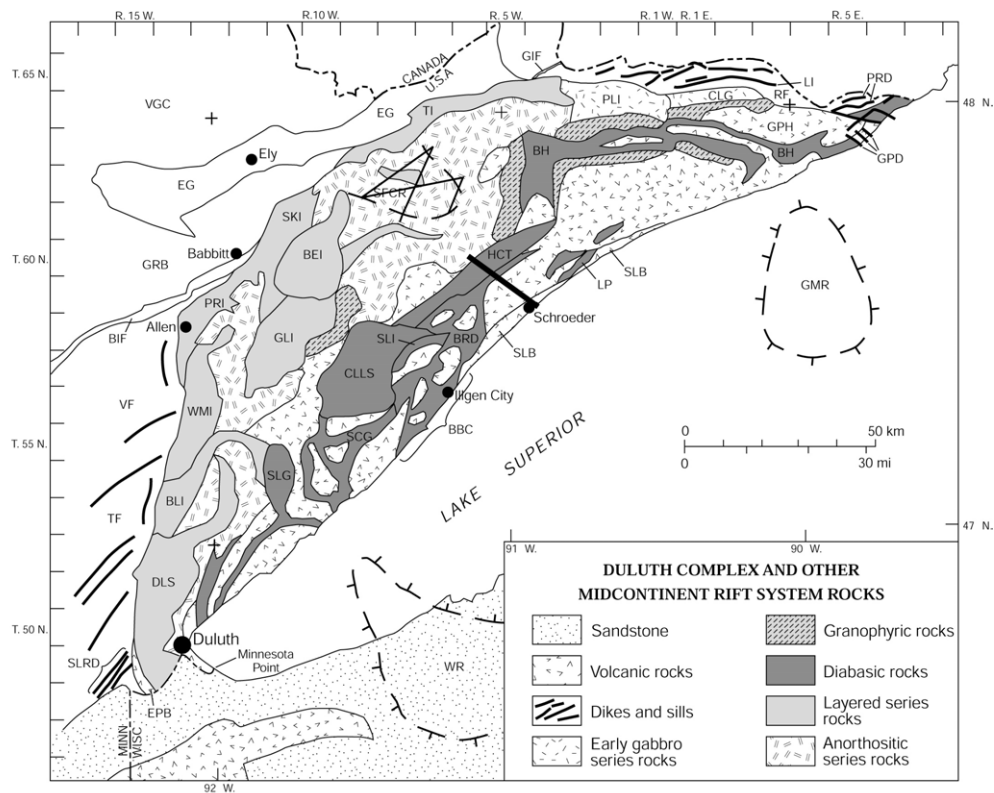
Generalized geology of northeastern Minnesota highlighting individual intrusions of the Duluth Complex. Layered series intrusions are:

- | | | |
|--------------------------------|--------------------------------|-------------------------------|
| BEIÑ Bald Eagle intrusion | BLIÑ Boulder Lake intrusion | DLSÑ Layered series at Duluth |
| GLIÑ Greenwood Lake intrusion | L1TÑ Lake One troctolite | OLIÑ Osier Lake intrusion |
| PRIÑ Partridge River intrusion | SKIÑ South Kawishiwi intrusion | TIÑ Tuscarora intrusion |
| WMIÑ Western Margin intrusion | WLIÑ Wilder Lake intrusion | |

Felsic Series intrusions:

- | | |
|---|------------------------------|
| BLÑ Beth Lake granophyre (and Wine Lake monzodiorite) | CLÑ Cucumber Lake granophyre |
| FBÑ Fairbanks/Brimson granophyres | IÑ Isabella granophyre |
| MWÑ Mt. Weber granophyre | MHÑ Misquah Hills granophyre |
| WLÑ Whitefish Lake granophyre | |

Fig.2 Generalized geology of the Duluth Complex and surrounding area (Miller and Severson, 2002).



Abbreviation of geological units: BEI—Bald Eagle intrusion, BIF—Biwabik Iron Formation (Paleoproterozoic), BLI—Boulder Lake intrusion, BBC—Beaver Bay Complex, BH—Brule Lake and Hovland gabbros, BRD—Beaver River diabase, CLG—Crocodile Lake gabbro, CLLS—Cloquet Lake layered series, DLS—Layered series at Duluth, EG—Ely Greenstone (Archean), EPB—Ely’s Peak basalts, GIF—Gunflint Iron Formation, GLI—Greenwood Lake intrusion GPH—Grand Portage basalts & Hovland lavas GRB—Giants Range batholith (Archean), GPD—Grand Portage dikes HCT—Houghtaling Creek troctolite LI—Logan intrusions, LP—Leveaux porphyritic diorite PLI—Poplar Lake intrusion PRD—Pigeon River diabase, PRI—Partridge River intrusion RF—Rove Formation (Paleoproterozoic) SCG—Silver Creek gabbro, SLRD—St. Louis River dikes SLB—Schroeder–Lutsen basalts SKI—South Kawishiwi intrusion, SLG—Sawmill Lake gabbro SLI—Sonju Lake intrusion TF—Thomson Formation (Paleoproterozoic), TI—Tuscarora intrusion VGC—Vermilion Granitic Complex (Archean) VF—Virginia Formation (Paleoproterozoic), WMI—Western Margin intrusion

Areas with heavy hatchures designate basement highs as follows: GMR—Grand Marais ridge SFCR—Schroeder–Forest Center ridge WR—White’s ridge

Fig.3 Generalized geology of the Duluth Complex and associated rocks with emphasis on each intrusions(Chandler, 2002).

1 – 1 – 2 . General Definition of the Duluth Complex

The Duluth Complex is defined as a more or less continuous mass of mafic to felsic plutonic rocks that extends in an arcuate fashion from Duluth nearly to Grand Portage. It is bounded by a footwall of predominantly Paleoproterozoic and Archean rocks, a hanging wall of largely mafic volcanic rocks and hypabyssal intrusions.

Defining the Duluth Complex more genetically, it is composed of multiple discrete intrusions of mafic to felsic tholeiitic magmas that were episodically emplaced into the base of a comagmatic volcanic edifice between 1108 and 1098 Ma. Within the nearly continuous mass of intrusive igneous rock forming the Duluth Complex, four general rock series are distinguished on the basis of age, dominant lithology, internal structure, and structural position within the complex.

Felsic series - Massive granophyric granite and smaller amounts of intermediate rock that occur as a semicontinuous mass of intrusions strung along the eastern and central roof zone of the complex emplaced during early stage magmatism (~1108 Ma).

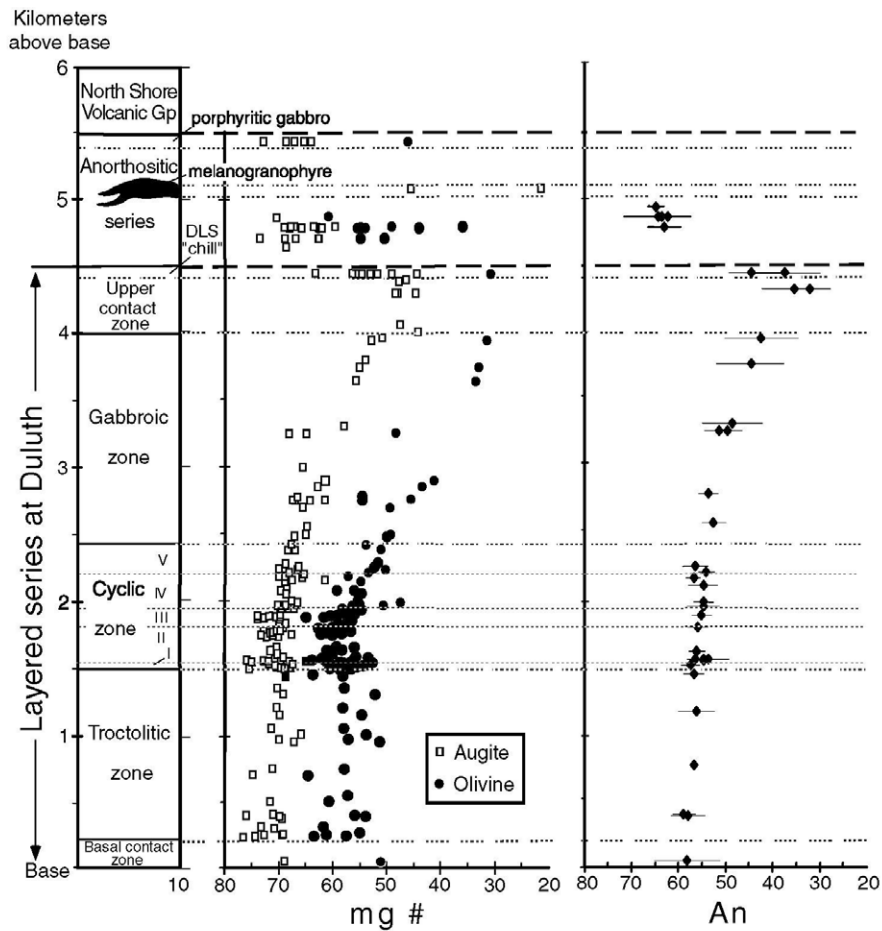
Early gabbro series - Layered sequences of dominantly gabbroic cumulates that occur along the northeastern contact of the Duluth Complex that were also emplaced during early stage magmatism (~1108 Ma).

Anorthositic series - A structurally complex suite of foliated, but rarely layered, plagioclase-rich gabbroic cumulates that was emplaced throughout the complex during main stage magmatism (~1099 Ma). The upper contact of the anorthositic series against volcanic or granophyric rocks is typically sharp and somewhat chilled.

Layered series - A suite of stratiform troctolitic to ferrogabbroic cumulates that comprises at least 11 variably differentiated mafic layered intrusions and occurs mostly along the base of the Duluth Complex. These intrusions were emplaced during main stage magmatism, but generally after the anorthositic series (~1099 Ma).

1 – 1 – 3 The main intrusions of the Duluth Complex

The layered series at Duluth is a well differentiated, 3- to 4.5-kilometer-thick, eastdipping, sheetlike mafic layered intrusion (Fig.4). Aeromagnetic data show that the layered series at Duluth has a north-south strike length of about 60 kilometers and eventually pinches out into the Boulder Lake intrusion. The hanging wall of the layered series at Duluth consists of olivine gabbroic and troctolitic anorthosite of the anorthositic series. The footwall of the layered series at Duluth at its southern end is composed of reversed polarity lavas of the Ely's Peak basalts.



Igneous stratigraphy and cryptic variation of olivine, pyroxene, and plagioclase composition through the layered series at Duluth. Mg values of olivine and pyroxene and An content of plagioclase are based on multiple microprobe analyses (error bars indicate one standard deviation).

Fig.4 Igneous stratigraphy and cryptic variation through the Layered Series at Duluth (Miller and Severson, 2002).

The layered series at Duluth is divided into five major zones based on dominant rock type. The basal contact zone is composed of coarse-grained, taxitic olivine gabbro and augite troctolite. The

lowermost cumulate sequence is the troctolite zone. It is 1- to 2-kilometers-thick, consists mostly of homogeneous foliated troctolitic cumulates(Fig.5A). The cyclic zone forms the medial section of the layered series at Duluth and is characterized by cyclical variations in cumulus mineralogy between troctolitic and gabbroic cumulates(Fig.5B). The persistent occurrence of gabbroic cumulates defines the gabbro zone. Gabbroic cumulates, in turn, grade upward into unlaminate (noncumulate) apatitic quartz ferromonzodiorite, which composes most of the upper contact zone. This quartz ferromonzodiorite complexly mixes with a fine-grained biotitic ilmenite ferrodiorite, which ultimately forms a "chilled" contact with anorthositic series rocks. A body of melanogranophyre that irregularly cuts through the anorthositic series probably represents the uppermost differentiate of the layered series at Duluth.

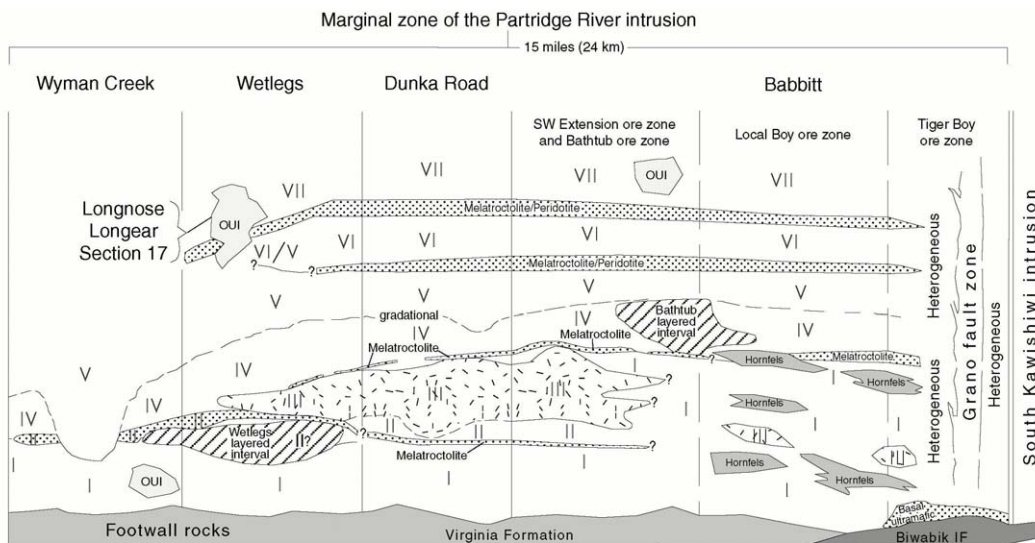
Although the layered series at Duluth is overall a well-differentiated intrusion, the repeated progression from troctolitic to gabbroic cumulates in the cyclic zone indicates that it did not fractionally crystallize as a closed system. The cyclic zone consists of at least five macrocycles within which troctolitic cumulates grade upward to gabbroic cumulates. Macrocycle boundaries are marked by the abrupt regression in the cumulus mineralogy from gabbroic back to troctolitic cumulates(Fig.5B). The gabbroic parts of the macrocycles commonly contain inclusions of anorthositic series rocks and the very uppermost parts of the macrocycles locally have discontinuous layers of fine-grained gabbroic adcumulate. Miller and Ripley (1996) suggested that the macrocyclic phase layering is predominantly related to devolatilization and decompression that attended magma venting events from a shallow chamber (less than 5-kilometers-deep), with magma recharge possibly having a secondary effect.



Fig.5 Photographs of
A)Troctolite zone(layered augite troctolite) and
B)Cyclic zone

The Partridge River intrusion has been studied in detail because it hosts at least four sub-economic copper-nickel deposits and at least seven potential Fe-Ti deposits. The Partridge River intrusion consists mainly of troctolitic cumulates that are exposed in an arc shaped area that extends from the Water Hen Fe-Ti deposit area to the Babbitt copper-nickel deposit. Miller and Ripley(1996) estimated that the Partridge River intrusion is 2.5-kilometers-thick. Its footwall includes the Paleoproterozoic Virginia Formation (slate and graywacke), and to a lesser extent, the Biwabik Iron Formation. The top of the Partridge River intrusion is in complex contact with anorthositic rocks, gabbroic rocks, and mafic volcanic. This assemblage of anorthositic, gabbroic, and hornfelsic rocks are also present as large inclusions within the interior of the Partridge River intrusion.

The basal 900 meters of the Partridge River intrusion are known in great detail from the abundance of exploration drill core. This marginal zone, consisting of varied troctolitic and gabbroic rock types, is subdivided into seven stratigraphic units that can be correlated over a strike length of 24 kilometers. All of these igneous units generally exhibit shallow dips (10o to 20°) to the southeast. The stratigraphy shown in Fig.6 is based on the re-logging of almost 700 drill holes, or 205 kilometers of drill core.



Generalized stratigraphy of the marginal zone of the Partridge River intrusion (modified from Severson, 1994). Stratigraphic relationships for the area between the Wyman Creek deposit and the Water Hen deposit are poorly understood and are not portrayed on this figure.

Fig.6 Generalized stratigraphy of the marginal zone of the Partridge River Intrusion (Miller and Severson, 2002).

Unit I:

Unit I is the dominant sulfide-bearing member of the Partridge River intrusion. It consists of a heterogeneous mixture of ophitic troctolitic to gabbroic rocks that contain abundant inclusions of hornfelsic sedimentary footwall rocks and minor thin, discontinuous layers of melatroctolite and peridotite. Noritic rocks are common at the basal contact and peripheral to sedimentary hornfels inclusions, probably due to contamination of the magma. An ultramafic interval, consisting of an oxide-bearing peridotite overlying an oxide-bearing pyroxenite, is present at the base of Unit I in areas where the Partridge River intrusion is in direct contact with the Biwabik Iron Formation .

Unit II :

This unit exhibits considerable variation from one copper-nickel deposit to the next. At the Dunka Road and Babbitt deposits, Unit II consists of homogeneous troctolitic rocks, with minor sulfide mineralization, and a fairly persistent basal ultramafic layer that separates Unit II from Unit I. The upper contact with Unit III is gradational at the Dunka Road deposit. At the Wetlegs deposit, Unit II is either a single ultramafic layer immediately beneath Unit III, and/or the Wetlegs layered interval, which consists of repeated, thin cyclic units that internally grade upward from ultramafic rock to troctolite. Still farther to the west at the Wyman Creek deposit, Unit II consists of a single ultramafic horizon that separates sulfide-bearing and heterogeneous troctolitic rocks of Unit I from homogeneous troctolitic rocks of Unit IV.

Unit III:

This unit consists of poikilitic leucotroctolite that commonly grades into poikilitic/ophitic augite troctolite. Because of its fine-grained texture and distinctive olivine oikocrysts that impart a mottled appearance, Unit III is a useful marker horizon, although it is absent from the Wyman Creek deposit and portions of the Babbitt deposit. Hornfelsed basalt inclusions, or roof rocks, are commonly associated with Unit III at the Dunka Road deposit. This relationship, and the highly gradational contact of Unit III with Unit II, suggest that Unit III may have formed as a roof cumulate during crystallization of Unit II.

Unit IV:

Homogeneous ophitic augite troctolite that contains a local basal unit of ultramafic rock. Unit IV typically exhibits a highly gradational upper contact with Unit V. A cyclic sequence of alternating troctolitic and ultramafic layers, termed the Bathtub layered interval, is present in the Bathtub ore zone of the Babbitt deposit.

Unit V:

Unit V consists of homogeneous, coarse-grained leucotroctolite. At the Wyman Creek deposit, drill hole relationships suggest that Unit V cuts downward into the lower units.

Units VI and VII:

Each of these units consist of homogeneous leucotroctolite that locally grades into ophitic augite troctolite; both also contain a fairly persistent ultramafic base. Several other similar units

(Unit VIII and up) are present above these two units, but they are not defined by drilling.

Oxide ultramafic intrusions (OUI):

Several late-stage pegmatitic plugs and vertical lenses of oxide-bearing ultramafic intrusions intrude the troctolitic rocks of the Partridge River intrusion. Severson and Hauck (1990) first used the acronym OUI to designate crosscutting pegmatitic bodies of peridotite, melatroctolite, melagabbro, and clinopyroxenite that contain a high percentage of coarse-grained oxides (15 to 100 percent). The OUI are spatially arranged along linear trends, suggesting that structural control was important to their genesis. In addition, the Longnose-Longear-Section 17 group of OUIs is positioned over a window in the basal contact where the Biwabik Iron Formation is the footwall rock. This relationship suggests a genetic link between assimilation of iron formation at the basal contact and formation of OUI along a coincident fault zone.



Wetlegs deposit, Unit I (D.H. W-8B): heterogeneous sulfide-bearing augite troctolite (Cu:0.25-0.34%, Ni:0.07-0.10%).



Wetlegs deposit, Unit II : abundant ultramafic layers in the Wetlegs area.

Fig.7 Photographs of A) Unit I, and B) Unit II of the Partridge River Intrusion.

Attributes of the more heterogeneous units (I and II) near the base of the Partridge River intrusion are interpreted to indicate rapid magma replenishment in a progressively developing magma chamber, and magmatic contamination from assimilated footwall rocks. The more homogeneous upper units (IV through VII), each floored by a persistent ultramafic layer, were probably emplaced later in a well-developed magma chamber. The ultramafic layers in the upper units, and abundant ultramafic layers of the Wetlegs and Bathtub layered intervals, probably represent the inception of episodic magma injection that crystallized more primitive ultramafic layers before mixing with the resident magma.

The South Kawishiwi intrusion hosts at least five subeconomic copper-nickel deposits and a potential platinum group element-copper-nickel deposit. The South Kawishiwi intrusion is dominantly composed of troctolitic cumulates that are exposed in an 8- by 32-kilometer arcuate band. Footwall rocks include the Virginia Formation in the Serpentine and Dunka Pit deposits, the Biwabik Iron Formation in the Dunka Pit and Birch Lake deposits, and the Archean Giants Range batholith in the Dunka Pit deposit north to the Spruce Road deposit. The presence of Biwabik Iron Formation as inclusions as far north as the Spruce Road deposit indicates that the majority of Paleoproterozoic units were assimilated and removed from the footwall during emplacement of the South Kawishiwi intrusion, leaving the Giants Range batholith as the dominant footwall rock type. Also present as inclusions in the Dunka Pit and Serpentine deposits are mafic volcanic hornfels (probably North Shore Volcanic Group) and quartz sandstone hornfels (probably equivalents to the Puckwunge and Nopeming Sandstones).

The South Kawishiwi intrusion consists of five major map units. These are, from the base upward:

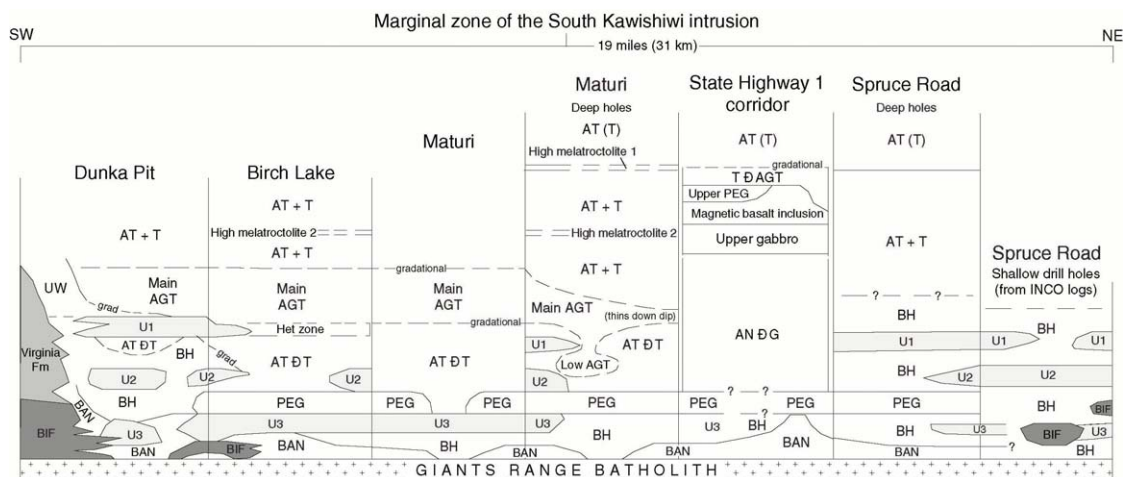
1. A basal contact zone that is a heterogeneous mix of sulfide-bearing troctolitic, gabbroic, and noritic rocks with abundant hornfels inclusions;
2. A thick unit of subophitic to ophitic augite troctolite that contains an internal ophitic olivine gabbro unit;
3. A discontinuous and localized layer of poikilitic leucotroctolite;
4. A thick, homogeneous sequence of ophitic troctolite;
5. An uppermost, thick sequence of homogeneous troctolite that contains numerous lensoidal layers and inclusions of anorthositic rocks.

Severson (1994) and Zanko and others (1994) further subdivided the marginal zone of the South Kawishiwi intrusion into 17 different lithostratigraphic units (Fig.8) that are present in over 180 drill holes over a strike length of 31 kilometers. Sulfide mineralization is confined to the BH, BAN, UW, and U3 units, and to a lesser extent the U1 and U2 units. Major marker horizons that are

correlated in drill holes include three horizons with abundant cyclic ultramafic layers (U1, U2, and U3 units) and a pegmatite-bearing unit (PEG unit). A large, anorthositic inclusion (less than 1-kilometer-thick) is intersected in six deep drill holes in the State Highway One corridor area.

The lowest units of the marginal zone are the most varied with respect to textures, rock type, and sulfide content. They are unevenly distributed along the strike length of the South Kawishiwi intrusion in a "compartmentalized" fashion, suggesting a complicated intrusive history (Fig. 8). The lowest units were emplaced early into several restricted magma chambers via repeated and close-spaced magmatic pulses. The U1, U2, and U3 units represent periods of rapid and continuous magma injection that crystallized more primitive ultramafic layers before mixing with the resident magma. The U3 unit is unique among the lower units in that it contains several massive oxide pods (titanomagnetite-rich) along its entire length. A spatial correspondence between the U3 unit and footwall iron-formation suggests that most of the massive oxide pods are iron-rich "restite" produced by the magmatic digestion of iron-formation. The U3 unit also contains the majority of the high platinum group element values that have been sampled within the South Kawishiwi intrusion.

The upper units of the South Kawishiwi intrusion are continuous throughout the intrusion. Their continuity and monotonous troctolitic composition suggest that they crystallized in a more quiescent and open magmatic system characterized by widely spaced magmatic pulses.



Generalized stratigraphy for the marginal zone of the South Kawishiwi intrusion (modified from Severson, 1994).

Fig.8 Generalized stratigraphy of the marginal zone of the South Kawishiwi Intrusion (Miller and Severson, 2002).

1 – 2. Mineral Deposits of the Duluth Complex

Increased demand and commodity prices for platinum group elements (PGEs) have recently reinvigorated exploration activity in the Duluth Complex and related Keweenaw intrusions in northeastern Minnesota. Known copper-nickel deposits along the base of the complex are being reassessed for their PGE concentrations, and intrusions stratigraphically higher in the Duluth and Beaver Bay Complexes are being evaluated using new models for stratiform PGE mineralization within well-differentiated tholeiitic systems.

1 – 2 – 1. Copper-Nickel-(PGE) Sulfide Mineralization

The fundamental characteristics of this style of mineralization involve sulfur contamination of mafic magmas by pre-Keweenaw footwall rocks, and mineralization occurring in the vicinity of the basal contact of mafic intrusions. The variants of this style of mineralization that may potentially exist in the Duluth Complex include:

- a) Disseminated copper-nickel-(PGE) sulfide mineralization in basal contact zones of mafic intrusions
- b) Massive sulfide mineralization at the basal intrusive contact and in footwall rocks
- c) Sulfide mineralization in major feeder zones (Noril'sk/Voisey's Bay-type; Naldrett, 1997)

Only disseminated copper-nickel sulfide mineralization and basal massive sulfide mineralization are presently known to occur in the Duluth Complex.

a) Disseminated Sulfide Mineralization

Large resources of low-grade copper-nickel sulfide ore that locally contain anomalous PGE concentrations are well documented by drilling in the basal zones of the South Kawishiwi and Partridge River intrusions (Fig.9, Table.1). At least nine subeconomic deposits have been delineated in the basal 100 to 300 meters of both intrusions. The mineralization consists predominantly of disseminated sulfides that collectively constitute over 4.4 billion tons of material averaging 0.66 percent copper and 0.2 percent nickel (Listerud and Meineke, 1977). Overall, the copper to nickel ratio averages 3.3:1, ranging from 2.4:1 to 4.11:1. PGE concentrations average about 10 parts per million platinum + palladium (recalculated to 100 percent sulfide), but may range as high as 50 parts per million platinum + palladium (recalculated to 100 percent sulfide) in associated stratabound zones, such as the Birch Lake and Dunka Road deposits. Sulfur isotope analyses consistently indicate that the source of sulfur was from the pelitic country-rocks of the Virginia Formation, which form much of the footwall to the Partridge River intrusion and locally to the South Kawishiwi intrusion.

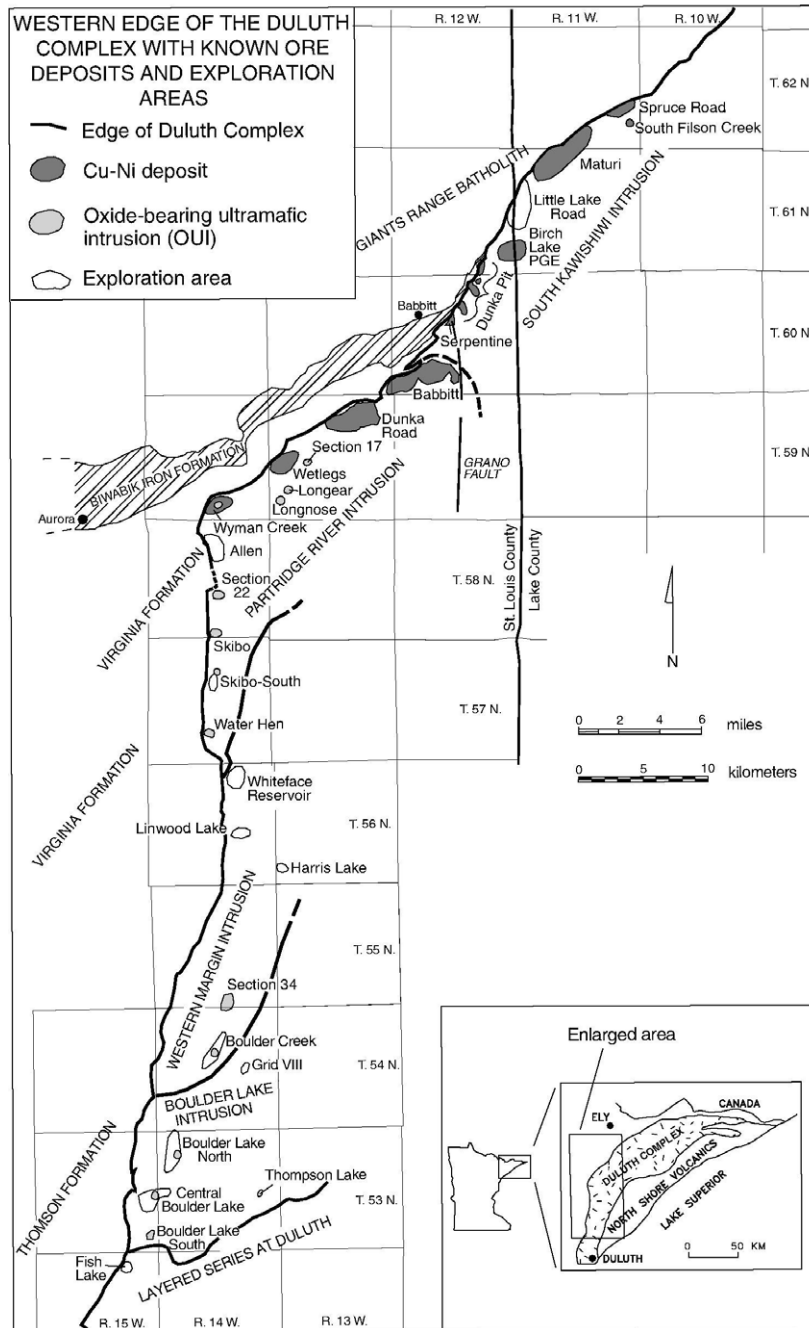


Fig.9 Location of the mineral deposits and exploration areas along the western base of the Duluth Complex(Miller at al., 2002).

Table 1 Calculated copper-nickel resources for various deposits along the western margin of the Duluth Complex(Miller et al., 2002).

(-) not available

Deposit (company)	Resources (in Mt)	Average Cu%	Average Ni%	Cu% cutoff	Cu/Ni ratio (x/1)
South Filson Creek (Hanna)	-	-	-	-	3.00
Spruce Road (INCO)	248	0.46	0.17	-	2.63
Maturi (INCO)	247 underground	0.50	0.19	-	2.40
Maturi Deep (Duval)	411	0.74	0.22	-	-
Dunka Pit (many)	272	0.25	-	-	-
Serpentine (Bear Cr., AMAX)	7	0.88	0.30	0.60	2.85
Babbitt/Minnamax (Bear Cr., AMAX)	364	0.84	0.19	0.60	4.11
Babbitt/Mesaba (Arimetco)	3300	0.46	0.12	0.38	-
Dunka Road (USS)	100 underground	0.77	0.24	-	3.43
Dunka Rd/NorthMet (PolyMet)	808	0.43	0.11	-	-
Wetlegs (Bear Creek, Exxon)	38	0.29	0.10	0.15	-
Wyman Creek (USS)	-	-	-	-	-
All deposits (Listerud and Meineke, 1977)	4400	0.66	0.20	0.50	3.30

The sulfide deposits are hosted by taxitic troctolitic to gabbroic rocks that contain abundant inclusions of the various footwall rock types. Within the Partridge River intrusion, the basal unit(Unit-I) hosts the vast majority of the disseminated sulfides. Similarly, mineralization within the South Kawishiwi intrusion is confined to the basal heterogeneous (BH), ultramafic 3 (U3), basal augite troctolite/norite (BAN), and updip wedge (UW) units.

The disseminated sulfide minerals occur as interstitial grains that make up between trace amounts and 10 percent of the rock by volume (visual estimation). The average sulfide mineral content is between 1 and 5 percent. Major sulfide minerals are pyrrhotite, chalcopyrite, cubanite, and pentlandite. Pyrrhotite is generally the dominant sulfide, especially closer to the basal contact. Chalcopyrite is generally the dominant copper-sulfide with variable amounts of cubanite. Also

present are minor amounts of bornite, talnakhite, chalcocite, digenite, mackinawite, valleriite, violarite, native copper, and platinum group minerals.

Although this mineralization type is described as being present within the basal units of the South Kawishiwi and Partridge River intrusions, the basal zones do not contain sulfides in all areas. Mineralized zones are extremely erratic in their spatial extent and ore grades. Zones that are barren of sulfides commonly interfinger with mineralized zones in a random pattern. This erratic pattern of mineralization, in part, mirrors the lithologic heterogeneity of the basal units. The only exception to this random mineralization pattern is the Maturi deposit (and its downdip eastward extension), where the upper portion of the basal heterogeneous (BH) unit consistently exhibits copper values in excess of 1.0 percent that gradually decrease with depth toward the basal contact. The change to more consistent mineralization at Maturi may be related to a thinning of the basal heterogeneous unit in the area and thus the sulfides are more restricted to a specific horizon.

Although the style of mineralization in all of the deposits is dominated by disseminated copper-nickel sulfides, differences occur between the deposits in copper-nickel and PGE grade, thickness, and tonnage. Mineralization profiles from single drill holes within each of the deposits of the South Kawishiwi intrusion are presented in Fig. 10. Peterson (2001) attributed the higher-grade, sheetlike mineralization of the Maturi and Maturi Extension deposits to confined magma flow of the lower South Kawishiwi intrusion beneath a large pillar of older anorthositic series rocks.

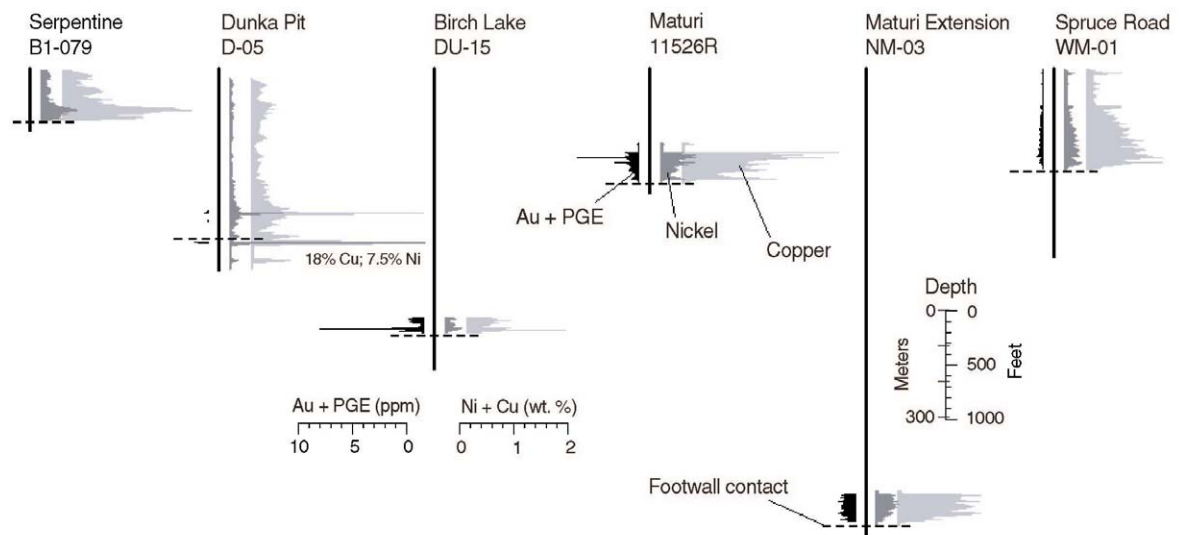


Fig.10 Assay from drill holes in each of the copper-nickel deposits of the South Kawishiwi Intrusion(Peterson, 2002).

Contrasting with this heterogeneity of rock types and mineralization, some internal PGE bearing sulfide zones within the lower units exhibit a stratabound relationship to the igneous stratigraphic section. For example, at the Dunka Road and Babbitt deposits, the top of Unit I immediately beneath an ultramafic layer commonly exhibits increased copper-PGE grades (Fig.11). These PGE-bearing zones are often copper enriched.

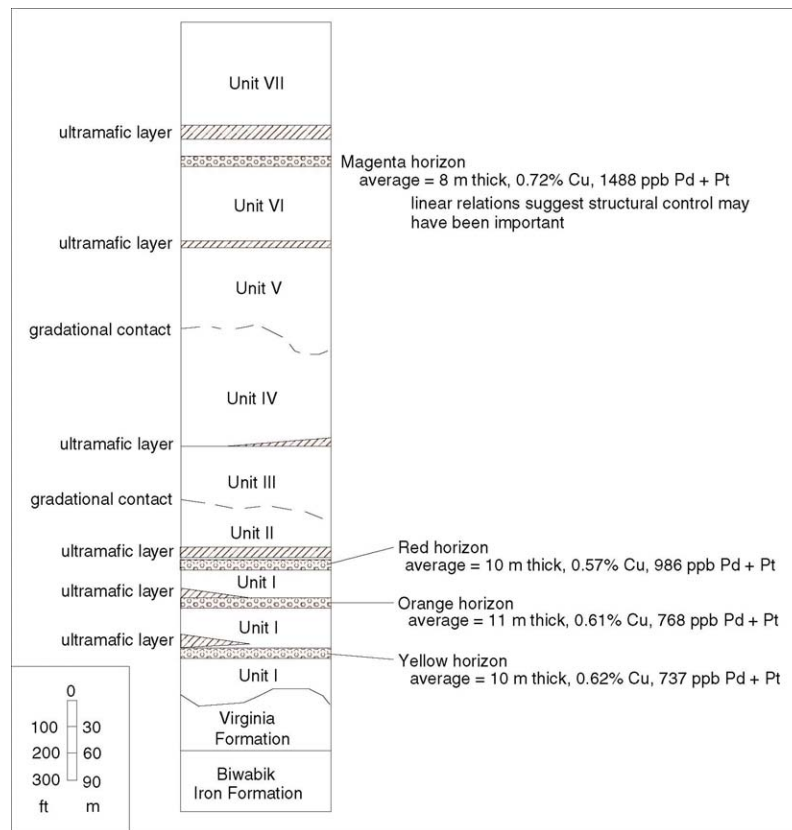


Fig.11 Distribution of PGE-bearing stratabound horizons relative to the igneous stratigraphy at the Dunka Road deposits(modified from Geerts, 1991). (Severson et al., 2002)

Two other stratabound copper-rich (\pm PGE) zones are present beneath discontinuous ultramafic horizons within the interior of Unit I at Dunka Road. At the Maturi deposit, a copper-enriched (and possibly PGE-enriched) zone also occurs immediately beneath an ultramafic unit (U3) at the top of the basal heterogeneous unit.

Copper-enrichment often occurs locally near fault zones as in the Wyman Creek, Wetlegs, and Dunka Road deposits close to the basal contact, and at the South Filson Creek deposit, 670 meters above the basal contact. The causes of copper-PGE enrichment along faults are not entirely clear,

but probably involve the remobilization and redeposition of copper and PGEs by circulating chlorine-rich hydrothermal fluids, possibly during the latest stages of magmatic crystallization. Evidence for such fluids is given by the presence of chlorine-rich brown liquid drops that commonly coat the surfaces of drill core after exposure to air. These drops form by a deliquescent process and are found on core from the Fish Lake area north to the Spruce Road deposit and the Gunflint Trail area. Analyses of the drops indicate high chlorine content values up to 3,000 parts per million. The drops are most common on core from the ultramafic layers (variably serpentinized) and the oxide ultramafic intrusions (OUI). They are also locally associated with disseminated sulfide zones, massive sulfides at the Local Boy ore zone of the Babbitt deposit, massive oxides and ultramafic layers from the U3 unit of the South Kawishiwi intrusion, pre-Duluth Complex sills (Logan-type sill and Cr-sill), and olivine-bearing portions of the metamorphosed footwall Biwabik Iron Formation. In addition to the liquid drops, chlorine-rich coatings on the core occur in a variety of colors and precipitate forms.

In summary, the disseminated sulfide mineralization at each of the copper-nickel deposits, with the exception of the Maturi deposit, can be classified as chaotic with diverse ore grades that are unevenly distributed throughout the basal mineralized units. This diverse nature makes it difficult to predict the overall spatial distribution of ore zones based on widely scattered drill holes and little to no outcrop. More in-fill drilling is needed in order to address ore controls. The cause of the erratic mineralization is probably related to a variety of factors that include:

- The nature of magma emplacement (turbulent versus quiescent)
- The mode of emplacement (along bedding planes in the footwall rocks versus overplating previous magma pulses)
- The volume of magmatic inputs (small incremental batches versus wide-spaced and large-volume batches)
- The total sulfide content of the magma prior to final emplacement
- The sulfide content of the footwall rocks (in-situ source)
- The secondary enrichment/depletion of copper and PGEs along faults

Magma pulses in the basal heterogeneous units appear to have formed dispersed sulfides that were unable to coalesce and form massive sulfides. Exceptions include pyrrhotite-rich massive sulfides at the Serpentine and Dunka Pit deposits; however, in these two cases a local sulfur source is also present in the footwall Virginia Formation. Most of the disseminated mineralization appears to have formed from a sulfide melt that interacted with low volumes of silicate magma, as indicated by low to moderate R-factors ranging from 50 to 3,000. The low R-factors are probably the result of close-spaced magma inputs of limited volume in progressively developing magma chambers accompanied by contamination from the footwall rocks. Conversely, the copper-PGE-enriched

zones beneath ultramafic horizons formed from a sulfide melt that apparently interacted with large volumes of silicate magma as indicated by high R-factors in the range of 5,000 to 17,000. The high R-factors of these zones are indicative of more turbulent magma conditions following a new magma input that crystallized the ultramafic layers. Overall, copper and PGE-enriched zones within the disseminated sulfide mineralization are associated with a dynamic changeover in the style of emplacement from close-spaced, low-volume, and contaminated pulses to wide-spaced, large-volume, and uncontaminated pulses. Some late hydrothermal secondary enrichment of copper and PGEs also took place adjacent to faults during later stages of emplacement.

b) Basal massive sulfide mineralization

In a few localized areas along the basal zones of the South Kawishiwi and Partridge River intrusions, semi-massive to massive sulfide mineralization is present at the basal contact. In most cases, the massive to semi-massive sulfide is proximal to either sulfide-rich footwall rocks or structures such as faults and pre-complex folds. Massive sulfide zones that are spatially related to sulfide-rich footwall rocks are intersected in scattered drill holes in the Dunka Pit, Babbitt, and Serpentine deposits. All of these massive sulfides are pyrrhotite-rich (with generally less than 2 percent copper) and are present at, or slightly above, the basal contact. In all cases, a pyrrhotite-rich member of the footwall Virginia Formation (bedded pyrrhotite unit) is located at the basal contact and is situated updip of the massive sulfide occurrences. This relationship suggests that the bedded pyrrhotite unit acted as a local sulfur source that generated a copper-poor, sulfide-rich melt that was concentrated downdip along the basal contact, via gravity settling.

The massive sulfide occurrence in the Local Boy ore zone of the Babbitt deposit is clearly structurally controlled. At this locality, the massive sulfide zones are copper-rich (generally 5 to 25 percent copper) and are situated along the axis of an anticline defined by the footwall rock units. The highest PGE values (11 parts per million palladium and 8 parts per million platinum) found to date within the Duluth Complex are associated with these structurally controlled, copper-rich, massive sulfides. The massive sulfides are almost exclusively hosted by the Virginia Formation, present as both inclusions above the basal contact and in the footwall rocks below the basal contact, while interfingering intrusive rocks are relatively barren of massive sulfide. These relationships, plus sulfide textures that are indicative of structural preparation, suggest that the massive sulfides were "injected" into the footwall rocks. Ripley (1986) and Severson and Barnes (1991) proposed that an immiscible sulfide melt, formed in an auxiliary magma chamber at depth, was injected into structurally prepared zones in the footwall rocks along the anticline to form the Local Boy ores. Late movement of chlorine-rich fluids along the axis of the anticline further redistributed and concentrated the PGEs. Recent studies indicated that there is an overall increase in the copper-PGE content of the massive sulfide in an east-to-west direction; this is possibly the result of fractional crystallization of immiscible sulfide melt as it migrated into the footwall rocks. In this scenario, the

north-south-trending Grano fault is inferred to be a potential feeder zone.

Structurally controlled veins and irregular pods of massive sulfide are locally present within granitic footwall rocks immediately beneath the South Kawishiwi intrusion. These occurrences are intersected in scattered holes that outline two northeast-trending belts (Fig.12). The linearity of the belts suggests they are fault controlled. One of these belts crudely aligns with the Birch Lake fault zone that trends through the Birch Lake PGE prospect. The massive sulfide veins were probably formed by the downward expulsion of a basal immiscible sulfide melt into fractured and faulted footwall rocks. The veins are moderately copper enriched due to fractional crystallization of the sulfide melt as it moved down through the footwall rocks.

The occurrence of local massive sulfide veins near and below the basal contact of the Duluth Complex is an indication that larger, potentially economic footwall massive sulfide deposits may yet be found.

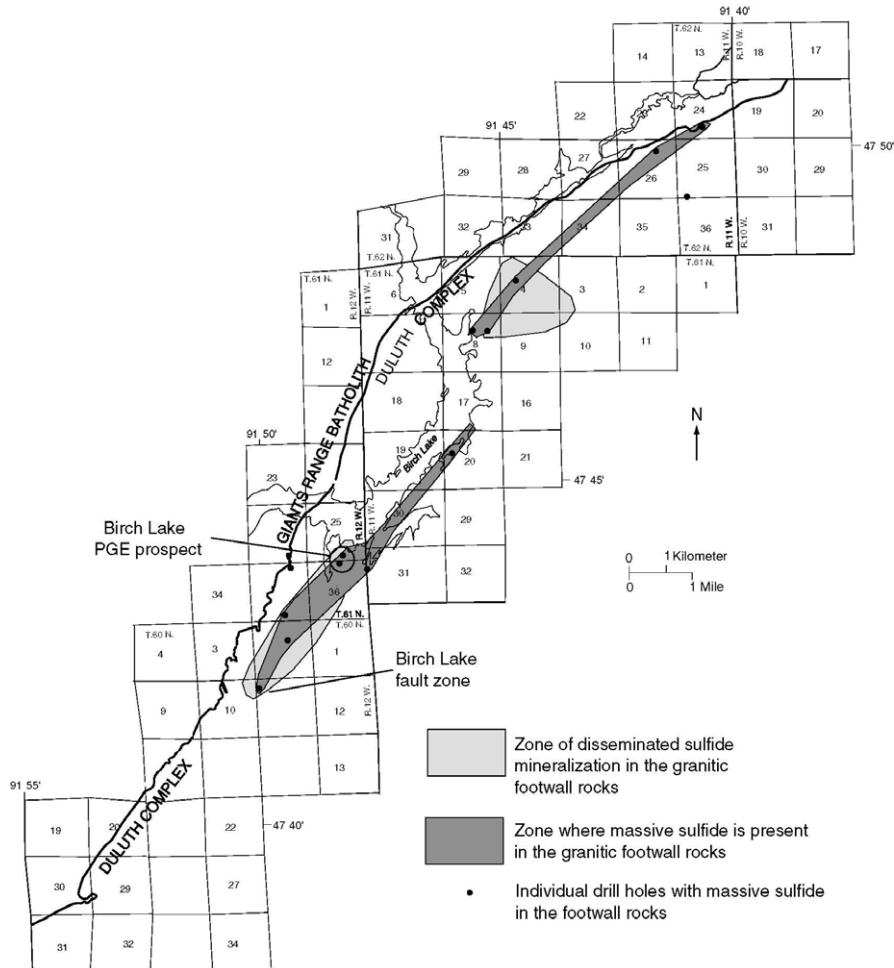


Fig.12 Linear distribution of massive sulfide, disseminated sulfide, and copper rich veins in the footwall rocks beneath the South Kawishiwi Intrusion (Severson et al., 2002).

c) Feeder Zone Sulfide Mineralization

Some of the attributes of the Duluth Complex copper-nickel-PGE sulfide deposits resemble those of deposits at Noril'sk, Russia and Voisey's Bay, Canada that are associated with sulfide mineralization in intrusive feeder zones. The common attributes include occurrence in shallow tholeiitic intrusions associated with plateau basalt volcanism, an external sedimentary source of sulfur, and openness to repeated magma influx and expulsion. A critical attribute of the high-grade Noril'sk, Talnakh and Voisey's Bay deposits, not yet positively identified in the Duluth Complex deposits, is the location of a magma conduit. A conduit that experienced repeated influxes of magma appears to be key to the formation of high grade copper-nickel-PGE deposits. Although this type of mineralization has not been recognized in the Duluth Complex, the conceptual model of magma flow has been established in the South Kiwishiwi and the Partridge River Intrusions and promising areas of the mineralization have been delineating.

One of the difficulties in evaluating the potential for feeder zone mineralization in the Duluth Complex is determining whether intrusions were fed one-by-one by local magma conduits or by master conduits that sequentially fed several intrusions.

1 – 2 – 2. Stratabound and Stratiform PGE Mineralization

PGE-enriched zones in the Duluth Complex and related intrusions can be classified into two categories, stratabound and stratiform. The term *stratabound* refers to PGE-enriched horizons that are restricted to a general stratigraphic unit; however, the PGE-enriched zones are often transgressive relative to the enclosing stratigraphy because they are associated with a wide variety of internal rock types that characterize the stratigraphic unit. The term *stratiform* denotes specific PGE-enriched horizons that are sheetlike in form, concordant, and strictly coextensive with laterally persistent igneous layers. The stratiform PGE horizons differ from the stratabound ones because they are consistently sulfide-poor and tend to occur at mid levels of well differentiated intrusions.

a) Stratabound PGE mineralization

Stratabound PGE-enriched horizons, with low to moderate sulfide concentrations (0.05 to 1.0 weight percent sulfur), are commonly associated with ultramafic layers in the Dunka Road, Babbitt, Wetlegs, and Birch Lake deposits, and occur at the extreme top of Unit I immediately beneath a laterally persistent ultramafic layer (Fig. 11). At Dunka Road, this stratabound horizon averages about 10-meters-thick and contains an average of 1.0 part per million palladium + platinum. Recent work suggested that the sulfur was largely derived from the mafic magma and that this stratabound horizon was formed as a result of magma mixing. Two similar stratabound PGE enriched horizons, related to laterally discontinuous ultramafic layers, occur toward the middle of Unit I at Dunka Road. To the west of Dunka Road, the stratabound PGE horizon at the top of Unit I is also present at the Wetlegs and Wyman Creek deposits. At the Dunka Road and the Wetlegs deposits, PGE-

enriched stratabound horizons also span across the ultramafic layer that separates Units VI and VII.

Another example of a PGE stratabound horizon is at the Birch Lake PGE prospect within the South Kawishiwi intrusion. There, PGE contents as high as 9 parts per million palladium + platinum, and Cr₂O₃ contents locally as high as 10 weight percent are associated with a wide variety of rock types within the U3 unit. The U3 unit consists of alternating troctolitic and ultramafic layers in which variably sulfide mineralized zones and discontinuous pods of Cr-bearing massive oxide both occur. As stated above, the massive oxide pods are interpreted based on empirical relationships to have been produced by assimilation and partial melting of the Biwabik Iron Formation. This oxide-rich partial melt may have initially acted as a trap that concentrated chromium and titanium, and through further assimilation and contamination of the magma, may have led to precipitation of PGEs. However, because the ultramafic layers of the U3 unit are interpreted to record new influxes of more primitive magma, magma mixing may have had a more significant effect on PGE mineralization. Furthermore, the presence of chlorine-rich drops on the surface of drill core from the Birch Lake area suggests that a hydrothermal model of concentrating the PGE could also be invoked. As depicted in Fig.13, ascending chlorine-rich hydrothermal fluids may have remobilized and further concentrated the PGE along the northeast trending Birch Lake fault.

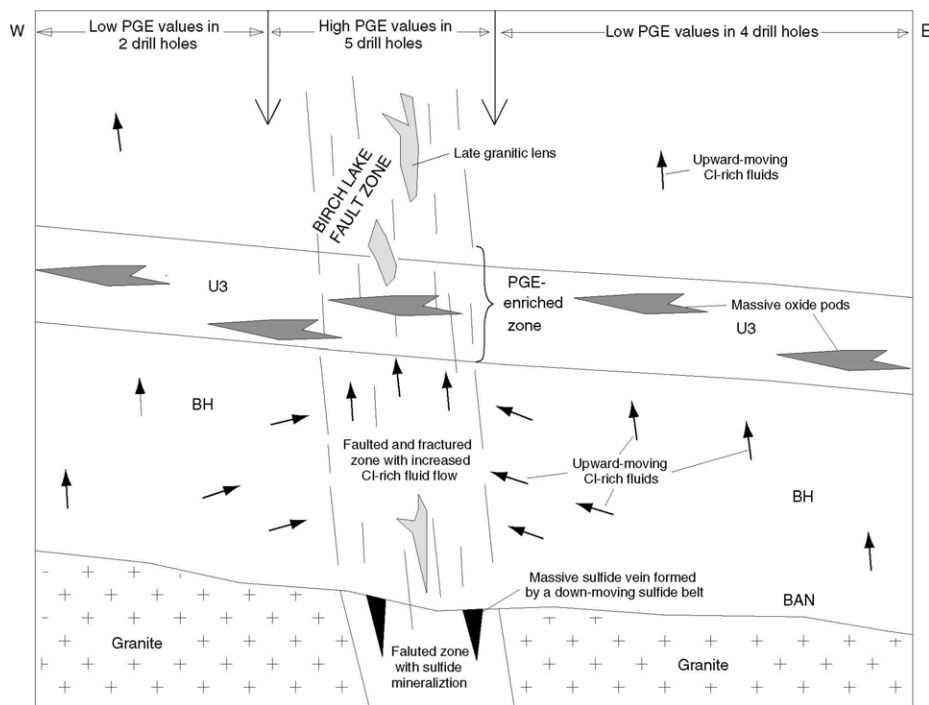
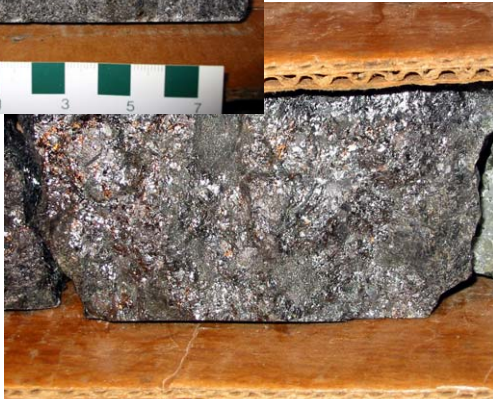


Fig.13 Schematic diagram showing the possible role of the Birch Lake fault for concentration PGE in U3 unit(Severson et al., 2002).



Fig.14A Photograph of Ni-Cu-PGE ore (disseminated ore).



Birch Lake area (D.H. BL-89-2): troctolite in the vicinity of 2595.5ft (2593-2597ft, Cu:0.71%, Ni:0.18%, Pd:1100ppb, Pt:520ppb)



Fig.14B Photograph of Ni-Cu-PGE ore (stratabound mineralization?).

Babbitt deposit (D.H. BI-120): melatroctolite in the vicinity of 1307ft (the top of Unit I), with native copper (1305-1315ft, Cu:0.72%, Ni:0.16%, Pd:1728ppb, Pt:837ppb).

Fig.14C Photograph of Ni-Cu-PGE ore(massive ore).

Babbitt deposit (D.H.10198): mixed norite and hornfels with massive sulfide in the vicinity of 21ft (20-25ft, Cu:14%, Ni:0.28%, Pd:1700ppb, Pt:260ppb, Au:250ppb).

b) Stratiform PGE mineralization

Recent discoveries of gold- and PGE-enriched horizons (or PGE reefs) in the upper part of the Skaergaard intrusion and related Tertiary intrusions of east Greenland have stimulated exploration for such deposits in other tholeiitic layered intrusions such as those comprising the Duluth Complex. Stratiform PGE mineralization in well-differentiated tholeiitic intrusions are similar to classic PGE reef deposits hosted by ultramafic-mafic complexes, such as the Bushveld and Stillwater Complexes, in that they occur as sulfide-poor (less than 1 weight percent), PGE-rich intervals that are several meters thick and are conformable with igneous layering. However, stratiform PGE mineralization in tholeiitic intrusions, termed Skaergaard-type PGE mineralization, differs from the classic PGE reefs because it is:

- Exclusively associated with mantle plume influenced, continental rift environments
- Of Middle Proterozoic age or younger
- Associated with aluminous, olivine tholeiitic parent magma compositions that experience Fenner-type crystallization differentiation
- Hosted by ferrogabbroic cumulate rocks
- Associated with copper-rich, nickel-poor sulfide
- Associated with significant gold that is stratigraphically offset above peak PGE concentrations

Although there is considerable disagreement about how classic PGE reefs formed, most believe that Skaergaard-type reefs are orthomagmatic (formation by the saturation, exsolution, and settling of sulfide melt from silicate magma).

The key requirements for the orthomagmatic formation of an economic PGE reef in a tholeiitic intrusion are:

- The parent magma must be initially sulfide undersaturated
- The parent magma must have a high initial PGE concentration and/or experience a considerable amount of fractional crystallization to build up noble metal concentrations prior to sulfide saturation
- The initial segregation of sulfide melt must be associated with a large R-factor (silicate/sulfide melt ratio)

These required conditions imply that the best chance for significant stratiform PGE mineralization in a tholeiitic intrusion should be at a mid- to upper-level horizon that marks the first "cumulus arrival" of immiscible sulfide melt. Given a well-differentiated, initially sulfide undersaturated tholeiitic intrusion, its ability to form an economic PGE reef depends on the manner by which sulfide becomes saturated in the magma. Mechanisms that may trigger sulfide saturation include:

Fractional crystallization - Because sulfur is an incompatible element in silicate minerals, sulfide will become enriched in differentiating magmas ultimately to the point of saturation and exsolution.

Magma recharge and mixing - Unlike variable sulfide solubility in ultramafic-mafic magmas that produce classic PGE reefs, sulfide solubility in differentiating tholeiitic magmas remains fairly constant over the course of fractional crystallization due to the offsetting effects of temperature and Fe enrichment. Therefore, mixing of magmas with comparable sulfide solubilities is unlikely to produce sulfide saturation or oversaturation.

Decompression due to magma venting - At shallow crustal depths, sulfide solubility appears to positively correlate with pressure. Consequently, in a magma that is close to sulfide saturation, decompression due to magma venting can potentially trigger system wide exsolution of sulfide melt.

Phase changes - The cumulus arrival of a mineral phase that abruptly lowers the iron content of the magma can potentially trigger sulfide saturation. Magnetite crystallization would be the most effective phase change. This was the apparent trigger for sulfide saturation in the Rincon del Tigre magma system of Bolivia (Prendergast, 2000).

Most intrusions of the Duluth and Beaver Bay Complexes formed from initially sulfide undersaturated magmas and show evidence of having undergone one or more of the above mentioned processes during crystallization. Two of these intrusions, the Sonju Lake intrusion and the layered series at Duluth, have been studied for their potential to host stratiform PGE mineralization, and subeconomic mineralization has been found in both areas. The Sonju Lake intrusion was a closed magmatic system and the layered series at Duluth was a more open magmatic system.

Stratiform PGE mineralization in the Sonju Lake intrusion

The Sonju Lake intrusion is a 1,200-meter-thick, shallow-dipping, sheetlike intrusion that forms part of the Beaver Bay Complex (Fig.3). Although its exposed strike-length is only about 3 kilometers, the Sonju Lake intrusion has a distinctive aeromagnetic signature that can be traced for at least 20 kilometers. The Sonju Lake intrusion is predominantly composed of well-laminated, mafic mesocumulates that define a unidirectionally differentiated sequence. Between a lower contact zone of fine grained melatroctolite cumulates (unit slmt - OP cumulate) and an upper zone of unlaminated olivine ferromonzodiorite (unit slmd), five units are distinguished on the basis of cumulus mineral assemblages. In ascending order these units are: sld - dunite (O), slt - troctolite (PO), slg - gabbro (POC), slfg - ferrogabbro (PCF ± O), and slad - apatite olivine ferrodiorite (PCFOA). This cumulus stratigraphy is complemented by a smooth cryptic variation in mineral compositions. The parent magma composition of the Sonju Lake intrusion is estimated to be an uncontaminated, moderately evolved olivine tholeiite. Cumulus stratigraphy and cryptic variation

suggests that the Sonju Lake intrusion formed by unidirectional fractional crystallization under nearly closed conditions.

To determine the possible location of a stratiform PGE reef in the Sonju Lake intrusion, a total of 67 outcrop and drill core samples that profile the intrusion were analyzed for platinum, palladium, gold, sulfur, copper, and other major and trace elements (Fig. 15). These data indicate that a narrow PGE-enriched interval exists within the ferrogabbro unit and lies immediately below the stratigraphic level marking the initial saturation of sulfide melt in the Sonju Lake intrusion magma. Initial saturation and segregation of sulfide melt is inferred to occur at the +322-meter level, where an increase in copper abundance is recognized. Copper content increases from consistently less than 100 parts per million below the +322-meter level to greater than 500 parts per million above, and then gradually decreases to below 100 parts per million in the upper part of the intrusion. This pattern is consistent with:

- Sulfide undersaturated conditions below the +322-meter level
- The saturation and cumulus segregation of sulfide melt at the +322-meter level
- The gradual depletion of copper abundance due to continued exsolution of sulfide melt above the +322-meter level

Nickel was strongly depleted in the magma at the time of inferred sulfide saturation, presumably by prolonged olivine crystallization.

Interestingly, the onset of sulfide saturation at the +322-meter level implied by the abrupt increase in copper abundance is not reflected in the sulfur data. Theoretically, a gabbro cumulate saturated in sulfide melt should contain a minimum of 400 parts per million sulfur based on sulfide solubility of 0.1 weight percent for a moderately evolved mafic magma. The sulfur abundance in cumulates below the +322 meter level are consistently less than the theoretical saturation benchmark. However, sulfur concentrations above the +322 meter level vary unsystematically between less than 100 and 500 parts per million, and only sustain greater than 400 parts per million sulfur concentrations above about +460 meters. Petrographic observations confirm that this zone of presumed sulfide saturation is locally poor in modal sulfide and where present, the sulfide is dominated by copper rich varieties (chalcopyrite, bornite, and digenite). A possible explanation for the high copper, low sulfur samples in the zone between +322 meters and +460 meters is that oxidizing deuteric fluids caused dissolution of sulfur without mobilizing copper. A similar oxidation/desulfurization process has been noted in the low-sulfide PGE reefs in eastern Greenland intrusions. The Sonju Lake intrusion displays several features that are consistent with such a process of oxidation and desulfurization:

- Moderate degrees of chloritic and uralitic alteration of pyroxene are common in ferrogabbroic cumulates of the ferrogabbro unit
- Copper-rich sulfides (chalcopyrite, bornite, and digenite) predominate in samples near the +322 meter level

- Secondary veins of pyrite with hematite occur locally above the +322-meter level
- If deuteric desulfurization is a widespread phenomenon in mafic layered intrusions, variations in copper abundance rather than total sulfur may be better indicators of magmatic sulfide saturation.

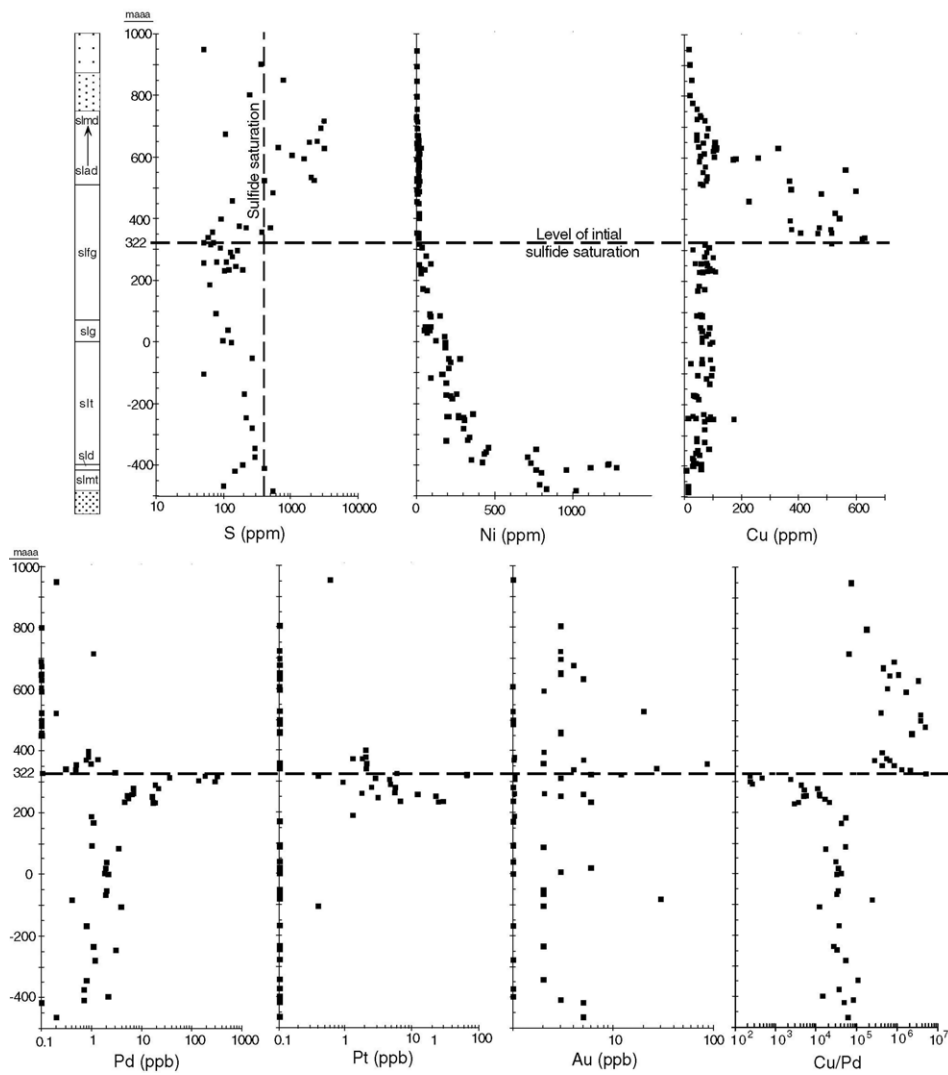


Fig.15 Chemostratigraphic variation in whole rock concentrations of S, Ni, Cu, Pd, Pt, Au, and Cu/Pd through the Sonju Lake intrusion (Severson et al., 2002).

There are no overt changes in lithology evident from field observations across the horizon of increased copper abundance (and PGE enrichment) that would suggest sulfide saturation was triggered by a significant perturbation to the magma system (such as magma recharge, assimilation, or venting). The only notable change observed across the horizon of initial sulfide saturation is the gradual changeover from titanomagnetite-dominated to ilmenite-dominated cumulus oxide phases. However, this also seems an unlikely triggering mechanism because such a change should have minimal effect on the FeO content of the magma and thus sulfide solubility. Rather, given the well-differentiated and closed nature of the Sonju Lake intrusion system, it is concluded that simple crystallization differentiation was the process primarily responsible for not only driving the magma to sulfide saturation, but also for triggering the initial sulfide exsolution event.

Upward from about the +230-meter level, palladium abundance increases regularly to a peak concentration of 320 parts per billion in the sample just below the first known sample to display an abrupt jump in copper to greater than 500 parts per million. This high palladium sample also yields the greatest platinum concentration (66 parts per billion), although a secondary platinum peak (less than 30 parts per billion) seems to exist about 90 meters below the sulfide saturation level. The highest gold concentration (85 parts per billion) occurs 35 meters above the level of initial sulfur saturation, although its variation is less systematic than palladium and platinum. In the first high concentration copper sample (greater than 500 parts per billion) and in most overlying samples, palladium and platinum abundances drop to below detection limits (less than 0.1 part per billion).

The efficiency with which PGEs were scavenged from the Sonju Lake intrusion magma with the "cumulus arrival" of sulfide melt is indicated by the 10^4 increase in copper/palladium observed across the +322-meter horizon. Changes in the copper/palladium ratio reflect the fact that the sulfide melt/silicate magma partition coefficient for palladium is several orders-of magnitude greater than that for copper.

When directly comparing the Sonju Lake and Skaergaard intrusions, not only are their petrologic similarities evident, but also when comparably scaled, their patterns of precious metal and sulfide mineralization are remarkably similar (Fig. 16). Their similar paragenetic sequences of cumulus minerals imply comparable closed-system crystallization differentiation of aluminous, olivine tholeiitic parent magmas. Both parent magmas were initially sulfide-undersaturated and became saturated at proportionally similar levels in their crystallization histories. Both intrusions were crystallizing a cumulus mineral assemblage of plagioclase + augite + iron oxide at the time of sulfide saturation. Both show peak PGE concentrations at or just below the level of initial sulfide saturation. Both show low and variable levels of palladium in the lower sulfide undersaturated part of the intrusions and almost complete and persistent palladium depletion above the sulfide saturation horizon. Finally, both show an upward displacement of peak gold concentration from the PGE peak.

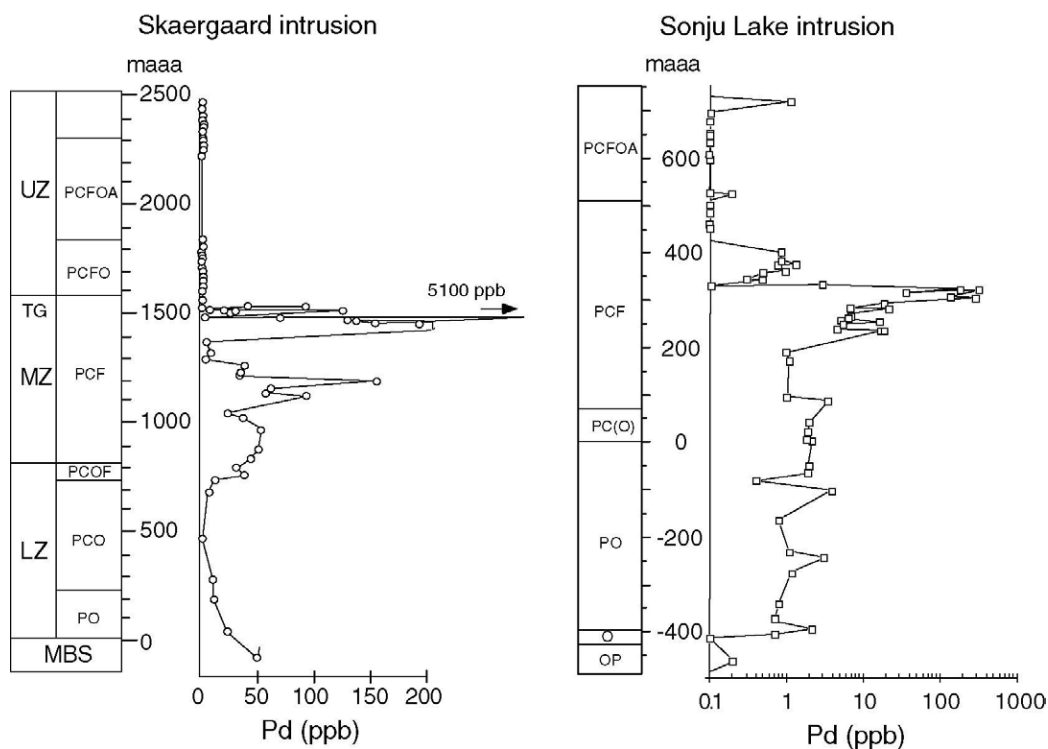


Fig.16 Comparison of generalized igneous stratigraphy and stratigraphic variation of Pd concentrations in the Skaergaard and Sonju Lake intrusions(Severson et al., 2002).

Stratiform PGE mineralization in the layered series at Duluth

The layered series at Duluth is a 3.5 to 5kilometer-thick, sheetlike, mafic layered intrusion, and the igneous stratigraphy of the layered series at Duluth is divided into five major zones. The repeated progression from troctolitic to gabbroic cumulates in the cyclic zone indicates that the layered series at Duluth, unlike the Sonju Lake intrusion, did not crystallize as a closed system. Based on these characteristics, Miller and Ripley (1996) suggested that the macrocyclic phase layering is predominantly related to devolatilization and decompression that attended magma venting events from a shallow (less than 5 kilometers) chamber, with magma recharge possibly having a secondary affect.

Geochemical data were acquired from 83 hand samples that profile the general stratigraphy of the layered series at Duluth. Most of the samples have platinum + palladium concentrations below 15 parts per billion, 14 samples have concentrations between 15 and 150 parts per billion, and five samples have concentrations greater than 150 parts per billion(Fig.17). Whole rock sulfur and copper concentrations imply that the layered series at Duluth magma was consistently sulfide

undersaturated during troctolite-zone crystallization, achieved intermittent saturation during crystallization of the cyclic zone, and was fairly consistently saturated as the gabbro zone accumulated. In contrast to the singular large increase in copper abundance and copper/palladium ratio observed in the Sonju Lake intrusion system (Fig.15), the layered series at Duluth shows erratic variability in these parameters (Fig.17). Such variability probably reflects the openness of the layered series at Duluth system to magmatic recharge.

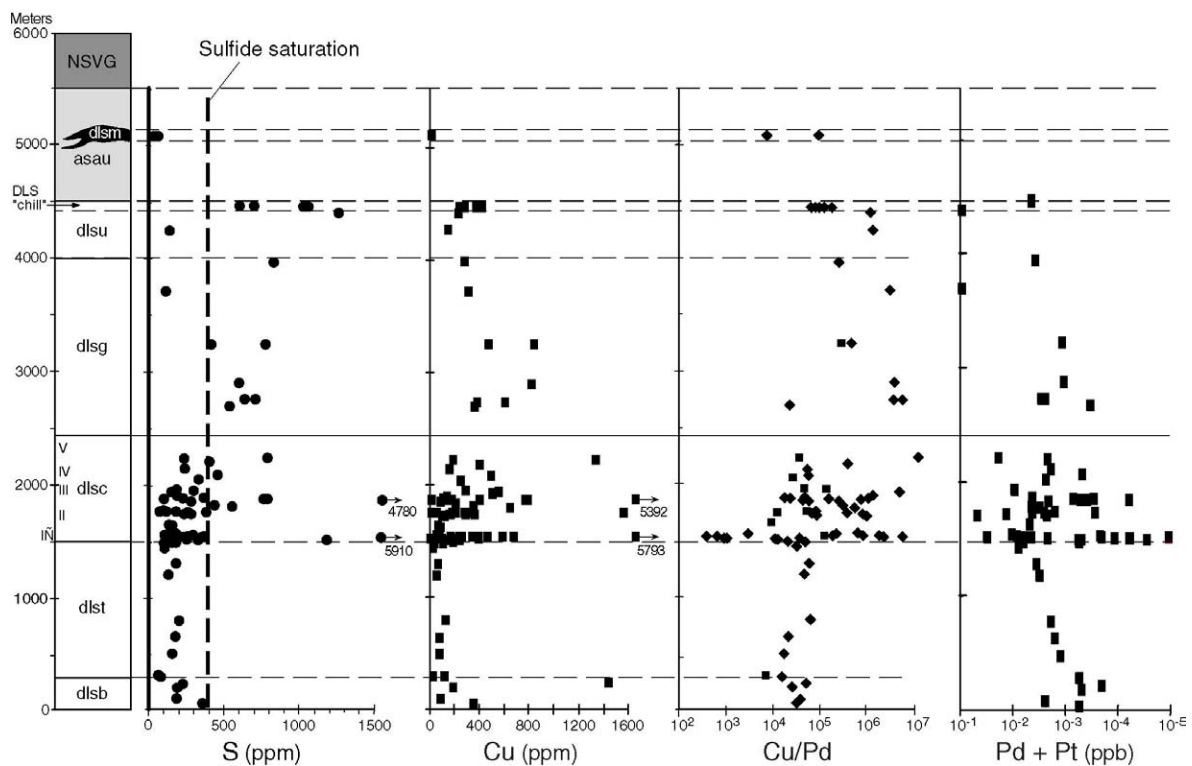


Fig.17 Stratigraphic variation in S, Cu, Cu/Pd, and Cu+Pd abundances through the Layered Series of Duluth(Severson et al., 2002).

1 – 3. Mineral Exploration in the Duluth Complex

1 – 3 – 1. Copper-Nickel Exploration activity at the Base of the Duluth Complex

Exploration for copper-nickel sulfide deposits along the base of the Duluth Complex began in 1948 and has continued intermittently. Between 1951 and 2000, at least 22 companies conducted drilling campaigns in 28 areas. During this period, over 1,700 holes totaling over 1.4 million feet (427 kilometers) of core have been drilled in the basal zone of the complex. At least ten copper nickel deposits, portrayed on Fig.9

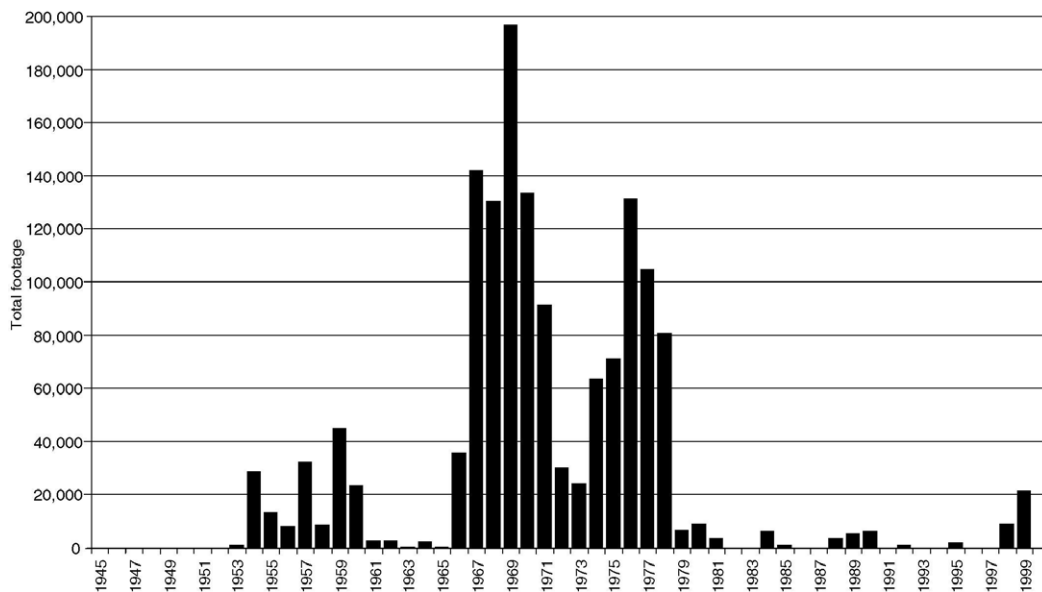


Fig.18 Histogram of annual exploration drilling footage along the basal contact of the Duluth Complex(Miller et al., 2002).

Outcrops that contain disseminated sulfide minerals were first noted at the base of the Duluth Complex by Grant(1899). However these occurrences were too sporadic and weakly mineralized to spark any exploration interest. More strongly mineralized rocks with copper stains were discovered in 1948 by prospector F.S. Childers at the area of present Spruce Road deposit. In 1951, a 188-foot (57-meter) exploration drill by Childers and partner R.V. Whiteside, situated within what is now referred to as the Maturi deposit, intersected disseminated sulfides in gabbroic rock that averaged 0.36 percent copper and 0.13 percent nickel.

This discovery sparked the exploration activity. In 1952, INCO and Bear Creek Mining(a

subsidiary of Kennecott Mining Company) conducted reconnaissance geological and geophysical surveys, which were beginning of the exploration activities by major mining companies. The drilling campaign commenced in 1954 is divided into two phases, the first phase(1954-1960), and the second phase(1966-1981).

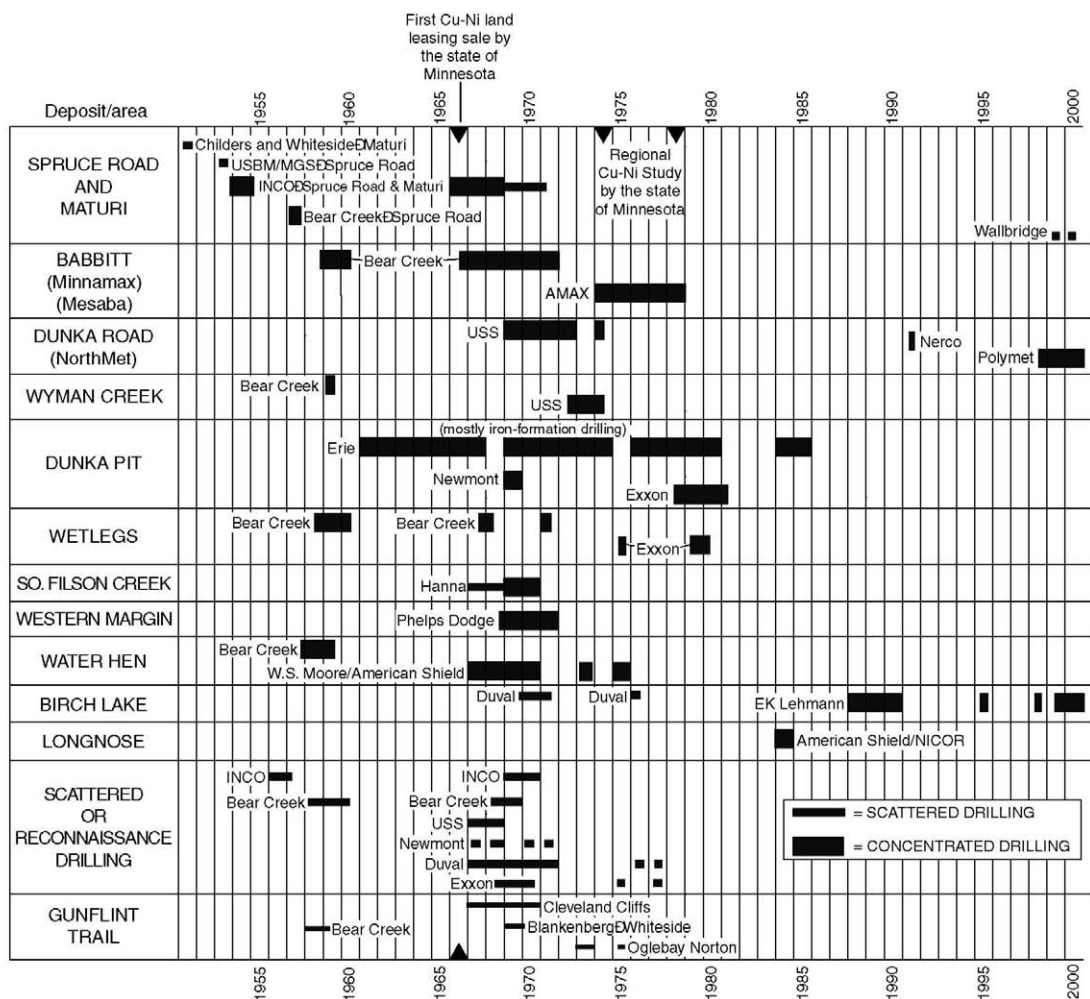


Fig.19 Chronology of drilling activities between 1950 and 2000 along the basal contact of the Duluth complex(Miller et al., 2002).

The first phase

INCO conducted drillings during 1954 and 1955 at the Spruce Road and Maturi deposits. A potentially large tonnage of low grade copper-nickel ore was indicated by this work. During 1956 and 1957, on the basis of geophysical surveys, INCO drilled reconnaissance holes into other areas.

However the results from these areas were disappointing.

Between 1957 and 1960, Bear Creek Mining conducted an extensive drilling program that targeted scattered geophysical anomalies, and the most interesting holes were located at the Babbitt Lake deposit.

After 1960, all drilling activities were suspended due to the low grade nature of the ore, coupled with low metal price, and unavailability of state-owned mineral lands.

The second phase

In 1966, regulations covering mining leases for copper and nickel on state lands were established. Leases were awarded to many companies and the second phase drilling activities began. Between 1967 and 1978, extensive drilling programs were conducted along the base of the Duluth Complex.

INCO completed a major drilling effort at its Spruce Road and Maturi deposits, and sunk a shaft at the Maturi deposit in 1966 and 1967. They decided that the mineralization at the Maturi deposit was too low grade to support an underground mine and concentrated their efforts at the Spruce Road deposit. In 1974, INCO announced their plan to establish an open-pit mine at the site.

Bear Creek Mining concentrated most of their effort at the Babbitt and Serpentine deposits between 1967 and 1971. After the activity of Bear Creek Mining, Amax renamed the deposit, "the Minnamax deposit", and conducted an extensive exploration program at the deposit that consisted of 228 surface drill holes (from 1974 to 1978), sinking a shaft with four drifts into massive sulfide mineralization near the basal contact at the Local Boy ore zone, and drilling 219 underground holes at Local Boy.

Between 1974 and 1978, in response to environment concerns regarding the impact of sulfide mining, "The Regional Copper and Nickel Study" was conducted by the state. Coincident with the study, the leasing of state lands was suspended between 1974 and 1981. The impact of the study on exploration activity in the Duluth Complex was significant. Development of both the Spruce Road and Babbitt deposits was delayed pending results of the study. By the end of the study in 1978, copper and nickel markets had weakened. As a result AMAX relinquished their land position at the Babbitt deposit and the Minnamax shaft was closed in 1981. INCO continues to maintain their leases, but has not pursued any further developments.

1 – 3 – 2 . PGEs exploration activity

During the early phases of drilling (prior to about 1980), all of the exploration companies recognized that the copper-nickel deposits had some potential for hosting precious metals, especially platinum group elements (PGEs). However, PGEs were rarely analyzed for in the early exploration period because PGE assays were too, and the assays were often inaccurate and not replicable. Based on limited PGE sampling, the companies assumed that the typical copper-nickel ore contained no more than a few hundred parts per billion combined platinum and palladium. In

1985, the Minnesota Department of Natural Resources and University of Minnesota (Minerals Resource Research Center [MRRC]) discovered significant values of about 9 parts per million combined platinum + palladium in the chromium bearing drill core from the Birch Lake area. A short time later, Morton and Hauck (1987) compiled all of the unpublished company PGE data for the Duluth Complex and reported the presence of significant PGE values, often associated with high copper contents. These discoveries sparked renewed interest in the copper-nickel deposits of the Duluth Complex as potential polymetallic deposits.

At least four companies are currently evaluating the polymetallic potential of five areas along the basal contact. PolyMet Mining Corporation has acquired the Dunka Road deposit and renamed it the NorthMet deposit. PolyMet has recently drilled over 90 holes (from 1998 to 2001) and is conducting pre-feasibility studies. They plan on using a hydrometallurgical process to treat the copper-nickel-cobalt-PGE-gold sulfide ore. Cominco American Incorporated has acquired the Babbitt deposit and has renamed it the Mesaba deposit. They also plan on using a hydrometallurgical process. Between 1988 and 2000, Lehmann drilled 20 surface holes to evaluate the PGE potential of the Birch Lake area at depths of 2,000 to 2,600 feet (610 to 792 meters) below the surface. In 1997, INCO began to renew their interest in the Maturi and Spruce Road deposits following the discovery of high-grade copper-PGE massive sulfide veins in the footwall rocks beneath the Sudbury Igneous Complex. In 1999, INCO's U.S. subsidiary, American Copper and Nickel Company, joint-ventured their properties with Wallbridge Mining Company Limited. Wallbridge has recently drilled two holes at Spruce Road to evaluate the high-grade vein potential, and one hole at Maturi to evaluate the PGE potential associated with the disseminated copper-nickel mineralization.

In addition to the recent activity along the basal contact of the Duluth Complex stated above, companies are showing interest in the possibilities of discovering PGE reef deposits at higher stratigraphic levels. Miller (1999) proposed that several intrusions within the Duluth and Beaver Bay Complexes appear to have petrologic attributes favorable for reef formation. Exploration companies that held mineral leases in the interior of the Duluth Complex and in the Beaver Bay Complex in 2001 include Falconbridge Limited, McVicar Minerals Limited, Polymet Mining Incorporated, and E.K. Lehmann and Associates.

APPENDIX 3

Report of
the 9th
International Platinum Symposium

*- Geology and PGE mineralization of
the Lac de Iles mine and others
in Ontario, CANADA -*

Lac des Iles

The mafic to ultramafic Archean Lac des Iles Intrusive Complex (2692Ma) (LDI-IC) is located approximately 85 kilometres north of the City of Thunder Bay in northwestern Ontario, CANADA. The LDI-IC is host to numerous zone of platinum group element (PGM) mineralization, including the Roby Zone (open-pit). The Roby Zone has been in commercial production since 1993 and as of December 31, 2001 has produced 629,272 ounces of palladium, 44,086 ounces of platinum, 42,073 ounces of gold, 11,139,923 pounds of copper and 7,765,647 pounds of nickel. As of December 31, 2001 measured and indicated resources for the 950 metres long by 815 metres wide Roby Zone deposit total 159 million tonnes grading 1.55g/t Pd, 0.17g/t Pt, 0.12g/t Au, 0.05 % Cu and 0.05 % Ni at a palladium cut-off grade of 0.7g/t.

Lac des Iles Intrusive Complex (LDI-IC) (Fig.1)

The LDI-IC is the largest of a series of mafic to ultramafic intrusions defining a circular pattern in the Lac des Iles area that is approximately 30 kilometres in diameters. The LDI-IC consists of three chambers, each a distinct lithological domain. Ultramafic intrusions (North Lac des Iles Intrusion) are centered on Lac des Iles, chaotic gabbroic intrusions (Mine Block Intrusion) occur immediately south of Lac des Iles, and a relatively homogeneous hornblende gabbro (Camp Lake Intrusion) is located southwest of Camp Lake.

North Lac des Iles Intrusion (NLDI-I)

The NLDI-I can be sub-divided into four domains based on lithological variations and structure.

- (1) massive to broadly layered, east-trending clinopyroxenite, with lesser websterite and minor olivine bearing units underlying the extreme north-northwest portion of The NLDI-I
- (2) massive to locally well layered north trending clinopyroxenite, websterite, thin olivine rich units and underlying gabbro in the northeastern portion of of The NLDI-I
- (3) east-trending massive gabbro, gabbronorite, variotextured gabbro/gabbronorite, and heterolithic gabbro/gabbronorite breccia interfingering with massive clinopyroxenite and minor websterite and olivine-bearing units in the south central portion of of The NLDI-I
- (4) clinopyroxenite with less abundant olivine-bearing units, websterite and gabbro underlying the southern portion of of The NLDI-I

Laterally, the most continuous unit is a 50m thick gabbro that can be traced over 2 kilometres on a northerly strike in the northeastern part of The NLDI-I. The most

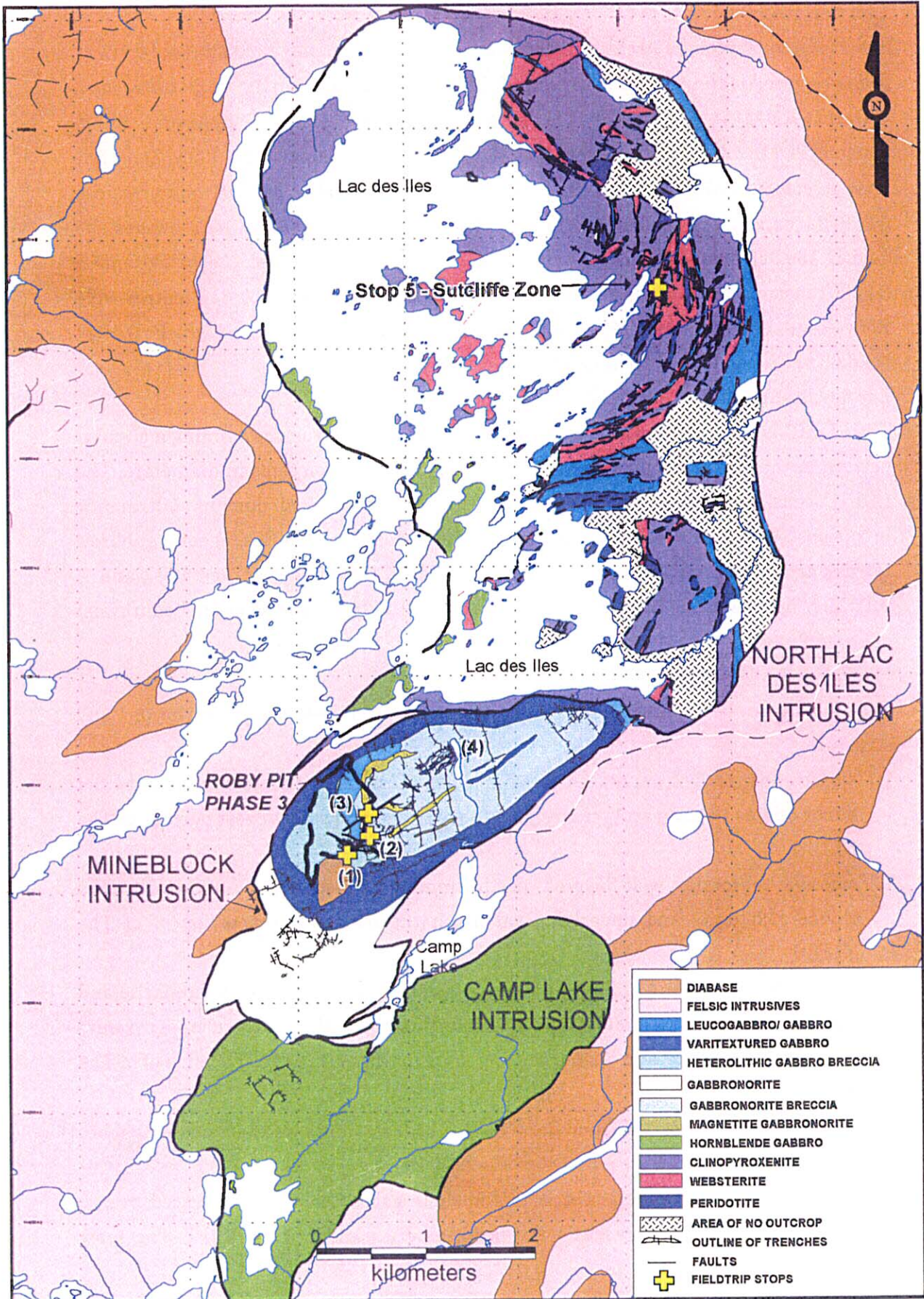


Fig. 1 Geology of the Lac des Iles Complex

common style of mineralization in the NLDI-I consists of disseminated chalcopyrite containing up to 1.0g/t Pd+Pt.

Mine Block Intrusion(MBI) (Fig.2)

MBI is texturally and compositionally complex. Its composition ranges from anorthosite to clinopyroxenite, leuco-gabbro to melanonorite and includes magnetite-rich gabbro. Textures include equigranular, fine-to coarse-grained, porphyritic, and pegmatitic, varitextured units and heterolithic gabbro breccia. These last three textural types are the most common host to PGE mineralization, including the Roby Zone. The most laterally continuous unit is a massive medium-grained gabbro, referred to as East Gabbro. It is adjacent to a varitextured gabbro "rim" to the west and more equigranular gabbro to the east. The varitextured rim is host to the Roby Zone palladium deposit where heterolithic gabbro breccia is common, occurring as pipes and pods, and large blocks (~60m) of varying composition.

Camp Lake Intrusion

Camp Lake Intrusion is dominantly a medium-grained hornblende gabbro, with lesser leucocratic, melanocratic, porphyritic, varitextured, and monolithic breccia. Minor PGE mineralization occurs along the northern edge of this intrusion.

Roby Zone (Fig.3,4,5,6)

The Roby Zone is a bulk mineable deposit with a minimum north to south length of 950 metres, a width of 815 metres and has been intersected at a depth of 1,000 metres by core drilling. The Roby Zone is comprised of three distinct ore types.

High Grade Ore(7.6% of volume), North Roby Ore(5.3%) and Breccia Ore(87.1%)

High Grade Ore

The High Grade Ore is mostly hosted by a portion of a 15 to 25 metres thick unit of pyroxenite/melanogabbro. It is located in the east-central portion of the Roby Zone, bounded by the PGE-barren East Gabbro hanging wall and Breccia Ore to the west. The High Grade Ore is confined to a 400-metres long segment of the pyroxenite, while the on strike extensions are barren to low grade.

Sulfide mineralization within the pyroxenite consists of 0.25% to 3.0% fine-grained disseminated sulfides with local net textured patches with up to 10% sulfide. The dominant sulfides are pyrite, pyrrhotite, chalcopyrite and pentlandite. Palladium and platinum mineralization within the High Grade Ore consists of predominantly fine-grained PGE sulfide, braggite and telluride minerals, merenskyite and kotulskite, which are present interstitially to cumulus grains and as inclusions within secondary silicates. The mean grade of the High Grade Ore is 7.89g/t Pd, with a maximum of

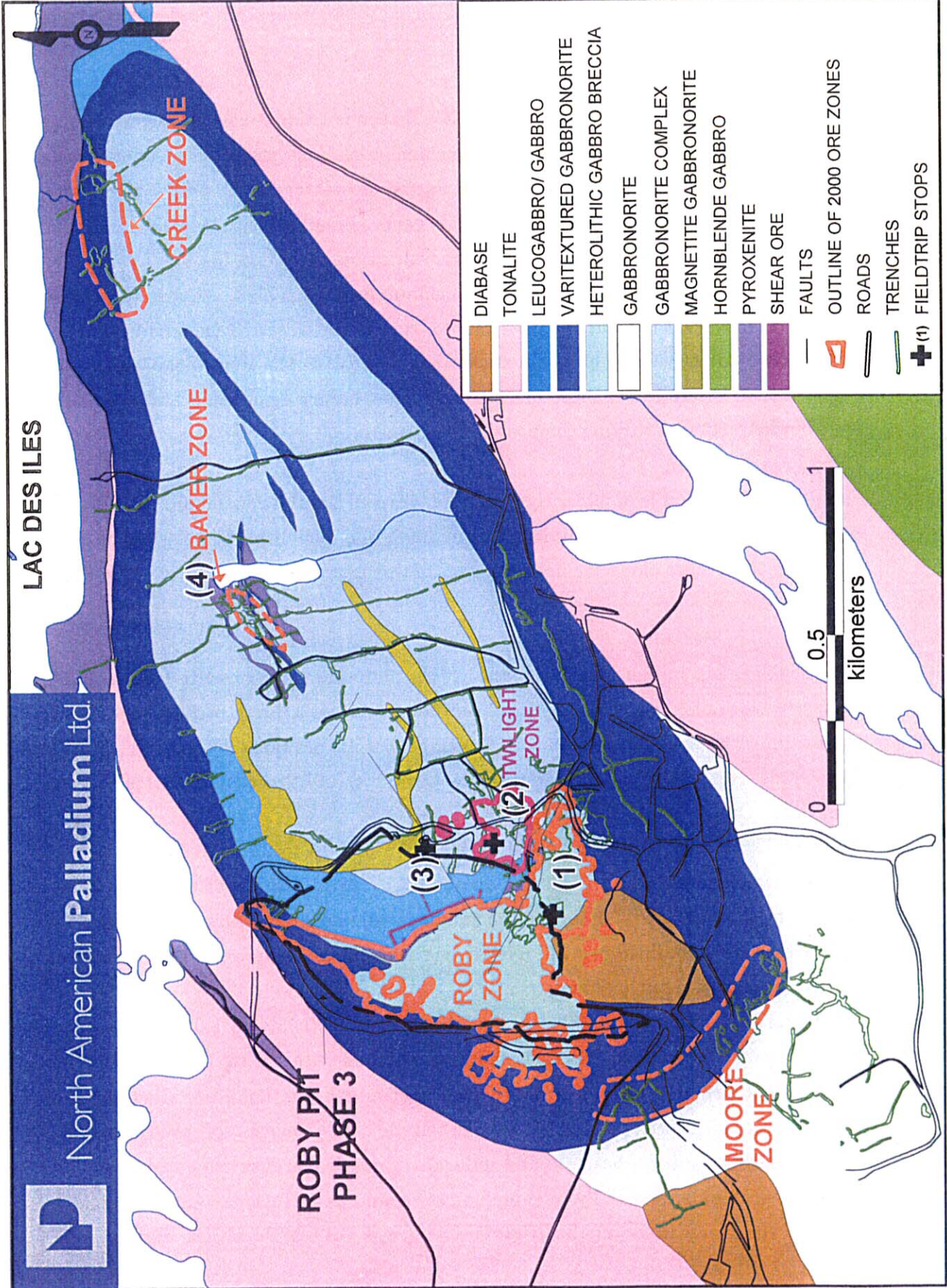


Fig. 2 Mineblock Intrusion Geology and Fieldtrip Stops

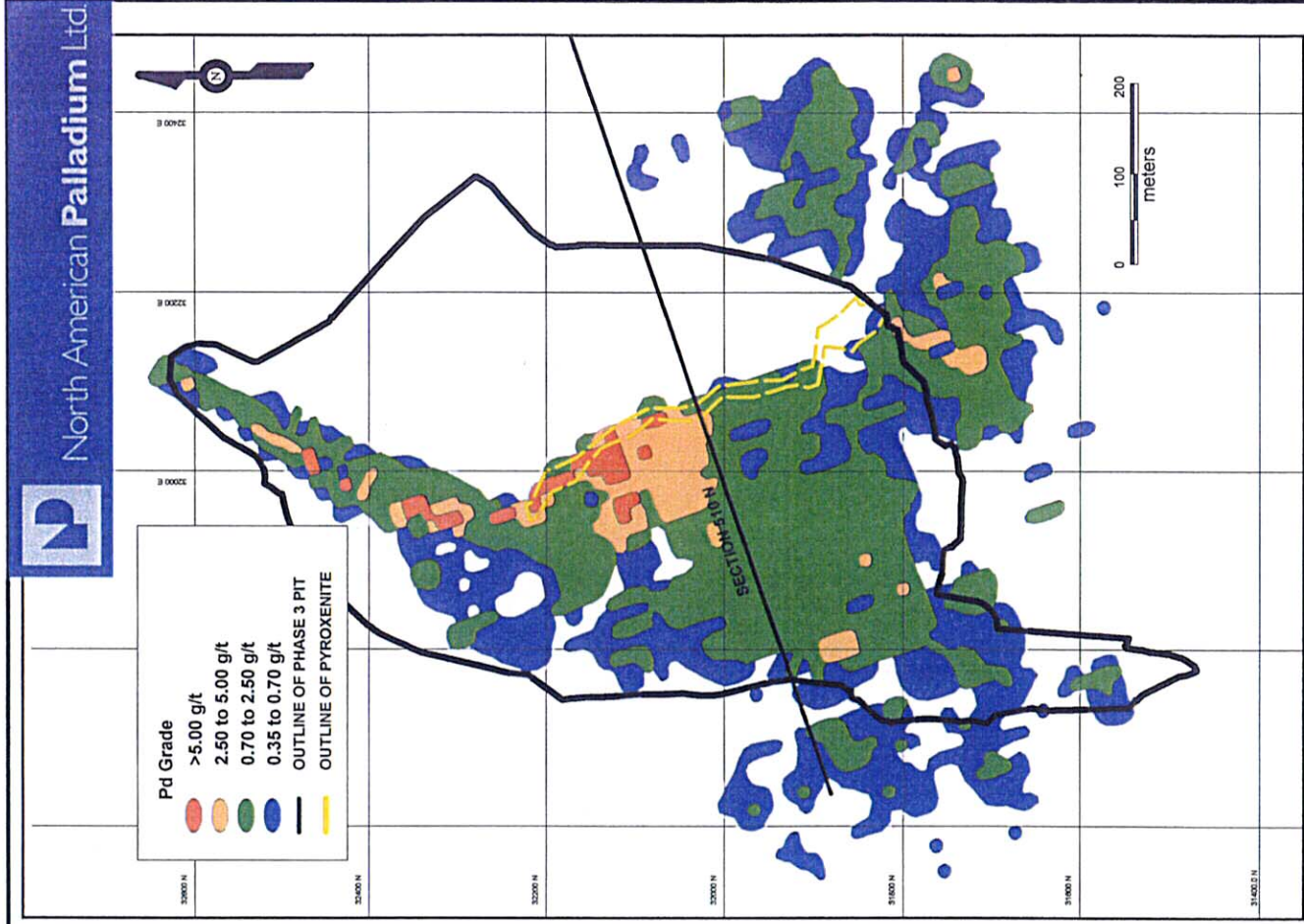


Fig. 4 ROBY ZONE Pd GRADE CONTOURS

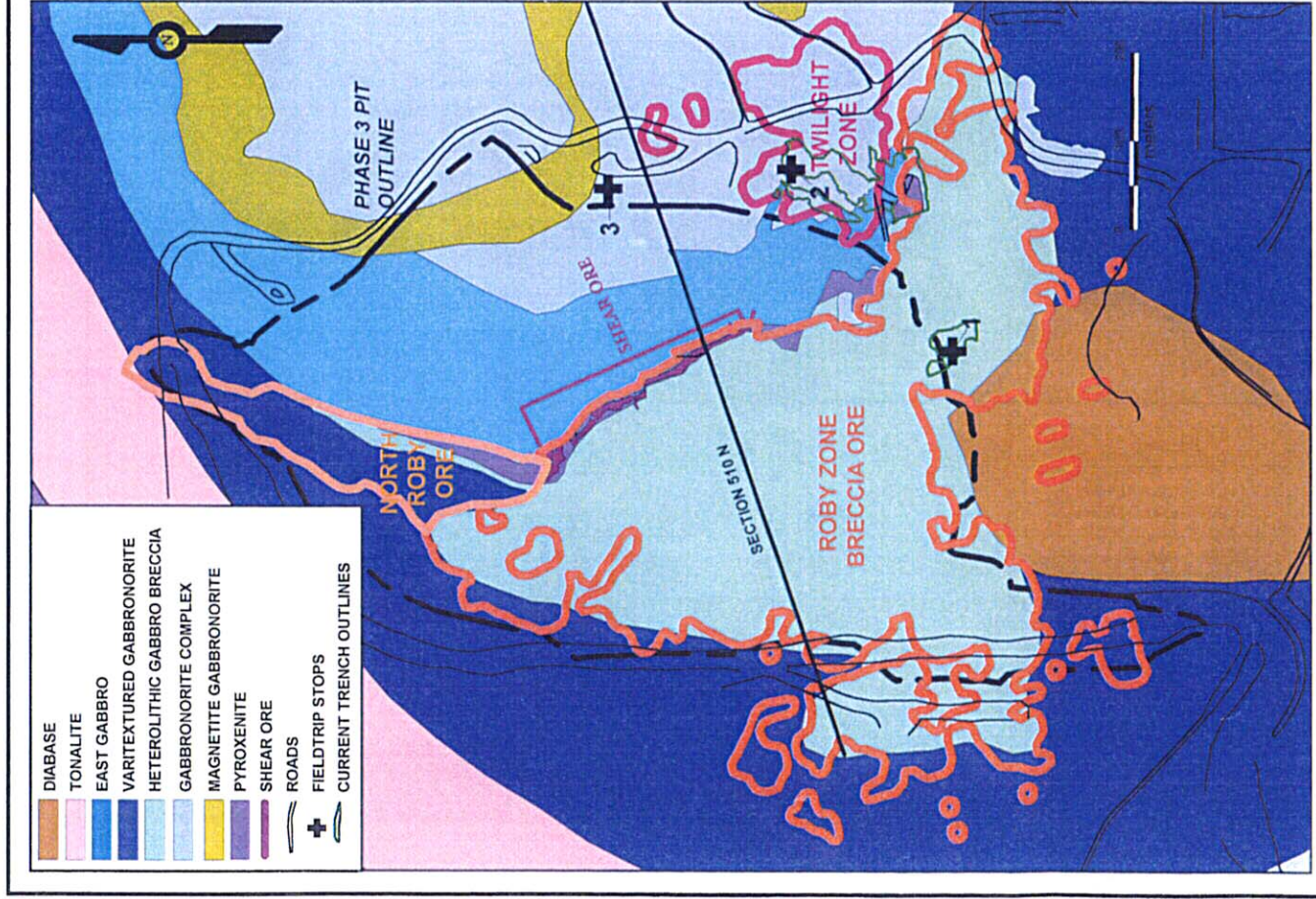


Fig. 3 ROBY ZONE GEOLOGY

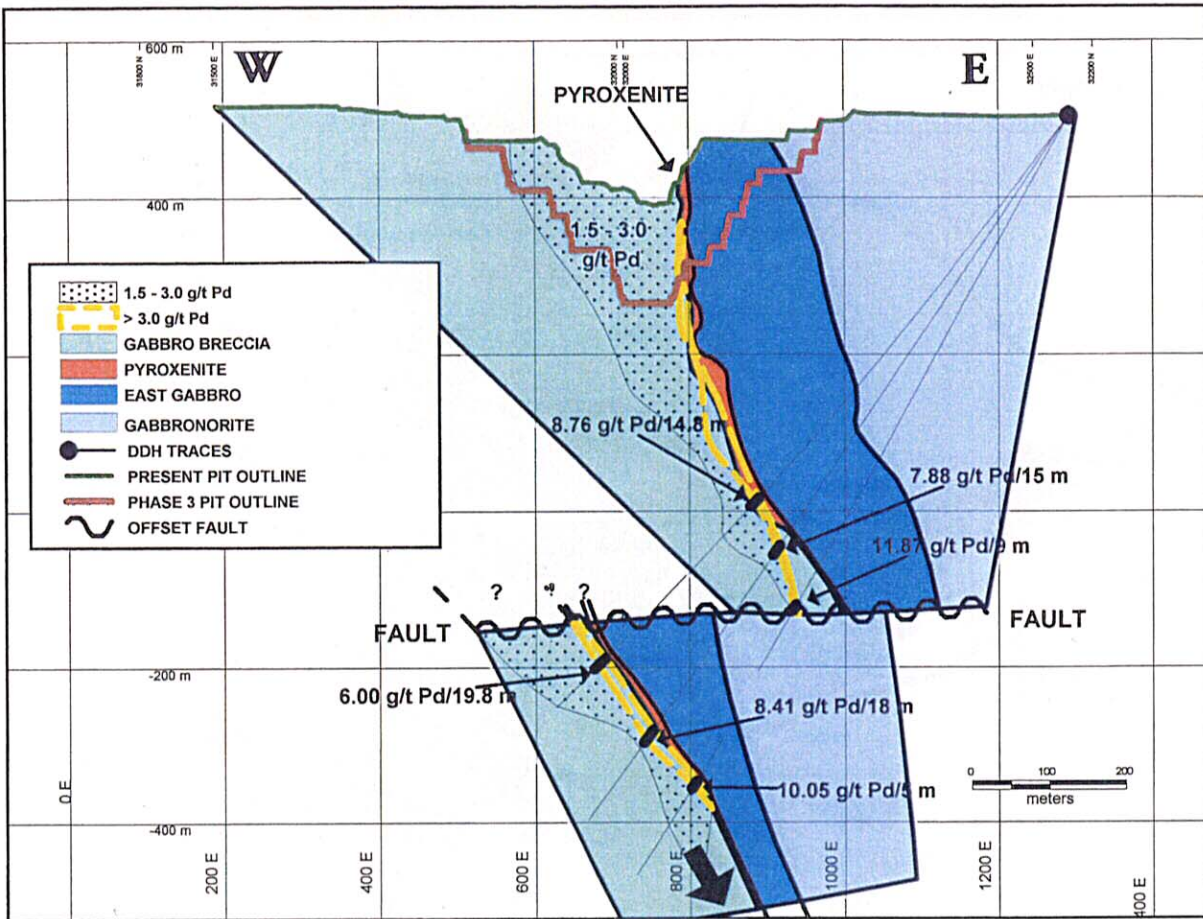


Fig. 5 SECTION 510 - GENERAL GEOLOGY

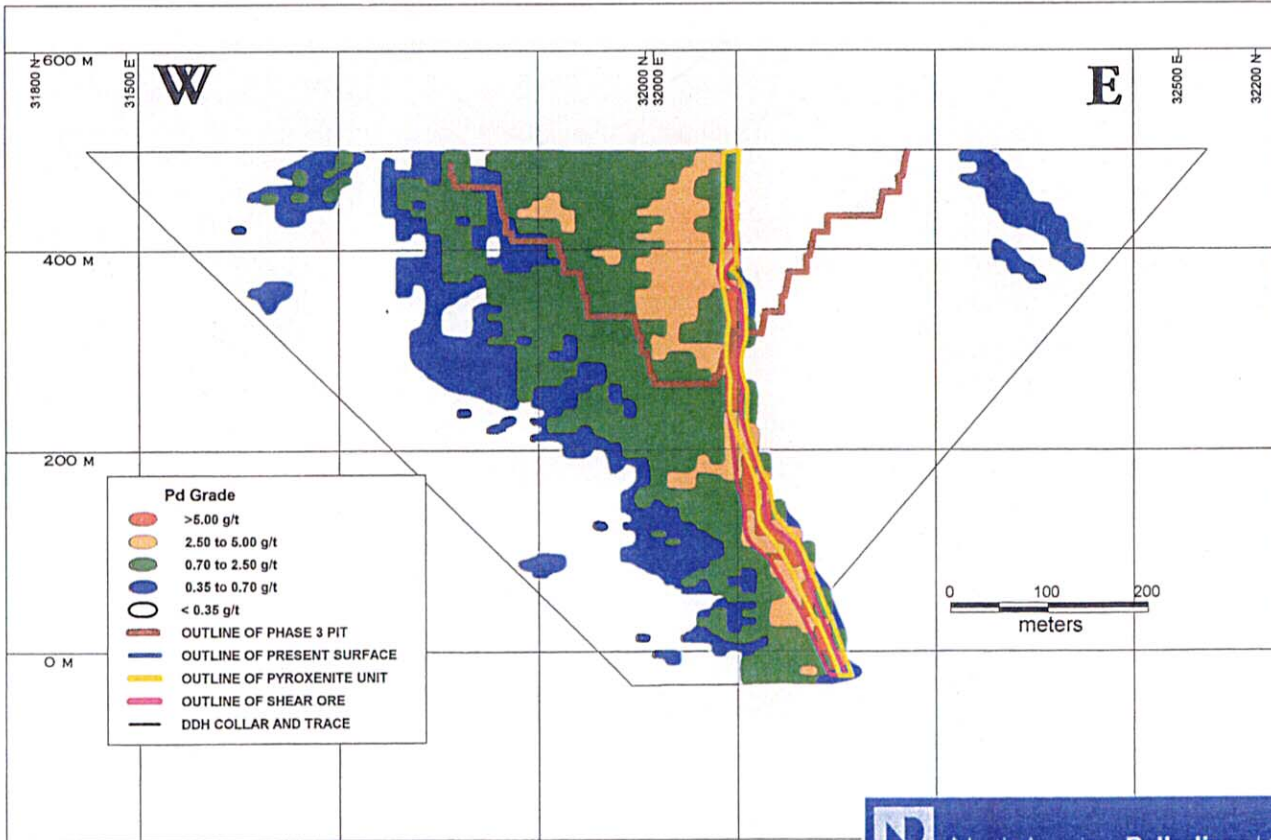


Fig. 6 SECTION 510 - Pd GRADES

55.95g/t Pd.

North Roby Ore

The North Roby Ore is a tabular zone that is 20 to 40 metres thick, and 200 metres long. Sulfide mineralization occurs primarily within gabbro and clinopyroxenite, ranging from trace amounts to 4%, and typically is less than 0.25%. The dominant sulfides are pyrrhotite, pentlandite, chalcopyrite and pyrite. Net-textured sulfides are common in gabbro and clinopyroxenite and occur locally within the gabbroic matrix of the heterolithic gabbro. PGE mineralization of the North Roby Ore, occurs primarily within gabbro and clinopyroxenite layers. The mean grade of the North Roby Ore is 1.7g/t Pd, and a maximum assay value of 39.7g/t Pd.

Breccia Ore

Southwest of the High Grade Ore, the central mass of the Breccia Ore is contained within a mineralized complex that measures 550 metres north-south by 350 metres wide. An arm of this mineralization has been traced for an additional 250 metres to the southeast. This complex contains lithologies that range in composition from clinopyroxenite to anorthosite to norite. Textures include equigranular fine to coarse grain, porphyritic, pegmatitic, varitextured and heterolithic gabbroic breccia, of which the latter three textures are host to the PGE mineralization. Sulfide mineralization within the Breccia Ore often comprises up to 5% of the rock and consists of intergrown chalcopyrite, pentlandite, pyrrhotite and pyrite. PGE mineralization within the Breccia Ore is associated with heterolithic gabbro breccia and accompanying varitextured gabbro. The mean grade of the Breccia Ore is 1.2g/t Pd, with a maximum of 36g/t Pd.

Twilight Zone

The Twilight Zone is a small pod of gabbroic breccia found on the eastern side of the Roby Pit. It is separated from the Roby Zone on its western boundary by the 50-70m thick East Gabbro and separated from the southeastern extension of the Roby Zone by a 100m thick barren, post-mineralization gabbro dike. The breccia is characterized by large, sub-round noritic clasts set within a matrix of melanogabbro.

Baker Zone

The Baker Zone is a lower grade deposit situated approximately 1km ENE of the Roby Zone Pit. The zone is situated in the northern/central portion of the Lac des Iles Complex and consists mostly of equigranular norite/gabbro that has been intruded by heterolithic melanogabbro breccia, melanogabbro, and a small amount of varitextured gabbro.

Keweenawan Supergroup

Mesoproterozoic intrusive, volcanic and minor sedimentary rocks associated with the Midcontinent Rift (MCR) comprise the Keweenawan Supergroup. On the northern margin of the MCR, Keweenawan rocks include a variety of intrusive rocks.

Copper-Nickel-Platinum Group Element Deposits Associated with the MCR

Rift-related intrusive rocks, such as the Duluth Complex in Minnesota, the Mellen Complex in Wisconsin, the Logan sills and Nipigon plate near Thunder Bay, as well as the alkalic-carbonatite complexes near Marathon, are most commonly exposed along the margins of the MCR. The presence of widespread disseminated Ni-Cu-sulphide mineralization in basal units of the the Duluth Complex, the Mellen Complex, the Coldwell alkalic complex, and in a number of gabbroic to peridotitic dykes and sills, shows that processes necessary for the formation of sulphide deposits were active in some intrusions locally in the MCR.

Intrusion / Lithology	Mineralization Style	Local Examples (Associated PGE-Mineralized Areas)	
<i>Reversed Magnetic Polarity???</i> < (1005-1102 Ma):			
Layered gabbro-anorthositic intrusions: Crystal Lake Gabbro	Disseminated and blebby sulphides in medium- to coarse-grained, varied-textured (a.k.a. 'taxitic') gabbro	Crystal Lake Gabbro (Great Lakes Nickel deposit)	
	Disseminated sulphides in Cr-spinel-bearing cumulate layers	(Cr-spinel-bearing, anorthositic gabbro above Cu-Ni deposit)	
Pine Point - Mount Mollie Gabbro	Disseminated, intergranular sulphides, fracture fillings	(Mount Mollie; Pine River)	
<i>Normal Magnetic Polarity</i> < (1005-1102 Ma):			
Pigeon River and Arrow River diabase dykes	Disseminated sulphides in medium- to coarse-grained gabbro and diabase	(Wallenius; Naomi Island; Jarvis Point)	
<i>Reversed Magnetic Polarity</i> (ca. 1108 Ma):			
Logan diabase sills, cone sheets	Sparse, disseminated sulphides	(Numerous)	
Nipigon gabbros			
Tholeiitic to alkaline complexes: Coldwell alkaline complex Killala Lake alkaline complex	Disseminated sulphides in medium- to coarse-grained, vari-textured (a.k.a. taxitic) gabbro	Coldwell	Killala Lake
		Two Duck Lake gabbro (Marathon deposit); Geordie Lake gabbro	Border gabbro (Sandspit - Killala)
	Disseminated sulphides in massive Fe-Ti-oxide cumulate layers	Eastern Border gabbro (Skipper Lake zone)	(Unknown)
Layered(?) picritic ultramafic intrusions	Disseminated sulphides and native metals in lherzolite, dunite, peridotite, etc.	Leckie Lake (Wolf Mountain); Hele Township; Eva-Kitto townships	

Exploring for PGM near the Lake Nipigon area (Fig.7,8,9)

The Lake Nipigon area near Thunder Bay where ultramafic rocks of the Nipigon embayment are being tested for their potential to host prolific Noril'sk-type nickel deposits. Many active companies are mentioned by this district.

Company's names are as follows.

North American Palladium Ltd.

New Millennium Metals Corp.

Platinum Group Metals Inc.

Jonpol Explorations Ltd and Wolfden Resources Inc.

East West Resources Corp.

Canadian Golden Dragon Resources Ltd.

McVicar Mineral Ltd.

Geomaque Explorations Ltd.

LEH Ventures Ltd.

Aurora Platinum Corp.

PGM Ventures Corp.

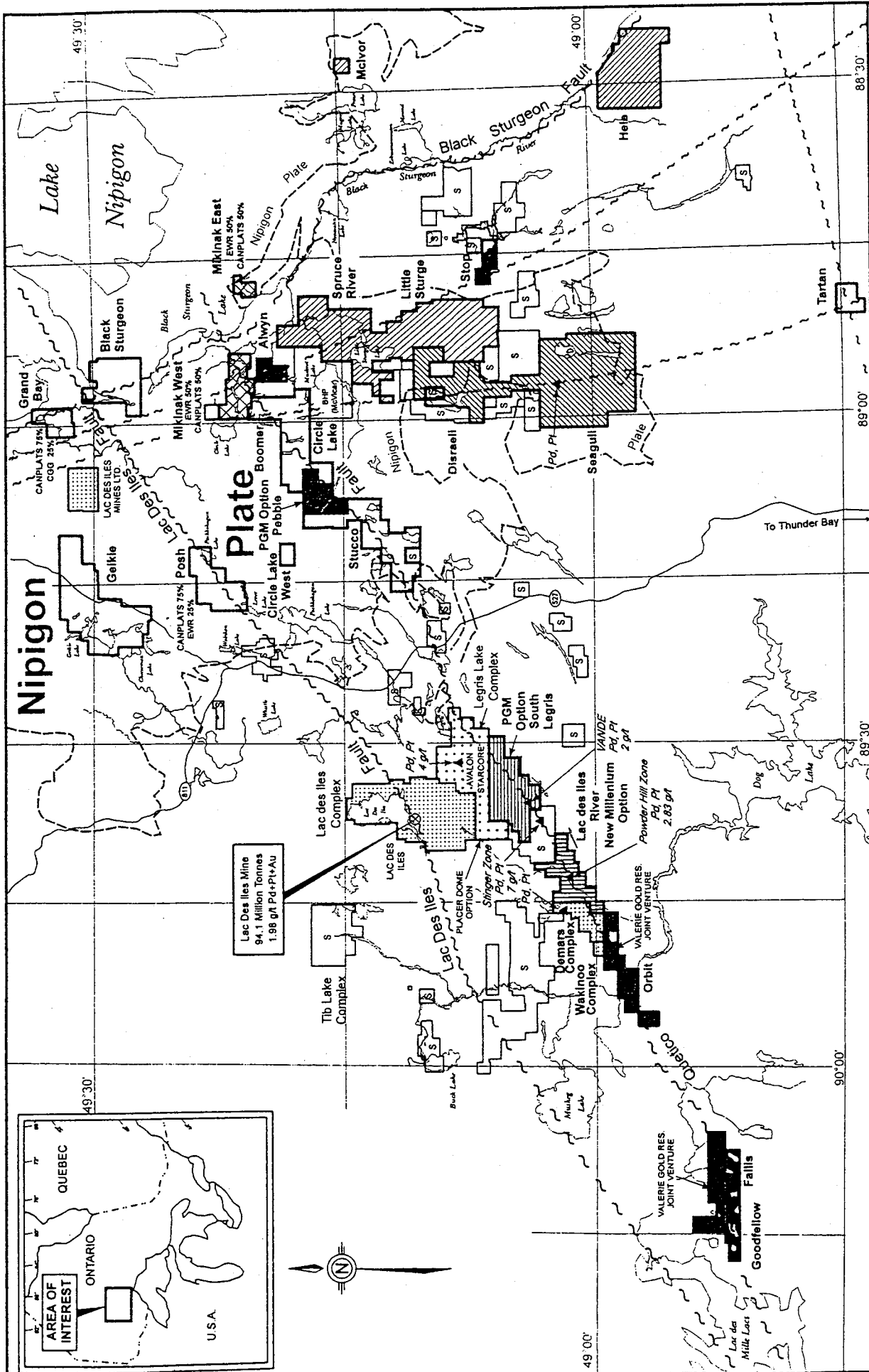
Spider Resources Inc.

Landore Resources Inc.

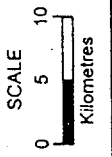
Canplats Resources Corp.

Hunter Dickinson Inc.

East West Resources Corp. and joint-venture partner Canadian Golden Dragon Resources Ltd. are exploring the Seagull and Disraeli intrusions for PGM. The project is 90km northeast of Thunder Bay, and forms part of the Wolf Mountain project. Diamond drilling has been conducted on magnetic anomaly in Seagull pluton, its conical phase measures 3km long 2km wide and is composed of an upper phase of pink granophyric gabbro and a pyroxenite-gabbro unit that contains a magnetite-rich layer with elevated PGE values (1 to 3g/t Pd+Pt). The sulphide zone with elevated PGE values ranges up to 16m in thickness and sometimes forms two layers containing more elevated values.



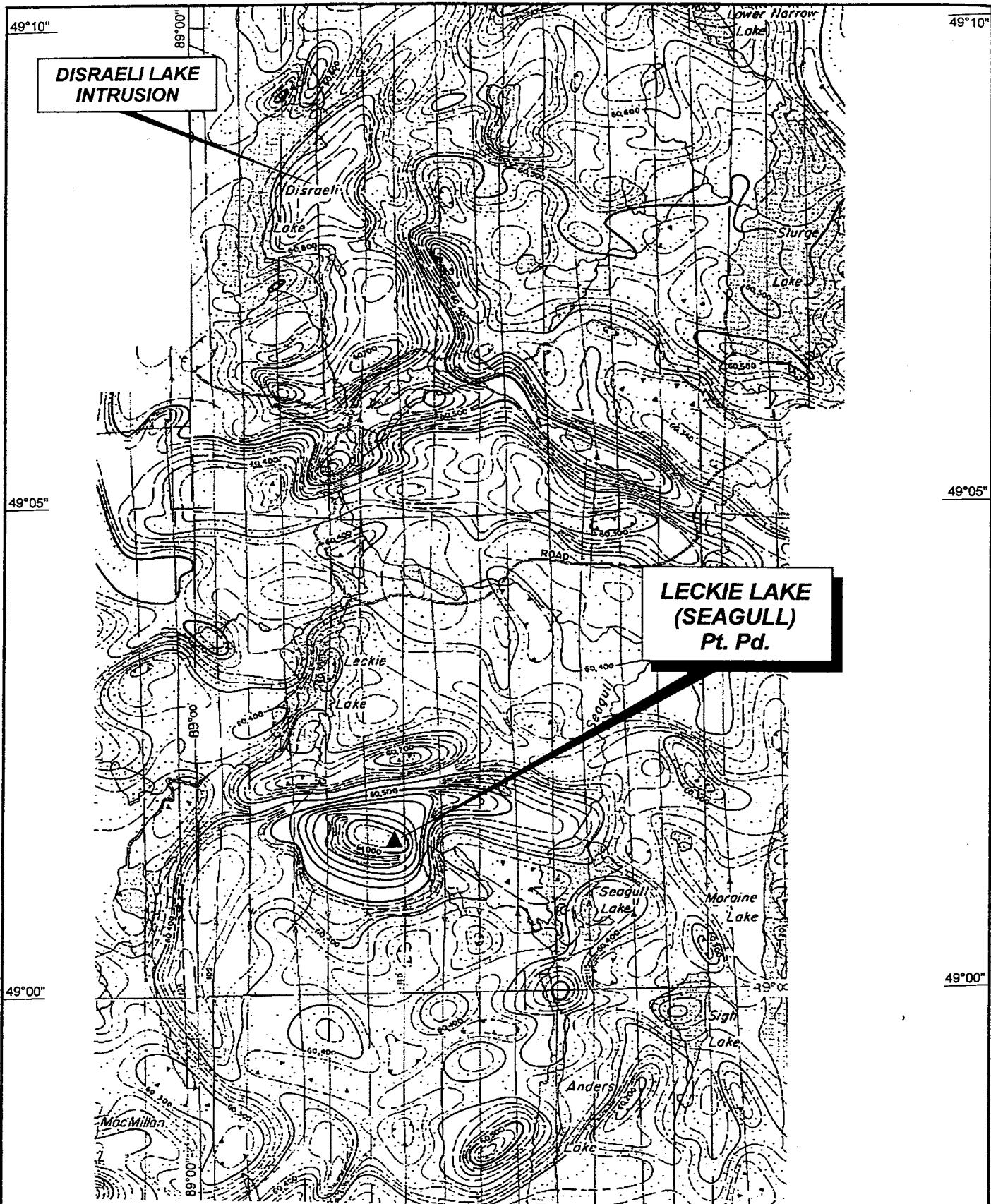
NOTE: This map is intended to provide a general overview of the location of the properties. It is not intended to be used as a legal document. The map is copyright by Robert S. Stephenson and the publisher may be held responsible in any form without the consent of Robert S. Stephenson.



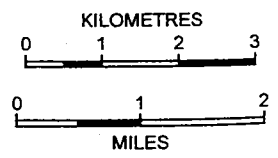
EAST WEST RESOURCE CORPORATION PGE PROPERTIES NIPIGON PLATE AND QUETICO BELT

ONTARIO

- | | |
|--|--|
| | LAC DES ILES MINES LTD. |
| | VALON VENTURES LTD. |
| | STARCORE RESOURCES LTD. |
| | STAKED |
| | Pd, Pl, Au (PGE) |
| | EAST WEST RESOURCE CORPORATION - 100% |
| | CANPLATS RESOURCES CORPORATION - 100% |
| | CANADIAN GOLDEN DRAGON RESOURCES LTD. - 100% |
| | EAST WEST RESOURCE CORPORATION - 30% |
| | CANADIAN GOLDEN DRAGON RESOURCES LTD. - 30% |
| | AVALON VENTURES LTD. - 40% |
| | EAST WEST RESOURCE CORPORATION - 50% |
| | CANADIAN GOLDEN DRAGON RESOURCES LTD. - 50% |
| | EAST WEST RESOURCE CORPORATION - 50% |
| | MAPLE MINERALS LTD. - 50% |
| | EAST WEST RESOURCE CORPORATION - 50% |
| | CANPLATS RESOURCES CORPORATION - 50% |



**Fig. 8 EAST WEST RESOURCE CORPORATION
DISRAELI LAKE INTRUSION
AND SEAGULL PLUTON
(Wolf Mountain)**



WOLF MOUNTAIN PROJECT

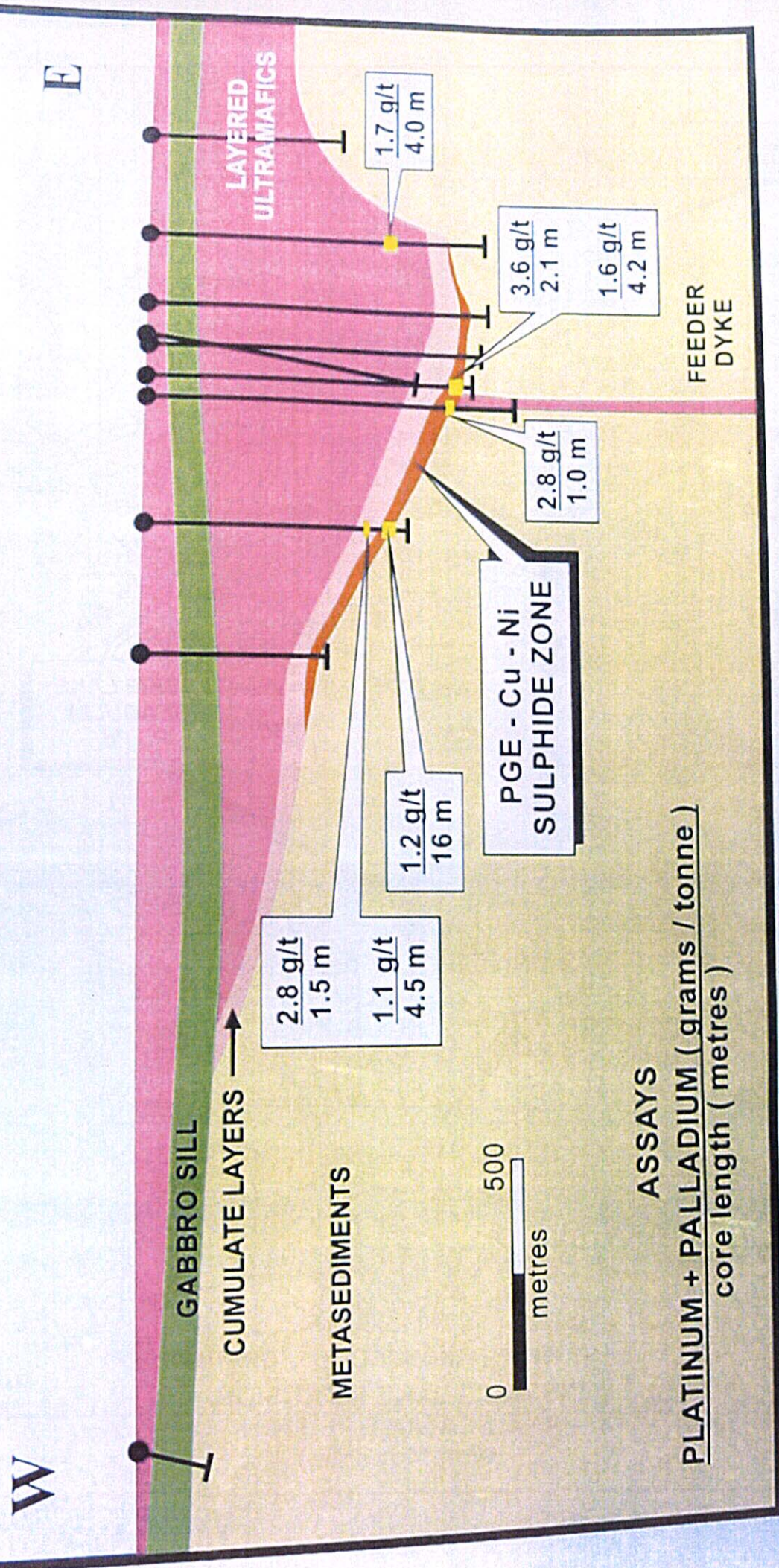


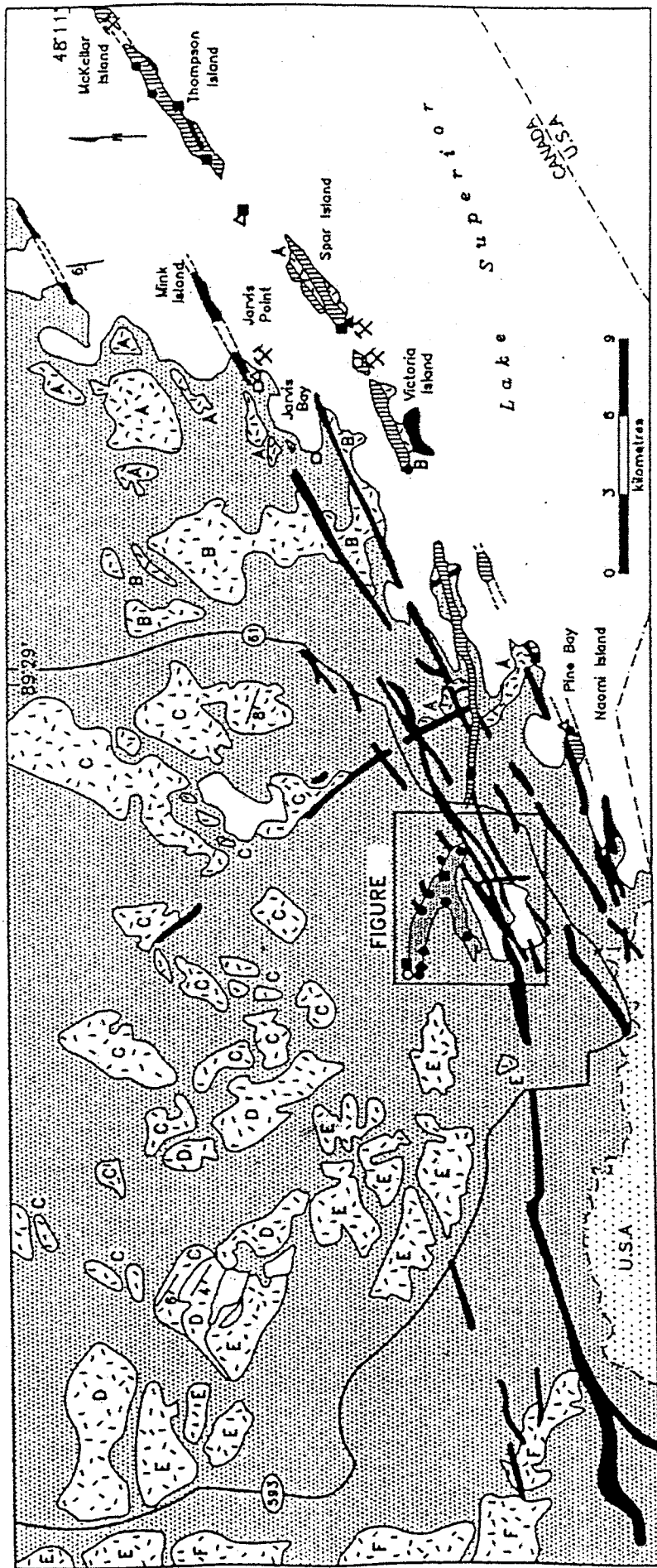
Fig. 9 Geological Cross section of the Seagull Intrusion (Wolf Mountain Project). The Noril'sk-type Cu-Ni-PGE target is indicated in red. Sulphides from this area display composite blebby textures and a high base metal-PGE to sulphur ratio similar to Noril'sk-type mineralization.

Crystal Lake gabbro (Fig.10,11)

Crystal Lake gabbro, which is Y-shaped in plan view, with a west-northwest-striking limb 5km long and an east-northeast-striking, southern limb 2.75km long. This intrusion was subdivided into four major, roughly equivalent, lithologic zone.

- (1) an upper zone(60 to 80m) of sulfide-barren troctolite, olivine gabbro and anorthositic gabbro
- (2) a middle zone(30 to 42m) of cyclic, layered anorthositic and olivine gabbro, Cr-spinel-bearing anorthosite, and olivine gabbro
- (3) a lower, unlayered zone(50m) of vari-textured gabbro and leucotroctolite, which hosts the bulk of the Cu-Ni-sulfide deposit
- (4) a basal zone(1 to 7m) of fine-grained, chilled melagabbro and hornfelsed country-rock xenoliths

The sulphides which comprise the Great Lakes Nickel deposit are disseminated, interstitial and included grains and droplets. Pyrrhotite, chalcopyrite, cubanite and pentlandite are the main sulphides. A grab sample of the lower adit area returned 116ppb/t Au, 688ppb/t Pd, 14ppb/t Pt, 13494ppm/t Cu, 4756ppm/t Ni and 255ppm/t Zn. This deposit reserves 45.6 million tons at a grade of 0.334% Cu and 0.183% Ni.



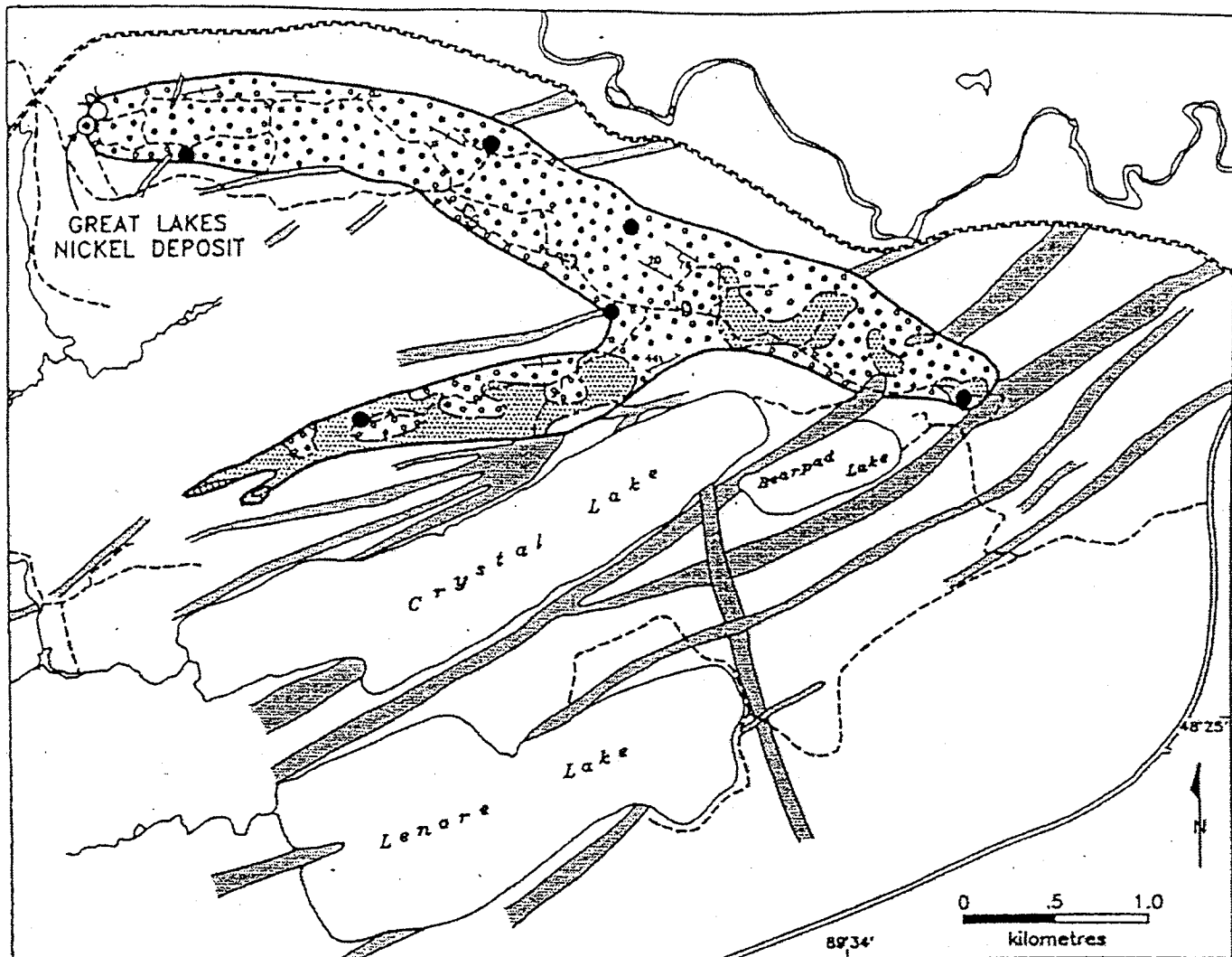
- Proterozoic Keweenaw**
- PINE RIVER-MOUNT MOLLIE INTRUSION (and associated composite dikes)
diorite, olivine gabbro, granophyre
 - CRYSTAL LAKE GABBRO
olivine gabbro, pegmatitic gabbro, anorthosite
 - PIGEON RIVER INTRUSIONS
olivine diabase dikes

- EARLY MAFIC INTRUSIONS
diabase sills and sheets
A, C-F = diabase
B = diabase + granophyre
stratigraphically up
- Anlimikite
- SEDIMENTARY ROCKS (ROYE FORMATION)
wacke, shale, siltstone, mudstone
strike and dip of sill

- Pt+Pd > 200 ppb
- Cr > 500 ppm
- Cu+Ni > 5000 ppm
- Pb > 500 ppm
- Zn > 100 ppm
- Ag > 5 ppm
- Au > 800 ppb
- Mines (Past Producers)

After Smith and Sutcliffe (1987)

Figure 1987). Geology and distribution of intrusive Keweenaw rocks of the Crystal Lake-Pigeon River Region (after Smith and Sutcliffe 1987).



Middle Proterozoic

Crystal Lake Gabbro

- olivine gabbro, troctolite
- gabbro, leucogabbro, pegmatitic gabbro, anorthosite
- diorite
- chilled gabbro

Pigeon River Dikes

- olivine diabase

Early Proterozoic

Rove Formation

- argillite, wacke

- layering (strike dip)
- foliation (strike dip)
- lithological contact
- adit
- roads

- sulphide mineralization with > 5000 ppm Cu+Ni
- sulphide mineralization with > 200 ppb Pt+Pd
- sulphide mineralization with > 500 ppm Cr

After Smith and Sutcliffe (1987)

Figure . Distribution of rock types in the Crystal Lake Intrusion (after Smith and Sutcliffe 1987).

Coldwell alkaline complex (Fig.12,13,14(1),(2),Tab.1)

A variety of Mesoproterozoic, Midcontinent Rift-related alkaline and carbonatitic rocks comprise several intrusive complexes on the north shore of Lake Superior. They include the Coldwell and Killala Lake alkaline complexes, the Prairie Lake carbonatite and numerous diatremes and related dykes in the vicinity of Dead Horse Creek. The Coldwell complex covers an area of approximately 580km². The complex comprises three, superimposed ring subcomplexes or magmatic centers that young progressively (Centers 1⇒3) to the southwest.

Center 1 : saturated alkaline rocks with oversaturated residue ; chiefly gabbro and Fe-rich augite syenite

Center 2 : miaskitic alkaline rocks with oversaturated residue ; chiefly alkaline biotite gabbro and nephline syenite

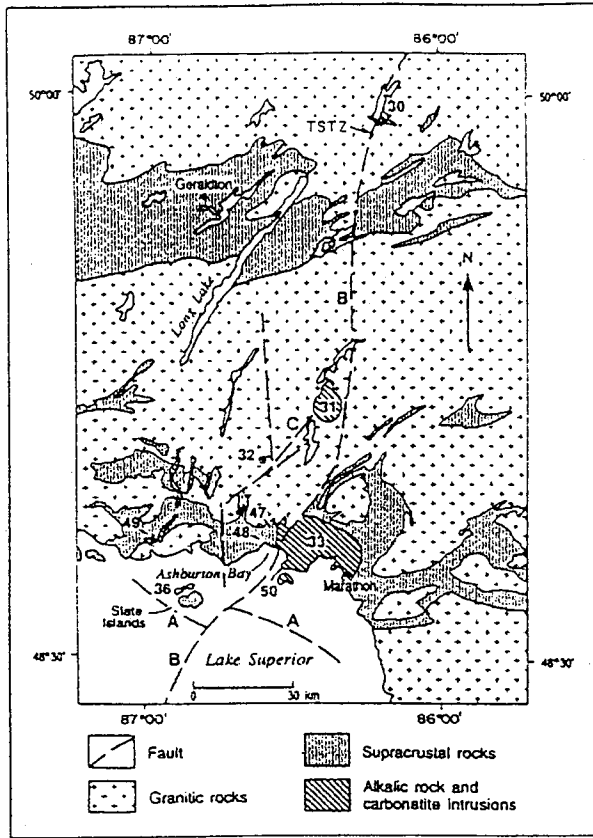
Center 3 : alkaline rocks with oversaturated residue ; chiefly syenite and quartz syenite

A variety lamprophyric and other dyke rocks occur within the complex.

Mafic(+ultramafic) intrusive rocks occur within Centers 1 and 2 and are tabulated, with their associated mineralized zones. Mineralized zones occur in the border gabbro at the eastern (Marathon deposit; Skipper Lake Zone) and western (Middleton occurrences) margins of the complex, and in its interior at Geordie Lake.

Marathon deposit

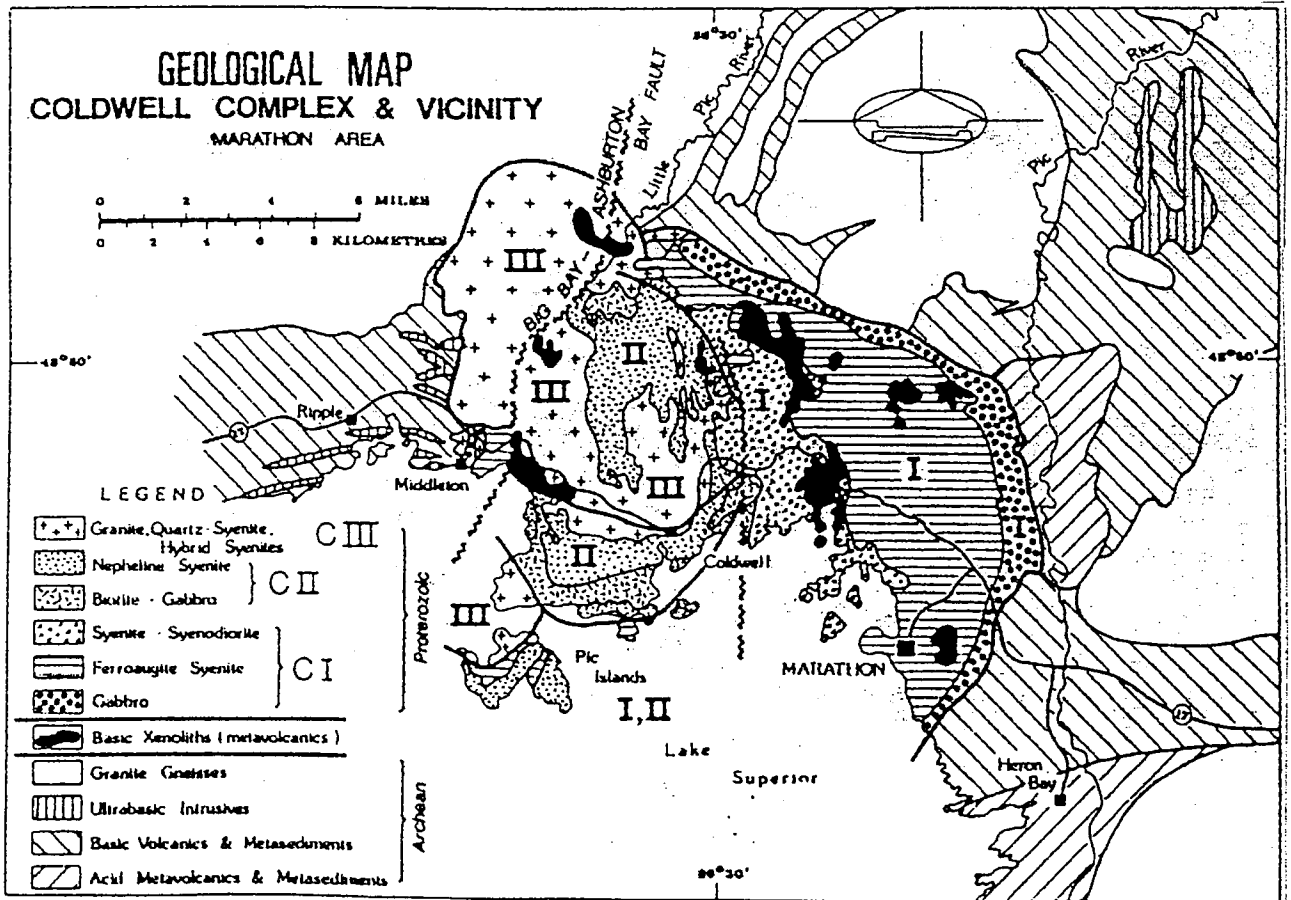
The mineralization has a north-south strike length of about 1.9 kilometers and dips 30° to 40° to the west, and has a true thickness ranging from 15 to 100m. The Southern Extension Zone incorporates two targets (BR Target and the RD Target) situated to the south of the known resource area (Main Zone) in the under-explored southern 2.3km of the Marathon Property. Grades of up to 4.70g/t Pt+Pd over 15.5m have been intersected in past drilling in the BR Target. The BR Target has a strike length of 225m. The RD Target is located 500m south of the BR Target. Several grab samples grading in the 2.0g/t Pt+Pd range have been found. The significant sulphide occurrences are restricted to the layered olivine gabbro (Eastern Border gabbro), with major concentrations as pods, lenses and interstitially, disseminated grains in the coarse basal phase at or near the contact with the Archean supracrustal rocks. Pyrrhotite and chalcopyrite are the principal sulphide minerals.



Regional geology (Sage 1991) in the vicinity of the Trans-Superior Tectonic Zone (TSTZ), extension of the Thiel Fault (B)

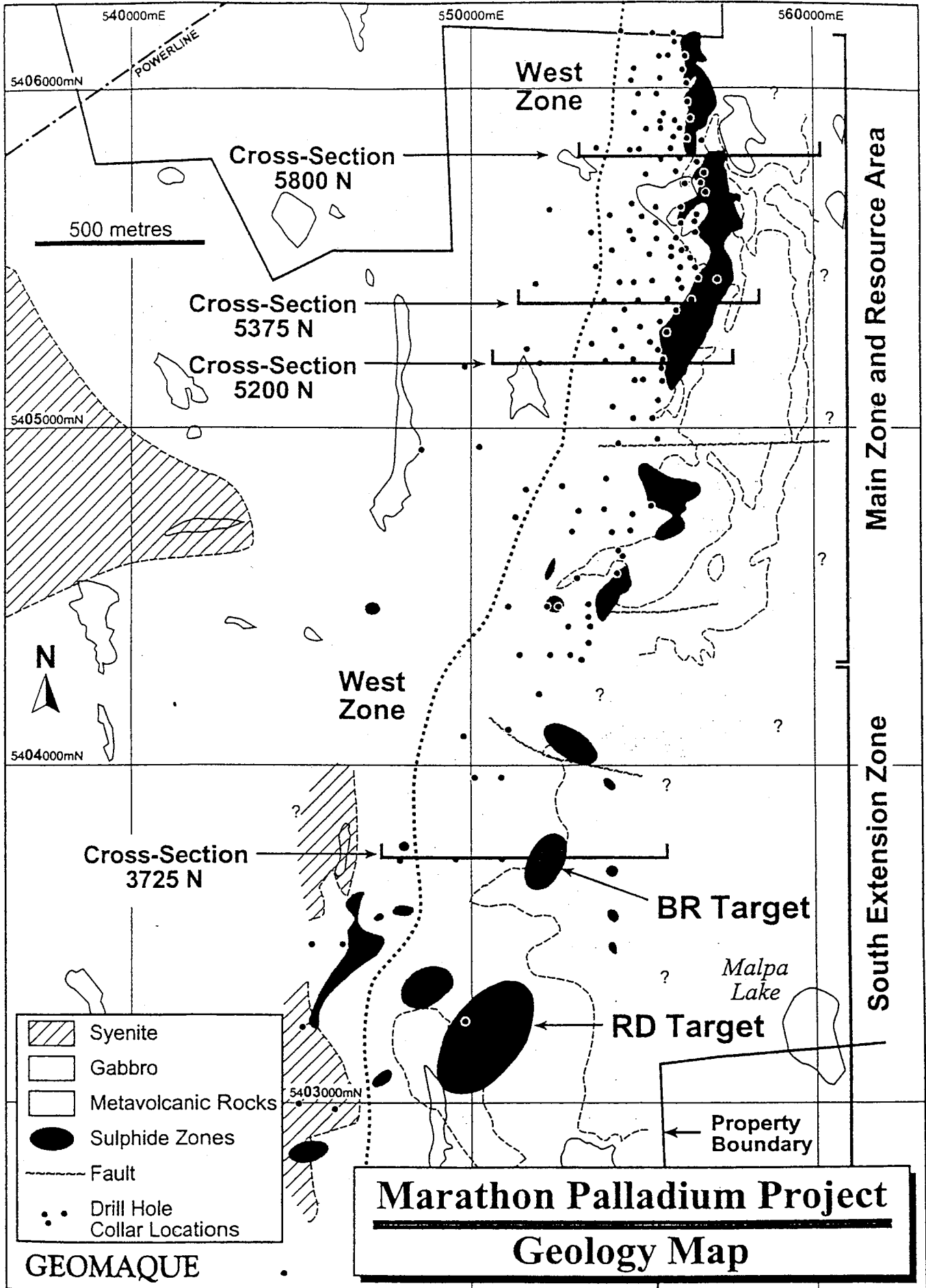
Key to numbering:
 30 - Chipman Lake fenites / carbonatite dykes
 31 - Killala Lake alkaline complex
 32 - Prairie Lake Carbonatite
 33 - Coldwell alkaline complex
 47 - Dead Horse Creek diatremes
 48 - McKellar Creek diatreme

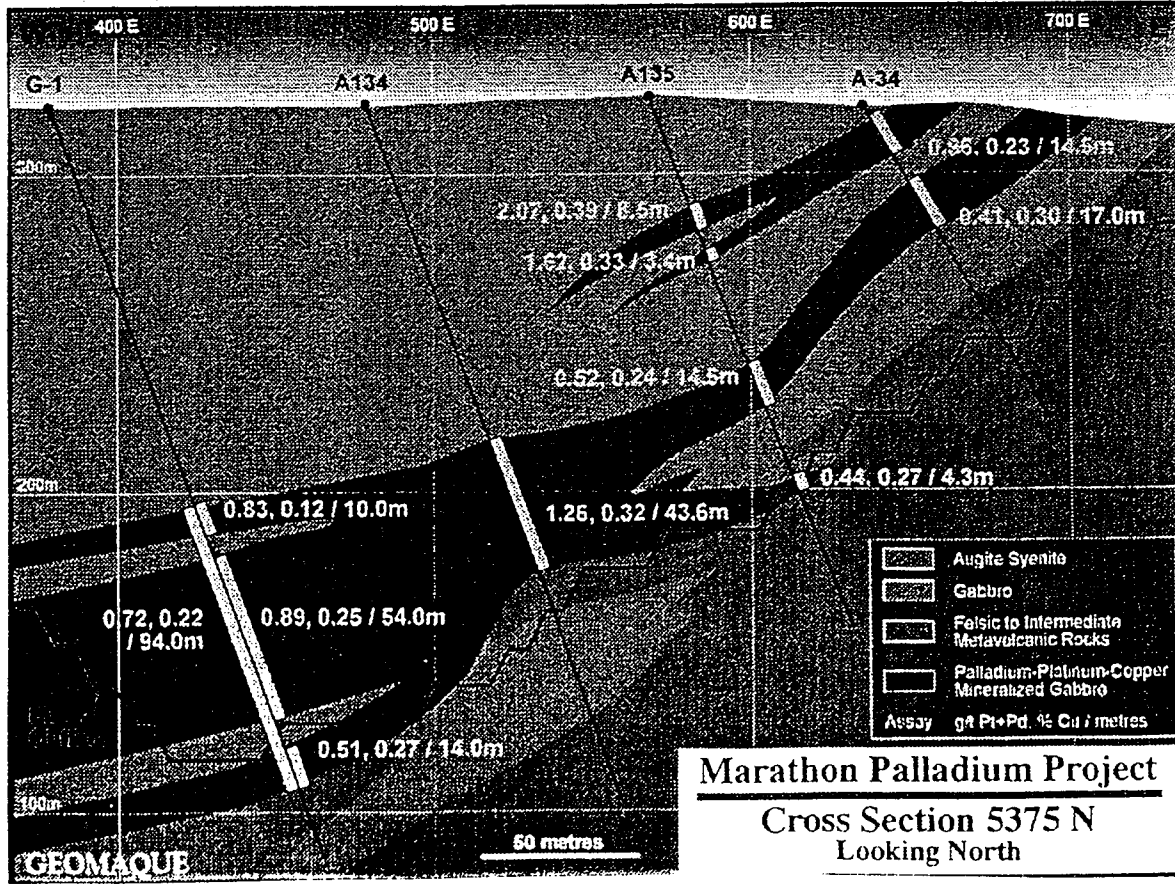
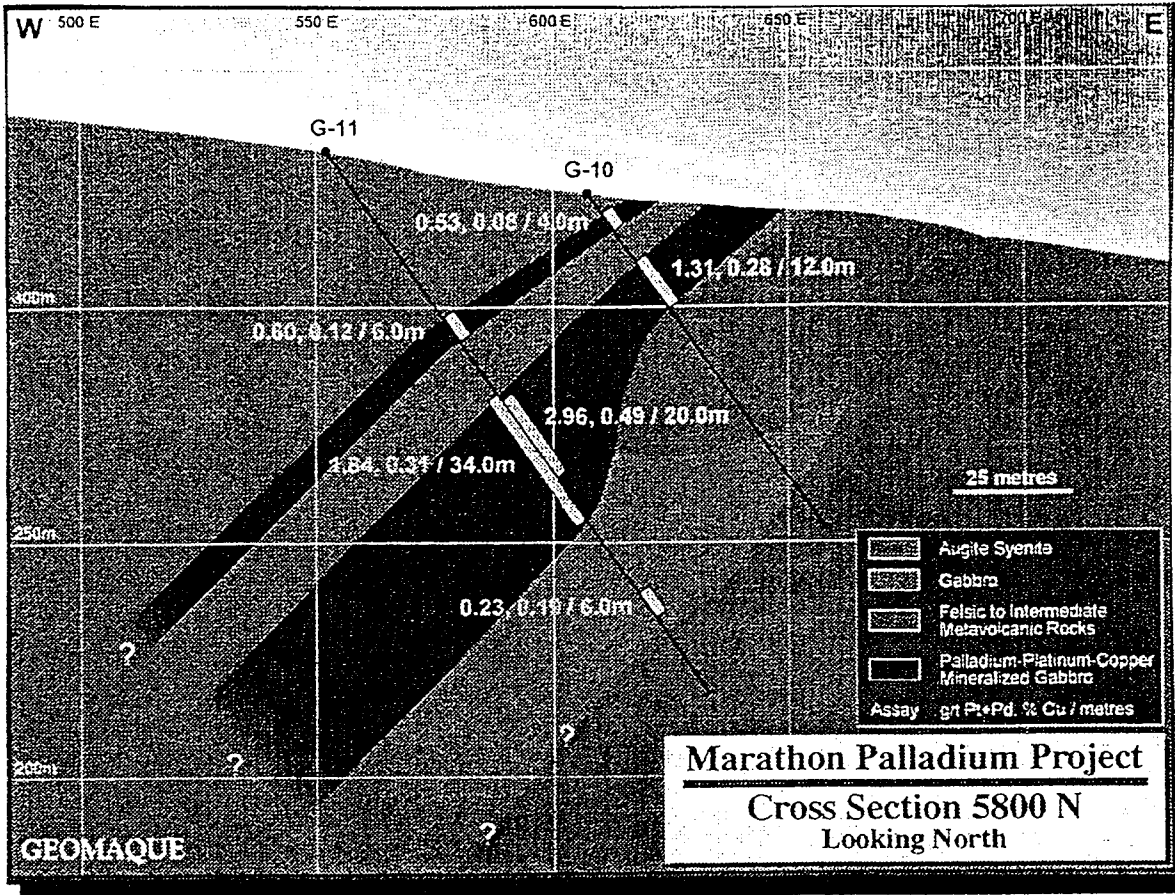
Generalized geology of the Coldwell alkaline complex after Mitchell and Platt (1994)

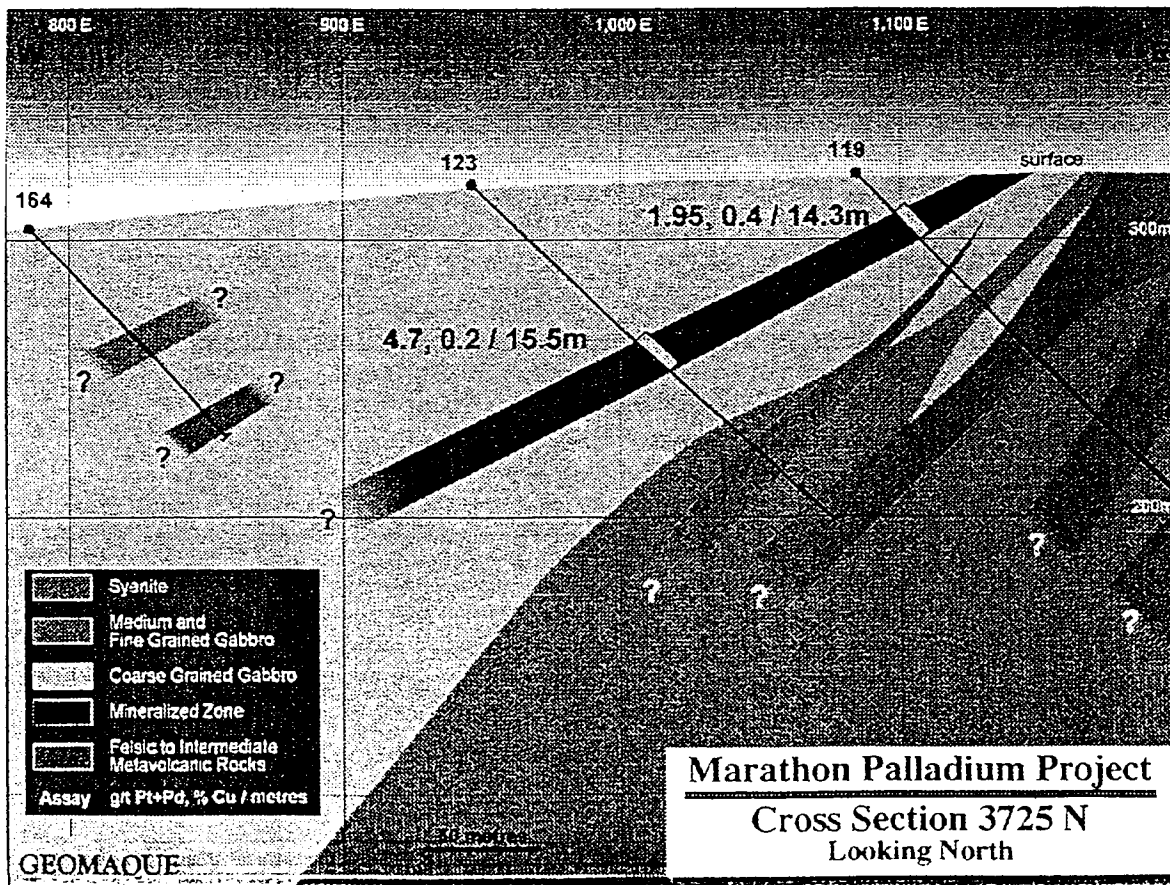
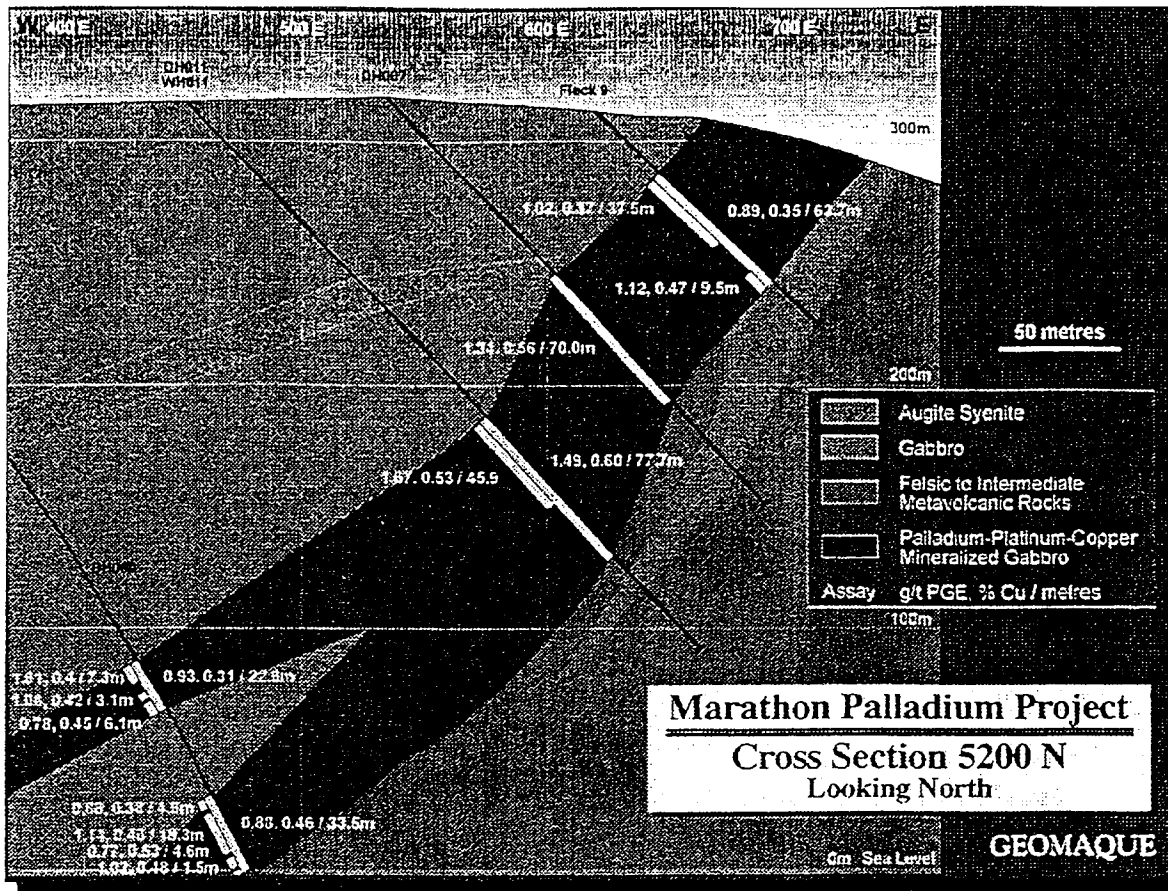


Intrusion (Center)	Lithologic Units	Reference(s)	
Eastern Gabbro (1)	layered gabbro cumulates (olivine gabbro, gabbro, troctolite, anorthositic gabbro)	Shaw (1994, 1997); Lum (1973)	
Western Gabbro (1)	Massive and layered series gabbro; olivine-bearing	Penczak (1992); Wilkinson (1983)	
Two Duck Lake (1)	Gabbronorite, olivine gabbronorite, olivine-bearing gabbro, leucogabbro	Shaw (1994, 1997)	
	Hornblende gabbro to monzodiorite; olivine ferrogabbro to ferrodiorite; olivine gabbro to diorite	Dahl <i>et al.</i> (1987)	
Malpas Lake (1)	Amphibole-bearing olivine gabbro	Shaw (1994, 1997)	
Geordie Lake (?)	Troctolite, olivine gabbro	Mulja (1989); MacTavish <i>et al.</i> (1987)	
Alkaline gabbro (2)	Biotite gabbro	Mitchell and Platt. (1982)	
	Biotite- and olivine-gabbro	Walker <i>et al.</i> (1993a)	
Deposit / Mineralized Zone	Grade / Significant Assays	Ore Mineralogy	Reference(s)
Marathon	37 million tonnes averaging 1.1 g/t Pd, 0.27 g/t Pt, 0.21 g/t Au 1.85 g/t Ag, 0.38% Cu and 0.032% Ni	Chalcopyrite ≥ pyrrhotite >> pentlandite > cubanite ≥ pyrite; sphalerite, hollingworthite, atokite-zvyaginsevite, sperrylite, Bi-kotulskite, michenerite, merenskyite, monceite, stibiopalladinite, paolovite, merteite II, palladoarsenide, unnamed (Pd ₅ As ₂), nickeline, majakite, aregentian gold	Ohnenstetter <i>et al.</i> (1991); Watkinson and Ohnenstetter (1992); Good and Crocket (1994a,b)
Geordie Lake	5.2 g/t Pd, 1.71% Cu, 244 ppb Pt, 373 ppb Au, 28.8 g/t Ag; drilling in 2000 yielded 1.164 g/t Pd, 0.054 g/t Pt and 0.070 g/t Au over 5.95 metres	Chalcopyrite, bornite, pyrite, millerite, siegenite, pentlandite, galena, chalcocite, melonite, hessite, unnamed (Ag ₃ Te ₂), altaite, kotulskite, merenskyite, michenerite, sopcheite, Pd-bismuthotelluride, paolovite, Pd-arsenide, guanglinite, Pd-antimonide, sperrylite, electrum, Pd _{1.6} As _{1.5} Ni, AgSb ₄	Mulja (1989); Mulja and Mitchell (1990, 1991)
Middleton		Chalcopyrite, pyrrhotite, pentlandite, sphalerite, pyrite	Penczak (1992)
Skipper Lake	average grade of 1.05 g/t Pd+Pt+Au over 12 m	Chalcopyrite, bornite, pentlandite, cobaltite, galena, chalcocite; telargpalite, polarite, kotulskite, taimyrite, merteite, zvyaginsevite, plumbopalladinite, majakite, tetraferroplatinum	MacTavish (2000)

CUT-OFF (g/t Pd)	MEASURED			INDICATED			INFERRED					
	TONNES	Pd (g/t)	Pt (g/t)	Cu (%)	TONNES	Pd (g/t)	Pt (g/t)	Cu (%)	TONNES	Pd (g/t)	Pt (g/t)	Cu (%)
0.001	24 602 220	0.698	0.182	0.264	46 450 710	0.621	0.168	0.243	51 637 320	0.460	0.130	0.218
0.04	16 963 200	0.912	0.229	0.317	29 387 070	0.658	0.221	0.306	24 245 100	0.749	0.194	0.302
0.8	8 649 000	1.215	0.296	0.357	13 613 805	1.163	0.287	0.347	7 940 340	1.135	0.279	0.386
1.2	3 357 765	1.588	0.391	0.389	4 614 560	1.523	0.378	0.377	2 374 290	1.633	0.378	0.424







APPENDIX 4

Description of rock and drill core samples

Sample No.	Depth		Longitude		Latitude		Altitude	Sample Type		Rock Name	Description	Mineral-ization	Weathering	Analysis method						
	m		D	M	S	D		M	S					Bulk	PGE	Ar-Ar	Thin	P-Thin	X-ray	EPMA
AS028B	---	50	13	44.9	27	46	41.5	1005	others	outcrop	phonolite									
AS028C	---	50	13	44.9	27	46	41.5	1005	others	outcrop	phonolite									
AS029	---	50	16	20.0	27	38	59.6	966	others	outcrop	carbonatite?									
AS030	---	50	13	50.4	27	32	45.4	866	others	outcrop	phonolite									
AS031	---	50	12	23.7	27	39	41.4	894	others	outcrop	kimberlite									
KM101	---	50	24	42.5	25	22	12.6	820	sediment	outcrop	shale	pyrite	---							
KM102	---	50	37	23.3	25	22	15.8	829	sediment	outcrop	shale	pyrite	---							
KM103	---	50	48	28.9	25	11	9.7		basic intrusion	outcrop	dolerite	---								
KM104	---	51	12	57.3	25	19	15.1		basic lava	outcrop	basalt	weak weathering								
KM105	---	51	20	58.6	25	21	50.0	1,107	basic lava	outcrop	basalt	weak weathering								
KM106	---	51	39	59.6	25	25	13.8	1,107	basic lava	outcrop	basalt	weak weathering								
KM107	---	52	9	41.8	25	41	35.6	814	basic lava	outcrop	basalt	pyrite	---						1	
KM108	---	52	18	33.0	25	17	43.7	852	basic lava	outcrop	basalt	---								
KM109	---	52	17	19.8	25	8	2.9	677	basic lava	outcrop	basalt	---								
KM110	---	52	4	7.7	24	52	19.4	984	basic lava	outcrop	basalt	weathering								
KM111	---	51	45	53.5	24	48	11.0	977	basic lava	outcrop	basalt	---								
KM112	---	51	49	7.6	24	40	13.5	955	basic lava	outcrop	basalt	---								
KM113	---	51	57	45.9	24	29	23.0	709	basic lava	outcrop	basalt	weak weathering								
KM114	---	52	13	30.8	24	29	46.2	855	basic lava	outcrop	basalt	weak weathering								
KM115	---	52	28	58.5	24	38	35.7		basic lava	outcrop	basalt	---								
KM116	---	52	45	35.5	24	36	44.5	531	basic lava	outcrop	basalt	weathering								
KM117	---	52	19	42.5	24	6	10.7	584	basic lava	outcrop	basalt	---							1	
KM118	---	52	6	15.3	24	18	36.8	435	basic lava	outcrop	basalt	weak weathering								
KM119	---	52	10	9.0	24	7	34.6	700	basic lava	outcrop	basalt-andesite	weak weathering								
KM120	---	52	9	14.8	24	7	6.4	682	basic lava	outcrop	basalt	weak weathering								
KM121	---	51	59	46.8	24	1	8.4	385	basic lava	outcrop	basalt	weak weathering							1	
KM122	---	51	51	9.3	23	51	41.3	419	basic lava	outcrop	basalt	weak weathering								
KM123	---	51	49	13.9	23	42	59.6	499	basic lava	outcrop	basalt	weak weathering								
KM124	---	52	1	54.0	23	31	33.7	411	basic lava	outcrop	basalt	weathering								
KM125	---	51	39	57.7	23	35	58.1	734	basic lava	outcrop	basalt	---								
KM126	---	51	38	34.8	23	46	48.2	590	basic lava	outcrop	basalt	---								
KM127	---	51	30	13.8	23	59	15.9	622	basic lava	outcrop	basalt	---								
KM128	---	51	19	43.7	23	59	31.1	807	basic lava	outcrop	basalt	---								
KM129	---	51	7	18.0	23	41	46.2	792	basic lava	outcrop	basalt-picroite?	weak weathering								
KM130A	---	51	6	30.1	23	39	15.1	686	basic lava	outcrop	basalt	weak weathering								

Description of rock and drill core samples

Sample No.	Depth		Longitude		Latitude		Altitude	Sample Type		Rock Name	Description	Mineral-ization	Weathering	Analysis method						
	m		D	M	S	D		M	S					Bulk	PGE	Ar-Ar	Thin	P-Thin	X-ray	EPMA
KMI30B	---	51	6	30.1	23	39	15.1	686	basic intrusion	dolerite (dyke/sill)	dark gray-black, coarse grained, fresh	---	---	1						
KMI31	---	51	7	52.1	23	29	42.4	612	basic lava	basalt?	pale gray (wk greenish), compact, medium grain	---	weathering						1	
KMI32	---	51	18	27.2	22	55	58.5	356	basic lava	basalt	dark gray (brownish stained), fine, wk weathered, compact	---	weak weathering							
KMI33	---	51	15	1.7	22	51	28.0	480	basic lava	basalt	pale gray-gray (partly greenish dot), fine, wk weathered	---	weak weathering							
KMI34	---	51	11	34.1	22	51	3.7	387	basic lava	basalt	greenish dark gray, wk weathered	---	weak weathering							
KMI35	---	50	56	59.5	22	57	49.1	354	basic lava	basalt	greenish dark gray, compact, partly black glass included	---	---							
KMI36	---	50	48	38.7	22	54	30.8	395	basic lava	basalt	greenish dark gray, compact, fine, wk weathered	---	weak weathering							
KMI37	---	50	44	1.1	23	5	7.1	487	basic lava	basalt	black, compact and fresh, partly fine black glass included	---	---						1	
KMI38	---	50	40	43.2	23	13	51.8	469	basic lava	basalt	black, compact and fresh, fine	---	---							
KMI39	---	51	0	39.9	23	15	38.5	419	basic lava	basalt	brownish gray, wk amygdaloidal, wk weathered, white clay(zeolite?) in amygdali.	---	weak weathering							
KMI40	---	50	42	17.3	23	29	2.8	753	basic lava	basalt	dark gray(wk greenish), small green dot included, wk weathered, compact, m. grained	---	weak weathering							
KMI41	---	50	27	23.8	23	24	45.0	482	basic lava	basalt	dark gray, m.grained, partly green dot included	---	---							
KMI42	---	50	50	23.2	23	20	19.5	508	basic lava	basalt	black-dark gray, fresh and compact, small black glass included	---	---						1	
KMI43	---	50	2	20.0	23	14	21.0	542	basic lava	basalt	dark gray, compact, pyroxene(1mm) →partly clay	---	---							
KMI44	---	50	15	22.8	23	4	3.9	490	basic lava	basalt	black-dark gray, compact and fresh, m.grained	---	---						1	
KMI45	---	51	20	45.2	23	19	34.0	634	basic lava	basalt	reddish-greenish gray, m weathered, greenish dot included	---	weathering							
KMI46	---	51	22	5.0	23	39	41.2	799	basic lava	basalt	brownish gray(wk greenish), wk weathered, m.grained	---	weak weathering							
KMI47	---	51	38	11.6	24	3	31.8	455	basic lava	basalt	dark gray, fine grained, fresh and compact	---	---							
KMI48	---	51	40	9.2	24	22	13.7	1,100	basic lava	basalt	black, compact, very fine grained	---	---						1	
KMI49	---	51	43	48.6	24	51	45.1	974	basic lava	basalt	dark gray-black, fine grained, fresh and compact, partly black glass included	---	---							
KMI50	---	51	32	15.9	25	5	19.6	1,141	basic lava	basalt	dark gray(wk greenish), compact, fine, mafic mineral →partly clay	---	weak weathering							
KMI51	---	48	25	55.4	22	47	28.4	639	basic lava	basalt	dark gray, medium grained, very weak weathered.	---	weak weathering							
KMI52	---	48	50	20.1	22	34	29.3	556	basic lava	basalt	dark gray, massive, few amygdal, very weak weathered	---	weak weathering							
KMI53	---	48	53	36.2	22	18	16.2	557	basic lava	basalt	dark gray, aphyric, massive, fine-m.grained	---	---						1	
KMI54	---	48	59	16.0	21	46	1.0	435	basic lava	basalt	ditto	---	---							
KMI55	---	49	4	43.9	21	36	16.9	386	basic lava	basalt	ditto	---	---							
KMI56	---	49	47	51.0	21	17	26.3	354	basic lava	basalt	greenish gray, part amygdaloidal, greenish chalcocopy included	---	weathering							
KMI57	---	50	9	33.9	21	16	37.9	351	basic lava	basalt	greenish gray, m.grained, massive, greenish celadonite and opal included	---	weathering							
KMI58	---	51	19	30.5	20	45	18.5	317	basic lava	basalt	greenish gray, weak-m weathered, greenish dot included	---	weathering							
KMI59	---	51	22	22.8	20	30	10.1	325	basic lava	basalt	reddish gray, m. weathered, partly greenish dot included	---	weathering							
KMI60	---	51	21	12.9	20	23	51.2	347	basic lava	basalt	greenish gray, m. weathered	---	weathering							
KMI61	---	50	54	13.7	20	18	2.9	328	basic lava	basalt	dark gray, massive and compact, fresh, a few amygdal included	---	---						1	
KMI62	---	50	10	45.9	20	27	48.9	409	basic lava	basalt	dark gray, fine grained, glassy, fresh	---	---							
KMI63	---	50	9	2.6	20	51	13.9	393	basic lava	basalt	ditto	---	---						1	
KN001	---	51	2	27.0	29	49	16.0	130	basic intrusion	dolerite	sill, dark greenish gray, fine grain(glassy), phenocryst: pyroxene=plagioclase, olivine?	---	---						1	

Description of rock and drill core samples

Sample No.	Depth		Longitude		Latitude		Altitude		Sample Type	Rock Name	Description	Mineral-ization	Weathering	Analysis method						
	m		D	M	S	D	M	S						m	Bulk	PGE	Ar-Ar	Thin	P-Thin	X-ray
KN136	---	49 45	38.9	25	0	36.5			others	diorite	phenocryst: Hb, Pyx, K-FI, Qtz., alteration min.:Epi	---	weathering							
KN137	---	49 42	49.5	25	2	24.2			others	schist	mica	---	weathering							
KN138	---	49 37	34.7	25	11	21.8			others	granite	megacrystal of feldspar (several mm to several cm), Hb, Pyx.	---	---							
KN139	---	49 35	59.3	25	14	6.5			others	schist	pale green mica schist/white quartz schist	---	---							
KN140	---	49 33	41.0	25	16	29.4			others	schist	mica	---	---							
KN141	---	48 24	26.7	23	31	41.7			basic intrusion	gabbro	sill, dark greenish grey, medium grain	---	---							
KN142	---	48 23	15.2	23	28	35.6			basic intrusion	gabbro	sill, dark greenish grey, medium grain, pyrite disseminated	pyrite	weathering							
KN143	---	48 17	22.8	23	29	30.9			basic intrusion	dolerite	dyke, dark greenish grey, coarse grain	---	weak weathering							
KN144	---	48 56	25.1	23	7	57.1			basic lava	basalt	lava, dark greenish grey, fine grain	---	weathering							
KN145	---	49 3	47.1	23	19	14.4			basic lava	basalt	lava, dark greenish grey, fine grain	---	weathering							
KN146	---	49 14	18.3	23	31	18.0			basic intrusion	gabbro	sill, dark greenish grey, coarse grain, pyrite disseminated	pyrite	weathering							
KN147	---	49 19	53.9	23	28	25.3			sediment	siltstone	Rio do Rasto Fm. White-pale green- greyish pink, ripple mark, bedding flat.	---	weathering							
KN148	---	49 20	49.0	23	28	3.5			basic intrusion	dolerite	dyke, dark greenish grey, fine grain dyke in coarse grain dyke	---	weathering							
KN149	---	49 21	13.1	23	27	42.5			basic intrusion	dolerite	dyke, dark greenish grey,	---	weathering							
KN150	---	49 31	27.7	23	23	29.3			basic intrusion	gabbro	sill, dark greenish grey, medium grain, pyrite disseminated	pyrite	weathering							
KN151	---	49 29	9.6	23	22	15.1			basic intrusion	dolerite	dyke, light greenish grey, porphyritic	---	weathering							
KN152	---	49 14	15.8	23	12	42.0			basic lava	basalt (Acidic Rock)	lava, pale brownish grey, fine grain. (Acidic Rock)	---	weathering							
KN153	---	49 20	16.7	23	14	3.7			basic lava	basalt	lava, greenish grey, medium grain,	---	weathering							
KN154	---	49 27	5.2	23	16	7.6			basic lava	basalt	lava, dark greenish grey, fine grain	---	weathering							
KN155	---	49 28	9.7	23	18	42.0			sediment	sandstone	Botucou Fm., red, fine grain,	---	weathering							
KN156	---	49 58	38.0	22	51	58.3			basic lava	basalt	lava, greenish grey, fine grain, celadonite in cavities and cracks,	---	weathering							
KN157	---	50 29	28.2	22	36	17.7			basic lava	basalt	lava, pale greenish grey, fine grain	---	weathering							
KN158	---	50 34	4.8	22	30	9.7			basic lava	basalt	lava, grey, fine grain,	---	weathering							
KN159	---	51 3	28.2	23	3	33.7			basic lava	basalt	lava, greenish grey, medium grain, celadonite in cavities and cracks,	---	weathering							
KN160	---	51 26	47.4	23	50	5.4			basic intrusion	dolerite	dyke in basalt lava, dark greenish grey, medium grain	---	weathering							
KN161	---	50 51	55.1	24	10	23.6			basic intrusion	dolerite	dyke, dark greenish grey, medium grain	---	weathering							
KN162	---	50 48	9.5	24	11	2.2			basic intrusion	gabbro	sill, dark greenish grey, medium grain, pyrite disseminated	sulfide	weathering							
KN163	---	50 46	34.1	24	6	33.8			basic intrusion	dolerite	dyke, dark greenish grey	---	weathering							
KN164	---	50 43	51.1	24	3	35.1			basic intrusion	gabbro	dyke, dark greenish grey, pyrite disseminated	sulfide	weathering							
KN165	---	50 42	4.3	23	58	4.3			basic intrusion	dolerite	dyke, dark greenish grey, coarse grain	---	weathering							
KN166	---	50 40	20.3	23	56	2.1			basic intrusion	dolerite	dyke, dark greenish grey, coarse grain	---	weathering							
KN167B	---	50 36	54.0	23	55	12.0			basic intrusion	gabbro	dyke/sill, dark greenish grey, coarse grain (pegmatic), sulfide disseminated	sulfide	weathering							
KN168	---	50 38	12.2	23	47	38.7			acidic lava	dacite	relic of lava in the Paleozoic sequence, phenocryst of Qtz	---	weathering							
KN169	---	50 49	17.5	24	27	47.2			basic intrusion	gabbro	sill, dark greenish grey, sulfide disseminated	sulfide	weathering							
KN170	---	50 51	35.3	24	25	11.4			basic intrusion	gabbro	sill, dark greenish grey, medium grain	---	weathering							

Description of rock and drill core samples

Sample No.	Depth		Longitude			Latitude			Altitude	Sample Type		Rock Name	Description	Mineralization	Weathering	Analysis method						
	m		D	M	S	D	M	S		D	M					S	Bulk	PGE	Ar-Ar	Thin	P-Thin	X-ray
KN171	---	---	50	55	11.8	24	24	29.7	894	outcrop	gabbro	dark greenish grey, sulfide disseminated (?)	sulfide	weathering	1							
KN172	---	---	50	49	58.9	24	32	3.2	992	basic intrusion	gabbro	sill, dark greenish grey, fine grain	---	weathering	1							
KN173	---	---	50	52	3.9	25	0	55.3	791	basic intrusion	dolerite	dike, dark greenish grey, fine grain	---	weathering	1							
KN174A	---	---	50	56	29.1	25	2	4.1	738	basic intrusion	gabbro	sill, dark greenish grey, fine grain	---	weathering	1							
KN174B	---	---	50	56	29.1	25	2	4.1	738	sediment	mudstone	laminated, weak horizontal	---	---								
KN175	---	---	50	42	26.3	25	22	34.3	882	basic intrusion	gabbro	sill, dark greenish grey, fine grain	---	weathering	1							
KN176	---	---	50	30	28.7	25	37	33.3	957	basic intrusion	gabbro	sill, dark greenish grey, fine grain	---	weathering	1							
TM101	---	---	50	43	14.6	16	57	56.6	666	basic intrusion	gabbro?	dike, grey, coarse grain(glassy), phenocryst: pyroxene(max 1cm)+plagioclase olivine?	---	---	1							
TM102	---	---	51	30	29.5	16	43	57.5	683	basic intrusion	gabbro?	sill, dark greenish grey, coarse grain(glassy), phenocryst: pyroxene+plagioclase(max 5mm)	---	---	1							
TM103	---	---	51	37	46.9	16	49	47.6	713	basic intrusion	gabbro?	sill, dark greenish grey, coarse grain(glassy), phenocryst: pyroxene+plagioclase	---	---	1							
TM104	---	---	51	43	16.2	16	54	18.2	753	basic intrusion	dolerite	sill, dark greenish grey, medium grain(glassy), phenocryst: pyroxene+plagioclase	---	---	1							
TM105	---	---	51	52	54.3	16	52	18.4	720	basic intrusion	gabbro?	sill, dark greenish grey, coarse grain(glassy), phenocryst: pyroxene+plagioclase, olivine?	---	---	1							
TM106	---	---	51	54	33.5	16	50	28.7	731	basic intrusion	dolerite	sill, dark grey, medium grain(glassy), phenocryst: pyroxene+plagioclase	---	---	1							
TM107	---	---	52	0	22.4	17	36	27.8	672	basic intrusion	dolerite	sill, dark grey, medium grain(glassy), phenocryst: pyroxene+plagioclase	---	---	1							
TM108	---	---	52	3	38.9	17	29	1.2	775	sediment	calcareous rock	pale grey, calcareous rock, pyrite dissemination.	pyrite	---								
TM109	---	---	52	3	38.9	17	29	1.2	775	basic intrusion	dolerite	dike-sill, dark grey, medium-coarse grain(glassy), phenocryst: pyroxene+plagioclase	---	---	1							
TM110	---	---	52	3	38.9	17	29	1.2	775	basic intrusion	basalt	dike-sill, grey, fine grain(glassy), phenocryst: plagioclase, pyrite film	pyrite	---	1							
TM111	---	---	52	3	21.2	17	28	27.5	768	basic intrusion	dolerite?	dike-sill, grey, medium grain(glassy), phenocryst: pyroxene+plagioclase, (py. diss?)	pyrite	---								
TM112	---	---	52	3	21.2	17	28	27.5	768	basic intrusion	dolerite?	sill, grey, medium-fine grain(glassy), phenocryst: pyroxene+plagioclase, olivine?	---	---	1							
TM113	---	---	51	44	58.9	17	55	8.4	651	basic lava	basalt?	pale purple-grey, fine grain(glassy), phenocryst: plagioclase	---	---	1							
TM114	---	---	51	47	40.7	17	59	12.4	660	basic lava	basalt	p-brownish grey, fine grain(glassy), phenocryst: pyroxene+plagioclase	---	---	1							
TM115	---	---	51	39	25.6	17	56	6.2	729	basic lava	basalt	p-greenish grey, compact, fine-medium grain(glassy), phenocryst: pyroxene+plagioclase	---	---	1							
TM116	---	---	50	48	20.0	17	42	42.9	669	basic lava	basalt	p-greenish grey, compact, medium-fine grain(glassy), phenocryst: pyroxene+plagioclase	---	---	1							
TM117	---	---	50	41	30.5	17	33	32.1	660	basic intrusion	gabbro?	sill-lake?, pale grey, partly redish, coarse grain (glassy), phenocryst: pyroxene (0.5cm)+plagioclase	---	---	1							
TM118	---	---	50	33	30.9	17	32	31.1	533	basic lava	basalt	dark grey, compact, fine grain(glassy), phenocryst: pyroxene+plagioclase	---	---	1							
TM119	---	---	50	32	29.4	17	36	52.3	530	basic lava	basalt	dark grey, compact, fine grain(glassy), phenocryst: pyroxene+plagioclase	---	---	1							
TM120	---	---	50	33	14.8	17	52	46.3	496	basic lava	basalt	p-grey, compact, fine-medium grain(glassy), phenocryst: pyroxene+plagioclase	---	---	1							
TM121	---	---	50	42	30.4	17	52	37.1	573	basic lava	basalt	dark grey, compact, fine grain(glassy), phenocryst: pyroxene+plagioclase, film	sulfide	---							1	
TM122	---	---	49	31	46.5	18	0	28.4	526	basic lava	basalt	dark greenish grey, medium-fine grain(glassy), phenocryst: pyroxene+plagioclase	---	---	1							
TM123	---	---	49	32	25.4	17	47	10.5	722	basic lava	basalt	dark grey, compact, fine grain(glassy), phenocryst: pyroxene+plagioclase	---	---	1							
TM124	---	---	49	23	24.5	17	57	37.4	668	basic lava	basalt	dark grey, compact, fine grain(glassy), phenocryst: pyroxene+plagioclase	---	---	1							
TM125	---	---	49	16	5.3	18	4	7.0	827	basic lava	basalt	dark grey, compact, fine grain(glassy), phenocryst: pyroxene+plagioclase	---	---	1							
TM126	---	---	49	18	38.8	18	24	2.5	678	basic lava	basalt	dark grey, compact, fine grain(glassy), phenocryst: pyroxene+plagioclase	---	---	1							
TM127	---	---	49	1	23.9	18	30	34.3	650	basic lava	basalt?	dark grey, fine-medium grain(glassy), phenocryst: pyroxene+plagioclase	---	---	1							
TM128	---	---	48	19	24.5	18	53	26.5	801	basic lava	basalt	dark grey, compact, fine grain(glassy), phenocryst: pyroxene+plagioclase	---	---	1							

Description of rock and drill core samples

Sample No.	Depth		Longitude		Latitude		Altitude	Sample Type		Rock Name	Description	Mineral-ization	Weathering	Analysis method						
	m		D	M	S	D		M	S					Bulk	PGE	Ar-Ar	Thin	P-Thin	X-ray	EPMA
TM129	---	47	55	12.1	19	41	58.8	747	basic lava	outcrop	basalt	dark grey, compact, fine grain(glassy), phenocryst: pyroxene+plagioclase	---	---	1	1				
TM130	---	47	52	49.5	19	50	21.5	724	basic lava	outcrop	basalt	dark grey, compact, fine grain(glassy), phenocryst: pyroxene+plagioclase	---	---	1	1				1
TM131	---	47	47	57.5	19	57	30.3	589	basic lava	outcrop	basalt	dark greenish grey, compact fine-medium grain(glassy), phenocryst: pyroxene+plagioclase	---	---	1					
TM132	---	47	47	19.5	20	3	37.7	602	basic lava	outcrop	basalt	dark grey, compact, fine grain(glassy), phenocryst: pyroxene+plagioclase, pyrite fine?	pyrite	---	---		1			
TM133	---	47	28	20.0	20	9	11.2	939	basic lava	outcrop	basalt	dark grey, compact, fine grain(glassy), phenocryst: pyroxene+plagioclase	---	---	1					
TM134	---	47	16	53.0	20	37	12.0	714	basic intrusion	outcrop	dolerite?	sill, dark greenish grey, fine-medium grain(glassy), phenocryst: pyroxene+plagioclase	---	---	1					
TM135	---	47	18	49.3	20	41	36.3	695	basic lava	outcrop	dolerite	p-grey, medium grain(glassy), phenocryst: pyroxene+plagioclase	---	---	1					
TM136	---	47	19	14.2	20	34	47.1	923	basic lava	outcrop	basalt	dark grey, compact, fine grain(glassy), phenocryst: pyroxene+plagioclase	---	---	1					
TM137	---	47	18	49.4	21	6	1.3	667	basic intrusion	outcrop	dolerite	sill, dark greenish grey, medium grain(glassy), phenocryst: pyroxene+plagioclase	---	---	1					
TM138	---	47	17	39.8	21	10	12.2	762	basic intrusion	outcrop	dolerite	sill, dark greenish grey, medium grain(glassy), phenocryst: pyroxene+plagioclase	---	---	1					
TM139	---	47	19	42.3	21	20	34.5	720	basic intrusion	outcrop	dolerite	sill, dark grey, medium grain(glassy), phenocryst: pyroxene+plagioclase	---	---	1					
TM140	---	47	16	34.9	21	44	26.4	814	basic intrusion	outcrop	dolerite	sill, dark grey, medium-coarse grain(glassy), phenocryst: pyroxene+plagioclase	---	---	1					
TM141	---	47	18	28.2	21	50	50.2	644	basic intrusion	outcrop	dolerite?	sill, dark greenish grey, medium-coarse grain(glassy), phenocryst: pyroxene+plagioclase	---	---	1					
TM142	---	47	22	7.0	21	55	31.7	554	basic intrusion	outcrop	gabbro?	sill, dark greenish grey, coarse grain(glassy), phenocryst: pyroxene+plagioclase	---	---	1					
TM143	---	47	19	1.1	22	12	29.4	733	basic intrusion	outcrop	gabbro?	sill, dark grey, medium-coarse grain (glassy), phenocryst: pyroxene+plagioclase	---	---	1					
TM144	---	47	28	29.5	22	34	20.0	621	basic intrusion	outcrop	dolerite	sill, dark greenish grey, medium-coarse grain(glassy), phenocryst: pyroxene+plagioclase	---	---	1					
TM145	---	47	15	10.4	22	29	21.1	683	basic intrusion	outcrop	gabbro?	sill, dark greenish grey, medium-coarse grain(glassy), phenocryst: pyroxene+plagioclase	---	---	1					
TM146	---	47	10	28.3	22	25	14.6	614	basic intrusion	outcrop	dolerite	sill, dark grey, medium grain(glassy), phenocryst: pyroxene+plagioclase	---	---	1					
TM147	---	47	11	34.0	22	36	28.3	607	basic intrusion	outcrop	gabbro?	sill, dark greenish grey, medium-coarse grain(glassy), phenocryst: pyroxene+plagioclase	---	---	1					
TM148	---	47	21	24.6	22	36	51.8	534	basic intrusion	outcrop	gabbro?	sill, dark greenish grey, coarse grain(glassy), phenocryst: pyroxene+plagioclase	---	---	1					
TM149	---	47	35	36.1	21	42	18.8	700	basic lava	outcrop	basalt	dark grey, compact, medium-fine grain(glassy), phenocryst: pyroxene+plagioclase	---	---	1					
TM150	---	47	42	3.0	21	23	3.5	728	basic intrusion	outcrop	dolerite	sill, dark greenish grey, medium grain(glassy), phenocryst: pyroxene+plagioclase	---	---	1					
TM151	---	47	45	11.0	21	16	8.3	713	basic lava	outcrop	basalt	purple grey, compact, fine grain, weathering	---	---	1					
TM152	---	54	44	57.2	20	27	14.3	514	basic lava	outcrop	basalt	grey, compact, medium grain, weathering	---	---	1					
TM153	---	55	6	38.5	20	38	16.1	226	basic lava	outcrop	basalt	grey, compact, medium grain, weathering	---	---	1					1
TM154	---	54	21	50.9	19	46	57.2	614	basic lava	outcrop	basalt	brown, medium grain, weathering	---	---	1					
TM155	---	54	22	36.9	19	17	45.7	577	basic lava	outcrop	basalt	brownish grey, medium grain, weathering	---	---	1					
TM156	---	53	23	19.2	19	49	38.1	392	basic lava	outcrop	basalt	dark greenish grey, fine grain(glassy), phenocryst: pyroxene+plagioclase	---	---	1					
TM157	---	54	31	52.1	20	59	53.5	428	basic lava	outcrop	basalt	brownish grey, medium grain, weathering	---	---	1					
TM158	---	55	26	10.8	21	11	1.9	551	basic lava	outcrop	basalt	brown, medium grain, weathering	---	---	1					
TM159	---	55	37	33.1	21	10	36.1	417	basic lava	outcrop	basalt	grey, medium grain, weathering, phenocryst: pyroxene+plagioclase	---	---	1					
TM160	---	55	49	17.2	21	29	3.1	448	basic lava	outcrop	basalt	purple-grey, medium grain, weathering	---	---	1					
TM161	---	54	59	13.6	21	37	26.6	318	basic lava	outcrop	basalt	pale purple grey, medium grain, weathering	---	---	1					
TM162	---	54	32	1.3	21	54	37.4	291	basic lava	outcrop	basalt	pale greenish grey, fine grain(glassy), phenocryst: pyroxene+plagioclase	---	---	1					1
TM163	---	56	7	46.7	22	9	12.2	360	basic lava	outcrop	basalt	dark greenish grey, medium grain(glassy), phenocryst: pyroxene+plagioclase	---	---	1					

Description of rock and drill core samples

Sample No.	Depeth		Longitude			Latitude			Altitude		Sample Type	Rock Name	Description	Mineral-ization	Weathering	Analysis method						
	m		D	M	S	D	M	S	m	Bulk						PGE	Ar-Ar	Thin	P-Thin	X-ray	EPMA	Nd-Sr Isotope
WW130		---	50	2	58.0	28	7	40.7			basic lava	basalt	lava, grey, medium grain.	---	---	1						
WW134		---	50	11	33.3	27	51	57.5			basic lava	basalt	lava, grey, medium grain.	---	---	1						
WW-161		---	55	5	45.8	30	10	36.3			basic intrusion	gabbro	sill, gabbro, host: Rosario Fm.	---	weak weathering	1						
WW-162		---	55	17	3.1	30	5	15.5			basic intrusion	gabbro	sill, gabbro with native copper	native copper	weathering	1						
WW-163		---	55	34	11.0	29	53	56.3			basic lava	andesitic basalt	andesitic basalt?	---	weathering	1						
WW-164		---	55	47	1.3	29	48	19.5			acidic lava	acidic rock	acidic flow	---	weathering	1						
WW-165		---	55	38	25.8	29	45	54.0			acidic lava	acidic rock	acidic flow	---	weathering	1						
WW-166		---	55	33	28.3	29	44	29.6			acidic lava	acidic rock	acidic flow	---	weathering	1						
WW-167		---	55	58	9.6	29	35	53.1			basic lava	basalt	basalt or andesitic basalt	---	weathering	1						
WW-174		---	55	38	9.7	29	58	10.9			acidic lava	acidic rock	acidic flow	---	weak weathering	1						
WW-175		---	55	34	37.2	30	9	34.4			acidic lava	acidic rock	acidic flow	---	weathering	1	1					
WW-177		---	55	52	11.9	29	55	47.8			acidic lava	acidic rock	acidic flow	---	weathering	1						
WW-181		---	56	36	8.9	29	41	37.4			acidic lava	acidic rock	lava, acidic or intermediate?	---	weathering	1						
WW-183A		---	56	35	20.0	29	57	24.0			basic lava	basalt	autobrecciated lava, near bottom of one flow	---	weak weathering	1						
WW-184		---	56	27	51.7	30	0	20.6			acidic lava	acidic rock	lava, acidic rock	---	weathering	1	1					
WW-191		---	55	33	47.6	30	50	28.1			acidic lava	acidic rock	lava, near contact between lava and sandstone	---	weak weathering	1						
WW-196		---	56	12	58.9	30	29	12.4			basic lava	basalt	lava, near contact between lava and sandstone (Botucatu Fm.)	---	weak weathering	1						
WW-197		---	56	23	59.5	30	24	27.5			basic lava	basalt	lava.	---	weathering	1						
WW-203		---	55	51	24.0	30	52	43.3			acidic lava	acidic rock	lava, acidic rock.	---	weathering	1						
WW-204		---	56	29	9.9	30	18	9.2			acidic lava	acidic rock	lava, acidic rock.	---	weathering	1						
WW-205		---	56	33	2.6	30	8	25.7			acidic lava	acidic rock	lava, acidic rock, flow texture.	---	weathering	1						
WW-209		---	57	6	56.0	29	53	1.3			acidic lava	acidic rock	lava, acidic rock.	---	weathering	1						
WW-210		---	56	59	22.9	29	44	53.3			acidic lava	acidic rock	lava, acidic rock.	---	weak weathering	1						
WW-219		---	55	5	42.4	29	2	37.1			acidic lava	acidic rock	lava, acidic rock.	---	weak weathering	1						
WW-221		---	54	49	24.3	29	15	0.3			acidic lava	acidic rock	lava, acidic rock.	---	weak weathering	1						
WW-222		---	53	20	35.0	29	43	17.6			basic intrusion	gabbro	sill, gabbro, in sediment (Rosario Fm.)	---	weak weathering	1						
WW-223		---	53	0	0.1	29	44	44.7			basic intrusion	gabbro	sill, gabbro.	---	weak weathering	1						
AC-028		---	53	13	20.6	26	45	26.4			basic lava	basalt	lava, greenish grey, fine grain.	---	---	1						
AC-034B		---	53	30	5.0	26	52	28.3			basic lava	dolerite	lava, greenish grey, coarse grain, phenocryst: pyroxene+plagioclase.	---	---	1						
AC-035		---	53	29	59.1	27	1	40.6			basic lava	basalt	lava, reddish grey, fine grain.	---	---	1						
ACR-119		---	51	36	27.2	27	31	5.8			acidic lava	basalt	lava, reddish grey, fine grain. (Acidic Rock)	---	---	1						
ACR-125		---	52	35	54.1	27	4	5.3			basic lava	basalt	lava, grey, fine grain, phenocryst: pyroxene+plagioclase.	---	---	1						
ADR-033		---	52	5	9.8	26	44	36.7			basic lava	basalt	lava, greenish grey, fine grain, weak weathered(hematite), rare amygdaloidal texture.	---	weak weathering	1						
YM105		---	54	57	28.8	28	33	51.4		181	basic lava	basalt	lava, massive, greenish grey, medium grain, phenocryst: plagioclase+pyroxene, amygdaloidal texture, weathering.	---	weak weathering	1						
YM108		---	55	15	46.4	28	11	44.0		130	basic lava	basalt	lava, massive, grey, fine grain, aphyritic, reddish dots (hematite).	---	weathering	1	1					

Description of rock and drill core samples

Sample No.	Depth		Longitude		Latitude		Altitude		Sample Type		Rock Name	Description	Mineral-ization	Weathering	Analysis method						
	m		D	M	S	D	M	S	m						Bulk	PGE	Ar-Ar	Thin	P-Thin	X-ray	EPMA
YM151	---	50	10	47.1	29	46	39.6	15	basic lava	outcrop	basalt (Acidic Rock)	lava, massive, greenish grey, medium grain, vesicle (common), weak weathering. (Acidic Rock)	---	weak weathering							
YM152	---	50	10	47.1	29	46	39.6	45	basic intrusion	outcrop	basalt (Acidic Rock)	feeder dyke?, black, glassy. (Acidic Rock)	---	---							
YM153	---	50	9	38.3	29	25	29.5	100	basic lava	outcrop	basalt	lava, massive, greenish grey, medium grain, phenocryst: plagioclase-olivine, a few vesicles, weak weathering.	---	weak weathering							
YM154	---	50	11	31.1	29	22	56.7	185	basic lava	outcrop	basalt	lava, massive, greenish grey, aphyric, a lot of vesicles, weak weathering.	---	weak weathering							
YM155	---	50	11	3.0	29	22	25.4	415	basic lava	outcrop	basalt	lava, massive, grey, fine grain, phenocryst: plagioclase-olivine, a few vesicles.	---	---							
YM156	---	50	10	52.0	29	22	31.2	500	basic intrusion	outcrop	basalt	sill?, massive, greenish grey, porphyritic?, phenocryst: plagioclase-olivine, including of fine grained sulfide.	sulfide	---							
YM157	---	50	10	48.9	29	22	11.6	610	basic lava	outcrop	basalt (Acidic Rock)	lava, massive, black, glassy. (Acidic Rock)	---	---							
YM158	---	50	10	18.4	29	21	18.1	770	basic lava	outcrop	basalt	lava, massive, reddish grey, a few vesicles, weak weathering.	---	weak weathering							
YM159	---	50	11	6.1	29	20	8.7	850	basic lava	outcrop	dacite	lava, low banding, grey, a few vesicles, weak weathering.	---	weak weathering							
YM160	---	50	11	6.1	29	20	8.7	850	basic lava	outcrop	basalt (Acidic Rock)	lava, massive, black, glassy, a few vesicles. (Acidic Rock)	---	---							
YM161	---	49	57	6.6	28	47	59.0	1,170	basic intrusion	outcrop	basalt	feeder dyke, fine pitch columnar joint, greenish grey, medium grain, phenocryst: plagioclase, weak weathering.	---	weak weathering							
YM162	---	49	56	57.1	28	48	1.9	1,090	basic lava	outcrop	basalt (Acidic Rock)	lava, massive, black, glassy, a few vesicles. (Acidic Rock)	---	---							
YM163	---	49	56	46.4	28	48	19.8	940	basic lava	outcrop	basalt	lava, massive, greenish grey, fine grain, including of fine grained native copper (rare), weak weathering.	native copper	weak weathering							
YM164	---	49	56	29.1	28	48	23.3	855	basic lava	outcrop	basalt	lava, massive, greenish grey, fine grain, weathering.	---	weathering							
YM165	---	49	56	4.2	28	48	17.0	750	basic lava	outcrop	basalt	lava, massive, reddish grey, fine grain, a few vesicles, fine grained native copper, weathering.	native copper	weathering							
YM166	---	49	55	46.1	28	48	26.0	630	basic lava	outcrop	basalt	lava, massive, grey, fine grain, weak weathering.	---	weak weathering							
YM167	---	49	55	27.9	28	48	12.8	440	basic lava	outcrop	basalt	lava, massive, grey, fine grain, a few vesicles, native copper (rare), weak weathering.	native copper	weak weathering							
YM168	---	49	55	13.8	28	48	27.4	370	basic lava	outcrop	basalt	lava, massive, grey, fine grain (aphyric), a few vesicles, weak weathering.	---	weak weathering							
YM169	---	49	22	18.4	28	45	1.3	130	basic intrusion	outcrop	basalt	sill, fine pitch columnar joint, fine grain, dots of sulfide along crack.	sulfide	---							
YM170	---	49	26	46.6	28	51	24.2	50	basic intrusion	outcrop	dolerite	sill, fine pitch columnar joint, grey, phenocryst: plagioclase-cpx-olivine, dots of sulfide (along crack).	sulfide	---							
YM171	---	49	26	46.6	28	51	24.2	60	basic intrusion	outcrop	dolerite	sill, fine pitch columnar joint, grey, phenocryst: plagioclase-cpx-olivine, dots of sulfide (along crack).	sulfide	---							
YM172	---	49	26	46.6	28	51	24.2	70	basic intrusion	outcrop	dolerite	sill, fine pitch columnar joint, grey, phenocryst: plagioclase-cpx-olivine, dissemination of sulfide.	sulfide	---							
YM173	---	49	26	46.6	28	51	24.2	100	basic intrusion	outcrop	gabro	sill, gabbroic part, coarse grain, plagioclase-cpx, dissemination of sulfide.	sulfide	---							
YM174	---	49	26	46.6	28	51	24.2	80	basic intrusion	outcrop	dolerite	sill, fine pitch columnar joint, grey, phenocryst: plagioclase-cpx-olivine, dissemination of sulfide.	sulfide	---							
YM175	---	49	26	46.6	28	51	24.2	100	basic intrusion	outcrop	dolerite	sill, fine pitch columnar joint, grey, phenocryst: plagioclase-cpx-olivine, dissemination of sulfide.	sulfide	---							
YM176	---	49	24	6.8	28	30	32.0	515	basic intrusion	outcrop	gabro	sill (gabbroic part), massive, grey, coarse grain, plagioclase-cpx, dissemination of sulfide.	sulfide	---							
YM177	---	49	12	55.1	28	30	16.7	168	basic intrusion	outcrop	basalt	dyke?, massive, grey, medium grain, phenocryst: plagioclase.	---	---							
YM178	---	49	31	38.0	28	23	38.5	760	basic lava	outcrop	basalt	lava, massive, medium grain, phenocryst: plagioclase, a few vesicles, weathering.	---	weathering							
YM179	---	49	32	13.7	28	23	34.9	905	basic lava	outcrop	basalt	lava, massive, fine grain, phenocryst: plagioclase-olivine?, a few vesicles, weak weathering.	---	weak weathering							
YM180	---	49	32	26.4	28	23	42.0	1,000	basic lava	outcrop	basalt	lava, massive, grey, fine-medium grain, including of vesicles, weak weathering.	---	weak weathering							
YM181	---	49	32	42.6	28	23	47.0	1,075	basic lava	outcrop	basalt	lava, massive, fine grain, including of vesicles, weak weathering.	---	weak weathering							
YM182	---	49	39	0.4	28	6	40.8	1,374	basic lava	outcrop	basalt	lava, massive, greenish grey, fine grain, phenocryst: cpx, including of vesicles, weak weathering.	---	weak weathering							
YM183	---	49	38	18.6	28	5	16.7	1,495	basic intrusion	outcrop	dolerite	dyke, fine pitch columnar joint, phenocryst: cpx-plagioclase, sulfide along crack.	sulfide	---							
YM184	---	49	36	29.4	28	1	51.8	1,275	basic lava	outcrop	basalt	lava, massive, grey, fine-medium grain, phenocryst: cpx-plagioclase, a few vesicles, weak weathering.	---	weak weathering							
YM185	---	49	34	16.0	27	54	42.4	1,195	basic intrusion	outcrop	dolerite	sill, massive, greenish grey, dissemination of sulfide, weak weathering.	sulfide	weak weathering							

Description of rock and drill core samples

Sample No.	Depth		Longitude				Latitude				Altitude	Sample Type		Rock Name	Description	Mineralization	Weathering	Analysis method								
	m		D	M	S	D	M	S	D	M		S	Bulk					PGE	Ar-Ar	Thin	P-Thin	X-ray	EPMA	Nd-Sr Isotope	S-Isotope	NRM
YM186	---	---	49	3	26.0	27	41	35.8	980		basic intrusion	outcrop	basalt	sil, massive, greenish grey, medium grain, phenocryst: plagioclase+cpx+olivine?, weak weathering.	---	weak weathering	1									
YM187	---	---	48	34	9.1	27	45	29.9	20		basic intrusion	outcrop	gabro (Acidic Rock)	dike, width: >10m, leucocratic, plagioclase+cpx+olivine, fine grained sulfide. (Acidic Rock)	sulfide	---	1									
YM188	---	---	48	33	38.1	27	47	38.5	25		basic intrusion	outcrop	gabro	dike, massive, width: 20m, merocratic, plagioclase+cpx+olivine, with fine grained sulfide.	sulfide	---	1	1								
YM189	---	---	48	30	18.9	27	43	24.4	22		basic intrusion	outcrop	gabro	dike, massive, width: 10m, NAOE, merocratic, plagioclase+cpx+olivine, with sulfide.	sulfide	---	1									
YM190	---	---	48	26	5.0	27	35	49.4	85		basic intrusion	outcrop	gabro	dike, massive, width: 3m, merocratic, plagioclase+cpx+olivine, with sulfide.	sulfide	---	1									
YM191	---	---	48	25	6.1	27	26	28.8	81		basic intrusion	outcrop	gabro	dike, massive, merocratic, width: >50m, plagioclase+cpx+olivine, with sulfide, weak weathering.	sulfide	weak weathering	1									
YM192	---	---	49	31	41.4	28	23	32.3	810		basic intrusion	outcrop	dolerite	dike, massive, reddish grey, coarse grain, phenocryst: plagioclase+cpx, a few vesicles, weathering.	---	weathering	1									
YM193	---	---	49	32	48.8	28	23	45.8	1,120		basic lava	outcrop	basalt	lava, massive, greenish grey, fine grain, a few vesicles, weathering.	---	weak weathering	1									
YM194	---	---	49	32	54.4	28	23	48.9	1,165		basic lava	outcrop	basalt	lava, massive, greenish grey, fine grain, a few vesicles, weathering.	---	weak weathering	1									
YM195	---	---	49	32	54.1	28	23	56.4	1,255		basic intrusion	outcrop	basalt	dike, width: >100m, black, glassy, weak weathering.	---	weak weathering	1									
YM197	---	---	49	33	3.1	28	23	57.1	1,330		basic lava	outcrop	basalt	lava, massive, greenish grey, fine grain, aphyritic, magnetite (common), a lot of vesicles, weathering.	---	weak weathering	1									
YM198	---	---	49	43	20.5	28	18	21.0	1,300		basic intrusion	outcrop	dolerite	dike, fine pitch columnar joint, grey, phenocryst: plagioclase+cpx+olivine.	---	weak weathering	1	1								
YM199	---	---	49	48	42.9	28	14	25.4	1,480		basic lava	outcrop	basalt	lava, grey, fine grain, phenocryst: plagioclase+cpx+olivine?, a lot of vesicles, weak weathering.	---	weak weathering	1									
YM200	---	---	50	31	7.5	27	44	3.4	785		basic lava	outcrop	basalt	lava, massive, grey, fine-medium grain, fine grained native copper (rare), weak weathering.	native copper	weak weathering	1									
YM201	---	---	50	23	38.2	27	26	34.9	872		basic intrusion	outcrop	basalt	sil or dike, grey, medium-course grain, phenocryst: plagioclase+cpx+olivine?, weak weathering.	---	weak weathering	1									
YM205	---	---	50	4	36.5	27	16	55.8	605		basic intrusion	outcrop	basalt	sil, fine pitched columnar joint, fine grain, a lot of sulfide along crack.	sulfide	---	1									
YM206	---	---	50	4	36.5	27	16	55.8	605		basic intrusion	outcrop	gabro	sil, gabbroic part of YM205, plagioclase+cpx+olivine, weak weathering.	---	weak weathering	1									
YM207	---	---	50	8	35.7	26	58	1.0	570		basic intrusion	outcrop	basalt	sil, thickness: 10-15m, fine pitched columnar joint, grey, medium grain, phenocryst: plagioclase+cpx+olivine, sulfide along crack.	sulfide	---	1									
YM208	---	---	50	25	59.7	26	56	10.0	1,070		basic lava	outcrop	basalt	lava, massive, grey, fine grain, a few vesicles, weak weathering.	---	weak weathering	1	1								
YM210	---	---	50	20	24.1	26	48	12.9	1,275		basic lava	outcrop	basalt	lava, massive, greenish grey, fine grain, weak weathering.	---	weak weathering	1	1								
YM211	---	---	50	32	15.6	26	4	22.2	755		basic intrusion	outcrop	dolerite	sil, thickness: 10-15m, massive, grey, porphyritic, phenocryst: cpx.	---	weak weathering	1									
YM212	---	---	50	32	15.6	26	4	22.2	755		basic intrusion	outcrop	basalt	sil (child margin), blackish green, fine grain, dissemination of sulfide, weathering. (Acidic Rock)	sulfide	---	1	1								
YM215	---	---	51	34	37.0	26	35	20.1	1,325		basic intrusion	outcrop	basalt	dike, massive, greenish grey, fine grain, a few vesicles, weak weathering. (Acidic Rock)	---	weak weathering	1									
YM218	---	---	52	24	30.1	26	46	58.1	710		basic intrusion	outcrop	dolerite	sil (lava?), massive, coarse grain, phenocryst: plagioclase+cpx.	---	weak weathering	1									
YM219	---	---	52	20	48.0	27	4	42.5	585		basic lava	outcrop	basalt	lava, massive, greenish grey, fine grain, weak weathering.	---	weak weathering	1									
YM222	---	---	50	42	23.0	26	55	26.2	1,010		basic lava	outcrop	basalt	lava, massive, greenish grey, fine grain, weathering.	---	weathering	1									
YM223	---	---	51	13	11.9	27	2	49.7	1,025		basic lava	outcrop	basalt	lava, massive, greenish grey, fine grain, weak weathering.	---	weak weathering	1	1								
YM225	---	---	51	34	26.5	27	15	32.5	600		basic lava	outcrop	dolerite	lava, massive, grey, coarse grain, phenocryst: plagioclase+cpx, dots of native copper, weak weathering.	native copper	weak weathering	1									
YM227	---	---	52	2	3.2	27	25	53.0	625		basic lava	outcrop	basalt	lava, massive, greenish grey, fine grain, a few vesicles, weak weathering.	---	weak weathering	1									
YM228	---	---	52	18	24.8	27	42	31.4	785		basic lava	outcrop	basalt?	lava, massive, greenish grey, green band (flow band?), fine grain.	---	weak weathering	1				1					
YM229	---	---	52	15	42.5	27	52	56.8	651		basic intrusion	outcrop	basalt	dike, fine pitched columnar joint, greenish grey, medium grain, weathering.	---	weathering	1									
YM230	---	---	51	48	8.3	27	55	58.2	800		basic lava	outcrop	basalt	lava, massive, greenish grey, fine grain, weak weathering.	---	weak weathering	1									
YM231	---	---	51	39	19.5	27	47	14.7	780		basic lava	outcrop	basalt	lava, massive, greenish grey, medium grain, phenocryst: olivine?, weak weathering.	---	weak weathering	1						1			
AT03-486.0	486.0	---	49	45	19.6	29	11	34.1	---		basic intrusion	drill core	drill core	native copper in cavity, chalcocitic quartz druse	native copper	---	weak weathering	1				1				
AT03-486.3	486.3	---							---		sediment	drill core	block shale	black brecciated shale with pyrite	pyrite	---	weak weathering	1				1				1

Description of rock and drill core samples

Sample No.	Depth		Longitude			Latitude			Altitude	Sample Type	Rock Name	Description	Mineral-ization	Weathering	Analysis method							
	m		D	M	S	D	M	S							D	M	S	Bulk	PGE	Ar-Ar	Thin	P-Thin
TC95-709.4	709.4									drill core	basalt	sill, black, fine			1							
TC95-725.0	725.0									drill core	basalt	sill, black, fine			1							
TC95-743.0	743.0									drill core	basalt	sill, black, fine			1							
TC95-756.3	756.3									drill core	basalt	sill, black, fine			1						1	
TC95-772.8	772.8									drill core	basalt	sill, very fine, chilled margin			1							
TC95-825.7	825.7									drill core	shale	black shale of Irati Fm.										
TC97-563.2	563.2	50	7	46.6	29	51	19.2			sediment	siltstone	pyrite rich silt stone, pyrite conc. In fracture	pyrite									
TC97-569.2	569.2									drill core	dolerite	sill, black, very fine, chilled margin			1							
TC97-590.0	590.0									drill core	dolerite	sill, black, fine, pyroxene(1mm) included			1							
TC97-602.0	602.0									drill core	dolerite	sill, black, fine, pyroxene(1mm) included			1							
TC97-615.0	615.0									drill core	dolerite	sill, black, fine, pyroxene(1mm) included			1							
TC97-625.0	625.0									drill core	dolerite	sill, black, fine, pyroxene(1mm) included			1							
TC97-630.0	630.0									drill core	dolerite	sill, black, fine, pyroxene(1mm) included			1							
TC97-650.0	650.0									drill core	dolerite	sill, black, fine, pyroxene(1mm) included			1							
TC97-664.0	664.0									drill core	dolerite	sill, black, very fine, possible chilled margin			1							
FP-02	125.20-125.40	50	35	43.6	23	50	15.2			drill core	dolerite	dark gray, dolerite			1							
FP-02	580.15-580.40									drill core	dolerite	dark gray, fine, dolerite, py-network	pyrite		1							
FP-02	582.80-582.95									drill core	dolerite	dark gray, fine, dolerite			1							
FP-02	586.15-586.35									drill core	dolerite	dark gray, dolerite, chilled margin, py-diss(dot)	pyrite		1							
FP-02	596.60-596.75									drill core	dolerite	dark gray, dolerite			1							
FP-02	604.60-604.80									drill core	dolerite	dark gray, coarse grained dolerite(gabbroic), py-diss	pyrite		1							
FP-02	605.55-605.70									drill core	dolerite	dark gray, coarse grained dolerite(gabbroic), py-diss	pyrite		1							
FP-09	24.75-24.95	50	57	20.0	24	32	43.0			drill core	dolerite	greenish gray, amygdaloidal, chilled margin(fine)			1							
FP-09	173.20-173.40									drill core	dolerite	dark gray, chilled margin (fine)			1							
FP-09	174.10-174.30									drill core	hornfels	white foot wall hornfels, py-net, native copper?	pyrite, native copper		2							
SP-21	129.80-130.00	50	30	26.1	23	46	58.7			drill core	dolerite	dark gray fine dolerite(chilled margin), py-diss	pyrite		1							
SP-33	74.50-74.70	50	27	54.4	23	46	29.4			drill core	dolerite	dark gray, fine, dolerite, wk py-diss	pyrite		1							
SP-51	171.40-172.05	50	30	18.7	23	46	37.2			drill core	shale	Irati Py-conc	pyrite		1							
SP-64	76.65-76.75	50	28	39.3	23	44	37.7			drill core	dolerite	black fine dolerite, py-diss (film)	pyrite		1							
SP-64	133.25-133.45									drill core	dolerite	black fine dolerite, py-diss (film), py-diss along crack	pyrite		1							
SP-64	138.15-138.30									drill core	shale	Irati Py-conc	pyrite		1							
SP-65	138.40-138.60	50	28	39.2	23	45	10.2			drill core	dolerite	black, fine and compact dolerite, chilled margin, py-diss along crack	pyrite		1							
SP-66	603.35-603.50	50	36	46.0	23	47	10.0			drill core	dolerite	black, dolerite			1							
SP-66	776.50-776.70									drill core	dolerite	dark gray, dolerite, py-diss(dot)	pyrite		1							
SP-66	777.50-777.65									drill core	dolerite	dark gray, fine, dolerite, py-diss along crack, chilled margin	pyrite		1							

Description of rock and drill core samples

Sample No.	Depth		Longitude			Latitude			Altitude	Sample Type	Rock Name	Description	Mineral-ization	Weathering	Analysis method																									
	m		D	M	S	D	M	S							m	Bulk	PGE	Ar-Ar	Thin	P-Thin	X-ray	EPMA	Nd-Sr Isotope	S-Isotope	NRM															
PP-02-SP	1240-1250											drill cuttings	basic lava		1																									
PP-02-SP	1280-1290											drill cuttings	basic lava		1																								1	
PP-02-SP	1320-1330											drill cuttings	basic lava		1																									
PP-02-SP	1380-1390											drill cuttings	basic lava		1																								1	
PP-02-SP	1420-1430											drill cuttings	basic lava		1																									
PP-02-SP	1460-1470											drill cuttings	basic lava		1																								1	
PP-02-SP	1500-1510											drill cuttings	basic lava		1																									
PP-02-SP	1540-1550											drill cuttings	basic lava		1																								1	
PP-02-SP	1590-1600											drill cuttings	basic lava		1																								1	
RP-03-SP	220-230		49	22	23.7							drill cuttings	basic lava		1																									
RP-03-SP	260-270											drill cuttings	basic lava		1																									
RP-03-SP	400-410											drill cuttings	basic lava		1																									
RP-03-SP	440-450											drill cuttings	basic lava		1																									
RP-03-SP	480-490											drill cuttings	basic lava		1																									
RP-03-SP	600-610											drill cuttings	basic lava		1																									
RP-03-SP	640-650											drill cuttings	basic lava		1																									
RP-03-SP	680-690											drill cuttings	basic lava		1																									
RP-03-SP	800-810											drill cuttings	basic lava		1																									
CA-85	432.20-432.70		50	40	46	29	43	50.9				drill core	basic intrusion		1																									
CA-85	473.00-473.50											drill core	basic intrusion		1																									
CA-85	522.50-522.95											drill core	basic intrusion		1																									
CA-85	637.40-637.90											drill core	basic intrusion		1																									
CA-85	666.15-666.65											drill core	basic intrusion		1																									
CA-86	496.25-496.75		50	33	39.7							drill core	basic intrusion		1																									
CA-86	516.50-517.00											drill core	basic intrusion		1																									
CA-86	548.75-549.25											drill core	basic intrusion		1																									

Analysis type

Bulk: Whole rock analysis (including of Au, Pt and Pd assay)

PGE: Measurement of Ir, Rh and Re contents (ICP)

Ar-Ar: ^{40}Ar - ^{39}Ar dating

Thin: Microscopic observation of thin section

P-thin: Microscopic observation of polished-thin section

X-ray: Powder X-ray diffraction

EPMA: Electron microprobe analysis

Nd-Sr isotope: Measurement of $^{143}\text{Nd}/^{144}\text{Nd}$ and $^{87}\text{Sr}/^{86}\text{Sr}$ isotopic ratioS-isotope: Measurement of $\delta^{34}\text{S}$ isotopic ratio

NRM: Measurement of natural remanent magnetism

APPENDIX 5

Microscopic observation of thin and polished-thin section for rock samples

(1/9)

Sample No.	Sample Type	Rock Name	Magma Type	Mg#	Minerals													Note							
					pl	am	opx	cpx	ol	qz	chl	ca	serp	mt	chr	sul	opaque pyrr		py	cp					
AS001	lava	cpx basalt	Gramado	72.3	⊙		•	○							△	•?									
AS002	lava	ol bearing cpx basalt	Paramapanema	60.6	⊙		•	○							△	•?								coarse basalt(or dolerite)	
AS003	lava	ol bearing cpx basalt	Ribeira	70.5	⊙		•	○	•						△	•?								contact with cpx dolerite	
AS004A	lava	ol bearing cpx basalt	Paramapanema	61.4	⊙		•	○	•						△	•?									
AS005	lava	ol bearing cpx basalt	Paramapanema	64.3	⊙		•	○	•						△	•?									
AS006	lava	ol bearing cpx basalt	Paramapanema	64.1	⊙		•	○	•						△	•?									
AS007	lava	ol bearing cpx basalt	Paramapanema	63.6	⊙		•	○	•						△	•?									
AS008	lava	ol cpx dolerite	Pitanga	66.7	⊙		•	○	•						△	•?									
AS009	lava	ol bearing cpx basalt	Paramapanema	59.5	⊙		•	○							△	•?								ferric hydroxide	
AS010	dyke	cpx dolerite	Paramapanema	61.0	⊙		•	○							△	•?									
AS011	dyke	cpx dolerite	Paramapanema	59.8	⊙		•	○							△	•?								ol(idingsite) bearing?	
AS013	dyke	cpx gabbro	Paramapanema?	52.1	⊙		•	○							△	•?									
AS014	sill	ol bearing cpx gabbro	unclassify	43.1	⊙		•	○	•						△	•?								needle shape cpx(1cm × 0,2mm)	
AS015A	dyke	ol bearing cpx dolerite	Acidic rock	23.1	⊙		•	○	•						△	•?									
AS015B	dyke	cpx basalt	Pitanga	55.2	⊙		•	○	•?						△	•?									
AS016A	dyke	cpx dolerite	Paramapanema-Ribeira	63.5	⊙		•	○							△	•?									
AS016B	dyke	cpx dolerite	Paramapanema-Ribeira	55.6	⊙		•	○							△	•?									
AS017	dyke	cpx dolerite	unclassify	47.8	⊙		•	○							△	•?									
AS018	sill	cpx dolerite	Paramapanema	60.4	⊙		•	○							△	•?									
AS021	lava	ol bearing cpx basalt	Gramado	66.0	⊙		•	○	•						△	•?								coarse basalt	
AS023	lava	cpx basalt	Gramado?	53.5	⊙		•	○							△	•?								partly brown glass matrix	
AS024A	lava	cpx basalt	Esmeralda	70.1	⊙		•	○							△	•?									
AS024B-1	dyke	cpx basalt (fragment)	unclassify	79.4	⊙		•	○							△	•?								intergranular texture	
AS024B-2	dyke	cpx ol basalt (matrix)	unclassify	79.4	⊙		•	○	△	•					△	•?								glass matrix with phenocrysts	
AS025	lava	weathered cpx basalt	unclassify	61.5	⊙		•	○							△	•?								partly glass matrix	
AS026	(Lajes district)	weathered basalt	kimberlite	80.5	⊙		•	△							△	•?								hyaloclastite	
AS027	(Lajes district)	cpx basalt	kimberlite	66.8	⊙		•	○							△	•?								ferric hydroxide	
AS028A	(Lajes district)	trachy andesite	phonolite	13.7	⊙			○							△	•								alkaline feldspar and alkaline amphibole	
KM104	lava	cpx basalt	Esmeralda	67.5	⊙										△	•								cryptocrystalline	
KM111	lava	ol-bearing cpx basalt	Paramapanema-Ribeira	67.0	⊙			△	○	•					△	•								intergranular texture, iddingsite	

Microscopic observation of thin and polished-thin section for rock samples

(5/9)

Sample No.	Sample Type	Rock Name	Magma Type	Mg#	Minerals										Note			
					pl	am	opx	cpx	ol	qz	chl	ca	serp	opaque				
											mt	chr	sul	pyrr	py	cp		
YM176	sill	cpx gabbro	Gramado?	40.9	☉													glass-cryptocrystalline matrix, native copper (3 grains)
YM179	lava	cpx basalt	Urubici	63.0	☉			△										
YM188	dyke	cpx dolerite	Urubici?	61.3	☉													
YM198	dyke	cpx basalt(-dolerite)	Esmeralda	66.9	☉													intergranular texture, iddingsite
YM211	sill	cpx basalt	Urubici	62.2	☉													intergranular texture, iddingsite
YM212	sill	cpx basalt	Urubici	63.2	☉													intersertal texture, iddingsite
YM228	lava	cpx basalt	Esmeralda	55.6	☉													intergranular texture
YM231	lava	cpx basalt	Esmeralda?	58.6	☉													intergranular texture
WW017	lava	aphilic andesite	Acidic rock	45.9	☉													iddingsite
WW024	lava	cpx basalt	Gramado	63.5	☉								•?					
WW026	lava	andesite	Acidic rock	42.9	☉			△					•?					
WW031	lava	cpx basalt	Esmeralda	57.7	☉								•?					
WW056	lava	cpx basalt	Parapanama	64.4	☉													ferric hydroxide (reddish)
WW068	lava	sandstone	Acidic rock	59.3	☉													medium grain size, rock fragment is absent
WW069	lava	cpx basalt	Acidic rock	49.5	☉													ferric hydroxide (reddish)
WW073	lava	cpx basalt	Gramado	48.3	☉			△										fine
WW076	lava	ol-cpx dolerite	Gramado	70.9	☉													
WW077	lava	ol-cpx dolerite	Gramado	73.1	☉													
WW083	lava	cpx basalt	Gramado	58.6	☉									•?				
WW092	lava	ol bearing cpx basalt	Esmeralda	69.1	☉									•?				ferric hydroxide (reddish)
WW095	lava	ol bearing cpx basalt	Esmeralda	65.5	☉									•?				ferric hydroxide (reddish)
WW099	lava	ol bearing cpx basalt	Gramado	75.1	☉													
WW117a	lava	cpx basalt	Gramado	65.6	☉													ferric hydroxide (reddish)
WW122	lava	ol bearing cpx basalt	Urubici	59.0	☉													
WW129	lava	cpx basalt	Gramado	63.7	☉													ferric hydroxide (reddish)
WW130	lava	cpx basalt	Esmeralda	71.0	☉													
WW134	lava	cpx basalt	Gramado	66.3	☉													
AC-028	lava	cpx basalt	Ribeira	70.0	☉													green clay mineral(malachite?)
AC-034B	lava	cpx dolerite	Parapanama	63.4	☉													green clay mineral(malachite?)
AC-035	lava	cpx basalt	Parapanama	59.4	☉													ferric hydroxide (reddish)

Microscopic observation of thin and polished-thin section for rock samples

Sampe No.	Sample Type	Rock Name	Magma Type	Mg#	Minerals										Note				
					pl	am	opx	cpx	ol	qz	chl	ca	serp	mt		chr	sul	pyrr	py
ACR-119	lava	cpx basalt	Acidic rock	43.6	☉				○						△				partly spinifex-like texture,
ACR-125	lava	cpx olivine dolerite	Pitanga	55.7	☉				△						△	•?			
ADR-033	lava	cpx basalt	Gramado	61.0	☉				○						△	•?			green clay mineral(malachite?)
5AT03-SC-486.3	sill (drill core)	basalt with shear plane	Gramado?	66.3	○				•						•	•?			
5AT03-SC-487.0	sill (drill core)	cpx basalt	Gramado	62.7	○				•						△	•?			needle shape opaque,
5AT03-SC-498.5	sill (drill core)	cpx gabbro	Gramado	50.6	☉				○						△	•?			
5AT03-SC-509.4	sill (drill core)	cpx gabbro	Gramado	48.9	☉				○						△	•?			
5AT03-SC-519.2	sill (drill core)	cpx gabbro	Gramado	52.7	☉				○						△	•?			
5AT03-SC-528.9	sill (drill core)	cpx gabbro	Gramado	51.1	☉				○						△	•?			
5AT03-SC-537.3	sill (drill core)	cpx gabbro	Gramado	50.5	☉				○	•					△	•?			
5AT03-SC-547.4	sill (drill core)	cpx basalt	Gramado	50.5	☉				○						△	•?			fine, quartz vein,
5AT03-SC-590.1	sill (drill core)	cpx basalt	Gramado	52.1	☉				○						△	•?			
5AT03-SC-600.0	sill (drill core)	cpx dolerite	Esmeralda	51.8	☉				○						△	•?			
5AT03-SC-607.9	sill (drill core)	cpx dolerite	Esmeralda	51.9	☉				○						△	•?			
5AT03-SC-616.9	sill (drill core)	cpx basalt	Esmeralda	51.9	☉				○						△	•?			
5AT08-SC-792.4	sill (drill core)	cpx basalt	Gramado	47.8	☉				○						△	•?			
5AT08-SC-802.85	sill (drill core)	cpx gabbro	Gramado	37.4	☉				○						△	•?			
5AT08-SC-815.1	sill (drill core)	cpx dolerite	Gramado	50.6	☉				○						△	•?			
5AT08-SC-825.0	sill (drill core)	cpx dolerite	Gramado	51.4	☉				○						△	•?			
5AT08-SC-834.4	sill (drill core)	cpx dolerite	Gramado	51.5	☉				○						△	•?			coarse
5AT08-SC-845.1	sill (drill core)	cpx dolerite	Gramado	50.9	☉				○						△	•?			
5AT08-SC-853.05	sill (drill core)	cpx basalt	Gramado	51.3	☉				○						△	•?			epidote-chlorite vein
5AT08-SC-925.3	sill (drill core)	cpx basalt	Gramado?	46.6	☉				○						△	•?			
5AT08-SC-936.5	sill (drill core)	cpx basalt	Gramado	45.6	☉				○						△	•?			
5AT08-SC-947.75	sill (drill core)	cpx basalt	Gramado	49.8	☉				○						△	•?			fine grained
TG07-235.0	sill (drill core)	basaltic andesite	Gramado?	39.6	☉				△	•					•				
TG07-250	sill (drill core)	basaltic andesite	Gramado?	39.1	☉				△	•					•				
TG07-270.0	sill (drill core)	basaltic andesite	Gramado?	37.2	☉				△	•					•				
TG114-272.0	sill (drill core)	cpx basalt	Gramado	71.0	☉				•						•				
TG114-275.0	sill (drill core)	cpx basalt	Gramado	71.0	☉				•						•				

Microscopic observation of thin and polished-thin section for rock samples

Sample No.	Sample Type	Rock Name	Magma Type	Mg#	Minerals												Note								
					pl	am	opx	cpx	ol	qz	chl	ca	serp	mt	chr	sul		pyrr	py	cp					
TG114-278.9	sill (drill core)	basalt with gabbro	Gramado?	78.7	☉		•	○				•						•	•?						mixed texture, contact part
TG114-283.8	sill (drill core)	olivine gabbro	Gramado?	88.8	☉		•	○			△							•	•?						
TG114-286.4	sill (drill core)	olivine gabbro	Gramado?	90.5	☉		•	○	☉			•						•							
TG114-289.9	sill (drill core)	olivine gabbro	Gramado?	88.6	☉		•	○	☉			•						•							
TG114-293.7	sill (drill core)	olivine gabbro	Gramado?	84.8	☉		•	○	○			•						•							
TG114-295.0	sill (drill core)	olivine gabbro	Gramado?	82.8	☉		•	○	•			△						•	•?						
TG228-725	sill (drill core)	cpx basalt	Gramado	61.1	☉		•	○				•						△		•?					
TG228-740	sill (drill core)	cpx dolerite	Gramado	60.7	☉		•	○				•						△		•?					
TG228-755	sill (drill core)	cpx gabbro	Paramanema	51.5	☉		•	○				•						△							
TG228-758.5	sill (drill core)	cpx gabbro	Gramado	63.1	☉		•	○				•						△							
TG228-770	sill (drill core)	cpx gabbro	Gramado	64.7	☉		•	○				•						△							
TG228-785	sill (drill core)	cpx dolerite	Gramado	61.8	☉		•	○				•						△							
TG228-800	sill (drill core)	cpx dolerite	Gramado	61.8	☉		•	○				•						△							
TG228-817	sill (drill core)	cpx dolerite	Gramado	61.6	☉		•	○				•						△							
TG228-830	sill (drill core)	cpx dolerite	Gramado	61.6	☉		•	○				•						△							
TG27-112.0	sill (drill core)	basaltic andesite	Acidic rock	41.8	☉				△	•		•						•							ol bearing cpx basalt
TG27-119.0	sill (drill core)	basaltic andesite	Gramado	48.0	☉				△			•						•							
TG27-138.5	sill (drill core)	basaltic andesite	Gramado	42.2	☉				△			•	△					•							fine, strongly altered
TG27-91.5	sill (drill core)	olivine gabbro	unclassify	79.7	☉				△	•		•						△	•?						
TG27-93.2	sill (drill core)	cpx basalt	unclassify	48.0	☉				△			•						•							fine, quartz-calcite veins
TG62-176.5	sill (drill core)	cpx basalt	Gramado	45.3	☉				○			•						•	?						
TG62-183.5	sill (drill core)	cpx basalt	Gramado	42.5	☉				○			•						•	?						
TG62-188.0	sill (drill core)	cpx basalt	Gramado	46.9	☉				○			•						•	?						
TG62-196.7	sill (drill core)	cpx basalt	Gramado	42.4	☉				○			•	○					•	•(pr?)						calcite vein with opaque
TG62-200.5	sill (drill core)	cpx basalt	Gramado	45.1	☉				○			•						•	?						
TG62-204.0	sill (drill core)	cpx basalt	Gramado	42.2	☉				○			•						•	?						
TG62-207.7	sill (drill core)	cpx basalt	Gramado	43.6	☉				○			•						•	?						two calcite veins with sulfide
TG62-211.4	sill (drill core)	cpx basalt	Gramado	42.5	☉				○			•						•	?						
TG62-215.4	sill (drill core)	Komatite	Gramado	41.6	•						•	△						•							spinifex texture
TG62-220.0	sill (drill core)	cpx basalt	Gramado	42.6	☉				○			•						•	?						

Microscopic observation of thin and polished-thin section for rock samples

Sample No.	Sample Type	Rock Name	Magma Type	Mg#	Minerals											Note										
					pl	am	opx	cpx	ol	qz	chl	ca	serp	mt	chr		sul	opaque pyrr	py	cp						
FP-02 (580.15-580.40)	sill (drill core)	cpx basalt	Parapananema?	57.9	⊙				⊙																intergranular texture	
FP-02 (586.15-586.35)	sill (drill core)	cpx basalt	Parapananema	58.0	⊙				○																	cryptocrystalline-interstitial texture
FP-02 (605.55-605.70)	sill (drill core)	cpx dolerite	Pitanga?	50.7	⊙				⊙																	
FP-09 (174.10-174.30)	sediment (drill core)	altered elastic rock (silt)	----	----	△																					
NF-09 (70.70-70.90)	sediment (drill core)	calcareous altered rock	----	----	△																					
SP-17 (731.00-731.20)	sill (drill core)	cpx basalt	Pitanga	56.6	⊙				○																	calcite vein cutted the pyrite vein; cryptocrystalline
SP-17 (734.40-734.55)	sediment (drill core)	silt stone	----	----																						pyrite vein; muscovite
SP-17 (735.30-735.45)	sediment (drill core)	calcareous rock	----	----																						clay minerals (△)
SP-21 (129.80-130.00)	sill (drill core)	weathered pyroxene basalt	Pitanga	57.9	○																					clay minerals (△), pyrite nodular (φ 5mm), pyrite vein
SP-66 (776.50-776.70)	sill (drill core)	cpx dolerite	Parapananema	61.3	⊙																					
SW-01 (553.15-553.35)	sill (drill core)	weathered basalt	Esmeralda?	52.5	⊙																					clay minerals (△)
SW-03 (234.00-234.20)	sill (drill core)	cpx basalt	Esmeralda	62.6	⊙																					

Legend

⊙: abundant; ○: common; △: minor; •: rare

pl: plagioclase, am: amphibole, opx: ortho pyroxene, cpx: clino pyroxene, ol: olivine

qtz: quartz, chl: chlorite, ca: carbonate mineral (mainly calcite), serp: serpentine

mt: magnetite, chr: chromite, sul: sulfide, pyrr: pyrrhotite, py: pyrite, cp: chalcopyrite, chr: chromite

APPENDIX 6

Geochemical grade assay for rocks, drill cores and drill cuttings samples

Table with 35 columns: Sample No., Detection Limit, SiO2, Al2O3, Fe2O3, MnO, MgO, CaO, Na2O, K2O, TiO2, P2O5, LOI, Total, Be, V, Cr, Co, Ga, Ge, As, Rb, Sr, Y, Zr, Nb, Mo, In, Sn, Sb, Cs, Ba, Sc. Each row represents a sample with numerical values for these elements, many in ppm or %.

Geochemical grade assay for rocks, drill cores and drill cuttings samples

Sample No.	La	Ce	Pr	Nd	Sm	Eu	Gd	Tb	Dy	Ho	Er	Tm	Yb	Lu	HF	Ta	W	Tl	Bi	Th	U	Ag	Cd	Cu	Ni	Pb	Zn	S	
	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	%	
Detection Limit	0.05	0.05	0.01	0.05	0.01	0.005	0.01	0.01	0.01	0.01	0.01	0.005	0.01	0.002	0.1	0.01	0.5	0.05	0.1	0.05	0.01	0.3	0.3	1	1	3	30	0.001	
AC-028	18.97	42.62	4.87	21.49	4.64	1.999	4.85	0.77	4.56	0.92	2.71	0.373	2.27	0.315	3.7	0.83	<0.5	0.12	<0.1	2.05	0.36	<0.3	0.6	214	70	<3	91	<0.001	
AC-034B	21.32	48.54	5.67	25.11	5.87	1.904	6.39	1.13	6.62	1.36	4.05	0.592	3.35	0.485	4.7	0.94	<0.5	0.12	<0.1	2.40	0.41	<0.3	<0.3	308	50	<3	117	<0.001	
AC-035	22.84	51.74	5.89	25.98	6.02	1.980	6.42	1.14	6.87	1.41	4.11	0.615	3.47	0.509	5.0	1.01	<0.5	0.14	<0.1	2.51	0.48	<0.3	<0.3	228	41	16	111	<0.001	
ACR-119	52.06	110.73	12.19	50.35	10.33	2.153	10.10	1.70	9.64	1.94	5.63	0.844	4.74	0.654	7.6	1.88	1.1	1.74	0.4	12.57	0.49	<0.3	<0.3	115	<1	12	86	<0.001	
ACR-125	37.32	85.89	9.69	42.27	9.05	2.888	8.68	1.40	7.75	1.51	4.25	0.602	3.38	0.474	6.9	1.86	<0.5	0.24	<0.1	3.94	0.79	<0.3	<0.3	199	22	8	126	0.004	
ADR-033	24.63	51.34	6.07	25.58	6.41	1.621	6.41	1.15	7.29	1.44	4.25	0.644	3.68	0.537	4.9	0.98	0.8	1.18	0.3	6.92	2.07	<0.3	<0.3	126	26	<3	96	<0.001	
AS001	16.32	35.23	4.03	17.05	3.75	1.145	4.02	0.70	4.16	0.85	2.55	0.371	2.37	0.349	3.1	0.50	<0.5	0.21	0.1	3.67	0.66	<0.3	<0.3	58	49	29	75	0.001	
AS002	21.49	47.44	5.56	23.97	5.54	1.847	6.15	1.05	6.25	1.28	3.77	0.565	3.47	0.504	4.8	0.97	<0.5	0.09	0.1	2.48	0.47	<0.3	<0.3	203	38	40	103	<0.001	
AS003	15.61	34.10	4.04	17.64	3.92	1.413	4.40	0.74	4.39	0.86	2.55	0.369	2.37	0.347	3.2	0.70	<0.5	0.05	<0.1	1.74	0.31	<0.3	<0.3	185	71	4	82	<0.001	
AS004A	20.53	45.22	5.17	22.77	5.39	1.739	5.77	0.98	6.07	1.20	3.55	0.547	3.42	0.496	4.3	0.85	<0.5	0.11	0.1	2.72	0.51	<0.3	<0.3	211	51	24	100	<0.001	
AS004B	20.55	45.34	5.25	22.86	5.31	1.743	5.75	1.00	6.07	1.24	3.60	0.537	3.26	0.500	4.3	0.85	<0.5	0.19	<0.1	2.76	0.54	<0.3	<0.3	264	42	7	103	0.005	
AS005	20.96	46.20	5.45	23.27	5.17	1.782	5.52	0.91	5.57	1.08	3.22	0.478	2.97	0.444	3.9	0.88	<0.5	0.05	<0.1	2.39	0.55	<0.3	<0.3	219	57	29	97	<0.001	
AS006	24.19	53.78	6.33	26.90	6.23	2.120	6.28	1.02	6.16	1.17	3.50	0.528	3.07	0.470	4.5	1.04	<0.5	0.09	0.2	2.55	0.53	<0.3	<0.3	226	56	13	96	0.005	
AS007	22.30	49.69	5.81	24.76	5.38	1.858	5.55	0.90	5.54	1.06	3.03	0.438	2.67	0.411	4.1	0.95	<0.5	0.09	<0.1	2.23	0.42	<0.3	<0.3	197	54	15	87	<0.001	
AS008	28.25	63.02	7.43	31.45	6.75	2.296	6.85	1.06	5.99	1.14	3.26	0.440	2.65	0.411	5.3	1.30	<0.5	0.07	0.1	2.63	0.53	<0.3	<0.3	209	73	<3	116	<0.001	
AS009	24.19	53.11	6.34	27.86	6.35	2.099	7.34	1.20	7.45	1.50	4.42	0.655	4.06	0.608	5.2	1.04	<0.5	0.13	0.1	2.47	0.46	<0.3	<0.3	264	44	<3	121	<0.001	
AS010	26.25	58.96	6.62	28.94	6.50	2.093	6.55	1.10	6.71	1.25	3.72	0.565	3.37	0.513	4.8	1.13	<0.5	0.30	0.3	2.83	0.57	<0.3	<0.3	260	44	4	117	0.366	
AS011	26.76	58.28	6.85	29.38	6.54	2.103	6.59	1.11	6.58	1.25	3.78	0.552	3.27	0.505	5.0	1.14	<0.5	0.09	0.1	2.98	0.59	<0.3	<0.3	268	43	<3	117	0.363	
AS013	31.29	69.47	8.13	34.55	7.73	2.508	8.14	1.33	7.89	1.54	4.60	0.675	4.13	0.586	5.9	1.38	<0.5	0.16	0.1	3.44	0.74	<0.3	<0.3	299	26	<3	133	0.037	
AS014	52.80	119.41	14.20	62.31	13.20	4.294	12.78	1.99	11.17	2.03	5.78	0.807	4.78	0.673	9.2	2.18	0.5	0.14	<0.1	4.68	0.96	<0.3	<0.3	79	1	11	155	0.133	
AS015A	75.99	166.51	19.06	79.14	16.29	4.620	15.38	2.41	13.98	2.63	7.63	1.067	6.52	0.980	15.3	3.11	0.7	0.27	0.1	8.00	1.62	<0.3	<0.3	91	<1	4	179	0.109	
AS015B	41.89	93.59	11.16	49.01	10.49	3.442	10.47	1.58	8.81	1.63	4.81	0.612	3.78	0.585	7.4	1.93	1.3	0.18	<0.1	3.65	0.71	<0.3	<0.3	45	7	<3	120	0.315	
AS016A	22.07	49.19	5.73	24.92	5.72	1.903	5.99	1.03	6.42	1.23	3.65	0.566	3.46	0.533	4.5	0.94	<0.5	0.09	<0.1	2.38	0.44	<0.3	<0.3	246	47	12	107	0.011	
AS016B	34.00	75.06	8.75	37.25	8.48	2.300	8.92	1.51	9.09	1.77	5.30	0.800	4.76	0.715	6.7	1.47	<0.5	0.10	<0.1	3.80	0.73	<0.3	<0.3	292	44	<3	137	0.011	
AS017	37.92	83.24	9.61	40.97	8.93	2.624	9.18	1.53	9.12	1.79	5.33	0.777	4.62	0.702	7.3	1.58	<0.5	0.15	0.2	4.51	0.88	<0.3	<0.3	387	21	<3	123	0.063	
AS018	23.86	52.89	6.21	27.00	6.10	2.060	6.29	1.07	6.27	1.24	3.73	0.548	3.34	0.520	4.6	1.06	<0.5	0.09	<0.1	2.64	0.54	<0.3	<0.3	255	49	7	116	0.082	
AS021	20.74	44.79	5.05	21.36	4.94	1.438	5.37	0.90	5.40	1.12	3.15	0.464	2.96	0.428	4.3	0.80	<0.5	0.15	0.2	2.41	0.41	<0.3	<0.3	112	41	8	93	0.013	
AS022A	39.72	91.39	11.40	55.13	14.69	4.882	19.55	3.37	19.48	3.80	10.76	1.459	8.33	1.212	5.0	2.98	1.0	0.21	0.2	4.41	0.93	<0.3	<0.3	111	144	67	<3	134	0.018
AS022B	17.16	33.91	3.61	15.65	3.78	1.377	4.79	0.89	5.59	1.23	3.82	0.648	4.04	0.632	7.0	2.41	0.8	0.15	0.1	6.26	2.41	0.7	0.9	276	78	24	138	0.246	
AS022C	11.87	23.82	2.30	8.89	1.60	0.572	1.47	0.23	1.16	0.21	0.59	0.084	0.54	0.083	1.8	1.26	<0.5	0.37	<0.1	1.77	0.94	<0.3	<0.3	14	63	4	<30	4.295	
AS022D	11.75	21.39	2.21	8.50	1.75	0.517	1.87	0.32	1.95	0.39	1.14	0.173	1.11	0.162	5.0	2.01	0.9	0.45	<0.1	6.57	5.84	0.3	0.6	136	63	15	<30	0.055	
AS023	23.52	51.34	5.86	26.23	6.26	1.962	7.31	1.30	7.74	1.59	4.56	0.687	4.20	0.608	5.2	1.70	<0.5	0.39	0.1	5.75	1.23	<0.3	<0.3	208	24	5	130	0.048	
AS024A	8.68	20.87	2.62	12.54	3.66	1.343	4.85	0.89	5.43	1.10	3.13	0.475	2.85	0.413	2.8	0.91	<0.5	0.08	<0.1	1.85	0.51	<0.3	<0.3	0.7	160	65	4	0.020	
AS024B	47.56	94.56	9.74	39.12	6.94	2.342	6.20	0.94	5.21	0.97	2.56	0.341	2.17	0.302	3.7	3.54	<0.5	0.29	<0.1	5.62	1.70	<0.3	<0.3	102	202	7	345	0.089	
AS025	43.93	98.88	11.44	50.46	10.68	3.655	10.66	1.50	8.12	1.45	3.75	0.464	2.99	0.393	8.7	2.46	<0.5	0.31	<0.1	4.67	1.00	0.4	0.7	143	54	7	135	0.012	
AS026	102.39	179.37	17.18	64.65	10.91	3.199	9.01	1.32	7.05	1.27	3.38	0.437	2.85	0.388	5.6	3.56	0.8	0.52	0.1	23.65	3.57	<0.3	<0.3	0.5	81	150	24	110	0.028
AS027	61.89	128.18	14.41	61.74	12.63	4.114	11.05	1.60	8.49	1.47	3.76	0.457	2.94	0.397	9.6	4.56	<0.5	0.11	<0.1	5.55	1.27	0.6	0.7	164	56	11	131	0.025	
AS028A	129.71	165.14	10.59	24.91	2.32	0.495	1.39	0.24	1.50	0.34	1.24	0.234	1.66	0.281	13.1	3.03	<0.5	1.55	0.3	40.88	10.61	1.2	<0.3	3	<1	46	127	0.006	
AS028B	132.27	163.95	10.45	24.49	2.37	0.552																							

Geochemical grade assay for rocks, drill cores and drill cuttings samples

(continued/8)

Sample No.	Pd	Pt	Ir	Rh	Re	Au
	ppb	ppb	ppb	ppb	ppb	ppb
Detection Limit	0.2	0.2	0.1	0.2	0.1	2
AC-028	16.5	11.9				4
AC-034B	19.1	5.9				6
AC-035	15.2	7.6				5
ACR-119	2.0	1.4				1
ACR-125	1.3	0.6				3
ADR-033	7.4	3.3				3
AS001	0.6	0.3				17
AS002	17.0	4.8				8
AS003	16.4	10.2				3
AS004A	22.3	9.3				6
AS004B	20.4	5.2				4
AS005	15.1	11.7				4
AS006	13.3	11.3				4
AS007	17.9	11.0				4
AS008	0.2	0.6				<2
AS009	16.6	4.3				9
AS010	16.0	13.1				5
AS011	14.7	11.5				4
AS013	20.9	9.8				5
AS014	<0.2	<0.2				<2
AS015A	<0.2	<0.2				<2
AS015B	<0.2	<0.2				<2
AS016A	15.6	5.2				4
AS016B	14.9	1.2				5
AS017	15.1	2.9				5
AS018	17.6	6.4				3
AS021	4.7	6.0				69
AS022A	8.3	8.5				4
AS022B	13.6	11.1				5
AS022C	1.9	2.4				3
AS022D	3.8	5.1				3
AS023	14.1	15.2				6
AS024A	11.6	6.0				3
AS024B	6.2	3.9				2
AS025	7.8	4.4				4
AS026	5.4	2.8				<2
AS027	7.1	6.4				4
AS028A	<0.2	<0.2				<2
AS028B	<0.2	<0.2				<2
KM103	16.6	12.8				4
KM104	15.4	11.1	0.2	1.0	0.7	17
KM105	5.8	5.1				2
KM106	<0.2	<0.2				<2
KM107	<0.2	<0.2				<2
KM108	11.7	11.1				4
KM109	4.2	1.3				<2
KM111	14.8	12.9				4
KM112	14.3	11.2				4
KM114	9.4	5.1				4
KM115	16.3	9.0				4
KM116	37.0	3.3	0.5	0.7	0.8	22
KM117	11.7	6.2				3
KM118	16.4	9.8				3
KM120	15.3	5.1	0.1	1.0	1.2	8
KM121	0.2	<0.2				3

Geochemical grade assay for rocks, drill cores and drill cuttings samples

Table with columns: Sample No., Detection Limit, and various chemical elements (SiO2, Al2O3, Fe2O3, MnO, MgO, CaO, Na2O, K2O, TiO2, P2O5, LOI, Total, Be, V, Cr, Co, Ga, Ge, As, Rb, Sr, Y, Zr, Nb, Mo, In, Sn, Sb, Cs, Ba, Sc) with their respective concentrations in ppm or %.

Geochemical grade assay for rocks, drill cores and drill cuttings samples

Table with 28 columns representing elements (La, Ce, Pr, Nd, Sm, Eu, Gd, Tb, Dy, Ho, Er, Tm, Yb, Lu, HF, Ta, W, Ti, Bi, Th, U, Ag, Cd, Cu, Ni, Pb, Zn, S) and rows for various sample IDs (KM122 to KN020) under the heading 'Sample No.' and 'Detection Limit'.

Geochemical grade assay for rocks, drill cores and drill cuttings samples

(2continued/8)

Sample No.	Pd	Pt	Ir	Rh	Re	Au
	ppb	ppb	ppb	ppb	ppb	ppb
Detection Limit	0.2	0.2	0.1	0.2	0.1	2
KM122	15.2	16.0				4
KM124	12.2	4.4				3
KM125	8.8	8.1				4
KM126	18.6	8.7				5
KM128	0.3	0.7				<2
KM129	3.9	3.3				<2
KM130A	5.6	2.9				<2
KM130B	15.4	8.8				4
KM131	14.3	7.0				4
KM133	17.3	6.7				6
KM134	12.4	6.0				5
KM135	15.3	6.0				5
KM136	20.0	5.9				7
KM137	15.2	8.5				4
KM138	1.9	1.7				<2
KM139	14.9	7.3				4
KM140	4.4	4.5				2
KM141	10.0	14.9				3
KM142	0.7	0.4				<2
KM143	9.7	7.4				4
KM144	16.4	4.4				4
KM145	11.9	8.1				4
KM146	15.1	5.0				4
KM147	0.4	<0.2				<2
KM148	0.5	0.4				<2
KM150	0.4	<0.2				<2
KM151	0.4	<0.2				<2
KM152	0.5	<0.2				<2
KM153	0.8	0.7				<2
KM154	3.2	4.3				<2
KM155	3.7	4.1				3
KM156	4.3	2.7				<2
KM157	18.4	10.7				5
KM158	10.7	3.3				5
KM159	11.2	7.3				15
KM160	17.9	12.8				4
KM161	16.8	7.3				4
KM162	18.1	5.9				5
KM163	16.0	7.5				4
KN001	0.8	0.7	<0.1	0.3	<0.1	2
KN002	0.3	1.2	<0.1	0.3	0.2	<1
KN003	8.4	8.5				3
KN005	10.1	4.6				3
KN006	7.4	3.0				3
KN009	13.5	12.0				4
KN010	9.6	3.0				3
KN011	6.3	2.7				1
KN012	11.4	4.5				3
KN014	0.2	0.2				<1
KN016	16.1	6.3				2
KN017	12.2	18.1				5
KN018	0.6	0.6				<1
KN019	18.5	14.6				6
KN020	17.5	12.1				6

Geochemical grade assay for rocks, drill cores and drill cuttings samples

Sample No. Detection Limit	SiO ₂		Al ₂ O ₃		Fe ₂ O ₃		MnO	MgO	CaO	Na ₂ O	K ₂ O	TitO ₂	P ₂ O ₅	LOI	Total	Be	V	Cr	Co	Ga	Ge	As	Rb	Sr	Y	Zr	Nb	Mo	In	Sn	Sb	Cs	Ba	Sc			
	%	ppm	%	ppm	%	ppm	%	ppm	%	ppm	%	ppm	%	ppm	%	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	
KN021	50.05	13.29	13.33	13.87	0.206	5.62	10.04	2.49	1.04	2.107	0.24	1.31	100.21	<1	433	100	44	21	1.5	<5	24	368	27.5	153	14.1	<2	<0.1	<1	<0.2	0.4	326	39	1				
KN022A	49.95	13.33	13.87	0.206	5.62	10.04	2.49	1.04	2.107	0.24	1.31	100.21	<1	433	100	44	21	1.5	<5	24	368	27.5	153	14.1	<2	<0.1	<1	<0.2	0.4	326	39	1					
KN022B	52.26	11.21	17.01	0.209	3.28	5.59	2.66	2.41	3.025	0.43	2.04	100.12	<1	298	<20	35	24	1.7	<5	83	326	41.6	287	25.2	<2	<0.1	<1	<0.2	0.3	556	32						
KN023A	50.28	13.32	14.10	0.201	5.67	9.95	2.52	1.02	2.145	0.24	1.06	100.50	<1	416	96	43	22	1.8	<5	24	363	28.5	160	14.8	<2	<0.1	<1	<0.2	0.4	339	38						
KN023B	50.01	13.37	17.67	0.220	3.78	7.31	2.63	1.60	3.509	0.40	1.81	100.34	<1	497	<20	40	25	1.7	<5	49	347	39.2	247	23.2	<2	<0.1	<1	<0.2	0.4	453	37						
KN024A	50.23	13.37	14.19	0.203	5.91	10.06	2.49	0.98	2.130	0.23	0.64	100.45	<1	441	101	45	22	1.9	<5	23	389	27.7	153	14.5	<2	<0.1	<1	<0.2	0.1	326	38						
KN024B	51.30	11.99	15.92	0.172	3.38	5.78	2.77	3.24	2.843	0.42	1.69	99.50	<1	397	<20	35	25	1.7	<5	161	367	42.9	259	23.3	<2	<0.1	<1	<0.2	0.8	483	33						
KN025A	49.73	13.32	13.95	0.190	5.98	10.18	2.34	0.80	1.810	0.21	1.63	100.12	<1	420	147	44	21	1.6	<5	15	321	28.1	127	10.8	<2	<0.1	<1	<0.2	0.1	265	39						
KN025B	54.24	10.36	13.21	0.344	3.61	6.99	2.73	1.99	3.045	0.37	1.20	100.28	<1	556	<20	35	22	1.7	<5	57	328	32.9	224	22.7	<2	<0.1	<1	<0.2	0.5	447	33						
KN026	49.93	13.31	14.14	0.205	5.81	10.19	2.45	0.98	2.132	0.24	0.82	100.10	<1	434	100	44	22	1.9	<5	21	368	28.0	155	14.6	<2	<0.1	<1	<0.2	0.1	324	39						
KN027	50.73	12.64	15.01	0.212	4.82	8.78	2.56	1.33	2.304	0.27	0.71	100.24	<1	420	52	41	22	1.5	<5	32	265	36.7	175	14.9	<2	<0.1	<1	0.3	0.2	311	41						
KN028	51.43	12.49	15.33	0.220	4.88	8.75	2.61	1.30	2.236	0.27	0.71	100.24	<1	420	53	41	22	1.8	<5	29	266	39.4	189	15.2	<2	<0.1	<1	0.3	0.2	335	42						
KN029	51.22	12.68	16.16	0.229	4.09	8.38	2.87	1.42	2.604	0.33	0.25	100.24	<1	489	41	38	22	2.1	<5	31	235	35.6	171	14.1	<2	<0.1	<1	0.3	0.3	341	41						
KN030	51.31	12.56	15.85	0.227	4.57	8.73	2.65	1.24	2.367	0.28	0.22	100.01	<1	450	27	40	21	1.7	<5	29	290	38.8	186	16.2	<2	<0.1	<1	<0.2	0.2	368	42						
KN031	50.66	13.56	13.20	0.204	6.31	10.29	2.43	0.87	1.660	0.18	0.86	100.22	<1	376	84	40	18	1.6	<5	19	275	27.1	127	10.3	<2	<0.1	<1	<0.2	0.2	243	43						
KN033	52.09	12.56	15.37	0.209	4.42	8.40	2.68	1.45	2.391	0.30	0.34	100.21	<1	443	26	38	22	1.7	<5	36	280	38.1	198	16.5	<2	<0.1	<1	<0.2	0.3	375	41						
KN034A	50.61	13.39	12.89	0.159	2.68	6.38	3.11	2.76	2.925	0.74	0.62	100.19	<1	366	72	22	18	0.6	<5	38	751	28.7	229	20.5	<2	<0.1	<1	<0.2	0.8	501	27						
KN035A	51.21	13.10	14.95	0.213	4.96	9.10	2.47	1.09	2.246	0.26	0.68	100.28	<1	423	54	39	21	1.6	<5	28	274	37.8	186	14.7	<2	<0.1	<1	<0.2	1.0	331	42						
KN036A	50.98	12.44	15.78	0.220	4.28	8.06	2.69	1.64	3.451	0.50	0.13	100.18	<1	458	21	38	25	1.7	<5	48	381	40.4	267	25.1	<2	<0.1	<1	<0.2	1.0	515	33						
KN036B	50.23	12.93	15.25	0.218	4.29	8.14	2.70	1.47	3.708	0.44	0.88	100.26	<1	509	20	39	25	1.2	<5	38	400	37.3	247	23.6	<2	<0.1	<1	<0.2	0.5	450	34						
KN037	49.99	14.83	14.86	0.196	3.69	8.90	2.83	1.30	2.775	0.36	0.58	100.31	<1	514	<20	36	23	1.7	<5	36	378	34.2	176	16.5	<2	<0.1	<1	<0.2	1.9	398	32						
KN038	53.12	12.59	14.65	0.221	3.13	7.02	3.10	1.74	3.171	0.46	1.27	100.49	<1	366	<20	20	14	<5	40	366	47.2	289	24.7	<2	<0.1	<1	<0.2	0.6	528	29							
KN039A	49.79	13.50	14.13	0.201	6.27	9.99	2.53	0.94	2.133	0.28	0.41	100.17	<1	396	135	42	20	1.7	<5	26	339	29.7	139	13.4	<2	<0.1	<1	<0.2	0.7	305	39						
KN039B	49.70	11.98	17.56	0.238	4.22	8.43	2.49	1.25	3.598	0.40	0.63	100.48	<1	592	<20	43	23	1.2	<5	35	340	40.4	212	19.9	<2	<0.1	<1	<0.2	0.9	409	40						
KN040A	50.37	12.99	14.77	0.208	5.70	9.64	2.41	1.04	2.358	0.39	0.51	100.40	<1	446	100	43	21	1.5	<5	32	340	32.7	165	15.6	<2	<0.1	<1	<0.2	0.8	355	39						
KN040B	50.02	13.17	14.64	0.204	5.64	9.67	2.48	0.96	2.372	0.32	0.75	100.22	<1	441	97	42	21	1.7	<5	29	338	32.4	165	15.6	<2	<0.1	<1	<0.2	1.0	303	38						
KN040C	46.94	11.66	19.22	0.238	5.22	9.09	2.32	1.02	3.871	0.29	0.40	100.26	<1	916	<20	54	23	1.7	<5	27	314	31.6	166	15.0	<2	<0.1	<1	<0.2	0.8	354	45						
KN042	52.81	13.95	12.56	0.183	5.62	9.14	2.72	1.27	1.111	0.13	0.93	99.98	<1	357	32	46	20	1.9	<5	45	201	22.9	105	10.4	<2	<0.1	<1	<0.2	1.4	239	44						
KN043	66.13	12.39	5.56	0.173	0.67	2.74	2.72	4.47	0.708	0.21	0.27	99.73	<1	27	<20	6	19	1.7	<5	195	98	53.8	295	25.0	<2	<0.1	<1	9	0.3	10.3	660	18					
KN044	52.38	13.98	11.39	0.165	5.85	9.60	2.20	0.82	1.118	0.14	1.72	99.36	<1	305	75	36	19	1.6	<5	18	192	25.8	132	8.4	<2	<0.2	55	<0.2	1.6	294	39						
KN045	54.60	13.74	12.35	0.175	4.59	8.02	2.65	1.76	1.322	0.20	0.96	100.41	<1	322	<20	37	21	1.8	<5	62	211	31.9	141	12.1	<2	<0.1	<1	<0.2	2.2	345	38						
KN046A	53.20	13.78	11.83	0.193	5.80	9.50	2.46	1.12	1.137	0.13	0.70	100.81	<1	311	31	38	19	1.8	<5	37	192	25.1	97	7.0	<2	<0.1	<1	<0.2	1.0	211	42						
KN047	49.56	12.57	15.74	0.214	4.31	8.37	2.68	1.45	3.472	0.55	0.81	99.71	<1	441	<20	37	23	1.8	<5	33	390	41.7	235	23.8	<2	<0.1	<1	<0.2	0.4	511	34						
KN048A	53.60	13.18	13.72	0.190	4.02	7.66	2.84	1.73	1.725	0.26	1.02	99.96	<1	393	<20	36	22	1.7	<5	59	232	39.5	187	17.3	<2	<0.1	<1	<0.2	1.6	457	37						
KN048B	51.20	11.94	13.74	0.155	2.22	5.01	3.01	2.61	1.800	0.40	1.30	99.39	<1	230	<20	27	23	2.0	<5	92	195	55.8	289	25.5	<2	<0.1	<1	<0.2	2.3	673	29						
KN049	49.86	13.47	13.87	0.208	5.78	9.84	2.43	0.90	1.926	0.22	0.94	99.44	<1	399	103	41	20	1.6	<5	19	317	28.8	137	12.7	<2	<0.1	<1	<0.2	0.2	300	41						
KN050	49.76	13.65	14.25	0.224	5.78	10.10	2.43	0.85	2.153	0.27	0.97	100.43	<1	415	116	41	20	1.6	<5	12	331	31.0	144	13.4	<2	<0.1	<1	<0.2	0.3	329	40						
KN051	50.81	12.73	15.76	0.209	4.89	8.72	2.45	1.35	2.313	0.28	0.52	100.04	<1	429	43	38	20	1.7	<5	38	272	37.3	164	14.2	<2	<0.1	<1	<0.2	0.4	332	43						
KN052	49.51	13.60	13.85	0.187	6.21	10.24	2.38	0.86	1.877	0.19	0.64	99.56	<1	408	118	43	20	1.8	<5	17	320	2															

Geochemical grade assay for rocks, drill cores and drill cuttings samples

Sample No.	La	Ce	Pr	Nd	Sm	Eu	Gd	Tb	Dy	Ho	Er	Tm	Yb	Lu	HF	Ta	W	Ti	Bi	Th	U	Ag	Cd	Cu	Ni	Pb	Zn	S
	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	%
	0.05	0.05	0.01	0.05	0.01	0.005	0.01	0.01	0.01	0.01	0.01	0.005	0.01	0.002	0.1	0.01	0.5	0.05	0.1	0.05	0.01	0.3	0.3	1	1	3	30	0.001
Detection Limit																												
KN021	22.56	51.31	5.83	25.12	5.56	1.873	5.74	0.90	5.45	1.07	3.14	0.440	2.62	0.381	4.3	1.00	<0.5	0.19	<0.1	2.36	0.45	<0.3	0.5	246	57	7	104	<0.001
KN022A	21.83	48.38	5.49	24.04	5.30	1.779	5.40	0.87	5.15	1.01	3.06	0.431	2.37	0.332	4.1	0.95	<0.5	0.14	<0.1	2.29	0.42	<0.3	0.9	234	55	<3	100	<0.001
KN022B	42.58	94.55	10.43	44.24	9.26	2.891	8.56	1.42	8.58	1.68	4.83	0.744	4.28	0.600	7.8	1.80	<0.5	0.21	<0.1	4.46	0.87	<0.3	0.5	273	10	<3	134	<0.001
KN023A	23.05	51.57	5.96	25.55	5.56	1.904	5.88	0.92	5.44	1.09	3.27	0.459	2.61	0.373	4.5	1.00	<0.5	0.14	<0.1	2.47	0.45	<0.3	0.6	239	49	8	95	<0.001
KN023B	35.69	81.17	9.13	38.92	8.24	2.696	8.33	1.33	7.68	1.56	4.36	0.626	3.63	0.517	6.7	1.59	<0.5	0.41	<0.1	3.69	0.73	<0.3	0.7	372	14	<3	137	<0.001
KN024	22.30	50.15	5.63	24.68	5.34	1.888	5.55	0.90	5.37	1.04	2.95	0.430	2.59	0.344	4.4	0.99	<0.5	0.11	<0.1	2.32	0.45	<0.3	0.6	226	55	<3	96	<0.001
KN024A	37.91	85.26	9.50	40.88	8.84	2.835	8.63	1.42	8.28	1.63	4.62	0.671	3.71	0.555	6.9	1.63	<0.5	0.34	<0.1	3.90	0.80	<0.3	0.3	232	14	<3	145	0.004
KN025A	18.12	40.33	4.64	20.77	4.81	1.700	5.06	0.87	5.27	1.05	3.13	0.459	2.45	0.372	3.6	0.73	<0.5	0.12	<0.1	2.19	0.50	<0.3	<0.3	227	64	8	96	<0.001
KN025B	32.95	74.66	8.47	35.85	7.48	2.382	7.31	1.15	6.62	1.25	3.61	0.514	2.79	0.433	6.5	1.61	<0.5	0.33	<0.1	3.33	0.54	<0.3	0.6	1315	16	<3	90	0.012
KN026	22.26	49.87	5.63	24.74	5.46	1.868	5.56	0.91	5.44	1.06	3.06	0.424	2.39	0.357	4.3	0.95	<0.5	0.12	<0.1	2.32	0.42	<0.3	0.8	241	53	<3	99	<0.001
KN027	23.08	53.12	6.12	26.11	6.20	2.011	6.52	1.13	6.88	1.41	4.08	0.603	3.25	0.484	4.8	1.01	<0.5	0.23	<0.1	2.55	0.50	<0.3	0.6	273	41	<3	117	<0.001
KN028	21.86	48.08	5.98	25.85	6.32	1.974	7.49	1.19	7.18	1.52	4.43	0.653	4.02	0.586	5.4	1.08	<0.5	0.15	<0.1	2.82	0.50	<0.3	0.5	269	40	<3	113	<0.001
KN029	27.72	59.09	7.47	31.91	6.22	2.246	8.72	1.35	8.29	1.72	4.93	0.715	4.38	0.610	6.4	1.29	<0.5	0.28	<0.1	3.35	0.55	<0.3	0.3	289	33	7	122	<0.001
KN030	21.06	44.63	5.60	24.07	5.86	1.854	6.81	1.10	6.75	1.41	4.08	0.602	3.81	0.523	5.1	1.01	<0.5	0.14	<0.1	2.31	0.57	<0.3	<0.3	240	36	15	114	<0.001
KN031	24.51	50.91	6.51	27.62	6.38	2.000	7.31	1.19	7.16	1.47	4.20	0.614	3.71	0.541	5.6	1.12	<0.5	0.16	<0.1	2.92	0.51	<0.3	1.4	268	32	18	125	<0.001
KN032	15.74	33.21	4.26	17.80	4.31	1.416	4.75	0.80	4.79	1.03	2.87	0.441	2.59	0.375	3.5	0.69	<0.5	0.09	<0.1	1.86	0.35	<0.3	0.3	172	54	<3	93	<0.001
KN033	24.74	51.06	6.50	27.62	6.39	2.058	6.91	1.17	6.98	1.45	4.16	0.627	3.77	0.518	5.4	1.16	<0.5	0.19	<0.1	2.86	0.50	<0.3	<0.3	185	27	13	120	<0.001
KN034A	31.15	64.92	8.45	36.43	8.03	2.729	7.71	1.14	6.00	1.07	2.85	0.366	2.28	0.302	6.4	1.71	<0.5	0.10	<0.1	3.36	0.73	<0.3	<0.3	111	66	<3	93	<0.001
KN034B	62.62	131.47	16.97	69.70	15.00	4.521	13.85	2.04	10.66	1.98	5.20	0.629	3.88	0.535	11.8	2.58	0.6	0.45	<0.1	7.23	1.56	<0.3	0.4	251	15	8	140	0.011
KN035A	22.24	48.71	5.94	26.16	6.20	1.950	6.76	1.14	7.05	1.45	4.22	0.608	3.71	0.529	5.2	1.01	<0.5	0.24	<0.1	2.71	0.49	0.4	0.9	256	43	<3	115	0.031
KN036A	35.71	75.77	9.82	41.33	9.04	2.908	9.39	1.40	7.99	1.60	4.36	0.584	3.64	0.523	7.5	1.71	<0.5	0.24	<0.1	4.45	0.84	<0.3	0.7	235	33	<3	133	0.037
KN036B	31.98	66.95	8.70	37.58	8.18	2.641	8.40	1.27	7.32	1.44	3.92	0.547	3.39	0.474	6.9	1.61	<0.5	0.18	<0.1	3.91	0.73	<0.3	0.7	243	33	5	131	0.686
KN037	26.38	55.18	6.98	29.17	6.49	2.275	6.74	1.11	6.64	1.34	3.89	0.572	3.38	0.500	5.3	1.20	<0.5	0.22	<0.1	3.23	0.63	<0.3	0.7	238	26	4	116	0.055
KN038	37.19	78.46	10.04	43.06	9.47	3.059	9.92	1.52	8.90	1.82	4.90	0.687	4.29	0.590	8.0	1.79	<0.5	0.23	<0.1	4.03	0.81	<0.3	0.6	124	4	<3	133	0.137
KN039A	20.62	43.07	5.46	23.51	5.28	1.778	5.63	0.91	5.48	1.10	3.21	0.465	2.92	0.408	4.1	0.92	<0.5	0.17	<0.1	2.43	0.51	<0.3	0.9	190	64	<3	96	0.051
KN039B	29.74	62.43	7.92	33.46	7.52	2.481	8.40	1.31	7.65	1.55	4.49	0.646	4.06	0.569	6.0	1.39	<0.5	0.15	<0.1	3.52	0.70	<0.3	0.3	348	21	<3	136	0.127
KN040A	24.40	50.50	6.39	28.02	6.17	2.006	6.90	1.04	6.30	1.28	3.68	0.529	3.23	0.452	4.7	1.11	<0.5	0.22	<0.1	2.89	0.58	<0.3	0.5	211	56	11	107	0.068
KN040B	24.43	50.24	6.36	26.94	6.13	2.004	6.70	1.03	6.32	1.27	3.63	0.524	3.24	0.451	4.8	1.03	<0.5	0.22	<0.1	2.87	0.57	<0.3	<0.3	210	54	8	107	0.104
KN040C	21.91	45.49	5.87	24.92	5.67	1.946	6.49	0.99	5.83	1.24	3.49	0.486	3.08	0.441	4.5	0.97	<0.5	0.12	<0.1	2.48	0.49	<0.3	0.5	416	43	<3	133	0.087
KN042	55.87	34.71	3.94	16.99	3.84	1.238	4.36	0.74	4.28	0.86	2.44	0.366	2.20	0.342	3.2	0.78	<0.5	0.28	0.1	4.77	0.98	<0.3	0.6	122	57	10	87	0.010
KN043	13.32	109.06	12.06	48.88	10.17	1.945	10.58	1.73	10.00	2.01	5.74	0.887	5.37	0.785	8.9	2.37	1.8	1.02	0.2	18.97	3.67	0.4	0.3	27	1	26	94	0.003
KN044	16.63	35.00	4.16	18.07	4.20	1.300	4.59	0.80	4.93	0.96	2.75	0.415	2.57	0.376	4.0	1.74	<0.5	0.08	0.2	4.94	1.13	<0.3	0.5	91	46	10	88	0.011
KN045	21.49	47.06	5.31	22.52	5.13	1.538	5.57	0.98	5.77	1.15	3.31	0.491	2.36	0.458	4.2	1.02	0.6	0.39	0.1	6.26	1.57	<0.3	0.6	65	31	8	95	0.009
KN046A	13.43	29.91	3.45	15.47	3.79	1.292	4.34	0.79	4.74	0.95	2.68	0.405	2.48	0.365	3.1	0.72	<0.5	0.15	<0.1	3.65	0.90	<0.3	0.6	127	32	9	87	0.011
KN047	36.52	81.58	9.45	41.45	8.87	2.977	8.84	1.41	8.21	1.54	4.14	0.597	3.66	0.552	6.9	1.89	<0.5	0.11	<0.1	4.43	0.87	0.4	0.8	146	27	4	125	0.032
KN048A	29.90	63.30	6.96	29.61	6.62	1.996	6.86	1.23	7.19	1.46	4.12	0.624	3.79	0.557	5.6	1.29	0.6	0.27	<0.1	7.17	1.55	<0.3	0.8	141	26	7	112	0.009
KN048B	45.31	96.46	10.81	45.13	9.67	2.615	10.51	1.74	10.43	2.08	5.85	0.887	5.33	0.804	8.5	1.81	1.3	0.54	<0.1	11.05	2.42	0.3	0.8	193	7	13	118	0.008
KN049	19.46	43.09	4.95	21.78	5.00	1.759	5.39	0.93	5.48	1.05	2.98	0.437	2.79	0.402	4.1	0.93	<0.5	0.06	<0.1	2.45	0.44	<0.3	0.7	210				

Geochemical grade assay for rocks, drill cores and drill cuttings samples

Sample No.	Pd	Pt	Ir	Rh	Re	Au
	ppb	ppb	ppb	ppb	ppb	ppb
Detection Limit	0.2	0.2	0.1	0.2	0.1	2
KN021	16.1	12.6				6
KN022A	17.8	12.9				5
KN022B	27.4	2.5				12
KN023A	12.9	11.8				5
KN023B	27.4	2.4				11
KN024A	21.9	13.4				6
KN024B	28.4	2.3				9
KN025A	18.4	16.8				6
KN025B	22.0	2.6				8
KN026	13.0	11.9				5
KN027	20.3	4.6				6
KN028	7.0	5.0				4
KN029	6.2	10.7				1
KN030	7.8	8.7				1
KN031	16.5	4.9				6
KN032	13.4	7.8				4
KN033	18.3	4.6				7
KN034A	5.3	7.9				3
KN034B	9.0	1.0				6
KN035A	15.1	6.7				6
KN036A	3.2	4.3				4
KN036B	3.8	4.4				1
KN037	7.3	6.8				6
KN038	<0.2	<0.2				<2
KN039A	5.6	5.7				3
KN039B	10.1	14.1				7
KN040A	6.5	6.4				3
KN040B	6.3	6.0				3
KN040C	14.1	14.7				9
KN042	1.0	3.6				<2
KN043	<0.2	0.2				<2
KN044	6.0	7.6				4
KN045	0.2	0.9				<2
KN046A	7.4	5.6				3
KN047	<0.2	0.2				<2
KN048A	3.4	5.2				2
KN048B	8.9	2.6				2
KN049	14.4	6.6				4
KN050	10.4	14.5				4
KN051	20.6	4.0				6
KN052	13.2	8.2				9
KN104	133.8	18.7	0.1	<0.2	1.4	41
KN105	35.7	6.0				6
KN108	31.6	10.5				3
KN110A	43.4	16.5	0.2	0.8	1.3	8
KN112	47.0	5.6	<0.1	0.7	0.7	6
KN115	2.3	3.5				<2
KN117	13.2	7.9				3
KN118	41.7	14.7				17
KN119	14.9	9.5				4
KN121	9.0	2.4				<2
KN123	24.2	3.9				6
KN124	24.8	8.0				6
KN125	15.3	11.1				4
KN126A	17.3	6.6				3
KN128	<0.2	<0.2				<2

Geochemical grade assay for rocks, drill cores and drill cuttings samples

Table with columns: Sample No., Detection Limit, SiO2, Al2O3, Fe2O3, MnO, MgO, CaO, Na2O, K2O, TiO2, P2O5, LOI, Total, Be, V, Cr, Co, Ga, Ge, As, Rb, Sr, Y, Zr, Nb, Mo, In, Sn, Sb, Cs, Ba, Sc. Rows include sample identifiers like KNI134, KNI141, KNI142, KNI143, KNI144, KNI145, KNI146, KNI148, KNI150, KNI151, KNI152, KNI153, KNI154, KNI155, KNI156, KNI157, KNI158, KNI159, KNI160, KNI161, KNI162, KNI163, KNI164, KNI165, KNI166, KNI167B, KNI168, KNI169, KNI170, KNI171, KNI172, KNI173, KNI174A, KNI175, KNI176, KNI177, KNI178, KNI179, KNI180, KNI181, KNI182, KNI183, KNI184, KNI185, KNI186, KNI187, KNI188, KNI189, KNI190, KNI191, KNI192, KNI193, KNI194, KNI195, KNI196, KNI197, KNI198, KNI199, KNI200.

Geochemical grade assay for rocks, drill cores and drill cuttings samples

(4continued/8)

Sample No.	Pd	Pt	Ir	Rh	Re	Au
	ppb	ppb	ppb	ppb	ppb	ppb
Detection Limit	0.2	0.2	0.1	0.2	0.1	2
KN134	0.3	<0.2				<2
KN141	7.7	6.9				2
KN142	<0.2	6.3				2
KN143	11.8	9.4				3
KN144	0.9	4.1				3
KN145	<0.2	<0.2				<2
KN146	<0.2	<0.2				<2
KN148	16.6	8.8				4
KN150	76.1	11.2	<0.1	<0.2	2.0	24
KN151	4.6	1.9				<2
KN152	1.3	<0.2				<2
KN153	0.7	<0.2				<2
KN154	1.0	<0.2				<2
KN156	2.4	3.1				<2
KN157	17.0	8.7				5
KN158	17.3	3.8				4
KN159	14.4	9.1				5
KN160	11.6	9.2				4
KN161	16.3	5.5				4
KN162	3.7	0.4				<2
KN163	<0.2	0.2				<2
KN164	13.9	11.1				3
KN165	0.4	0.4				<2
KN166	<0.2	<0.2				<2
KN167B	<0.2	0.7				<2
KN168	1.8	<0.2				3
KN169	12.9	5.3				4
KN170	14.9	6.3				5
KN171	14.5	6.0				4
KN172	<0.2	<0.2				<2
KN173	59.3	13.6	<0.1	0.6	0.7	17
KN174A	16.7	6.0				4
KN175	6.6	4.3				3
KN176	16.0	6.2				4
TM101	3.6	2.9				<2
TM102	6.7	4.6				<2
TM103	13.0	17.0				6
TM104	5.6	5.8				<2
TM105	3.0	1.6				<2
TM106	3.7	3.9				<2
TM107	19.4	5.0				6
TM109	7.7	10.0				7
TM110	8.6	11.0				4
TM111	6.5	9.5				4
TM112	6.7	9.8				7
TM113	1.0	0.6				<2
TM114	0.3	0.7				<2
TM115	3.1	3.9				<2
TM116	1.3	1.7				<2
TM117	2.8	3.4				<2
TM118	3.4	3.8				<2
TM119	0.2	<0.2				<2
TM120	9.1	2.3				3

Geochemical grade assay for rocks, drill cores and drill cuttings samples

Table with 35 columns (Sample No., SiO2, Al2O3, Fe2O3, MnO, MgO, CaO, Na2O, K2O, TiO2, P2O5, LOI, Total, Be, V, Cr, Co, Ga, Ge, As, Rb, Sr, Y, Zr, Nb, Mo, In, Sn, Sb, Cs, Ba, Sc) and 17 rows of data (TM121 to TM173). Each cell contains numerical values or '<math>$'$' symbols.

Geochemical grade assay for rocks, drill cores and drill cuttings samples

(5continued/8)

Sample No.	La	Ce	Pr	Nd	Sm	Eu	Gd	Tb	Dy	Ho	Er	Tm	Yb	Lu	HF	Ta	W	Ti	Bi	Th	U	Ag	Cd	Cu	Ni	Pb	Zn	S
	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	%
	0.05	0.05	0.01	0.05	0.01	0.005	0.01	0.01	0.01	0.01	0.01	0.005	0.01	0.002	0.1	0.01	0.5	0.05	0.1	0.05	0.01	0.3	0.3	1	1	3	30	0.001
TM121	29.37	63.32	8.21	35.83	8.16	2.776	7.96	1.43	6.83	1.25	3.60	0.484	2.90	0.408	5.9	1.43	<0.5	0.12	0.1	2.79	0.59	0.6	0.3	102	16	15	105	0.041
TM122	40.97	88.78	11.13	47.43	10.36	3.247	9.83	1.47	7.87	1.44	4.28	0.545	3.39	0.463	7.5	1.88	0.6	0.19	<0.1	5.07	1.25	0.8	1.0	153	29	13	107	0.007
TM123	33.99	74.54	9.62	42.76	9.72	3.246	9.59	1.46	7.94	1.44	4.16	0.552	3.31	0.486	7.1	1.57	<0.5	0.14	<0.1	3.22	0.69	0.8	0.7	50	<1	7	112	0.041
TM124	28.22	61.48	8.03	35.35	8.02	2.760	8.09	1.22	6.59	1.19	3.45	0.452	2.68	0.369	5.8	1.32	<0.5	0.15	<0.1	2.55	0.64	1.0	11.9	145	48	<3	92	0.036
TM125	44.58	99.01	11.70	51.54	10.98	3.680	11.01	1.71	9.47	1.82	4.92	0.694	4.15	0.575	8.4	1.96	<0.5	0.17	<0.1	4.44	1.03	1.3	11.9	66	<1	4	104	0.008
TM126	41.64	91.14	11.47	50.20	11.06	3.570	10.88	1.63	8.64	1.54	4.45	0.578	3.57	0.489	7.8	2.01	<0.5	0.13	<0.1	3.86	0.82	0.9	13.9	116	8	8	94	0.048
TM127	37.77	81.63	10.33	45.93	10.35	3.366	10.31	1.57	8.39	1.52	4.39	0.571	3.48	0.482	7.2	1.66	<0.5	0.14	<0.1	3.81	0.87	1.2	13.1	45	<1	<3	92	0.012
TM128	30.21	66.34	8.44	36.87	8.33	2.709	8.01	1.25	6.80	1.25	3.71	0.490	2.95	0.407	6.2	1.40	<0.5	0.09	<0.1	2.82	0.60	0.7	11.7	135	29	12	84	0.015
TM129	30.92	68.06	8.61	38.18	8.56	2.823	8.29	1.28	6.92	1.29	3.73	0.483	3.00	0.412	6.1	1.42	<0.5	0.11	<0.1	2.81	0.63	0.9	11.6	143	25	11	82	0.026
TM130	29.16	63.83	8.15	35.56	8.08	2.688	8.12	1.25	6.77	1.24	3.61	0.472	2.94	0.401	6.0	1.37	<0.5	0.07	<0.1	2.72	0.56	0.7	11.2	142	33	11	82	0.024
TM131	32.78	73.13	9.45	42.78	9.61	3.153	9.19	1.44	7.75	1.41	4.08	0.544	3.23	0.440	7.2	1.55	<0.5	0.11	<0.1	3.25	0.69	0.7	12.4	43	<1	13	89	0.031
TM132	35.78	77.87	9.99	44.10	10.17	3.258	9.44	1.48	7.88	1.43	4.23	0.539	3.21	0.445	7.2	1.61	<0.5	0.16	<0.1	3.23	0.68	0.9	13.3	40	<1	4	90	0.053
TM133	40.96	89.73	11.40	49.85	11.09	3.574	10.41	1.56	8.30	1.49	4.33	0.557	3.39	0.468	7.7	1.86	<0.5	0.12	<0.1	3.60	0.77	1.3	14.2	130	<1	16	100	0.055
TM134	33.42	71.85	9.06	40.02	9.05	2.872	8.76	1.38	7.50	1.40	4.15	0.566	3.40	0.471	6.7	1.58	<0.5	0.09	<0.1	3.10	0.65	0.5	13.3	136	39	23	89	0.024
TM135	44.30	94.11	11.98	51.14	11.06	3.493	10.96	1.69	8.75	1.65	4.73	0.633	3.75	0.507	8.1	1.91	<0.5	0.13	<0.1	3.73	0.67	0.6	12.1	188	<1	22	100	0.026
TM136	45.65	96.37	12.29	51.59	10.77	3.465	10.16	1.50	7.69	1.41	4.07	0.510	3.13	0.420	7.2	1.82	<0.5	0.21	0.1	4.17	0.93	0.9	12.2	72	18	21	87	0.021
TM137	50.64	107.23	13.64	57.20	12.21	3.886	11.91	1.84	9.70	1.74	5.02	0.683	4.21	0.573	8.9	1.99	<0.5	0.16	<0.1	4.21	0.85	1.2	12.4	168	<1	6	102	0.028
TM138	38.01	81.07	10.38	43.52	9.34	3.073	9.46	1.44	7.71	1.38	4.08	0.551	3.24	0.447	6.9	1.54	<0.5	0.11	<0.1	3.09	0.63	0.4	13.9	154	19	16	89	0.022
TM139	52.54	110.90	14.03	60.48	12.80	3.955	12.44	1.90	10.09	1.85	5.29	0.718	4.44	0.585	9.4	2.00	<0.5	0.16	<0.1	4.57	0.93	1.7	11.4	157	<1	19	100	0.084
TM140	40.76	85.12	10.81	45.79	9.91	3.210	9.89	1.58	8.41	1.59	4.82	0.632	3.82	0.529	6.7	1.61	<0.5	0.12	<0.1	3.44	0.69	0.9	13.2	87	<1	18	100	0.096
TM141	31.01	64.68	8.23	35.64	7.90	2.657	8.08	1.30	6.77	1.27	3.63	0.486	2.96	0.425	5.5	1.25	<0.5	0.10	<0.1	2.52	0.52	0.8	11.2	125	107	19	84	0.049
TM142	28.33	60.48	7.69	33.96	7.47	2.477	7.63	1.18	6.39	1.17	3.19	0.449	2.74	0.376	5.5	1.32	<0.5	0.10	<0.1	2.78	0.58	0.9	11.7	137	39	18	84	0.063
TM143	224.70	445.11	49.26	187.37	31.02	9.298	25.58	2.96	12.65	1.98	5.14	0.649	2.98	0.367	13.9	15.19	1.1	0.16	<0.1	21.92	4.88	1.1	11.8	70	2	16	55	0.017
TM144	40.03	82.06	10.29	43.39	9.19	3.041	9.16	1.43	7.66	1.46	4.21	0.568	3.45	0.480	6.3	1.63	<0.5	0.14	<0.1	3.16	0.62	0.8	13.8	63	<1	15	100	0.067
TM145	45.95	95.64	11.99	50.97	11.07	3.484	11.04	1.73	9.33	1.72	5.13	0.709	4.28	0.580	7.3	1.81	<0.5	0.22	<0.1	3.88	0.79	1.1	14.4	76	<1	19	100	0.095
TM146	33.51	69.83	8.83	37.43	7.98	2.663	8.17	1.28	6.64	1.24	3.52	0.481	2.91	0.401	5.5	1.45	<0.5	0.11	<0.1	3.71	0.56	0.6	13.2	58	11	21	93	0.069
TM147	33.27	68.83	8.74	37.07	8.07	2.693	7.94	1.28	6.88	1.29	3.77	0.514	3.18	0.442	5.8	1.48	<0.5	0.09	<0.1	2.82	0.56	0.7	14.5	154	11	7	96	0.093
TM148	43.47	90.09	11.34	46.94	9.86	3.263	9.89	1.53	8.20	1.52	4.40	0.611	3.64	0.509	7.1	1.86	<0.5	0.10	0.1	3.66	0.74	0.7	11.0	16	<1	<3	99	0.053
TM149	44.86	93.13	11.76	49.40	10.54	3.347	10.55	1.65	8.74	1.65	4.70	0.630	3.80	0.520	7.4	1.89	<0.5	0.09	<0.1	3.76	0.75	0.8	13.4	65	<1	13	100	0.041
TM150	37.03	79.24	10.24	43.16	9.27	2.975	8.94	1.37	7.33	1.34	3.82	0.488	3.04	0.419	7.1	1.56	<0.5	0.13	<0.1	3.17	0.67	0.9	12.7	85	<1	4	97	0.008
TM151	44.02	93.14	12.03	51.06	10.84	3.525	10.55	1.64	8.65	1.58	4.43	0.588	3.59	0.487	7.9	1.94	<0.5	0.11	<0.1	3.60	0.75	0.9	14.4	133	<1	12	106	0.063
TM152	42.38	90.37	11.44	50.02	10.93	3.502	10.96	1.68	9.11	1.70	4.84	0.657	4.03	0.551	7.6	1.97	<0.5	0.13	<0.1	3.72	0.69	1.0	13.3	130	<1	20	112	0.009
TM153	25.63	53.68	7.15	32.12	7.44	2.594	7.48	1.14	6.28	1.16	3.27	0.427	2.66	0.361	5.2	1.08	<0.5	0.11	<0.1	2.08	0.45	0.6	10.2	228	53	14	97	0.010
TM154	16.90	46.50	4.42	19.18	4.69	1.492	4.90	0.93	5.70	1.17	3.61	0.548	3.53	0.494	6.4	1.32	<0.5	0.21	<0.1	2.67	0.64	1.1	18.7	313	44	23	159	0.001
TM155	18.38	36.98	4.76	20.77	4.99	1.637	5.46	0.93	5.54	1.07	3.24	0.477	2.97	0.416	3.8	0.73	<0.5	0.14	<0.1	1.79	0.37	0.5	12.9	186	57	4	207	0.010
TM156	17.99	37.40	4.78	20.93	4.94	1.708	5.45	0.95	5.54	1.10	3.29	0.482	2.98	0.430	3.8	0.73	<0.5	<0.05	<0.1	1.83	0.37	0.6	11.0	184	44	11	79	0.005
TM157	22.93	43.68	5.95	25.73	6.19	2.059	6.71	1.14	6.79	1.33	4.03	0.583	3.58	0.513	4.2	0.83	<0.5	0.09	<0.1	1.98	0.37	0.6	11.4	177	23	21	90	0.005
TM158	37.31	76.39	9.70	41.11	9.19	2.877	9.26	1.53	8.56	1.66	4.89	0.704	4.27	0.609	6.6	1.60	<0.5	0.14	<0.1	3.68	0.80	0.9	12.6	110	<1	12	103	0.008
TM159	27.03	57.50	7.18	31.39	6.96	2.360	6.85	1.10	6.08	1.13	3.25	0.428	2.70	0.381														

Geochemical grade assay for rocks, drill cores and drill cuttings samples

(5continued/8)

Sample No.	Pd	Pt	Ir	Rh	Re	Au
	ppb	ppb	ppb	ppb	ppb	ppb
Detection Limit	0.2	0.2	0.1	0.2	0.1	2
TM121	2.8	2.6				<2
TM122	5.7	6.9				2
TM123	1.6	0.9				<2
TM124	5.0	6.3				<2
TM125	3.3	2.3				<2
TM126	22.8	4.1	0.2	<0.2	0.6	8
TM127	0.4	0.8				<2
TM128	5.9	9.2				2
TM129	17.7	10.1	0.3	0.7	0.7	2
TM130	8.1	9.5	1.8	0.7	0.6	4
TM131	0.3	0.8				<2
TM132	0.7	0.5				<2
TM133	0.5	0.7				<2
TM134	9.0	9.2				3
TM135	8.0	11.2				3
TM136	0.7	1.1				<2
TM137	0.4	4.7				<2
TM138	7.5	9.9				3
TM139	<0.2	2.5				<2
TM140	2.0	2.4				<2
TM141	6.1	7.9				<2
TM142	11.6	9.3				27
TM143	2.8	2.5				<2
TM144	1.7	1.7				<2
TM145	4.0	1.4				<2
TM146	1.8	1.3				<2
TM147	1.6	2.9				<2
TM148	1.5	1.2				<2
TM149	0.8	1.7				<2
TM150	0.3	0.9				<2
TM151	0.4	0.9				<2
TM152	0.5	1.9				<2
TM153	10.3	12.1	0.1	0.3	0.6	3
TM154	27.4	8.7				6
TM155	21.7	6.2				3
TM156	15.1	10.6				3
TM157	10.8	4.6				<2
TM158	<0.2	0.6				<2
TM159	1.6	2.6				<2
TM160	1.4	2.8				<2
TM161	9.3	5.7				3
TM162	6.5	3.5				6
TM163	9.0	13.8				4
TM164	9.4	1.7	<0.1	0.5	1.3	2
TM165	0.3	<0.2				<2
TM166	15.5	5.4				3
TM167	18.8	5.8				5
TM168	7.8	6.2				3
TM169	1.2	2.8				<2
TM170	4.1	3.9				3
TM171	18.5	6.5				<2
TM172	10.3	7.8				3
TM173	4.8	7.3				<2

Geochemical grade assay for rocks, drill cores and drill cuttings samples

Sample No.	SiO ₂ %	Al ₂ O ₃ %	Fe ₂ O ₃ %	MnO %	MgO %	CaO %	Na ₂ O %	K ₂ O %	TiO ₂ %	P ₂ O ₅ %	LOI %	Total %	Be ppm	V ppm	Cr ppm	Co ppm	Ga ppm	Ge ppm	As ppm	Rb ppm	Sr ppm	Y ppm	Zr ppm	Nb ppm	Mo ppm	In ppm	Sn ppm	Sb ppm	Cs ppm	Ba ppm	Sc ppm
Detection Limit	0.01	0.01	0.01	0.001	0.01	0.01	0.01	0.01	0.001	0.01	0.01	0.01	1	5	20	1	1	0.5	5	1	2	0.5	1	0.2	2	0.1	1	0.2	0.1	3	1
TM174	50.57	11.85	15.92	0.224	5.47	8.30	2.32	1.11	2.203	0.23	1.62	99.83	<1	405	102	45	18	1.7	<5	20	235	61.9	137	12.9	<2	<0.1	1	<0.2	0.2	351	42
TM175	49.30	12.78	14.38	0.199	5.54	9.00	2.45	0.58	2.093	0.25	3.39	99.96	<1	419	62	40	19	1.5	<5	16	284	33.4	145	13.8	<2	<0.1	1	<0.2	0.3	290	41
TM176	52.16	12.49	15.37	0.214	4.42	8.47	2.57	1.05	2.284	0.28	0.60	99.78	<1	433	33	40	20	1.5	<5	23	283	38.7	172	15.5	<2	<0.1	1	<0.2	0.2	289	41
TM177	52.04	11.83	15.15	0.213	5.52	8.94	2.55	1.08	2.081	0.24	0.60	100.23	<1	401	90	42	18	1.7	<5	22	218	35.3	148	13.1	<2	<0.1	1	<0.2	0.2	289	45
TM178	49.79	12.83	15.24	0.232	4.99	8.82	2.64	1.30	2.339	0.27	1.28	99.91	1	357	53	40	19	1.6	<5	26	272	36.3	153	14.2	<2	<0.1	1	<0.2	0.2	356	42
TM179	51.37	13.08	14.32	0.219	5.34	9.31	2.46	0.94	2.009	0.24	0.99	100.27	<1	392	73	42	18	1.5	<5	18	259	32.0	132	12.6	<2	<0.1	1	<0.2	0.1	298	42
WW017	66.40	13.24	6.38	0.108	1.16	2.30	2.68	4.19	0.968	0.28	2.79	100.50	3	95	<20	13	22	2.0	<5	247	113	49.6	273	21.6	<2	<0.1	5	0.3	10.1	893	19
WW024	52.52	13.73	13.17	0.185	4.92	8.46	2.53	1.39	1.397	0.19	1.44	99.93	1	341	26	40	22	1.7	<5	43	225	31.9	158	12.7	<2	<0.1	2	0.3	1.3	331	37
WW026	67.34	12.53	6.45	0.099	1.04	2.72	3.21	4.26	1.014	0.31	1.48	100.44	3	78	<20	10	18	1.7	<5	182	129	43.8	255	20.5	<2	<0.1	4	0.3	5.9	628	17
WW031	52.75	12.98	14.34	0.219	4.20	8.11	2.72	1.41	1.695	0.21	1.06	99.70	<1	413	<20	38	21	1.8	<5	64	172	37.0	164	9.8	<2	<0.1	1	<0.2	2.9	299	40
WW056	49.66	13.05	14.30	0.206	5.54	9.80	2.28	0.94	2.180	0.27	2.11	100.34	<1	428	96	40	20	1.6	<5	40	323	28.6	147	13.4	<2	<0.1	1	<0.2	4.1	324	39
WW068	69.33	13.94	3.74	0.069	1.17	2.40	3.64	2.84	0.456	0.14	2.62	100.32	2	52	62	9	16	1.2	<5	74	476	18.8	155	9.1	<2	<0.1	1	<0.2	0.9	1210	8
WW069	65.83	12.75	6.81	0.136	1.43	3.48	3.15	3.95	1.037	0.29	1.57	100.43	3	101	<20	13	19	1.5	<5	180	142	37.5	243	19.9	<2	<0.1	4	<0.2	7.0	573	19
WW073	54.69	12.25	15.44	0.200	3.09	6.82	2.75	2.38	1.872	0.27	0.60	100.37	1	479	<20	38	21	1.9	<5	145	153	41.6	172	12.2	<2	<0.1	2	0.3	9.3	329	39
WW076	50.90	14.71	10.67	0.150	5.58	9.89	2.21	1.23	1.181	0.17	1.46	98.86	<1	257	148	38	20	1.4	<5	42	215	25.1	135	9.7	<2	<0.1	1	<0.2	0.7	285	35
WW077	52.88	14.71	10.64	0.155	6.19	9.16	2.38	1.48	1.044	0.16	1.57	100.37	<1	232	209	38	17	1.7	<5	48	195	21.7	122	8.2	<2	<0.1	1	<0.2	0.9	309	35
WW083	53.03	12.99	14.48	0.199	4.40	8.41	2.79	1.35	1.686	0.22	0.69	100.24	<1	415	<20	38	21	1.6	<5	48	191	102.0	158	10.9	<2	<0.1	1	0.7	1.0	331	39
WW092	51.04	13.55	13.12	0.213	6.30	10.45	2.41	0.50	1.195	0.13	1.47	100.37	<1	351	30	42	18	1.7	<5	16	161	27.1	95	5.8	<2	<0.1	1	<0.2	0.7	140	42
WW095	49.51	13.20	14.35	0.196	5.85	9.61	2.20	0.20	1.427	0.14	3.75	100.53	<1	406	49	45	20	1.6	<5	6	161	30.6	105	5.5	<2	<0.1	1	<0.2	0.4	80	42
WW099	53.57	14.64	10.03	0.163	6.49	8.76	2.22	1.61	1.059	0.17	1.72	100.43	<1	244	181	37	19	1.5	<5	60	202	22.9	130	8.7	<2	<0.1	1	<0.2	1.5	343	34
WW117A	53.12	13.78	12.23	0.180	5.00	8.91	2.43	1.32	1.289	0.18	1.49	99.92	1	351	40	42	21	1.8	<5	84	234	35.2	130	9.8	<2	<0.1	2	<0.2	4.2	716	37
WW122	50.48	13.02	13.13	0.165	4.06	7.89	2.50	1.85	1.426	0.60	2.22	100.04	2	337	26	32	26	1.6	<5	28	778	36.1	340	30.8	<2	<0.1	3	<0.2	0.7	673	22
WW129	54.35	14.04	11.80	0.206	4.44	8.41	2.71	1.17	1.429	0.19	1.62	100.38	1	368	20	42	22	1.7	<5	62	234	37.1	155	13.1	<2	<0.1	3	<0.2	2.3	387	37
WW130	51.71	14.08	11.64	0.206	6.10	10.60	2.53	0.50	1.320	0.12	0.07	98.86	<1	361	71	34	18	1.2	<5	18	158	27.3	89	5.0	<2	<0.1	1	<0.2	0.4	118	42
WW134	55.47	13.76	11.47	0.178	4.84	7.64	2.96	1.68	1.092	0.15	0.85	100.99	<1	288	<20	38	19	1.9	<5	71	216	29.2	132	9.4	<2	<0.1	1	<0.2	2.0	368	39
WW161	51.70	15.23	11.19	0.150	5.81	9.82	2.22	1.24	1.061	0.17	1.35	99.93	1	252	71	43	18	1.7	<5	38	236	23.8	126	8.7	<2	<0.1	1	<0.2	0.8	330	36
WW162	55.55	12.43	12.97	0.200	2.25	5.84	2.66	2.55	1.784	0.24	1.71	99.18	2	375	<20	40	21	1.7	<5	99	190	40.5	235	16.5	<2	<0.1	2	<0.2	2.9	487	34
WW163	57.08	12.70	13.64	0.181	2.24	6.00	2.75	2.59	1.776	0.25	0.80	100.00	2	365	<20	38	22	1.9	<5	99	190	42.7	245	17.9	<2	<0.1	2	<0.2	2.8	508	35
WW164	51.69	14.09	10.89	0.168	6.36	9.99	2.26	1.08	1.061	0.15	1.33	99.07	1	278	112	40	17	1.6	<5	24	197	22.5	107	8.8	<2	<0.1	1	<0.2	0.5	243	37
WW165	56.85	12.31	12.85	0.163	2.26	5.57	2.71	2.58	1.744	0.26	1.48	98.79	2	398	<20	34	21	1.9	7	102	178	39.9	219	16.4	<2	<0.1	3	<0.2	4.5	500	33
WW166	57.81	12.32	13.15	0.166	2.36	5.60	2.59	2.69	1.726	0.24	1.25	99.90	2	326	<20	35	20	1.6	<5	96	175	39.9	229	16.6	<2	<0.1	2	<0.2	2.5	465	33
WW167	52.66	13.56	13.21	0.222	5.20	8.83	2.52	0.95	1.318	0.19	2.12	99.77	1	320	30	39	19	1.4	<5	36	212	28.7	140	11.4	<2	<0.1	1	<0.2	1.9	332	36
WW174	53.01	13.21	13.18	0.184	5.00	8.85	2.07	0.81	1.285	0.16	2.30	100.05	1	378	22	48	19	1.7	<5	41	207	29.2	129	9.8	<2	<0.1	1	<0.2	1.1	325	42
WW175	52.39	14.29	13.30	0.194	6.78	10.22	2.01	0.88	0.882	0.10	1.20	99.86	<1	283	72	46	17	1.6	<5	32	185	18.3	83	6.5	<2	<0.1	1	<0.2	2.3	499	34
WW177	56.96	12.39	13.30	0.194	2.33	5.84	2.60	2.80	1.704	0.24	0.92	99.27	2	350	<20	37	22	1.8	<5	102	186	43.1	235	16.5	<2	<0.1	2	<0.2	0.7	282	40
WW181	52.45	14.53	10.84	0.173	6.59	10.20	2.31	1.02	0.962	0.11	1.05	100.22	<1	280	96	42	20	1.7	<5	30	191	21.7	95	7.1	<2	<0.1	1	<0.2	2.8	311	36
WW183A	52.62	13.60	12.52	0.197	5.07	8.81	2.66	0.94	1.358	0.18	1.42	99.38	1	322	32	40	20	1.7	<5	52	222	29.9	127	10.5	<2	<0.1	1	<0.2	0.8	209	39
WW184	52.97	13.77	11.37	0.204	6.19	9.92	2.38	1.05	1.091	0.13	1.06	100.13	<1	306	56	41	18	1.7	<5	35	187	22.5	98	6.9	<2	<0.1	1	<0.2	0.8	197	40
WW191	55.62	14.56	10.19	0.150	3.87	7.77	2.38	2.23	1.275	0.20	1.93	100.19	2	228	61	31	19	1.6	<5	70	204	29.4	193	13.6	<2	<0.1	2	<0.2	0.6	455	32
WW196	56.92	14.51	13.43	0.180	2.43	6.01	2.65	2.52	1.698	0.24	1.36	99.95	2	339	<20	37	22	1.7	<5	93	193	41.1	234	16.8	<2	<0.1	2	<0.2	2.3	505	34
WW197	52.52	14.11	11.18	0.199	6.40	10.34	2.24	1.08	1.069	0.13	0.90	100.17	1	298	87	40	18	1.8	<5	33	198	22.5	108	7.8	<2	<0.1	1	<0.2	0.9	232	39
WW203	56.85	12.33	13.61	0.154	2.31	4.78	2.72	2.74	1.778	0.24																					

Geochemical grade assay for rocks, drill cores and drill cuttings samples

(6continued/8)

Sample No.	La	Ce	Pr	Nd	Sm	Eu	Gd	Tb	Dy	Ho	Er	Tm	Yb	Lu	HF	Ta	W	Ti	Bi	Th	U	Ag	Cd	Cu	Ni	Pb	Zn	S	
	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	%
	Detection Limit																												
TM174	31.03	39.36	7.53	34.32	7.83	2.730	9.91	1.61	9.58	1.96	5.86	0.838	4.95	0.735	4.1	0.78	1.8	0.08	<0.1	1.88	0.41	0.5	1.5	227	59	4	128	0.014	
TM175	20.74	42.95	5.30	23.08	5.38	1.779	5.85	1.03	6.07	1.17	3.47	0.494	3.07	0.443	4.3	0.84	<0.5	<0.05	<0.1	2.01	0.18	0.6	<0.3	166	29	16	105	0.017	
TM176	22.94	47.94	6.06	26.47	6.12	2.010	6.67	1.16	6.82	1.35	4.05	0.587	3.64	0.537	4.7	0.96	<0.5	<0.05	0.2	2.37	0.68	0.6	1.4	156	14	16	114	0.015	
TM177	19.03	40.19	5.04	22.42	5.53	1.754	6.28	1.07	6.38	1.27	3.85	0.574	3.45	0.500	4.3	0.82	<0.5	0.09	<0.1	2.00	0.39	0.5	1.1	229	46	12	108	0.014	
TM178	21.14	44.44	5.66	24.67	5.83	1.902	6.30	1.10	6.72	1.32	3.99	0.578	3.58	0.526	4.4	0.90	<0.5	0.15	<0.1	2.26	0.55	0.8	0.5	711	23	8	114	0.015	
TM179	18.32	38.71	4.86	21.64	5.07	1.726	5.68	0.97	5.72	1.14	3.47	0.515	3.19	0.462	3.9	0.76	<0.5	0.06	<0.1	1.92	0.39	0.6	1.2	256	33	8	102	0.011	
WW017	47.93	104.87	11.58	45.69	9.58	2.085	9.22	1.54	9.33	1.79	5.33	0.788	4.84	0.719	7.3	1.97	1.3	1.46	0.7	12.98	4.40	<0.3	<0.3	61	2	19	89	<0.001	
WW024	51.96	51.96	5.93	24.16	5.43	1.665	5.71	0.99	5.93	1.22	3.58	0.532	3.24	0.487	4.2	0.87	<0.5	0.15	0.2	5.36	1.16	<0.3	0.5	176	33	<3	100	<0.001	
WW026	41.87	88.14	9.56	39.14	8.28	1.782	8.21	1.37	7.83	1.58	4.65	0.683	3.70	0.552	7.4	1.83	1.1	1.34	1.0	12.26	3.76	<0.3	0.4	126	1	18	85	<0.001	
WW031	19.50	42.41	5.03	22.83	5.85	1.703	6.38	1.14	6.77	1.39	4.10	0.607	3.37	0.500	4.6	0.66	<0.5	0.53	0.2	4.31	0.99	<0.3	0.4	179	28	6	107	<0.001	
WW056	20.35	45.76	5.34	23.37	5.22	1.755	5.47	0.91	5.37	1.10	3.14	0.468	2.54	0.383	4.2	0.90	<0.5	0.46	<0.1	2.42	0.50	<0.3	<0.3	265	54	<3	105	<0.001	
WW068	43.97	81.25	8.21	29.90	5.01	1.457	3.98	0.57	3.28	0.63	1.96	0.280	1.60	0.233	4.0	0.70	<0.5	0.71	0.1	6.46	0.71	<0.3	<0.3	12	25	23	49	0.004	
WW069	39.81	85.18	9.31	37.34	7.99	1.680	7.40	1.24	7.22	1.38	4.19	0.599	3.36	0.489	6.9	1.77	1.2	1.44	0.3	12.30	4.27	<0.3	<0.3	55	5	26	87	0.003	
WW073	23.85	51.34	6.00	25.58	6.30	1.884	6.98	1.25	7.76	1.58	4.70	0.707	3.92	0.595	5.1	0.98	0.7	0.91	0.2	6.18	1.75	<0.3	<0.3	250	23	10	126	<0.001	
WW076	18.13	40.69	4.57	19.32	4.47	1.307	4.61	0.77	4.61	0.94	2.80	0.403	2.24	0.351	3.8	0.63	<0.5	0.31	<0.1	3.90	0.63	<0.3	0.3	80	86	6	84	<0.001	
WW077	16.62	36.61	4.11	17.53	3.95	1.162	4.08	0.69	4.04	0.81	2.41	0.344	1.98	0.284	3.4	0.58	<0.5	0.42	0.1	3.46	0.85	<0.3	<0.3	151	103	<3	82	<0.001	
WW083	31.52	45.89	7.68	34.16	9.21	3.174	12.77	2.42	15.81	3.41	10.09	1.391	6.83	0.943	4.5	0.70	<0.5	0.36	<0.1	4.30	0.90	<0.3	<0.3	167	30	<3	113	<0.001	
WW092	10.66	24.37	2.94	13.26	3.62	1.231	4.31	0.78	4.94	1.04	2.95	0.441	2.60	0.363	2.9	0.42	<0.5	0.50	<0.1	2.44	0.69	<0.3	<0.3	166	48	<3	88	<0.001	
WW095	10.79	26.22	3.15	14.81	4.10	1.423	5.00	0.92	5.92	1.17	3.45	0.523	2.92	0.420	3.2	0.41	<0.5	0.20	<0.1	2.39	0.60	<0.3	0.6	217	43	<3	100	<0.001	
WW099	19.71	44.13	4.85	20.43	4.58	1.266	4.52	0.76	4.37	0.90	2.62	0.390	2.24	0.301	3.9	0.66	<0.5	0.54	<0.1	4.49	0.99	<0.3	0.4	63	117	9	78	0.002	
WW117A	20.72	44.41	5.16	22.45	5.30	1.644	6.07	1.05	6.21	1.26	3.72	0.539	3.12	0.428	3.7	0.72	<0.5	2.38	0.1	4.37	1.23	<0.3	<0.3	239	41	8	100	<0.001	
WW122	47.79	108.98	12.48	54.97	11.73	3.722	10.24	1.94	8.13	1.41	3.78	0.473	2.79	0.349	9.2	2.31	<0.5	0.22	<0.1	4.56	1.03	<0.3	<0.3	226	40	7	115	<0.001	
WW129	25.76	55.77	6.17	26.00	5.98	1.693	6.26	1.10	6.77	1.36	3.78	0.587	3.38	0.503	4.3	0.85	<0.5	0.95	<0.1	5.64	1.25	<0.3	<0.3	183	31	8	102	<0.001	
WW130	8.53	20.97	2.60	12.38	3.67	1.281	4.39	0.84	5.25	1.09	3.19	0.453	2.63	0.366	2.7	0.37	<0.5	0.14	<0.1	1.64	0.48	<0.3	<0.3	186	57	<3	94	0.012	
WW134	20.62	45.05	4.96	21.00	4.69	1.332	4.90	0.81	4.94	0.97	2.94	0.428	2.34	0.341	3.9	0.70	<0.5	0.77	0.1	5.17	1.19	<0.3	<0.3	88	40	8	93	<0.001	
WW161	23.85	47.62	5.72	23.66	5.05	1.504	5.12	0.85	4.93	0.86	2.92	0.422	2.58	0.364	3.3	0.51	<0.5	0.23	<0.1	2.97	0.59	<0.3	<0.3	106	270	7	87	<0.001	
WW162	45.84	89.46	10.54	42.51	9.03	2.176	8.59	1.44	8.16	1.58	4.79	0.699	4.21	0.593	6.0	1.11	0.7	0.49	<0.1	7.71	1.66	0.4	<0.3	92	8	15	131	<0.001	
WW163	45.59	89.52	10.59	43.27	9.19	2.232	8.64	1.44	8.42	1.62	5.02	0.709	4.34	0.604	6.3	1.11	0.9	0.51	<0.1	7.83	1.69	0.4	<0.3	94	<1	13	136	<0.001	
WW164	19.73	38.58	4.67	19.82	4.47	1.361	4.63	0.79	4.73	0.91	2.88	0.403	2.43	0.382	2.9	0.46	<0.5	0.15	<0.1	2.88	0.60	<0.3	<0.3	110	155	<3	88	<0.001	
WW165	43.11	80.97	9.84	39.72	8.69	2.162	8.70	1.45	8.19	1.56	4.84	0.706	4.23	0.593	5.8	1.20	1.2	0.58	0.3	7.21	1.91	<0.3	<0.3	109	32	13	124	<0.001	
WW166	43.64	85.97	10.22	41.21	8.77	2.159	8.44	1.43	8.18	1.58	4.87	0.708	4.22	0.614	6.1	1.09	0.7	0.49	0.2	7.51	1.50	<0.3	<0.3	77	6	10	126	<0.001	
WW167	27.05	53.17	6.38	26.47	5.90	1.748	6.10	1.03	6.00	1.16	3.52	0.514	3.10	0.434	3.8	0.68	<0.5	0.30	<0.1	3.98	0.86	<0.3	<0.3	153	123	10	99	<0.001	
WW174	26.18	49.58	6.41	26.97	6.06	1.819	6.53	1.06	6.10	1.19	3.54	0.497	3.08	0.424	3.5	0.61	<0.5	0.47	<0.1	3.72	0.86	<0.3	<0.3	156	150	11	108	<0.001	
WW175	14.38	29.58	3.63	15.23	3.61	1.143	3.64	0.65	3.79	0.74	2.25	0.328	1.93	0.280	2.3	0.39	<0.5	0.18	<0.1	2.36	0.58	<0.3	<0.3	172	261	<3	82	<0.001	
WW177	42.03	82.93	9.96	39.73	8.56	2.073	8.12	1.39	8.26	1.58	4.83	0.724	4.24	0.602	6.0	1.10	0.8	0.45	<0.1	8.03	1.68	0.4	<0.3	76	7	13	139	<0.001	
WW181	18.49	30.04	4.50	18.96	4.28	1.361	4.51	0.79	4.47	0.84	2.60	0.361	2.14	0.296	2.5	0.40	<0.5	0.14	<0.1	2.35	0.49	<0.3	<0.3	71	179	4	85	<0.001	
WW183A	19.75	41.03	5.17	20.51	4.72	1.509	5.03	0.92	5.34	1.16	3.00	0.472	2.79	0.410	3.5	0.67	<0.5	0.53	0.2	4.94	1.18	<0.3	<0.3	157	130	5	104	<0.001	
WW184	14.57	30.05	3.78	16.09	4.00	1.283	4.26	0.78	4.54	0.87	2.63	0.386	2.32	0.323	2.7	0.43	<0.5	0.22	<0.1	2.51	0.61	<0.3	<0.3	142	118	<3	86	<0.001	
WW191	35.65	71.08	8.48	33.71	7.03	1.745	6.52	1.07	6.06	1.16	3.61	0.513	3.17	0.445	5.1	0.87	<0.5	0.35	<0.1	7.24									

Geochemical grade assay for rocks, drill cores and drill cuttings samples

(6continued/8)

Sample No.	Pd ppb	Pt ppb	Ir ppb	Rh ppb	Re ppb	Au ppb
TM174	0.2	0.2	0.1	0.2	0.1	2
TM175	4.7	7.3				3
TM176	23.3	6.8				2
TM177	14.1	4.4				3
TM178	10.1	5.2				5
TM179	18.2	3.2				5
WW017	14.6	4.8				4
WW024	1.4	1.3				1
WW026	11.7	9.9				3
WW031	1.7	1.2				3
WW056	8.8	6.9				3
WW068	14.3	11.3				5
WW069	0.2	0.1				<1
WW073	0.4	1.5				1
WW076	4.8	8.8				1
WW077	6.9	4.9				16
WW083	5.6	5.7				2
WW092	9.2	3.8				4
WW095	8.7	2.3				2
WW099	8.8	4.8				3
WW117A	1.1	2.5				2
WW122	12.8	10.6				3
WW130	6.9	3.7				5
WW134	11.5	9.7				3
WW161	12.1	5.5				3
WW162	13.7	12.6	0.2	1.0	0.8	3
WW163	8.3	5.5				3
WW164	0.5	1.3				<2
WW165	0.8	1.9				<2
WW166	6.3	5.7				<2
WW167	0.3	0.4				<2
WW174	0.4	1.2				<2
WW175	9.4	8.9				3
WW177	1.7	2.5				5
WW181	33.4	11.8	<0.1	<0.2	0.6	<2
WW183A	0.8	1.6				<2
WW184	2.8	5.2				2
WW191	10.7	8.4				3
WW196	20.6	6.1	0.3	0.8	1.2	7
WW197	2.7	3.2				5
WW203	0.7	1.3				3
WW204	3.2	6.0				6
WW205	0.3	0.7				<2
WW209	0.4	0.6				<2
WW210	3.4	4.5				3
WW219	10.3	9.1				3
WW221	10.3	10.1				3
WW222	0.5	0.2				<2
WW223	0.5	<0.2				<2
YM105	5.4	4.7				<2
YM108	5.4	8.6				3
YM109	12.1	14.7				3
YM114	16.2	7.8	0.1	0.2	0.4	5
	11.9	10.7				3
	8.8	3.8	<0.1	0.6	0.9	<2

Geochemical grade assay for rocks, drill cores and drill cuttings samples

Sample No.	SiO ₂	Al ₂ O ₃	Fe ₂ O ₃	MnO	MgO	CaO	Na ₂ O	K ₂ O	TiO ₂	P ₂ O ₅	LOI	Total	Be	V	Cr	Co	Ga	Ge	As	Rb	Sr	Y	Zr	Nb	Mo	In	Sn	Sb	Cs	Ba	Sc	
Detection Limit	%	%	%	%	%	%	%	%	%	%	%	%	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm
YM115	52.51	12.55	16.01	0.225	4.12	8.20	2.65	1.20	1.890	0.22	0.73	100.29	1	493	<20	46	22	1.5	<5	45	176	42.9	163	10.2	<2	<0.1	2	<0.2	1.5	289	41	
YM117	50.16	12.54	15.65	0.222	4.81	8.68	2.67	1.20	2.319	0.28	1.49	100.01	1	443	52	44	21	1.6	<5	28	263	39.6	181	13.1	<2	<0.1	1	<0.2	0.3	334	42	
YM118	49.36	13.07	14.52	0.263	5.35	9.91	2.39	0.90	2.385	0.31	1.75	100.17	1	414	106	41	20	1.5	<5	13	339	30.7	157	13.3	<2	<0.1	<1	<0.2	0.2	321	37	
YM119	50.17	13.03	13.09	0.215	6.17	10.10	2.34	0.84	1.756	0.20	1.88	100.13	<1	384	87	43	19	1.7	<5	16	283	27.9	130	11.3	<2	<0.1	<1	<0.2	0.1	252	44	
YM120	50.61	12.79	13.64	0.179	5.81	9.87	2.33	1.50	1.721	0.19	1.58	100.22	1	388	102	44	19	1.6	<5	48	266	28.6	131	10.4	<2	<0.1	<1	<0.2	0.5	262	44	
YM121	51.20	12.54	15.16	0.225	4.74	8.64	2.55	1.48	2.330	0.28	0.74	100.03	1	446	50	42	20	1.6	<5	39	263	38.0	179	14.3	<2	<0.1	1	<0.2	0.4	322	42	
YM122	51.35	12.65	15.27	0.202	4.88	8.78	2.49	1.27	2.329	0.28	0.74	100.23	1	414	46	40	20	1.6	<5	26	248	35.8	170	13.6	<2	<0.1	<1	<0.2	0.2	305	42	
YM124	49.40	12.11	15.58	0.214	4.91	8.51	2.63	1.42	3.374	0.47	1.45	100.06	1	447	<20	42	23	2.0	<5	29	415	35.2	231	19.2	<2	<0.1	2	<0.2	0.4	410	34	
YM125	51.06	12.61	15.19	0.221	4.64	8.73	2.56	1.28	2.444	0.27	1.05	100.04	1	445	70	43	21	1.6	<5	26	248	39.5	192	15.4	<2	<0.1	2	<0.2	0.2	343	41	
YM126	51.04	12.78	14.95	0.219	4.93	8.79	2.40	1.32	2.315	0.27	1.36	100.36	1	430	61	42	20	1.2	<5	31	245	38.3	184	14.3	<2	<0.1	<1	<0.2	1.0	332	41	
YM127	51.83	12.42	15.28	0.230	4.65	8.66	2.43	1.38	2.277	0.26	0.97	100.39	1	447	53	43	21	1.6	<5	31	247	38.1	173	14.1	<2	<0.1	1	<0.2	0.5	321	41	
YM128	51.31	12.46	15.20	0.202	4.75	8.61	2.52	1.25	2.260	0.27	0.87	99.70	1	430	55	39	20	1.5	<5	29	242	38.3	173	14.0	<2	<0.1	2	<0.2	0.4	315	41	
YM129	49.44	11.94	15.97	0.207	5.53	8.77	2.57	1.08	2.267	0.25	1.81	99.84	1	454	72	46	18	1.6	<5	15	218	35.1	159	13.0	<2	<0.1	2	<0.2	<0.1	277	44	
YM130	50.43	13.37	13.61	0.202	6.00	9.99	2.38	0.89	1.843	0.22	1.34	100.27	<1	437	150	44	19	1.6	<5	19	304	29.3	129	11.0	<2	<0.1	1	<0.2	0.2	282	40	
YM131	50.90	12.38	15.65	0.216	4.64	8.67	2.67	1.26	2.364	0.27	0.98	99.99	<1	467	60	42	20	1.9	<5	28	248	39.5	176	14.2	<2	<0.1	2	<0.2	0.3	327	42	
YM132	50.77	13.28	14.08	0.207	5.71	9.80	2.51	0.96	1.960	0.24	0.78	100.31	<1	408	97	42	19	1.5	<5	19	306	28.0	138	12.0	<2	<0.1	1	<0.2	0.1	297	40	
YM133	50.35	12.98	13.20	0.192	6.13	9.85	2.35	1.07	1.633	0.19	1.68	99.63	<1	377	93	42	17	1.6	<5	19	242	26.2	117	10.0	<2	<0.1	1	<0.2	0.2	233	43	
YM134	51.19	12.69	15.22	0.206	5.05	9.00	2.58	1.17	2.225	0.27	0.75	100.35	1	440	54	40	20	1.7	<5	24	262	36.6	164	13.5	<2	<0.1	2	<0.2	0.2	312	42	
YM135	50.82	12.52	15.56	0.231	4.72	8.84	2.60	1.20	2.336	0.34	0.96	100.14	1	456	69	42	21	1.5	<5	30	257	41.0	184	12.6	<2	<0.1	1	0.9	0.4	363	41	
YM136	49.81	12.51	15.22	0.235	5.03	9.06	2.36	0.81	2.261	0.27	1.71	99.28	<1	433	55	40	19	1.6	<5	14	240	36.1	164	12.9	<2	<0.1	2	<0.2	0.8	302	41	
YM137	49.18	13.40	13.22	0.200	6.44	10.31	2.32	0.92	1.887	0.21	1.21	99.30	<1	402	153	47	19	1.4	<5	18	367	25.1	133	12.6	<2	<0.1	1	<0.2	0.2	290	37	
YM138	48.97	12.84	14.73	0.185	4.62	8.84	2.63	1.26	3.139	0.43	1.73	99.36	1	460	26	40	22	1.5	<5	24	425	31.3	199	17.7	<2	<0.1	2	<0.2	0.2	378	32	
YM139	53.92	11.92	14.48	0.157	2.80	5.82	2.91	2.40	2.670	0.79	2.25	100.11	3	222	<20	26	27	1.7	<5	51	438	54.1	378	27.8	<2	<0.1	3	<0.2	0.5	657	23	
YM140	50.17	12.53	14.90	0.250	5.34	8.92	2.37	1.06	2.148	0.26	1.46	99.40	<1	416	76	41	19	1.6	<5	20	240	36.8	173	14.6	<2	<0.1	2	<0.2	0.2	315	41	
YM141	49.91	12.17	15.77	0.226	4.38	8.46	2.64	1.38	3.555	0.53	1.27	100.27	2	461	21	39	22	1.3	<5	26	413	35.8	229	20.9	<2	<0.1	2	<0.2	0.5	430	32	
YM142	51.18	12.43	15.16	0.215	4.69	8.51	2.65	1.37	2.195	0.27	1.35	99.91	1	457	56	41	20	1.7	<5	23	236	37.5	170	14.7	<2	<0.1	2	<0.2	0.1	304	40	
YM143	50.56	12.42	15.42	0.214	4.88	8.75	2.39	1.17	2.219	0.27	1.98	100.27	1	429	56	42	21	1.5	<5	32	250	38.2	170	12.0	<2	<0.1	1	<0.2	0.4	320	40	
YM144	50.91	12.53	15.17	0.207	4.09	8.19	2.73	1.65	3.133	0.50	1.07	100.17	2	368	38	36	22	1.6	<5	33	401	39.8	256	21.9	<2	<0.1	2	<0.2	0.3	507	31	
YM145	51.10	12.51	15.27	0.223	4.19	8.10	2.73	1.66	3.121	0.49	0.96	100.35	2	414	39	37	24	1.7	<5	34	414	41.9	267	23.5	<2	<0.1	2	<0.2	0.3	522	30	
YM146	49.60	13.59	14.06	0.209	5.66	10.39	2.48	0.33	1.434	0.14	1.55	99.23	<1	357	4	43	18	1.7	<5	4	148	29.0	86	6.3	<2	<0.1	1	<0.2	0.2	118	41	
YM147	53.22	12.70	15.36	0.218	4.20	8.04	2.68	1.24	1.730	0.21	0.65	100.25	1	439	32	44	21	1.5	<5	37	182	39.0	146	8.9	<2	<0.1	2	<0.2	0.8	273	38	
YM148	51.35	12.73	15.31	0.232	4.84	8.89	2.50	1.17	2.182	0.28	0.86	100.33	1	421	67	43	21	1.5	<5	36	258	40.3	183	13.9	<2	<0.1	2	<0.2	0.8	327	41	
YM150	50.89	13.32	13.38	0.259	5.07	8.24	2.37	1.65	1.498	0.21	2.81	99.70	2	330	24	39	21	1.5	<5	81	241	36.7	165	11.1	<2	<0.1	1	<0.2	1.8	340	37	
YM151	63.30	13.65	6.85	0.100	2.20	3.62	2.53	3.76	1.074	0.23	2.89	100.00	2	140	32	16	20	1.5	<5	132	157	34.6	200	16.2	<2	<0.1	<1	<0.2	1.6	711	19	
YM152	62.63	13.83	7.33	0.111	2.19	5.19	2.24	2.45	1.055	0.24	3.14	100.41	2	147	33	18	20	1.5	<5	62	365	34.6	240	16.3	<2	<0.1	2	<0.2	3.7	674	20	
YM153	52.73	15.33	10.56	0.164	5.08	7.78	2.81	2.34	1.170	0.20	2.01	100.17	2	229	119	39	20	1.6	<5	62	286	29.3	167	13.2	<2	<0.1	1	<0.2	0.4	413	33	
YM154	54.73	12.79	13.71	0.177	3.28	7.49	2.09	2.09	1.615	0.21	1.68	99.87	1	420	<20	42	20	1.7	<5	44	197	34.0	183	12.5	<2	<0.1	2	<0.2	1.2	434	36	
YM155	52.90	14.07	11.79	0.181	5.26	9.10	2.58	0.85	1.195	0.17	1.86	99.95	1	261	36	42	20	1.7	<5	35	244	30.5	150	11.0	<2	<0.1	1	<0.2	1.2	388	37	
YM156	50.61	14.73	11.96	0.158	5.41	9.41	2.57	1.22	1.776	0.22	1.35	99.40	1	304	150	42	23	1.5	<5	35	327	29.0	169	12.0	<2	<0.1	1	<0.2	1.0	333	26	
YM157	66.18	13.67	12.45	0.184	4.40	7.86	2.60	1.76	1.401	0.28	0.83	99.98	4	47	<20	10	19	1.6	<5	172	145	42.6	253	18.5	<2	<0.1	5	<0.2	8.6	641	16	
YM158	54.54	12.61	12.45	0.201	4.40	7.86	2.60	1.76	1.401	0.28	0.83	99.98	4	47	<20	10	19	1.6	<5	172	145	42.6	253	18.5	<2	<0.1	5	<0.2	8.6	641	16	
YM159	67.97	12.41	5.65	0.103	1.32	2.74	3.07	3.99	0.892	0.28	1.35	99.77	4	63	<20	10	18	1.3	<5	164	120	39.4	231	19.9	<2	<0.1	5	<0.2	5.1	600	17	
YM160																																

Geochemical grade assay for rocks, drill cores and drill cuttings samples

Sample No.	La	Ce	Pr	Nd	Sm	Eu	Gd	Tb	Dy	Ho	Er	Tm	Yb	Lu	HF	Ta	W	Ti	Bi	Th	U	Ag	Cd	Cu	Ni	Pb	Zn	S
	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm
YM115	19.09	40.83	5.07	22.00	6.07	1.871	6.66	1.25	7.46	1.51	4.56	0.684	4.18	0.603	4.4	0.70	0.5	0.21	<0.1	4.56	1.14	0.8	13.8	200	9	16	30	0.001
YM117	24.30	50.82	6.27	26.96	6.50	2.080	6.76	1.19	7.08	1.42	4.24	0.620	3.73	0.536	4.7	0.86	<0.5	0.11	<0.1	2.48	0.50	0.7	14.6	302	27	10	101	0.011
YM118	22.93	48.27	5.92	25.18	5.91	1.995	5.92	1.00	5.71	1.13	3.35	0.482	2.87	0.415	4.0	0.86	<0.5	0.10	<0.1	2.56	0.51	0.7	12.2	227	48	18	85	0.010
YM119	17.79	37.96	4.63	20.06	4.89	1.667	4.81	0.86	5.09	1.02	3.01	0.443	2.63	0.386	3.3	1.25	<0.5	<0.05	<0.1	1.91	0.31	0.6	10.7	182	42	20	77	0.010
YM120	17.22	36.15	4.37	19.05	4.75	1.605	4.74	0.84	5.04	1.01	3.00	0.447	2.73	0.391	3.4	0.60	<0.5	0.09	<0.1	1.75	0.37	0.5	11.2	78	43	11	79	0.011
YM121	24.07	50.47	6.19	26.65	6.48	2.041	6.60	1.13	6.73	1.34	4.00	0.601	3.65	0.525	4.7	0.89	3.2	0.06	<0.1	3.46	0.42	0.8	13.1	252	25	13	90	0.007
YM122	23.53	49.25	6.07	26.25	6.27	2.036	6.27	1.11	6.66	1.33	4.06	0.595	3.49	0.535	4.6	0.86	<0.5	0.07	<0.1	2.51	0.40	0.8	15.1	231	24	9	97	0.006
YM124	30.50	65.84	8.31	35.03	8.23	2.741	7.54	1.23	6.73	1.26	3.56	0.494	2.88	0.419	5.5	1.30	<0.5	0.09	<0.1	2.66	0.44	0.9	13.8	85	13	9	101	0.010
YM125	25.49	53.74	6.69	28.53	6.88	2.155	6.79	1.18	7.12	1.44	4.25	0.640	3.73	0.562	4.9	0.94	<0.5	0.07	<0.1	2.66	0.50	0.7	13.0	192	22	21	94	0.010
YM126	24.68	51.81	6.46	27.57	6.87	2.095	6.85	1.18	6.93	1.39	4.07	0.630	3.76	0.547	4.8	0.91	<0.5	<0.05	<0.1	2.55	0.39	1.3	10.7	254	31	18	92	0.008
YM127	23.56	49.70	6.18	26.45	6.69	2.112	6.61	1.19	7.08	1.41	4.26	0.621	3.74	0.561	4.8	0.94	<0.5	0.06	<0.1	2.44	0.44	0.7	13.5	238	22	14	100	0.011
YM128	20.65	44.78	5.43	24.35	5.93	1.979	6.73	1.09	6.78	1.34	4.02	0.629	3.72	0.542	4.8	0.86	<0.5	0.09	<0.1	2.47	0.49	0.8	14.2	255	20	16	98	0.011
YM129	19.36	41.97	5.00	22.05	5.40	1.808	6.44	1.06	6.42	1.29	3.83	0.571	3.56	0.507	4.4	0.79	<0.5	<0.05	<0.1	2.38	0.89	0.7	13.5	214	44	5	100	0.008
YM130	16.82	36.46	4.37	19.53	4.61	1.651	5.24	0.83	5.16	1.01	3.03	0.445	2.72	0.396	3.4	0.64	<0.5	0.08	<0.1	2.27	0.45	0.6	11.2	209	61	13	84	0.010
YM131	21.50	46.14	5.52	25.11	6.09	1.978	6.88	1.11	6.77	1.35	4.01	0.616	3.58	0.529	4.6	0.87	<0.5	0.09	<0.1	2.46	0.55	0.7	15.4	104	23	15	101	0.009
YM132	18.82	40.33	4.82	21.31	4.97	1.718	5.35	0.86	5.17	1.04	3.04	0.453	2.85	0.385	3.8	0.75	<0.5	0.07	<0.1	2.14	0.42	0.6	12.3	208	49	26	89	0.009
YM133	15.74	33.99	4.07	18.22	4.37	1.542	4.85	0.82	4.91	0.96	2.82	0.426	2.61	0.383	3.3	0.61	<0.5	0.05	<0.1	1.86	0.36	0.4	12.6	153	51	3	77	0.009
YM134	21.43	47.06	5.66	24.72	5.98	2.000	6.66	1.10	6.54	1.30	3.91	0.581	3.82	0.517	4.5	0.86	<0.5	0.09	<0.1	2.58	0.50	0.8	14.0	231	21	8	97	0.009
YM135	23.64	50.13	6.48	27.02	6.26	2.082	6.82	1.22	7.05	1.57	4.05	0.632	3.62	0.559	4.9	0.86	<0.5	0.07	<0.1	2.70	0.58	0.8	13.8	109	22	16	99	0.009
YM136	21.09	46.26	5.54	24.18	5.96	1.967	6.66	1.08	6.65	1.32	3.90	0.591	3.54	0.508	4.6	0.81	<0.5	<0.05	<0.1	2.45	0.45	0.7	11.6	252	23	19	102	0.008
YM137	19.41	41.59	4.87	21.18	4.82	1.686	4.96	0.78	4.59	0.87	2.60	0.379	2.22	0.326	3.5	0.76	<0.5	0.09	<0.1	2.07	0.41	0.5	11.9	166	81	6	83	0.010
YM138	28.79	62.76	7.69	34.15	7.40	2.579	7.51	1.15	6.52	1.19	3.39	0.475	2.79	0.389	5.2	1.23	<0.5	0.09	<0.1	2.76	0.56	0.9	12.1	119	23	11	105	0.011
YM139	49.08	111.28	14.61	60.94	13.20	4.175	12.43	1.99	10.35	2.10	5.08	0.705	4.26	0.602	9.7	2.02	<0.5	0.20	<0.1	4.74	1.05	1.6	13.2	76	<1	5	127	0.010
YM140	21.81	47.57	5.71	24.98	6.00	1.965	6.59	1.09	6.53	1.28	3.83	0.578	3.56	0.521	4.5	0.82	<0.5	0.06	<0.1	2.44	0.47	0.7	13.3	213	34	16	100	0.007
YM141	32.51	70.01	8.58	37.78	8.46	2.861	8.37	1.30	7.41	1.33	3.77	0.524	3.15	0.432	6.0	1.40	<0.5	0.08	<0.1	2.99	0.62	0.8	14.7	82	11	9	112	0.019
YM142	21.72	46.93	5.68	24.85	5.94	2.007	6.66	1.12	6.71	1.32	3.90	0.593	3.56	0.528	4.5	0.84	<0.5	0.06	<0.1	2.44	0.50	0.7	13.7	213	19	14	101	0.008
YM143	21.19	45.64	5.54	24.24	5.79	1.979	6.61	1.13	6.81	1.37	4.04	0.595	3.67	0.522	4.6	0.87	<0.5	0.09	<0.1	2.53	0.50	0.7	12.8	273	18	10	100	0.007
YM144	35.72	75.20	9.07	39.60	8.64	2.871	8.74	1.30	7.29	1.36	3.95	0.559	3.43	0.461	6.7	1.56	<0.5	0.10	<0.1	3.75	0.80	1.1	14.1	206	9	12	111	0.006
YM145	36.13	78.01	9.34	40.73	8.86	2.914	8.94	1.37	7.63	1.43	4.16	0.590	3.41	0.497	7.0	1.61	<0.5	0.11	<0.1	3.91	0.79	1.0	13.8	204	16	8	116	0.005
YM146	6.93	16.32	2.21	10.91	3.11	1.244	4.50	0.80	5.02	1.01	3.00	0.459	2.62	0.389	2.5	0.31	<0.5	<0.05	<0.1	1.26	0.39	0.5	11.5	170	34	29	85	0.011
YM147	17.51	38.20	4.65	20.40	5.46	1.775	6.39	1.10	6.77	1.38	4.07	0.620	3.73	0.547	4.1	0.58	<0.5	0.21	<0.1	4.33	1.03	0.5	12.0	166	12	16	102	0.004
YM148	22.10	47.66	5.80	25.43	6.05	2.035	6.96	1.13	6.94	1.37	4.06	0.615	3.81	0.563	5.0	0.90	<0.5	0.07	<0.1	2.66	0.52	0.7	14.3	236	29	15	106	0.009
YM150	24.32	50.47	6.10	25.50	5.95	1.741	6.44	1.05	6.31	1.27	3.74	0.570	3.48	0.488	4.6	0.64	<0.5	0.31	<0.1	5.53	0.93	0.8	12.6	129	5	5	93	0.014
YM151	42.29	87.26	10.27	38.12	7.58	1.680	6.84	1.10	6.12	1.30	3.39	0.515	3.09	0.448	5.3	1.09	0.9	0.12	<0.1	11.52	2.08	1.0	7.2	20	<1	32	82	0.004
YM152	40.09	82.83	9.24	36.91	7.63	1.674	7.17	1.08	6.21	1.16	3.53	0.527	3.15	0.442	6.4	1.02	0.6	0.34	0.1	12.34	2.04	1.0	6.0	23	6	32	80	0.025
YM153	24.48	50.38	5.80	24.42	5.31	1.524	5.33	0.87	5.10	0.97	2.92	0.441	2.70	0.380	4.4	0.78	<0.5	0.35	<0.1	5.85	1.00	0.8	8.4	90	53	15	79	0.014
YM154	26.52	56.03	6.58	27.02	6.17	1.669	6.29	1.03	6.00	1.17	3.46	0.521	3.20	0.455	5.0	0.84	<0.5	0.17	0.2	7.33	1.84	0.8	12.3	295	13	33	98	0.017
YM155	22.19	45.98	5.32	22.59	5.04	1.528	5.39	0.89	5.32	1.03	3.08	0.463	2.82	0.413	4.0	0.56	<0.5	0.24	<0.1	4.89	0.80	0.6	10.8	82	26	18	86	0.016
YM156	21.14	45.70	5.67	24.58	5.70	1.840	5.96	0.91	5.23	0.97	2.85	0.399	2.40	0.337	4.5	0.75	<0.5	0.17	<0.1	4.24	0.72	0.7	11.0	108	78	4	85	0.010
YM157	43.10	89.12	9.96	39.89	8.56	1.752	8.18	1.30	7.47	1.42	4.32	0.649	3.92	0.545	7.0	1.65	1.6	0.89	0.8	15.71	5.50	1.2	5.1	20	<1	31	83	0.009
YM158	21.92	45.65	5.29	22.56	5.40	1.563	5.77	0.93	5.62	1.10	3.34	0.503	3.05	0.434	4.2	0.76	0.7	0.32	<0.1	6.13	1.65	<0.3	<0.3	148	91	8	106	<0.001
YM159	41.51	85.03	10.18	37.41	7.45	1.564	7.03	1.22	7.05	1.47	3.76	0.581	3.56	0.521	6.2	1.71	1.4	0.86	0.1	13.98	4.62	0.6	<0.3	45	1	21	84	<0.001
YM160	38.78	80.71	9.22	36.90	7.74	1.643	7.82	1.22	7.03	1.32	4.00	0.593	3.59	0.508	6.7	1.64	0.9	1.25	0.7	14.45	5.12	0.5	<0.3	51	1	18	84	<0.001
YM161	21.36	44.24	5.53	23.92	6.22	1.853	7.24	1.21	7.48	1.51	4.40	0.680	4.05	0.581	4.6	0.69	1.0	0.36	0.2	5.23	1.38	0.4	<0.3	213	86	10	131	<0.001
YM162	40.66	81.58	9.48	38.19	7.95	1.668	7.88	1.23	6.96	1.35	3.93	0.599	3.77	0.509	6.6	1.54	1.5	1.										

Geochemical grade assay for rocks, drill cores and drill cuttings samples

Sample No.	Pd ppb	Pt ppb	Ir ppb	Rh ppb	Re ppb	Au ppb
YM115	0.2	0.2	0.1	0.2	0.1	2
YM117	12.9	12.4	0.1	0.7	0.8	3
YM118	18.9	6.5				5
YM119	10.3	10.3				5
YM120	41.1	5.8	0.5	0.8	0.4	4
YM121	8.7	5.3				<2
YM122	18.1	3.8				5
YM124	17.3	3.6				4
YM125	<0.2	0.2				<2
YM126	23.3	4.8	<0.1	0.5	0.6	4
YM127	10.6	6.6				<2
YM128	14.2	7.6				7
YM129	17.1	7.3				4
YM130	16.7	5.4				4
YM131	24.2	16.1	0.2	0.4	0.6	7
YM132	9.0	5.3				9
YM133	13.6	6.8				2
YM134	20.0	6.0				5
YM135	18.1	4.4				5
YM136	13.6	6.1				5
YM137	15.0	7.2				5
YM138	14.0	22.8				4
YM139	0.3	0.4				<2
YM140	<0.2	<0.2				<2
YM141	13.3	7.0				6
YM142	<0.2	<0.2				<2
YM143	7.3	7.1				<2
YM144	17.3	8.2				5
YM145	0.8	0.9				<2
YM146	0.6	0.9				<2
YM147	13.0	5.5	0.1	0.3	1.4	6
YM148	11.9	14.9				4
YM150	10.8	6.1				2
YM151	8.1	4.5				2
YM152	0.2	0.3				<2
YM153	0.4	0.7				<2
YM154	2.5	2.0				<2
YM155	1.1	4.2				3
YM156	0.9	1.2				<2
YM157	5.4	6.6				2
YM158	0.5	<0.2				<2
YM159	6.7	8.2				<2
YM160	9.8	2.7				4
YM161	1.1	0.9				<2
YM162	9.1	12.2				2
YM163	1.4	1.3				<2
YM164	11.9	10.9				2
YM165	2.7	2.6				<2
YM166	7.0	8.2				3
YM167	4.5	3.0				<2
YM168	1.3	1.7				2
YM169	16.2	12.3	0.6	0.8	1.0	4
YM170	8.2	2.3				3
YM170	9.6	9.9				3

Geochemical grade assay for rocks, drill cores and drill cuttings samples

Sample No.	SiO ₂ %	Al ₂ O ₃ %	Fe ₂ O ₃ %	MnO %	MgO %	CaO %	Na ₂ O %	K ₂ O %	TiO ₂ %	P ₂ O ₅ %	LOI %	Total %	Be ppm	V ppm	Cr ppm	Co ppm	Ga ppm	Ge ppm	As ppm	Rb ppm	Sr ppm	Y ppm	Zr ppm	Nb ppm	Mo ppm	In ppm	Sn ppm	Sb ppm	Cs ppm	Ba ppm	Sc ppm		
Detection Limit	0.01	0.01	0.01	0.001	0.01	0.01	0.01	0.01	0.001	0.01	0.01		1	5	20	1	1	0.5	5	1	2	0.5	1	0.2	2	0.1	1	0.2	0.1	3	1		
YM171	53.45	12.09	15.76	0.216	3.48	7.55	2.88	1.48	1.892	0.26	0.82	99.87	1	549	<20	46	22	1.6	<5	55	168	43.0	169	12.1	<2	<0.1	2	<0.2	2.5	295	39		
YM172	53.65	12.23	15.60	0.211	3.57	7.70	2.84	1.42	1.849	0.25	0.58	99.92	1	539	<20	44	22	1.4	<5	54	176	41.6	166	12.0	<2	<0.1	2	<0.2	2.6	295	39		
YM173	57.08	11.90	15.04	0.206	1.91	6.13	3.17	1.98	1.831	0.37	0.79	100.41	2	510	<20	34	24	1.7	<5	84	182	58.9	248	16.8	<2	<0.1	3	<0.2	5.8	435	31		
YM174	54.15	12.13	15.50	0.209	3.28	7.40	2.90	1.61	1.791	0.29	0.64	99.89	1	202	<20	44	23	1.8	<5	62	182	46.0	182	12.8	<2	<0.1	2	<0.2	3.3	322	38		
YM175	53.88	12.22	15.60	0.209	3.34	7.74	2.80	1.61	1.861	0.27	0.90	100.43	1	533	<20	44	23	1.6	<5	58	186	43.7	173	12.5	<2	<0.1	2	<0.2	2.3	311	38		
YM176	55.62	11.41	15.95	0.215	2.37	6.49	2.81	1.84	1.871	0.35	1.18	100.12	2	287	<20	37	22	1.5	<5	70	158	62.2	242	15.7	<2	<0.1	3	0.4	3.7	359	35		
YM177	53.92	12.66	15.53	0.216	3.75	7.81	2.83	1.40	1.754	0.23	-0.01	100.10	1	484	27	40	21	1.5	<5	53	166	44.0	165	11.2	<2	<0.1	2	<0.2	2.8	279	38		
YM178	52.82	14.36	12.46	0.195	4.48	8.58	2.60	1.31	1.343	0.20	1.32	99.67	1	261	29	42	21	1.5	<5	39	258	32.2	178	13.7	<2	<0.1	<1	<0.2	0.6	462	36		
YM179	52.72	12.96	11.82	0.147	4.32	7.71	2.61	1.93	3.386	0.49	1.86	99.95	2	291	65	32	25	1.5	<5	29	571	34.2	281	23.8	<2	<0.1	2	<0.2	0.2	688	22		
YM180	50.75	14.98	9.57	0.152	8.13	10.89	2.04	0.60	0.778	0.09	2.07	100.04	<1	333	278	42	18	1.5	<5	14	189	17.1	70	7.1	<2	<0.1	<1	<0.2	<0.1	146	37		
YM181	49.34	12.75	12.78	0.156	4.93	8.84	2.49	0.69	3.629	0.49	3.24	99.33	2	315	93	36	25	1.5	<5	9	781	34.8	268	24.1	<2	<0.1	2	<0.2	0.5	485	24		
YM182	50.63	12.84	13.49	0.183	4.36	8.02	2.58	1.61	3.404	0.49	1.21	99.72	2	339	37	39	27	1.5	<5	33	613	39.1	289	26.6	<2	<0.1	2	<0.2	0.9	636	27		
YM183	50.80	13.61	13.68	0.205	6.01	10.27	2.59	0.57	1.374	0.15	0.33	99.58	<1	352	67	45	21	1.7	<5	20	165	30.5	97	8.0	<2	<0.1	<1	3.2	0.5	129	40		
YM184	49.67	12.98	12.53	0.178	5.28	7.45	3.55	1.63	3.253	0.45	2.16	99.13	2	295	89	37	25	1.7	<5	36	616	32.6	257	22.1	<2	<0.1	2	<0.2	0.5	561	24		
YM185	51.10	13.81	13.56	0.206	6.01	10.51	2.51	0.59	1.392	0.15	0.38	100.21	<1	334	66	42	20	1.7	<5	16	172	29.4	90	7.1	<2	<0.1	<1	<0.2	0.8	122	40		
YM186	51.67	14.02	11.98	0.184	6.86	9.77	2.54	1.08	1.039	0.11	0.69	99.95	<1	293	74	43	19	1.5	<5	45	204	24.8	82	6.0	<2	<0.1	<1	<0.2	1.1	248	42		
YM187	77.42	12.02	1.38	0.022	0.08	0.59	3.27	4.70	0.086	0.02	0.55	100.14	3	<5	<20	<1	20	0.9	<5	27	6	66.6	17	66.6	128	18.7	<2	<0.1	3	<0.2	3.3	25	4
YM188	51.78	13.10	12.75	0.159	4.33	7.85	2.87	1.69	3.575	0.50	0.94	99.56	2	307	43	35	27	1.8	<5	39	697	38.1	293	24.2	<2	<0.1	2	<0.2	0.8	645	24		
YM189	52.84	12.94	12.16	0.161	3.95	7.48	3.03	2.05	3.415	0.51	0.57	99.12	2	274	59	32	27	1.9	<5	67	692	39.1	326	29.3	<2	<0.1	2	<0.2	3.3	741	22		
YM190	51.64	13.06	12.52	0.156	4.63	8.26	2.77	1.70	3.588	0.48	0.58	99.37	2	311	73	36	27	1.5	<5	44	752	36.9	304	28.0	<2	<0.1	2	<0.2	2.2	618	24		
YM191	52.33	12.87	12.92	0.159	3.96	7.58	2.58	1.92	3.586	0.55	0.50	99.56	2	308	54	32	26	1.6	<5	40	693	39.3	313	27.2	<2	<0.1	2	<0.2	0.7	687	23		
YM192	53.09	14.17	11.61	0.189	5.29	9.00	2.51	1.40	1.148	0.17	1.11	99.68	1	255	27	40	21	1.6	<5	40	232	28.5	146	13.0	<2	<0.1	<1	<0.2	0.6	415	37		
YM193	53.67	13.04	13.01	0.180	3.81	7.72	2.67	1.47	1.666	0.25	2.07	99.54	2	375	<20	39	20	1.5	<5	41	282	32.4	171	14.7	<2	<0.1	2	<0.2	1.6	383	34		
YM194	54.09	13.39	12.24	0.174	3.32	6.64	2.56	2.16	1.666	0.29	3.08	99.62	2	307	22	35	21	1.6	<5	61	195	39.0	224	17.3	<2	<0.1	2	<0.2	2.3	463	31		
YM195	52.27	13.36	13.23	0.198	4.67	8.60	2.60	1.14	1.570	0.23	1.77	99.63	1	364	37	43	21	1.6	<5	39	244	32.0	177	15.3	<2	<0.1	1	<0.2	1.4	372	34		
YM196	52.26	13.36	12.49	0.198	4.89	8.34	2.64	1.41	1.344	0.18	1.58	99.69	1	357	<20	43	19	1.9	<5	62	197	29.8	140	11.2	<2	<0.1	1	<0.2	1.7	301	35		
YM198	51.03	13.70	13.86	0.207	6.00	10.52	2.53	0.54	1.397	0.15	0.21	100.13	<1	378	65	48	19	1.4	<5	16	156	29.3	98	6.6	<2	<0.1	<1	<0.2	0.5	114	42		
YM199	50.44	14.22	10.91	0.162	6.99	11.13	2.24	0.40	1.010	0.12	1.50	99.12	<1	299	219	45	17	1.6	<5	20	178	20.7	86	6.4	<2	<0.1	1	<0.2	2.8	138	38		
YM200	52.94	14.15	12.15	0.170	5.15	9.07	2.36	1.01	1.257	0.19	1.98	100.42	1	275	28	44	20	1.6	<5	27	241	29.7	164	11.4	<2	<0.1	1	<0.2	0.6	390	38		
YM201	50.78	13.02	12.41	0.159	4.71	8.42	2.88	1.71	3.501	0.48	1.02	99.09	2	324	65	37	23	1.7	<5	33	672	34.0	290	22.6	<2	<0.1	2	<0.2	0.3	596	24		
YM205	51.78	13.89	11.81	0.183	6.82	10.15	2.42	1.10	0.992	0.11	0.34	99.60	<1	296	94	44	17	1.5	<5	46	292	23.5	91	7.8	<2	<0.1	<1	<0.2	1.3	201	42		
YM206	50.45	13.80	13.65	0.207	5.96	10.87	2.48	0.41	1.365	0.15	0.71	100.06	<1	372	66	46	19	1.5	<5	13	176	29.0	93	6.2	<2	<0.1	1	<0.2	0.5	154	41		
YM207	53.04	12.52	15.99	0.228	4.07	8.09	2.74	1.31	1.766	0.22	0.20	100.17	<1	476	<20	45	20	1.8	<5	48	180	41.5	162	10.1	<2	<0.1	2	<0.2	2.0	256	41		
YM208	51.81	12.77	12.38	0.162	4.91	8.56	2.63	1.74	3.443	0.45	0.11	98.96	2	313	99	36	23	1.8	<5	31	671	33.0	256	20.1	<2	<0.1	2	<0.2	0.4	495	24		
YM210	52.13	13.00	12.55	0.155	4.82	8.18	2.82	1.87	3.647	0.46	0.45	100.07	2	317	79	36	23	1.6	<5	37	642	33.6	268	23.8	<2	<0.1	2	<0.2	0.4	567	25		
YM211	51.93	12.92	12.92	0.160	4.57	8.18	2.57	1.76	3.724	0.51	0.95	100.21	2	314	69	37	24	1.5	<5	31	981	37.0	300	27.0	<2	<0.1	2	<0.2	1.1	630	24		
YM212	50.70	13.06	12.56	0.131	4.62	8.02	2.66	1.92	3.818	0.51	2.01	100.03	2	328	66	38	25	1.4	<5	42	754	38.7	306	28.2	<2	<0.1	2	<0.2	0.9	682	24		
YM215	68.39	12.00	5.09	0.061	0.92	1.68	2.61	4.96	0.729	0.22	1.94	98.60	4	19	<20	6	17	1.8	8	215	86	55.2	291	23.8	<2	<0.1	6	0.4	9.3	651	17		
YM218	49.34	12.75	14.97	0.219	4.89	8.71	2.60	1.27	3.572	0.48	1.23	100.03	1	421	31	42	21	1.6	<5	30	410	35.6	214	22.9	<2	<0.1	1	<0.2	0.4	461	33		
YM219	51.59	13.48	13.73	0.232	5.70	9.97	2.56	0.99	1.371	0.17	0.55	100.35	<1	369	39	45	19	1.6	<5	30	183	30.1	118	10.5	<2	<0.1	<1	<0.2	0.4	740	43		
YM222	50.63	12.57	13.05	0.165	4.05	7.28	2.69	2.53	4.014	0.60	1.02	98.61	2	314	41	35	24	1.8	<5	58	735	37.6	329	30.4	<2	<0.1	3	<0.2	0.4	282	22		
YM223	50.58	13.01	13.82	0.159	4.43	8.00	2.71	1.98	4.055	0.5																							

Geochemical grade assay for rocks, drill cores and drill cuttings samples

(8continued/8)

Sample No.	Pd ppb	Pt ppb	Ir ppb	Rh ppb	Re ppb	Au ppb
YM171	9.2	10.7			0.2	2
YM172	9.4	12.2			0.1	4
YM173	14.9	0.9			0.2	4
YM174	10.2	8.2			0.1	6
YM175	14.5	10.3			0.2	4
YM176	13.2	1.1			0.1	5
YM177	10.5	11.8			0.2	4
YM178	4.5	5.2			0.1	3
YM179	5.6	3.6			0.1	2
YM180	7.5	6.8			0.1	2
YM181	8.2	5.6			0.1	3
YM182	8.6	6.8	0.2	0.6	0.2	3
YM183	10.3	5.7			0.2	4
YM184	7.3	5.8	0.3	1.0	0.9	3
YM185	10.2	5.6			0.9	3
YM186	6.0	5.9			0.9	3
YM187	<0.2	<0.2			0.9	<2
YM188	8.0	5.3			0.9	<2
YM189	4.8	4.6			0.9	3
YM190	7.1	6.3			0.9	2
YM191	6.9	5.5			0.9	3
YM192	0.4	1.1			0.9	4
YM193	9.0	4.8			0.9	<2
YM194	3.6	2.9			0.9	5
YM195	8.1	8.6			0.9	<2
YM197	11.3	12.0			0.9	3
YM198	11.6	5.9			0.9	3
YM199	9.6	11.4			0.9	2
YM200	1.4	2.0			0.9	<2
YM201	4.1	3.8			0.9	3
YM205	6.0	6.0			0.9	<2
YM206	10.1	6.5			0.9	3
YM207	11.5	11.8			0.9	3
YM208	10.3	4.3	0.3	0.7	0.3	2
YM210	7.4	6.3	0.2	0.5	0.8	<2
YM211	6.5	3.7			0.8	3
YM212	6.8	4.0			0.8	3
YM215	0.4	0.2			0.8	<2
YM218	1.2	0.4			0.8	<2
YM219	9.7	2.8			0.8	<2
YM222	4.8	4.4			0.8	<2
YM223	7.1	7.3	0.2	0.7	1.0	2
YM225	<0.2	<0.2			1.0	3
YM227	6.2	6.4			1.0	<2
YM228	10.1	9.7			1.0	4
YM229	14.0	11.8			1.0	4
YM230	12.2	10.3			1.0	4
YM231	7.6	7.6			1.0	<2

APPENDIX 7

Geochemical grade assay for pan concentrated samples collected in the Lajes districts

Sample No.	Ag	Cd	Cu	Mn	Mo	Ni	Pb	Zn	S	Pd	Pt
	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	%	ppb	ppb
AS027P	< 0.2	< 0.5	68	318	< 2	45	7	68	0.006	1.0	1.0
AS022P-1	< 0.2	< 0.5	79	1236	6	51	6	75	3.966	<1.0	4.0
AS022P-2	< 0.2	< 0.5	75	758	< 2	109	6	66	9.976	3.0	3.0
AS022P-3	< 0.2	< 0.5	130	922	< 2	105	5	50	10.026	3.0	3.0
AS022P-4	< 0.2	< 0.5	53	5397	< 2	1161	14	86	3.192	5.0	6.0

APPENDIX 8

Geochemical grade assay for stream sediments collected in the Lomba Grande district

Sample No.	UTM_E	UTM_N	Latitude	Longitude	Location	Ag ppm	Cd ppm	Cu ppm	Mn ppm	Mo ppm	Ni ppm	Pb ppm	Zn ppm	Al %	As ppm	Ba ppm	Be ppm	Bi ppm	Ca %	Co ppm	Cr ppm	Fe %	K %	Mg %	Na %	P %	Sb ppm	Sc ppm	Sn ppm
CK-S-001	495450	6698500	-51.0471	-29.8439	Lomba Grande	<0.2	<0.5	8	133	<2	7	6	8	0.53	<10	67	<1	<10	0.02	6	11	1.06	0.04	0.04	0.02	0.006	<10	3	<10
CK-S-002	494900	6699000	-51.0528	-29.8394	Lomba Grande	<0.2	<0.5	7	82	<2	4	6	6	0.38	<10	43	<1	<10	0.04	2	7	0.58	0.03	0.05	0.02	0.004	<10	1	<10
CK-S-003	492450	6698150	-51.0782	-29.8470	Lomba Grande	<0.2	<0.5	5	58	<2	3	6	8	0.37	<10	40	<1	<10	0.03	2	8	0.62	0.03	0.03	0.02	0.004	<10	1	<10
CK-S-004	493000	6699100	-51.0725	-29.8385	Lomba Grande	<0.2	<0.5	5	79	<2	3	6	9	0.24	<10	37	<1	<10	0.02	2	7	0.58	0.03	0.03	0.02	0.004	<10	1	<10
CK-S-005	494800	6698550	-51.0538	-29.8344	Lomba Grande	<0.2	<0.5	7	141	<2	4	7	10	0.26	<10	40	<1	<10	0.02	4	21	1.36	0.03	0.03	0.01	0.006	<10	1	<10
CK-S-006	496000	6699500	-51.0518	-29.8335	Lomba Grande	<0.2	<0.5	10	362	<2	24	6	13	0.44	<10	57	<1	<10	0.15	9	27	1.86	0.02	0.32	0.03	0.011	<10	2	<10
CK-S-007	495700	6700200	-51.0445	-29.8286	Lomba Grande	<0.2	<0.5	8	159	<2	11	7	12	0.25	<10	37	<1	<10	0.07	5	16	1.81	0.01	0.14	0.02	0.009	<10	1	<10
CK-S-008	495850	6700400	-51.0430	-29.8267	Lomba Grande	<0.2	<0.5	26	799	<2	227	4	34	1.50	<10	62	<1	<10	0.09	46	95	4.87	0.02	3.88	0.16	0.011	<10	7	<10
CK-S-009	494250	6700350	-51.0595	-29.8272	Lomba Grande	<0.2	<0.5	5	95	<2	4	7	8	0.33	<10	42	<1	<10	0.03	2	9	1.17	0.03	0.06	0.02	0.005	<10	1	<10
CK-S-010	494150	6701550	-51.0605	-29.8164	Lomba Grande	<0.2	<0.5	9	204	<2	7	12	19	0.71	<10	96	<1	<10	0.05	7	21	2.08	0.03	0.06	0.02	0.023	<10	2	<10
CK-S-011	493200	6700600	-51.0704	-29.8249	Lomba Grande	<0.2	<0.5	7	586	<2	3	11	13	0.75	<10	114	<1	<10	0.04	5	10	1.34	0.05	0.04	0.02	0.007	<10	5	<10
CK-S-012	493100	6701400	-51.0714	-29.8177	Lomba Grande	<0.2	<0.5	9	94	<2	4	11	15	1.42	<10	130	<1	<10	0.04	2	9	1.28	0.04	0.05	0.02	0.010	<10	3	<10
CK-S-013	492200	6702000	-51.0807	-29.8123	Lomba Grande	<0.2	<0.5	5	147	<2	3	7	10	0.69	<10	79	<1	<10	0.03	2	7	0.91	0.03	0.03	0.02	0.006	<10	1	<10
CK-S-014	498150	6698500	-51.0192	-29.8439	Lomba Grande	<0.2	<0.5	5	260	<2	6	4	7	0.29	<10	42	<1	<10	0.04	5	12	0.60	0.02	0.04	0.02	0.006	<10	1	<10
CK-S-015	497700	6698500	-51.0238	-29.8439	Lomba Grande	<0.2	<0.5	4	76	<2	4	5	6	0.18	<10	39	<1	<10	0.02	3	8	0.71	0.02	0.04	0.02	0.005	<10	1	<10
CK-S-016	493300	6699550	-51.0072	-29.8344	Lomba Grande	<0.2	<0.5	30	475	<2	36	13	26	2.23	<10	121	<1	<10	0.15	19	146	5.77	0.04	0.14	0.03	0.014	<10	18	<10
CK-S-017	498800	6699350	-51.0124	-29.8362	Lomba Grande	<0.2	<0.5	9	254	<2	9	6	10	0.31	<10	44	<1	<10	0.07	5	17	1.03	0.02	0.10	0.02	0.008	<10	2	<10
CK-S-018	499100	6699650	-51.0093	-29.8335	Lomba Grande	<0.2	<0.5	4	78	<2	3	4	6	0.16	<10	25	<1	<10	0.03	2	7	0.41	0.01	0.02	0.02	0.003	<10	1	<10
CK-S-019	499550	6701100	-51.0047	-29.8204	Lomba Grande	<0.2	<0.5	4	58	<2	2	5	5	0.23	<10	27	<1	<10	0.02	2	7	0.58	0.02	0.03	0.02	0.003	<10	1	<10
CK-S-020	498200	6703000	-51.0186	-29.8033	Lomba Grande	<0.2	<0.5	6	145	<2	5	6	13	0.33	<10	101	<1	<10	0.06	4	14	1.08	0.08	0.07	0.02	0.008	<10	1	<10
CK-S-021	499050	6703100	-51.0098	-29.8024	Lomba Grande	<0.2	<0.5	6	86	<2	5	7	12	0.43	<10	48	<1	<10	0.06	3	15	1.08	0.06	0.07	0.03	0.008	<10	1	<10
CK-S-022	498700	6704000	-51.0155	-29.7943	Lomba Grande	<0.2	<0.5	6	83	<2	4	5	9	0.27	<10	58	<1	<10	0.02	3	15	1.24	0.04	0.04	0.02	0.008	<10	1	<10
CK-S-023	498500	6704000	-51.0155	-29.7943	Lomba Grande	<0.2	<0.5	5	181	<2	3	6	7	0.27	<10	46	<1	<10	0.03	3	6	0.52	0.04	0.03	0.02	0.006	<10	1	<10
CK-S-024	498650	6704650	-51.0140	-29.7884	Lomba Grande	<0.2	<0.5	5	178	<2	4	6	9	0.39	<10	59	<1	<10	0.08	3	8	0.89	0.04	0.05	0.03	0.012	<10	1	<10
CK-S-025	498650	6705000	-51.0140	-29.7852	Lomba Grande	<0.2	<0.5	4	58	<2	2	4	7	0.20	<10	31	<1	<10	0.02	2	7	0.71	0.03	0.03	0.02	0.005	<10	1	<10
CK-S-026	498300	6705200	-51.0176	-29.7834	Lomba Grande	<0.2	<0.5	5	78	<2	3	5	5	0.27	<10	33	<1	<10	0.03	2	6	0.46	0.02	0.03	0.02	0.003	<10	1	<10
CK-S-027	499150	6705850	-51.0088	-29.7780	Lomba Grande	<0.2	<0.5	5	85	<2	2	6	10	0.31	<10	53	<1	<10	0.04	2	5	0.62	0.03	0.04	0.02	0.009	<10	1	<10
CK-S-028	496450	6706100	-51.0367	-29.7753	Lomba Grande	<0.2	<0.5	4	45	<2	3	4	5	0.19	<10	23	<1	<10	0.04	1	10	0.86	<0.01	0.03	0.02	0.003	<10	1	<10
CK-S-029	494200	6705600	-51.0620	-29.7798	Lomba Grande	<0.2	<0.5	8	201	<2	4	4	4	0.20	<10	41	<1	<10	0.06	4	6	0.52	0.01	0.04	0.03	0.006	<10	1	<10
CK-S-030	496900	6703400	-51.0321	-29.7997	Lomba Grande	<0.2	<0.5	11	266	<2	15	5	13	0.69	<10	62	<1	<10	0.31	6	25	1.21	0.03	0.17	0.07	0.009	<10	3	<10
CK-S-032	496100	6704400	-51.0404	-29.7907	Lomba Grande	<0.2	<0.5	5	36	<2	3	5	6	0.28	<10	30	<1	<10	0.02	<1	7	0.45	0.03	0.02	0.03	0.004	<10	1	<10
CK-S-033	495250	6704300	-51.0491	-29.7915	Lomba Grande	<0.2	<0.5	8	102	<2	7	5	11	0.87	<10	42	<1	<10	0.31	4	15	1.76	0.02	0.14	0.09	0.011	<10	3	<10
CK-S-034	494100	6702900	-51.0611	-29.8042	Lomba Grande	<0.2	<0.5	13	199	<2	9	10	23	0.97	<10	80	<1	<10	0.05	6	20	1.85	0.04	0.05	0.03	0.010	<10	3	<10
CK-S-035	493800	6704000	-51.0642	-29.7943	Lomba Grande	<0.2	<0.5	3	27	<2	<1	3	1	0.17	<10	13	<1	<10	0.01	<1	1	0.10	0.01	<0.01	0.03	0.001	<10	1	<10
CK-S-036	494750	6703700	-51.0543	-29.7970	Lomba Grande	<0.2	<0.5	5	108	<2	4	4	7	0.31	<10	39	<1	<10	0.04	3	6	0.54	0.02	0.05	0.03	0.005	<10	1	<10
CK-S-037	494900	6703700	-51.0528	-29.7970	Lomba Grande	<0.2	<0.5	4	137	<2	4	4	6	0.38	<10	41	<1	<10	0.08	2	7	0.55	0.02	0.06	0.04	0.006	<10	1	<10
CK-S-038	495300	6702750	-51.0486	-29.8055	Lomba Grande	<0.2	<0.5	7	324	<2	10	5	10	0.57	<10	64	<1	<10	0.17	5	13	1.03	0.03	0.17	0.05	0.010	<10	2	<10
CK-S-039	493350	6704900	-51.0688	-29.7861	Lomba Grande	<0.2	<0.5	6	64	<2	2	6	6	0.46	<10	30	<1	<10	0.04	1	4	0.65	0.02	0.02	0.03	0.004	<10	1	<10
CK-S-040	493200	6704900	-51.0704	-29.7861	Lomba Grande	<0.2	<0.5	3	18	<2	2	4	12	0.20	<10	20	<1	<10	0.04	<1	2	0.23	0.01	0.01	0.03	0.004	<10	1	<10
CK-S-041	492300	6703450	-51.0797	-29.7992	Lomba Grande	<0.2	<0.5	4	175	<2	5	12	10	0.14	<10	87	<1	<10	0.01	4	19	2.52	0.01	0.01	0.03	0.007	<10	1	<10
CK-S-042	503750	6705700	-50.9612	-29.7789	Lomba Grande	<0.2	<0.5	4	195	<2	6	7	16	0.66	<10	62	<1	<10	0.06	5	11	1.19	0.04	0.06	0.03	0.012	<10	2	<10
CK-S-043	502850	6705300	-50.9705	-29.7825	Lomba Grande	<0.2	<0.5	5	176	<2	5	9	19	0.83	<10	49	<1	<10	0.05	5	13	1.29	0.04	0.04	0.03	0.012	<10	2	<10
CK-S-044	502900	6705450	-50.9700	-29.7812	Lomba Grande	<0.2	<0.5	2	104	<2	3	3	8	0.26	<10	36	<1	<10	0.03	2	5	0.68	0.02	0.02	0.02	0.009	<10	1	<10
CK-S-045	502400	6705350	-50.9752	-29.7821	Lomba Grande	<0.2	<0.5	6	989	<2	5	7	18	0.51	<10	86	<1	<10	0.07	10	8	1.15	0.03	0.04	0.02	0.018	<10	1	<10
CK-S-046	502700	6703000	-50.9721	-29.8033	Lomba Grande	<0.2	<0.5	6	303	<2	11	9	16	0.40	<10	72	<1	<10	0.09	6	14	0.95	0.03	0.07	0.02	0.010	<10	1	<10
CK-S-047	502900	6703050	-50.9700	-29.8028																									

Geochemical grade assay for stream sediments collected in the Lomba Grande district

(1 continuerd/4)

Sample No.	Sr ppm	Ti %	V ppm	W ppm	Y ppm	Zr ppm	S %
CK-S-001	15	0.01	20	<10	5	5	0.036
CK-S-002	7	0.03	15	<10	3	4	0.011
CK-S-003	5	0.02	15	<10	3	5	0.007
CK-S-004	6	0.02	12	<10	3	2	0.007
CK-S-005	7	0.03	28	<10	4	4	0.007
CK-S-006	9	0.03	33	<10	3	4	0.007
CK-S-007	5	0.03	41	<10	2	4	0.009
CK-S-008	50	0.06	53	<10	4	12	0.006
CK-S-009	7	0.02	25	<10	2	3	0.004
CK-S-010	13	0.05	57	<10	4	4	0.009
CK-S-011	10	0.03	33	<10	6	4	0.006
CK-S-012	11	0.03	28	<10	3	11	0.007
CK-S-013	7	0.02	20	<10	2	4	0.005
CK-S-014	7	0.02	12	<10	2	2	0.003
CK-S-015	7	0.02	13	<10	2	2	0.004
CK-S-016	11	0.07	113	<10	8	29	0.008
CK-S-017	5	0.02	22	<10	3	4	0.006
CK-S-018	3	0.01	9	<10	1	1	0.003
CK-S-019	5	0.02	14	<10	2	2	0.003
CK-S-020	15	0.03	23	<10	3	4	0.004
CK-S-021	15	0.03	24	<10	4	4	0.004
CK-S-022	11	0.06	28	<10	2	4	0.004
CK-S-023	12	0.01	8	<10	2	2	0.005
CK-S-024	18	0.02	11	<10	2	2	0.004
CK-S-025	8	0.02	14	<10	1	2	0.003
CK-S-026	9	<0.01	9	<10	2	2	0.003
CK-S-027	9	0.01	11	<10	2	1	0.005
CK-S-028	4	0.02	27	<10	1	4	0.003
CK-S-029	4	0.01	10	<10	2	2	0.005
CK-S-030	17	0.05	30	<10	4	5	0.004
CK-S-032	7	0.02	13	<10	3	2	0.004
CK-S-033	18	0.04	29	<10	3	9	0.004
CK-S-034	18	0.05	46	<10	4	6	0.007
CK-S-035	3	<0.01	4	<10	<1	2	0.002
CK-S-036	8	0.02	13	<10	2	2	0.004
CK-S-037	8	0.02	12	<10	2	2	0.003
CK-S-038	12	0.02	18	<10	3	3	0.004
CK-S-039	5	<0.01	16	<10	1	6	0.005
CK-S-040	4	<0.01	7	<10	2	2	0.004
CK-S-041	19	0.04	49	<10	1	8	0.003
CK-S-042	11	0.03	28	<10	2	3	0.005
CK-S-043	8	0.03	29	<10	2	6	0.009
CK-S-044	6	0.02	13	<10	1	2	0.004
CK-S-045	10	0.02	15	<10	3	5	0.009
CK-S-046	12	0.01	20	<10	3	2	0.007
CK-S-047	10	0.02	25	<10	2	4	0.003
CK-S-048	8	0.02	21	<10	2	3	0.004
CK-S-049	11	0.05	51	<10	4	7	0.011
CK-S-050	12	0.05	43	<10	4	8	0.007
CK-S-051	6	0.04	42	<10	2	5	0.004
CK-S-052	13	0.02	22	<10	2	4	0.008
CK-S-053	9	0.04	30	<10	3	5	0.006
CK-S-054	6	0.04	29	<10	2	4	0.006

Geochemical grade assay for stream sediments collected in the Lomba Grande district

(2/4)

Sample No.	UTM_E	UTM_N	Latitude	Longitude	Location	Ag ppm	Cd ppm	Cu ppm	Mn ppm	Mo ppm	Ni ppm	Pb ppm	Zn ppm	Al %	As ppm	Ba ppm	Be ppm	Bi ppm	Ca %	Co ppm	Cr ppm	Fe %	K %	Mg %	Na %	P %	Sb ppm	Sc ppm	Sn ppm
CK-S-055	504750	6703850	-50.9509	-29.7956	Lomba Grande	<0.2	<0.5	2	97	<2	3	6	9	0.24	<10	<10	<1	<10	0.02	3	7	0.65	<0.01	0.01	0.01	0.005	<10	<1	<10
CK-S-056	506300	6704050	-50.9348	-29.7938	Lomba Grande	<0.2	<0.5	3	159	<2	1	6	10	0.34	<10	<10	<1	<10	0.04	2	4	0.68	<0.01	0.02	0.02	0.007	<10	<1	<10
CK-S-057	505700	6703800	-50.9410	-29.7961	Lomba Grande	<0.2	<0.5	9	192	<2	3	8	18	0.54	<10	<10	<1	<10	0.12	4	8	1.00	0.02	0.06	0.02	0.011	<10	<2	<10
CK-S-058	506700	6703200	-50.9307	-29.8015	Lomba Grande	<0.2	<0.5	4	44	<2	2	5	12	0.30	<10	<10	<1	<10	0.05	1	4	0.45	<0.01	0.03	0.02	0.007	<10	<1	<10
CK-S-059	506450	6703150	-50.9333	-29.8019	Lomba Grande	<0.2	<0.5	5	134	<2	5	6	17	0.66	<10	<10	<1	<10	0.08	3	8	0.98	0.02	0.08	0.02	0.016	<10	<2	<10
CK-S-060	507150	6702000	-50.9260	-29.8123	Lomba Grande	<0.2	<0.5	3	72	<2	2	4	10	0.28	<10	<10	<1	<10	0.03	2	4	0.53	<0.01	0.02	0.02	0.008	<10	<1	<10
CK-S-061	506650	6701700	-50.9312	-29.8150	Lomba Grande	<0.2	<0.5	7	62	<2	1	4	9	0.24	<10	<10	<1	<10	0.03	1	4	0.42	0.01	0.02	0.01	0.004	<10	<1	<10
CK-S-062	506700	6701450	-50.9307	-29.8173	Lomba Grande	<0.2	<0.5	6	48	<2	2	6	14	0.21	<10	<10	<1	<10	0.02	1	3	0.29	0.01	0.01	0.01	0.004	<10	<1	<10
CK-S-063	505800	6702150	-50.9400	-29.8109	Lomba Grande	<0.2	<0.5	5	163	<2	2	6	15	0.43	<10	<10	<1	<10	0.04	2	8	0.96	0.01	0.03	0.02	0.006	<10	<1	<10
CK-S-064	504850	6702500	-50.9498	-29.8078	Lomba Grande	<0.2	<0.5	9	348	<2	8	9	22	0.98	<10	<10	<1	<10	0.38	6	17	1.22	0.02	0.13	0.07	0.013	<10	<2	<10
CK-S-065	503900	6702950	-50.9596	-29.8037	Lomba Grande	<0.2	<0.5	4	129	<2	2	5	12	0.31	<10	<10	<1	<10	0.03	2	5	0.52	0.01	0.02	0.01	0.006	<10	<1	<10
CK-S-066	503900	6702100	-50.9596	-29.8114	Lomba Grande	<0.2	<0.5	4	100	<2	1	6	10	0.28	<10	<10	<1	<10	0.02	2	7	0.93	0.01	0.02	<0.01	0.004	<10	<1	<10
CK-S-067	503900	6701100	-50.9596	-29.8204	Lomba Grande	<0.2	<0.5	6	101	<2	3	7	13	0.21	<10	<10	<1	<10	0.03	2	9	0.99	0.01	0.02	0.01	0.004	<10	<1	<10
CK-S-068	505500	6701250	-50.9431	-29.8191	Lomba Grande	<0.2	<0.5	10	168	<2	7	9	17	0.53	<10	<10	<1	<10	0.23	4	16	1.38	0.02	0.10	0.04	0.008	<10	<2	<10
CK-S-069	505100	6700550	-50.9472	-29.8254	Lomba Grande	<0.2	<0.5	4	104	<2	4	7	12	0.23	<10	<10	<1	<10	0.08	2	12	1.27	0.01	0.04	0.02	0.008	<10	<1	<10
CK-S-070	504900	6699650	-50.9493	-29.8335	Lomba Grande	<0.2	<0.5	4	161	<2	3	7	13	0.31	<10	<10	<1	<10	0.08	2	10	0.82	0.03	0.05	0.02	0.008	<10	<1	<10
CK-S-071	504900	6699650	-50.9493	-29.8344	Lomba Grande	<0.2	<0.5	5	351	<2	3	7	16	0.23	<10	<10	<1	<10	0.05	5	9	1.02	0.03	0.04	0.01	0.016	<10	<1	<10
CK-S-072	504850	6698950	-50.9503	-29.8398	Lomba Grande	<0.2	<0.5	6	350	<2	2	8	17	0.23	<10	<10	<1	<10	0.07	4	6	0.59	0.02	0.04	0.01	0.007	<10	<1	<10
CK-S-073	504850	6698950	-50.9498	-29.8371	Lomba Grande	<0.2	<0.5	7	610	<2	3	6	18	0.25	<10	<10	<1	<10	0.07	5	6	0.88	0.02	0.05	0.01	0.010	<10	<1	<10
CK-S-074	504850	6698250	-50.9498	-29.8371	Lomba Grande	<0.2	<0.5	7	610	<2	3	6	18	0.25	<10	<10	<1	<10	0.07	5	6	0.88	0.02	0.05	0.01	0.010	<10	<1	<10
CK-S-075	502050	6700200	-50.9788	-29.8286	Lomba Grande	<0.2	<0.5	5	118	<2	4	8	14	0.19	<10	<10	<1	<10	0.03	3	15	1.34	0.02	0.03	0.01	0.006	<10	<1	<10
CK-S-076	502550	6699650	-50.9736	-29.8344	Lomba Grande	<0.2	<0.5	5	289	<2	4	7	18	0.30	<10	<10	<1	<10	0.08	4	11	0.81	0.07	0.08	0.01	0.010	<10	<1	<10
CK-S-077	502600	6701150	-50.9731	-29.8200	Lomba Grande	<0.2	<0.5	3	183	<2	3	6	12	0.15	<10	<10	<1	<10	0.04	3	7	0.59	0.02	0.03	0.01	0.007	<10	<1	<10
CK-S-078	503100	6700350	-50.9679	-29.8272	Lomba Grande	<0.2	<0.5	2	238	<2	2	5	10	0.13	<10	<10	<1	<10	0.04	2	9	0.64	0.02	0.03	0.01	0.006	<10	<1	<10
CK-S-079	503000	6700200	-50.9689	-29.8286	Lomba Grande	<0.2	<0.5	14	126	<2	4	7	19	0.13	<10	<10	<1	<10	0.04	2	8	0.69	0.02	0.03	0.01	0.008	<10	<1	<10
CK-S-080	503400	6700100	-50.9648	-29.8295	Lomba Grande	<0.2	<0.5	5	239	<2	3	6	15	0.13	<10	<10	<1	<10	0.05	3	5	0.43	0.02	0.03	0.01	0.006	<10	<1	<10
CK-S-081	503400	6698900	-50.9648	-29.8403	Lomba Grande	<0.2	<0.5	4	171	<2	3	6	16	0.14	<10	<10	<1	<10	0.05	3	9	0.71	0.03	0.05	0.01	0.009	<10	<1	<10
CK-S-082	503900	6700150	-50.9596	-29.8290	Lomba Grande	<0.2	<0.5	3	61	<2	2	4	10	0.17	<10	<10	<1	<10	0.02	1	5	0.65	<0.01	0.01	<0.01	0.004	<10	<1	<10
CK-S-083	503850	6700050	-50.9601	-29.8299	Lomba Grande	<0.2	<0.5	4	217	<2	2	4	12	0.10	<10	<10	<1	<10	0.04	2	5	0.40	0.02	0.02	<0.01	0.005	<10	<1	<10
CK-S-084	503950	6699800	-50.9591	-29.8322	Lomba Grande	<0.2	<0.5	6	186	<2	4	10	28	0.11	<10	<10	<1	<10	0.08	3	6	0.62	0.02	0.03	0.02	0.009	<10	<1	<10
CK-S-085	502700	6698100	-50.9720	-29.8475	Lomba Grande	<0.2	<0.5	4	72	<2	4	10	21	0.31	<10	<10	<1	<10	0.07	2	6	0.53	0.02	0.05	0.02	0.010	<10	<1	<10
CK-S-086	501750	6696900	-50.9819	-29.8403	Lomba Grande	<0.2	<0.5	4	292	<2	5	9	23	0.37	<10	<10	<1	<10	0.12	4	11	1.02	0.08	0.10	0.02	0.013	<10	<1	<10
CK-S-087	500700	6698400	-50.9928	-29.8448	Lomba Grande	<0.2	<0.5	13	485	<2	20	9	27	0.95	<10	<10	<1	<10	0.22	10	49	1.95	0.03	0.14	0.04	0.012	<10	<5	<10
CK-S-088	501000	6697550	-50.9886	-29.8525	Lomba Grande	<0.2	<0.5	1	35	<2	1	6	8	0.17	<10	<10	<1	<10	0.01	1	4	0.41	0.01	0.02	0.02	0.005	<10	<1	<10
CK-S-089	501050	6700200	-50.9891	-29.8286	Lomba Grande	<0.2	<0.5	1	46	<2	1	6	8	0.18	<10	<10	<1	<10	0.02	1	7	0.46	0.02	0.02	0.02	0.004	<10	<1	<10
CK-S-090	501450	6701400	-50.9850	-29.8177	Lomba Grande	<0.2	<0.5	3	121	<2	2	8	12	0.20	<10	<10	<1	<10	0.03	3	8	0.62	0.02	0.02	0.02	0.006	<10	<1	<10
CK-S-091	501200	6701600	-50.9876	-29.8159	Lomba Grande	<0.2	<0.5	3	182	<2	3	9	15	0.32	<10	<10	<1	<10	0.04	4	8	0.88	0.02	0.03	0.02	0.009	<10	<1	<10
CK-S-092	501100	6701100	-50.9886	-29.8204	Lomba Grande	<0.2	<0.5	3	98	<2	7	8	11	0.22	<10	<10	<1	<10	0.05	2	8	0.80	0.03	0.04	0.02	0.008	<10	<1	<10
CK-S-093	500050	6700950	-50.9895	-29.8218	Lomba Grande	<0.2	<0.5	3	113	<2	2	6	11	0.15	<10	<10	<1	<10	0.02	3	6	0.59	0.01	0.02	0.02	0.004	<10	<1	<10
CK-S-094	502250	6694300	-50.9767	-29.8818	Lomba Grande	<0.2	<0.5	1	80	<2	1	5	10	0.16	<10	<10	<1	<10	0.04	2	3	0.48	<0.01	0.01	0.01	0.012	<10	<1	<10
CK-S-095	500200	6693250	-50.9879	-29.8913	Lomba Grande	<0.2	<0.5	6	62	<2	3	7	16	0.17	<10	<10	<1	<10	0.05	2	4	0.32	0.02	0.03	0.02	0.008	<10	<1	<10
CK-S-096	501700	6691650	-50.9824	-29.9057	Lomba Grande	<0.2	<0.5	6	220	<2	4	13	56	0.71	<10	<10	<1	<10	0.10	6	10	1.52	0.08	0.10	0.02	0.018	<10	<2	<10
CK-S-097	513500	6704750	-50.8603	-29.7874	Lomba Grande	<0.2	<0.5	2	178	<2	3	5	11	0.29	<10	<10	<1	<10	0.04	3	7	0.66	0.02	0.03	0.02	0.005	<10	<1	<10
CK-S-098	513400	6704200	-50.8614	-29.7924	Lomba Grande	<0.2	<0.5	6	203	<2	7	8	21	0.57	<10	<10	<1	<10	0.12	7	17	1.59	0.03	0.09	0.03	0.012	<10	<3	<10
CK-S-100	511500	6702450	-50.8810	-29.8082	Lomba Grande	<0.2	<0.5	<1	405	<2	2	4	11	0.24	<10	<10	<1	<10	0.02	4	5	0.71							

Geochemical grade assay for stream sediments collected in the Lomba Grande district

(2continuerd/4)

Sample No.	Sr ppm	Ti %	V ppm	W ppm	Y ppm	Zr ppm	S %
CK-S-055	2	0.09	27	<10	2	3	0.006
CK-S-056	5	0.01	12	<10	2	2	0.006
CK-S-057	8	0.02	22	<10	2	3	0.008
CK-S-058	4	0.01	11	<10	1	2	0.004
CK-S-059	6	0.02	23	<10	2	3	0.007
CK-S-061	3	0.01	13	<10	1	2	0.004
CK-S-062	3	0.01	10	<10	1	1	0.003
CK-S-063	3	<0.01	7	<10	<1	1	0.004
CK-S-064	5	0.02	23	<10	2	3	0.004
CK-S-065	21	0.03	29	<10	8	4	0.008
CK-S-066	4	0.01	12	<10	2	2	0.005
CK-S-067	3	0.01	25	<10	1	2	0.003
CK-S-068	3	0.02	26	<10	1	2	0.004
CK-S-069	14	0.03	37	<10	4	7	0.004
CK-S-070	5	0.02	34	<10	2	4	0.002
CK-S-071	6	0.01	17	<10	2	4	0.005
CK-S-072	5	0.01	13	<10	2	3	0.005
CK-S-073	5	<0.01	11	<10	2	2	0.004
CK-S-074	6	0.01	10	<10	2	2	0.005
CK-S-075	6	0.02	32	<10	2	3	0.003
CK-S-076	8	0.01	12	<10	3	2	0.005
CK-S-077	6	0.01	13	<10	2	1	0.003
CK-S-078	5	<0.01	13	<10	1	1	0.002
CK-S-079	3	0.01	12	<10	1	2	0.005
CK-S-080	3	<0.01	7	<10	1	2	0.004
CK-S-081	4	<0.01	12	<10	2	2	0.003
CK-S-082	2	0.01	16	<10	<1	2	0.002
CK-S-083	3	<0.01	6	<10	<1	1	0.002
CK-S-084	5	0.01	10	<10	2	2	0.005
CK-S-085	5	0.01	12	<10	3	1	0.010
CK-S-086	11	0.01	16	<10	4	3	0.004
CK-S-087	15	0.05	54	<10	5	10	0.005
CK-S-088	3	0.01	8	<10	1	1	0.004
CK-S-089	3	0.01	10	<10	2	3	0.002
CK-S-090	4	0.01	14	<10	2	2	0.004
CK-S-091	6	0.02	18	<10	2	2	0.006
CK-S-092	5	<0.01	14	<10	2	3	0.004
CK-S-093	3	0.01	12	<10	1	2	0.002
CK-S-094	3	<0.01	11	<10	1	2	0.004
CK-S-095	4	<0.01	9	<10	2	2	0.004
CK-S-096	11	0.01	26	<10	4	6	0.005
CK-S-097	5	0.02	15	<10	1	2	0.003
CK-S-098	11	0.04	45	<10	3	5	0.006
CK-S-100	6	<0.01	8	<10	2	2	0.005
CK-S-102	17	0.08	96	<10	5	11	0.005
CK-S-103	25	0.10	112	<10	10	20	0.008
CK-S-104	22	0.03	115	<10	12	13	0.012
CK-S-105	22	0.02	48	<10	4	10	0.011
CK-S-106	9	<0.01	8	<10	2	2	0.003
CK-S-107	11	0.01	17	<10	3	4	0.005
CK-S-108	12	0.01	16	<10	3	3	0.004
CK-S-109	13	0.01	14	<10	3	3	0.003
CK-S-110	5	0.01	13	<10	3	3	0.002

Geochemical grade assay for stream sediments collected in the Lomba Grande district

Sample No.	UTM_E	UTM_N	Latitude	Longitude	Location	Ag	Cd	Cu	Mn	Mo	Ni	Pb	Zn	Al	As	Ba	Be	Bi	Ca	Co	Cr	Fe	K	Mg	Na	P	Sb	Sc	Sn
CK-S-111	512250	6696900	-50.8732	-29.8583	Lomba Grande	<0.2	<0.5	6	207	<2	5	9	16	0.17	<10	53	<1	<10	0.05	5	18	1.35	0.04	0.04	0.02	0.006	<10	<1	<10
CK-S-112	511850	6697650	-50.8773	-29.8515	Lomba Grande	<0.2	<0.5	13	188	<2	3	8	16	0.13	<10	44	<1	<10	0.04	4	11	0.74	0.03	0.03	0.01	0.007	<10	<1	<10
CK-S-113	511050	6697450	-50.8856	-29.8533	Lomba Grande	<0.2	<0.5	2	30	<2	1	5	8	0.14	<10	22	<1	<10	0.02	1	5	0.39	0.01	0.01	0.01	0.002	<10	<1	<10
CK-S-114	510850	6696700	-50.8877	-29.8601	Lomba Grande	<0.2	<0.5	3	122	<2	5	8	13	0.15	<10	47	<1	<10	0.02	4	23	2.97	0.03	0.02	0.02	0.004	<10	<1	<10
CK-S-115	510750	6696450	-50.8887	-29.8624	Lomba Grande	<0.2	<0.5	3	90	<2	3	4	11	0.21	<10	43	<1	<10	0.03	2	5	0.53	0.04	0.03	0.02	0.005	<10	<1	<10
CK-S-116	510150	6697000	-50.8949	-29.8574	Lomba Grande	<0.2	<0.5	3	162	<2	5	7	12	0.29	<10	52	<1	<10	0.04	3	9	0.61	0.04	0.04	0.02	0.007	<10	<1	<10
CK-S-117	509600	6696150	-50.9006	-29.8651	Lomba Grande	<0.2	<0.5	3	84	<2	<1	4	8	0.15	<10	21	<1	<10	0.06	<1	3	0.63	0.09	0.01	0.03	0.003	<10	<1	<10
CK-S-118	508850	6696350	-50.9084	-29.8615	Lomba Grande	<0.2	<0.5	2	73	<2	3	6	15	0.29	<10	54	<1	<10	0.07	3	4	0.59	0.04	0.06	0.01	0.006	<10	<1	<10
CK-S-119	511400	6696950	-50.8820	-29.8669	Lomba Grande	<0.2	<0.5	3	62	<2	2	7	12	0.12	<10	39	<1	<10	0.02	2	7	0.43	0.03	0.02	0.01	0.005	<10	<1	<10
CK-S-120	510950	6694250	-50.8866	-29.8822	Lomba Grande	<0.2	<0.5	2	87	<2	1	7	11	0.15	<10	43	<1	<10	0.04	2	3	0.47	0.02	0.02	0.01	0.003	<10	<1	<10
CK-S-121	510100	6695550	-50.8954	-29.8705	Lomba Grande	<0.2	<0.5	2	77	<2	2	5	12	0.13	<10	33	<1	<10	0.03	3	10	0.52	0.02	0.02	0.01	0.004	<10	<1	<10
CK-S-122	509550	6695000	-50.9011	-29.8755	Lomba Grande	<0.2	<0.5	4	22	<2	<1	6	14	0.09	<10	23	<1	<10	0.01	1	4	0.29	0.01	<0.01	0.01	0.004	<10	<1	<10
CK-S-123	514950	6699950	-50.8463	-29.8307	Lomba Grande	<0.2	<0.5	5	53	<2	1	7	14	0.19	<10	36	<1	<10	0.03	<1	6	0.41	0.01	0.02	0.01	0.003	<10	<1	<10
CK-S-124	514250	6696900	-50.8525	-29.8339	Lomba Grande	<0.2	<0.5	4	72	<2	1	7	13	0.09	<10	31	<1	<10	0.02	2	7	0.44	0.01	0.01	<0.01	0.003	<10	<1	<10
CK-S-125	514900	6696900	-50.8458	-29.8388	Lomba Grande	<0.2	<0.5	4	77	<2	2	7	14	0.37	<10	54	<1	<10	0.04	2	9	0.86	0.03	0.04	0.02	0.005	<10	<1	<10
CK-S-126	513350	6696900	-50.8618	-29.8393	Lomba Grande	<0.2	<0.5	3	141	<2	4	7	15	0.36	<10	45	<1	<10	0.09	3	14	1.09	0.03	0.05	0.02	0.007	<10	<1	<10
CK-S-127	515700	6696400	-50.8375	-29.8357	Lomba Grande	<0.2	<0.5	2	48	<2	1	5	12	0.16	<10	35	<1	<10	0.02	1	5	0.32	0.02	0.02	0.01	0.004	<10	<1	<10
CK-S-128	515600	6696200	-50.8385	-29.8375	Lomba Grande	<0.2	<0.5	7	66	<2	2	7	17	0.24	<10	46	<1	<10	0.05	2	6	0.51	0.02	0.03	0.01	0.006	<10	<1	<10
CK-S-129	516300	6698250	-50.8313	-29.8461	Lomba Grande	<0.2	<0.5	3	168	<2	3	5	17	0.21	<10	51	<1	<10	0.05	4	6	0.57	0.03	0.05	0.01	0.006	<10	<1	<10
CK-S-130	517050	6701050	-50.9239	-29.8209	Lomba Grande	<0.2	<0.5	2	71	<2	2	5	12	0.23	<10	22	<1	<10	0.02	2	8	0.84	<0.01	0.02	0.01	0.004	<10	<1	<10
CK-S-131	507450	6601100	-50.9229	-29.8204	Lomba Grande	<0.2	<0.5	2	48	<2	1	5	11	0.15	<10	17	<1	<10	0.02	1	3	0.37	<0.01	0.02	<0.01	0.005	<10	<1	<10
CK-S-132	507700	6700300	-50.9203	-29.8276	Lomba Grande	<0.2	<0.5	3	130	<2	2	4	13	0.18	<10	29	<1	<10	0.04	2	5	0.65	<0.01	0.02	<0.01	0.007	<10	<1	<10
CK-S-133	508450	6701250	-50.9125	-29.8191	Lomba Grande	<0.2	<0.5	5	142	<2	1	5	14	0.12	<10	26	<1	<10	0.02	2	4	0.48	<0.01	0.01	<0.01	0.004	<10	<1	<10
CK-S-134	517550	6700450	-50.8184	-29.8262	Lomba Grande	<0.2	<0.5	3	51	<2	<1	5	13	0.06	<10	19	<1	<10	0.03	1	1	0.22	<0.01	0.01	<0.01	0.003	<10	<1	<10
CK-S-135	516550	6700350	-50.8287	-29.8271	Lomba Grande	<0.2	<0.5	4	130	<2	4	4	15	0.31	<10	38	<1	<10	0.10	3	5	0.46	0.02	0.06	0.02	0.006	<10	<1	<10
CK-S-136	516750	6699250	-50.8266	-29.8370	Lomba Grande	<0.2	<0.5	6	88	<2	8	11	21	0.77	<10	84	<1	<10	0.20	5	10	1.47	0.02	0.17	0.03	0.011	<10	<1	<10
CK-S-137	516650	6698900	-50.8276	-29.8402	Lomba Grande	<0.2	<0.5	2	56	<2	2	4	11	0.13	<10	31	<1	<10	0.02	1	4	0.27	0.01	0.02	<0.01	0.004	<10	<1	<10
CK-S-138	517500	6698250	-50.8188	-29.8460	Lomba Grande	<0.2	<0.5	6	98	<2	1	8	15	0.11	<10	26	<1	<10	0.02	<1	3	0.24	0.01	0.02	<0.01	0.004	<10	<1	<10
CK-S-139	517400	6698050	-50.8199	-29.8478	Lomba Grande	<0.2	<0.5	6	50	<2	2	8	17	0.10	<10	26	<1	<10	0.03	1	3	0.29	0.01	0.02	<0.01	0.004	<10	<1	<10
CK-S-140	517900	6698500	-50.8147	-29.8347	Lomba Grande	<0.2	<0.5	5	132	<2	1	6	16	0.16	<10	34	<1	<10	0.03	2	3	0.46	0.02	0.03	<0.01	0.008	<10	<1	<10
CK-S-141	518900	6697800	-50.8043	-29.8501	Lomba Grande	<0.2	<0.5	4	130	<2	<1	6	16	0.13	<10	32	<1	<10	0.03	2	2	0.37	0.01	0.02	<0.01	0.008	<10	<1	<10
CK-S-142	517350	6696050	-50.8204	-29.8659	Lomba Grande	<0.2	<0.5	3	51	<2	<1	6	14	0.14	<10	20	<1	<10	0.02	1	2	0.52	<0.01	0.01	<0.01	0.013	<10	<1	<10
CK-S-143	515500	6698400	-50.8395	-29.8447	Lomba Grande	<0.2	<0.5	2	77	<2	1	5	12	0.07	<10	28	<1	<10	0.02	1	3	0.27	<0.01	0.02	<0.01	0.003	<10	<1	<10
CK-S-144	515600	6698450	-50.8385	-29.8443	Lomba Grande	<0.2	<0.5	4	41	<2	<1	7	15	0.07	<10	26	<1	<10	0.03	1	2	0.28	<0.01	0.02	<0.01	0.003	<10	<1	<10
CK-S-145	516550	6696650	-50.8286	-29.8605	Lomba Grande	<0.2	<0.5	4	20	<2	3	5	14	0.09	<10	11	<1	<10	0.03	<1	4	0.25	<0.01	0.03	0.01	0.003	<10	<1	<10
CK-S-146	513800	6696850	-50.8571	-29.8677	Lomba Grande	<0.2	<0.5	2	41	<2	2	7	15	0.15	<10	29	<1	<10	0.02	1	4	0.40	0.02	0.02	0.01	0.004	<10	<1	<10
CK-S-147	514050	6696850	-50.8545	-29.8677	Lomba Grande	<0.2	<0.5	4	27	<2	<1	5	12	0.10	<10	21	<1	<10	<0.01	1	3	0.33	<0.01	<0.01	0.01	0.003	<10	<1	<10
CK-S-148	512700	6695450	-50.8685	-29.8894	Lomba Grande	<0.2	<0.5	3	37	<2	2	4	12	0.09	<10	28	<1	<10	0.02	1	3	0.28	0.01	0.01	0.01	0.003	<10	<1	<10
CK-S-149	515550	6694200	-50.8392	-29.8826	Lomba Grande	<0.2	<0.5	25	36	<2	7	6	18	0.07	<10	26	<1	<10	0.01	2	4	0.43	0.01	0.01	<0.01	0.009	<10	<1	<10
CK-S-150	515550	6694200	-50.8390	-29.8826	Lomba Grande	<0.2	<0.5	10	25	<2	4	6	<1	0.06	<10	11	<1	<10	37.17	<1	1	0.33	<0.01	<0.01	<0.01	0.005	<10	<1	<10
CK-S-151	515400	6694000	-50.8405	-29.8844	Lomba Grande	<0.2	<0.5	8	15	<2	3	6	12	0.06	<10	20	<1	<10	<0.01	<1	1	0.22	<0.01	<0.01	0.01	0.005	<10	<1	<10
CK-S-152	514450	6692450	-50.8503	-29.8984	Lomba Grande	<0.2	<0.5	5	29	<2	2	4	11	0.06	<10	12	<1	<10	0.02	<1	1	0.19	<0.01	<0.01	0.01	0.003	<10	<1	<10
CK-S-153	514250	6692150	-50.8545	-29.9011	Lomba Grande	<0.2	<0.5	5	29	<2	<1	4	11	0.06	<10	9	<1	<10	0.01	<1	2	0.31	<0.01	<0.01	<0.01	0.007	<10	<1	<10
CK-S-154	512850	6690900	-50.8669	-29.9124	Lomba Grande	<0.2	<0.5	3	67	<2	1	4	11	0.04	<10	15	<1	<10	0.02	2	1	0.41	<0.01	<0.01	0.01	0.004	<10	<1	<10
CK-S-157	515500	6693000	-50.8395	-29.8934	Lomba Grande	<0.2	<0.5	4	167	<2	2	8	12	0.10	<10	37	<1	<10	0.03	3	4	0.79	<0.01	0.01	0.01	0.007			

Geochemical grade assay for stream sediments collected in the Lomba Grande district

(3continuerd/4)

Sample No.	Sr ppm	Ti %	V ppm	W ppm	Y ppm	Zr ppm	S %
CK-S-111	6	0.01	32	<10	3	3	0.004
CK-S-112	4	0.01	17	<10	2	2	0.006
CK-S-113	3	<0.01	12	<10	2	2	0.003
CK-S-114	7	0.04	74	<10	2	7	0.002
CK-S-115	7	<0.01	8	<10	2	2	0.004
CK-S-116	9	0.01	11	<10	3	2	0.003
CK-S-117	6	<0.01	9	<10	2	2	0.003
CK-S-118	8	<0.01	14	<10	6	4	0.002
CK-S-119	4	<0.01	11	<10	5	2	0.003
CK-S-120	3	<0.01	9	<10	3	3	0.003
CK-S-121	3	<0.01	11	<10	2	2	0.003
CK-S-122	1	<0.01	7	<10	2	1	0.004
CK-S-123	3	<0.01	14	<10	2	3	0.002
CK-S-124	2	<0.01	11	<10	1	1	0.004
CK-S-125	8	<0.01	22	<10	4	5	0.003
CK-S-126	9	0.02	24	<10	2	3	0.004
CK-S-127	5	<0.01	7	<10	1	1	0.003
CK-S-128	6	<0.01	12	<10	4	3	0.006
CK-S-129	5	<0.01	11	<10	3	2	0.003
CK-S-130	3	0.02	24	<10	1	2	0.002
CK-S-131	2	<0.01	9	<10	<1	1	0.003
CK-S-132	3	<0.01	14	<10	1	2	0.004
CK-S-133	2	<0.01	11	<10	1	1	0.004
CK-S-134	2	<0.01	4	<10	<1	<1	0.003
CK-S-135	7	0.01	10	<10	2	2	0.008
CK-S-136	16	0.02	27	<10	5	11	0.006
CK-S-137	3	<0.01	6	<10	2	<1	0.003
CK-S-138	3	<0.01	5	<10	1	<1	0.004
CK-S-139	3	<0.01	6	<10	2	1	0.003
CK-S-140	3	<0.01	7	<10	2	1	0.004
CK-S-141	3	<0.01	4	<10	1	<1	0.004
CK-S-142	2	<0.01	15	<10	1	1	0.004
CK-S-143	2	<0.01	6	<10	1	<1	0.002
CK-S-144	2	<0.01	7	<10	2	1	0.002
CK-S-145	2	<0.01	8	<10	<1	1	0.003
CK-S-146	3	<0.01	13	<10	2	2	0.003
CK-S-147	2	<0.01	7	<10	2	2	0.003
CK-S-148	3	<0.01	7	<10	2	1	0.003
CK-S-149	2	<0.01	6	<10	1	1	0.006
CK-S-150	1	<0.01	8	<10	<1	1	<0.001
CK-S-151	1	<0.01	4	<10	1	<1	0.003
CK-S-152	1	<0.01	5	<10	<1	<1	0.003
CK-S-153	1	<0.01	13	<10	<1	1	0.004
CK-S-154	1	<0.01	10	<10	<1	1	0.003
CK-S-157	3	<0.01	25	<10	3	3	0.005
CK-S-158	2	<0.01	6	<10	1	2	0.003
CK-S-159	1	<0.01	8	<10	<1	2	0.002
CK-S-161	10	0.03	94	<10	3	27	0.004
CK-S-162	6	0.02	62	<10	2	16	0.004
CK-S-163	2	<0.01	18	<10	1	3	0.004
CK-S-164	3	<0.01	21	<10	2	5	0.004
CK-S-165	3	0.02	32	<10	1	6	0.008
CK-S-166	3	0.01	11	<10	2	3	0.004

Geochemical grade assay for stream sediments collected in the Lomba Grande district

(4/4)

Sample No.	UTM_E	UTM_N	Latitude	Longitude	Location	Ag ppm	Cd ppm	Cu ppm	Mn ppm	Mo ppm	Ni ppm	Pb ppm	Zn ppm	Al %	As ppm	Ba ppm	Be ppm	Bi ppm	Ca %	Co ppm	Cr ppm	Fe %	K %	Mg %	Na %	P %	Sb ppm	Sc ppm	Sn ppm
CK-S-167	509600	6692850	-50.9006	-29.8949	Lomba Grande	<0.2	<0.5	2	57	<2	3	6	8	0.18	<10	27	<1	<10	0.02	2	3	0.27	0.02	0.02	0.02	0.005	<10	<1	<10
CK-S-168	508300	6689550	-50.9140	-29.9246	Lomba Grande	<0.2	<0.5	10	347	<2	7	12	43	0.86	<10	56	<1	<10	0.09	10	15	2.45	0.03	0.09	0.02	0.010	<10	5	<10
CK-S-169	508250	6689800	-50.9145	-29.9224	Lomba Grande	<0.2	<0.5	20	100	<2	8	7	21	0.50	12	29	<1	<10	0.26	4	9	0.82	0.02	0.10	0.05	0.016	<10	1	<10
CK-S-170	506550	6690350	-50.9321	-29.9174	Lomba Grande	<0.2	<0.5	2	81	<2	3	5	12	0.24	<10	19	<1	<10	0.05	2	6	0.74	0.01	0.02	0.03	0.005	<10	1	<10
CK-S-172	503150	6690450	-50.9674	-29.9165	Lomba Grande	<0.2	<0.5	2	48	<2	<1	7	15	0.22	<10	30	<1	<10	0.04	2	3	0.33	0.04	0.04	0.02	0.004	<10	<1	<10
CK-S-173	502950	6690250	-50.9694	-29.9184	Lomba Grande	<0.2	<0.5	4	156	<2	2	9	19	0.34	<10	66	<1	<10	0.06	4	6	0.73	0.03	0.05	0.02	0.006	<10	1	<10
CK-S-174	503200	6691950	-50.9669	-29.9030	Lomba Grande	<0.2	<0.5	2	39	<2	<1	5	14	0.13	<10	23	<1	<10	0.04	1	2	0.27	0.02	0.03	0.01	0.003	<10	<1	<10
CK-S-175	505400	6692600	-50.9441	-29.8971	Lomba Grande	<0.2	<0.5	6	139	<2	6	9	21	0.67	<10	58	<1	<10	0.28	4	12	0.89	0.07	0.14	0.04	0.008	<10	2	<10
CK-S-176	505950	6690700	-50.9384	-29.9143	Lomba Grande	<0.2	<0.5	7	80	<2	9	8	20	0.46	<10	39	<1	<10	0.11	3	14	0.63	0.04	0.10	0.02	0.006	<10	1	<10
CK-S-177	508150	6692950	-50.9156	-29.8940	Lomba Grande	<0.2	<0.5	3	137	<2	2	6	18	0.38	<10	39	<1	<10	0.06	3	6	0.58	0.04	0.05	0.02	0.010	<10	1	<10
CK-S-178	504650	6695800	-50.9519	-29.8683	Lomba Grande	<0.2	<0.5	<1	80	<2	2	5	10	0.32	<10	31	<1	<10	0.04	2	4	0.38	0.04	0.04	0.02	0.006	<10	<1	<10
CK-S-179	507900	6694450	-50.9182	-29.8804	Lomba Grande	<0.2	<0.5	2	178	<2	3	6	14	0.42	<10	41	<1	<10	0.04	3	6	0.44	0.06	0.06	0.02	0.009	<10	1	<10
CK-S-180	506100	6695350	-50.9368	-29.8723	Lomba Grande	<0.2	<0.5	<1	402	<2	3	8	17	0.53	<10	61	<1	<10	0.07	5	7	0.65	0.07	0.07	0.02	0.013	<10	1	<10
CK-S-181	504550	6696400	-50.9529	-29.8628	Lomba Grande	<0.2	<0.5	<1	163	<2	4	7	15	0.45	<10	43	<1	<10	0.06	3	6	0.56	0.10	0.09	0.02	0.010	<10	1	<10
CK-S-182	491600	6698850	-51.0870	-29.8407	Lomba Grande	<0.2	<0.5	10	55	<2	3	14	32	0.28	<10	27	<1	<10	0.05	1	4	0.35	0.03	0.03	0.02	0.005	<10	<1	<10

Geochemical grade assay for stream sediments collected in the Lomba Grande district

(4continuerd/4)

Sample No.	Sr ppm	Ti %	V ppm	W ppm	Y ppm	Zr ppm	S %
CK-S-167	3	<0.01	6	<10	2	2	0.003
CK-S-168	8	0.05	66	<10	3	15	0.003
CK-S-169	10	0.03	24	<10	2	6	0.011
CK-S-170	3	0.04	31	<10	1	6	0.002
CK-S-172	3	<0.01	8	<10	2	2	0.003
CK-S-173	6	<0.01	16	<10	3	4	0.005
CK-S-174	3	<0.01	6	<10	1	2	0.002
CK-S-175	12	0.02	23	<10	4	7	0.003
CK-S-176	7	0.02	18	<10	3	4	0.003
CK-S-177	7	0.01	11	<10	3	2	0.005
CK-S-178	6	<0.01	8	<10	3	3	0.003
CK-S-179	7	0.01	9	<10	4	2	0.006
CK-S-180	10	<0.01	10	<10	3	2	0.006
CK-S-181	10	<0.01	7	<10	3	3	0.004
CK-S-182	6	<0.01	8	<10	1	2	0.003

Geochemical grade assay for stream sediments collected in the São Gabriel district

Sample No.	UTM_E	UTM_N	Latitude	Longitude	Location	Ag	Cd	Cu	Mn	Mo	Ni	Pb	Zn	Al	As	Ba	Be	Bi	Ca	Co	Cr	Fe	K	Mg	Na	P	Sb	Sc	Sn
SJ-001			-29.80532	-55.78368	São Gabriel	<0.2	<0.5	21	320	<2	11	8	26	0.61	<10	62	<1	<10	0.18	8	10	1.48	0.02	0.08	0.04	0.007	<10	5	<10
SJ-002			-29.57620	-55.98628	São Gabriel	<0.2	<0.5	45	1731	4	22	15	53	2.04	<10	225	1	<10	0.70	44	27	4.19	0.05	0.27	0.08	0.016	<10	15	<10
SJ-003			-29.72000	-55.85834	São Gabriel	<0.2	<0.5	72	1553	3	36	8	65	2.95	<10	262	1	<10	1.11	40	54	4.17	0.05	0.44	0.11	0.006	<10	24	<10
SJ-004			-29.72481	-56.05511	São Gabriel	<0.2	<0.5	69	1823	3	32	12	81	3.22	<10	270	2	<10	1.00	37	45	4.06	0.05	0.42	0.08	0.008	<10	25	<10
SJ-005			-29.81949	-55.77008	São Gabriel	<0.2	<0.5	18	307	<2	17	7	28	0.89	<10	79	<1	<10	0.17	9	11	1.40	0.02	0.09	0.04	0.016	<10	6	<10
SJ-006			-29.97298	-55.74703	São Gabriel	<0.2	<0.5	38	563	3	16	<2	47	1.55	<10	105	<1	<10	0.53	16	19	2.92	0.03	0.23	0.06	0.008	<10	14	<10
SJ-007			-29.95093	-55.59867	São Gabriel	<0.2	<0.5	16	99	3	9	7	26	0.80	<10	56	<1	<10	0.12	5	12	1.18	0.02	0.07	0.03	0.010	<10	5	<10
SJ-008			-30.15953	-55.57694	São Gabriel	<0.2	<0.5	40	569	<2	16	9	56	2.22	<10	216	1	<10	0.50	18	15	3.54	0.04	0.17	0.03	0.028	<10	17	<10
SJ-009			-30.07010	-55.77653	São Gabriel	<0.2	<0.5	61	1433	<2	25	7	68	2.26	<10	221	1	<10	0.60	38	25	5.28	0.05	0.32	0.06	0.012	<10	21	<10
SJ-010			-30.02345	-55.86850	São Gabriel	<0.2	<0.5	44	1289	<2	19	6	55	1.85	<10	193	1	<10	0.67	31	23	3.80	0.03	0.27	0.07	0.016	<10	16	<10
SJ-011			-29.92994	-55.86998	São Gabriel	<0.2	<0.5	53	830	5	20	8	58	2.04	<10	169	1	<10	0.80	20	28	4.13	0.05	0.31	0.05	0.011	<10	19	<10
SJ-012			-29.73854	-56.25440	São Gabriel	<0.2	<0.5	53	861	5	25	11	52	2.00	<10	162	<1	<10	1.22	33	32	4.53	0.06	0.46	0.18	0.020	<10	21	<10
SJ-013			-29.78532	-56.36606	São Gabriel	<0.2	<0.5	64	1600	2	25	12	71	2.56	<10	248	1	<10	1.00	32	29	4.32	0.06	0.37	0.09	0.022	<10	20	<10
SJ-014			-29.71876	-56.67583	São Gabriel	<0.2	<0.5	85	1631	3	30	11	71	3.09	<10	336	2	<10	1.24	33	35	4.86	0.06	0.48	0.13	0.024	<10	23	<10
SJ-015			-29.74362	-56.69124	São Gabriel	<0.2	<0.5	71	1903	4	28	15	77	3.03	<10	369	2	<10	0.98	39	26	5.85	0.07	0.44	0.07	0.024	<10	24	<10
SJ-016			-29.91547	-56.04875	São Gabriel	<0.2	<0.5	87	2779	<2	44	3	54	3.34	<10	342	1	<10	1.14	55	46	5.15	0.04	0.47	0.17	0.022	<10	22	<10
SJ-017			-30.16495	-55.91664	São Gabriel	<0.2	<0.5	82	1620	<2	34	7	81	3.25	<10	282	2	<10	0.89	42	47	6.53	0.07	0.41	0.08	0.016	<10	28	<10
SJ-018			-30.24133	-55.89169	São Gabriel	<0.2	<0.5	73	2431	6	25	16	86	2.56	<10	354	2	<10	0.94	52	20	6.92	0.08	0.34	0.09	0.031	<10	22	<10
SJ-019			-30.55491	-55.93910	São Gabriel	<0.2	<0.5	84	3307	5	33	8	78	2.87	<10	589	2	<10	1.03	70	32	6.51	0.06	0.40	0.10	0.021	<10	23	<10
SJ-020			-30.49425	-56.10691	São Gabriel	<0.2	<0.5	58	1056	4	19	8	60	2.16	<10	292	2	<10	0.78	27	19	4.39	0.06	0.38	0.08	0.022	<10	17	<10
SJ-021			-30.49425	-56.20988	São Gabriel	<0.2	<0.5	56	1332	<2	22	4	72	2.31	<10	232	2	<10	0.72	26	22	4.66	0.05	0.32	0.04	0.021	<10	22	<10
SJ-022			-30.45125	-56.30272	São Gabriel	<0.2	<0.5	24	717	2	10	9	34	0.93	<10	139	<1	<10	0.20	16	9	2.40	0.02	0.09	0.02	0.014	<10	9	<10
SJ-023			-30.38882	-56.43548	São Gabriel	<0.2	<0.5	63	1461	5	23	11	134	2.23	<10	328	1	<10	1.30	34	22	4.01	0.04	0.39	0.11	0.021	<10	14	<10
SJ-024			-30.24397	-56.30597	São Gabriel	<0.2	<0.5	74	1800	6	30	13	94	2.83	<10	257	2	<10	1.43	34	36	4.41	0.05	0.40	0.09	0.041	<10	19	<10
SJ-025			-30.62689	-55.68289	São Gabriel	<0.2	<0.5	59	1463	<2	24	7	77	2.83	<10	302	2	<10	0.71	35	26	5.84	0.07	0.36	0.04	0.020	<10	26	<10
SJ-026			-30.45639	-55.71204	São Gabriel	<0.2	<0.5	64	1558	4	26	3	72	2.17	<10	257	2	<10	0.68	49	32	5.60	0.04	0.30	0.07	0.018	<10	20	<10
SJ-027			-30.44213	-55.74472	São Gabriel	<0.2	<0.5	84	1558	4	26	3	72	2.17	<10	257	2	<10	1.22	65	65	6.79	0.05	0.44	0.11	0.018	<10	30	<10
SJ-028			-30.87885	-55.85689	São Gabriel	<0.2	<0.5	50	1159	<2	12	14	88	2.55	<10	304	3	<10	0.63	32	11	6.69	0.07	0.22	0.03	0.044	<10	24	<10
SJ-029			-30.91931	-55.95766	São Gabriel	<0.2	<0.5	63	1214	<2	17	12	94	2.52	<10	281	3	<10	0.52	33	26	7.80	0.08	0.32	0.04	0.036	<10	25	<10
SJ-030			-30.82867	-55.84183	São Gabriel	<0.2	<0.5	56	1566	2	15	20	112	2.54	<10	452	3	<10	0.62	29	10	5.47	0.08	0.15	0.02	0.063	<10	23	<10
SJ-031			-30.25256	-56.50440	São Gabriel	<0.2	<0.5	51	1116	5	19	13	57	1.98	<10	231	1	<10	0.73	25	27	3.90	0.05	0.33	0.06	0.016	<10	17	<10
SJ-032			-30.19128	-56.53511	São Gabriel	<0.2	<0.5	4	63	<2	3	3	7	0.09	<10	36	<1	<10	0.03	2	1	0.36	<0.01	0.01	0.01	0.003	<10	<1	<10
SJ-033			-30.14058	-56.55071	São Gabriel	<0.2	<0.5	67	1926	3	24	10	72	2.13	<10	316	2	<10	0.89	33	30	3.86	0.04	0.33	0.04	0.031	<10	18	<10
SJ-034			-30.12144	-56.42154	São Gabriel	<0.2	<0.5	68	2087	3	30	11	83	2.83	<10	316	2	<10	1.05	35	35	4.66	0.06	0.38	0.06	0.031	<10	22	<10
SJ-035			-30.01502	-56.82671	São Gabriel	<0.2	<0.5	48	1655	2	17	12	59	1.77	<10	420	1	<10	0.68	32	16	3.27	0.05	0.26	0.04	0.035	<10	12	<10
SJ-036			-30.08952	-56.76924	São Gabriel	<0.2	<0.5	59	3329	3	26	20	62	2.48	<10	608	2	<10	0.99	45	21	4.06	0.06	0.39	0.06	0.018	<10	16	<10
SJ-037			-29.88367	-57.11540	São Gabriel	<0.2	<0.5	39	1070	<2	19	11	50	1.75	<10	214	1	<10	0.76	22	22	2.66	0.03	0.25	0.04	0.023	<10	11	<10
SJ-038			-29.75201	-56.99299	São Gabriel	<0.2	<0.5	70	1748	6	30	4	75	3.15	<10	426	2	<10	0.94	38	26	4.85	0.06	0.39	0.05	0.019	<10	21	<10
SJ-039			-29.65353	-56.89146	São Gabriel	<0.2	<0.5	37	836	4	22	17	64	2.59	<10	166	1	<10	0.73	19	28	3.30	0.06	0.37	0.03	0.009	<10	15	<10
SJ-040			-29.49741	-56.59528	São Gabriel	<0.2	<0.5	67	1461	6	26	11	84	3.11	<10	231	2	<10	0.63	21	34	4.88	0.06	0.42	0.03	0.008	<10	25	<10
SJ-041			-29.46838	-56.47515	São Gabriel	<0.2	<0.5	22	920	2	9	6	26	0.62	<10	125	<1	<10	0.24	15	9	1.34	0.02	0.10	0.01	0.011	<10	5	<10
SJ-042			-29.46535	-56.29649	São Gabriel	<0.2	<0.5	22	920	2	9	6	26	0.62	<10	125	<1	<10	0.29	46	13	8.67	0.03	0.08	0.02	0.026	<10	26	<10
LA-S-001	690295	6836812	-28.56159	-55.05425	São Gabriel	<0.2	<0.5	97	1798	3	21	16	82	1.61	<10	245	2	<10	0.30	51	8	10.03	0.04	0.10	0.02	0.026	<10	28	<10
LA-S-002	688984	6837330	-28.57562	-54.96554	São Gabriel	<0.2	<0.5	130	1953	4	20	12	80	1.93	<10	203	3	<10	0.90	51	13	9.81	0.06	0.16	0.03	0.028	<10	30	<10
LA-S-003	701552	6832278	-28.62079	-54.93841	São Gabriel	<0.2	<0.5	120	1233	<2	24	16	73	2.74	<10	204	2	<10	0.35	42	14	9.81	0.06	0.16	0.03	0.028	<10	30	<10
LA-S-004	646828	6844846	-28.51474	-55.49958	São Gabriel	<0.2	<0.5	94	3764	3	31	18	62	2.16	<10	599	2	<10	1.47	82	35	6.85	0.02	0.31	0.05	0.016	<10	21	<10
LA-S-005	683088	6847434	-28.48680	-55.12960	São Gabriel	<0.2	<0.5	34	756	3	10	23	38	0.62	<10	96	<1	<10	0.06	20	17	3.84	<0.01	0.04	0.02	0.004	<10	11	<10

Geochemical grade assay for stream sediments collected in the São Gabriel district

Sample No.	Sr ppm	Ti %	V ppm	W ppm	Y ppm	Zr ppm	S %
SJ-001	25	0.10	75	<10	6	24	0.009
SJ-002	52	0.08	130	<10	20	36	0.010
SJ-003	94	0.01	34	<10	31	24	0.009
SJ-004	81	<0.01	25	<10	35	24	0.011
SJ-005	16	0.04	57	<10	6	13	0.013
SJ-006	55	0.09	117	<10	16	31	0.009
SJ-007	12	0.04	45	<10	4	13	0.012
SJ-008	33	0.02	106	<10	23	30	0.017
SJ-009	48	0.12	214	<10	26	52	0.012
SJ-010	65	0.13	159	<10	20	66	0.010
SJ-011	51	0.07	168	<10	24	70	0.014
SJ-012	73	0.18	155	<10	16	86	0.005
SJ-013	71	0.14	157	<10	26	53	0.024
SJ-014	81	0.11	179	<10	28	55	0.026
SJ-015	84	0.11	219	<10	30	66	0.023
SJ-016	59	0.11	196	<10	19	60	0.029
SJ-017	87	0.16	273	<10	35	75	0.017
SJ-018	76	0.20	353	<10	30	81	0.029
SJ-019	108	0.18	307	<10	31	78	0.027
SJ-020	70	0.13	181	<10	23	73	0.017
SJ-021	47	0.09	182	<10	31	56	0.021
SJ-022	16	0.04	102	<10	15	29	0.019
SJ-023	42	0.09	158	<10	19	58	0.022
SJ-024	72	0.10	169	<10	28	46	0.072
SJ-025	129	0.17	258	<10	38	76	0.022
SJ-026	102	0.24	316	<10	23	87	0.011
SJ-027	98	0.18	302	<10	24	88	0.017
SJ-028	84	0.24	300	<10	49	77	0.043
SJ-029	58	0.34	380	<10	39	97	0.015
SJ-030	60	0.19	258	<10	67	49	0.066
SJ-031	46	0.11	171	<10	24	47	0.021
SJ-032	2	<0.01	14	<10	2	4	0.003
SJ-033	50	0.05	154	<10	27	42	0.033
SJ-034	59	0.09	179	<10	30	46	0.037
SJ-035	50	0.09	125	<10	22	42	0.023
SJ-036	60	0.11	170	<10	30	49	0.025
SJ-037	60	0.12	170	<10	24	70	0.027
SJ-038	34	0.07	107	<10	15	31	0.032
SJ-039	56	0.13	190	<10	26	74	0.021
SJ-040	36	0.07	106	<10	16	59	0.015
SJ-041	42	0.05	159	<10	36	46	0.024
SJ-042	14	<0.01	57	<10	10	15	0.014
LA-S-001	25	0.19	406	<10	45	86	0.012
LA-S-002	23	0.37	547	<10	36	112	0.014
LA-S-003	30	0.31	469	<10	34	121	0.010
LA-S-004	71	0.29	340	<10	40	82	0.007
LA-S-005	6	0.49	201	<10	10	75	0.006
LA-S-006	23	0.30	333	<10	25	90	0.010
LA-S-007	23	0.42	300	<10	20	80	0.005
LA-S-008	19	0.17	330	<10	22	74	0.014
LA-S-009	7	0.54	638	<10	8	91	0.004
LA-S-010	25	0.33	420	<10	32	91	0.012
LA-S-011	8	0.39	485	<10	8	73	0.005
LA-S-012	28	0.35	499	<10	21	96	0.005

Geochemical grade assay for stream sediments collected in the São Gabriel district

(2/7)

Sample No.	UTM_E	UTM_N	Latitude	Longitude	Location	Ag ppm	Cd ppm	Cu ppm	Mn ppm	Mo ppm	Ni ppm	Pb ppm	Zn ppm	Al %	As ppm	Ba ppm	Be ppm	Bi ppm	Ca %	Co ppm	Cr ppm	Fe %	K %	Mg %	Na %	P %	Sb ppm	Sc ppm	Sn ppm
LA-S-013	678715	6865756	-28.32210	-55.17709	São Gabriel	<0.2	<0.5	105	1926	4	26	14	61	2.61	<10	231	2	12	0.38	44	26	7.95	0.02	0.11	0.03	0.025	<10	32	<10
LA-S-014	670601	6875187	-28.23809	-55.26120	São Gabriel	<0.2	<0.5	61	1075	3	21	9	49	1.90	<10	141	1	19	0.26	32	19	5.62	0.03	0.12	0.03	0.015	<10	17	<10
LA-S-015	692292	6888533	-28.11470	-55.04242	São Gabriel	<0.2	<0.5	105	1142	<2	37	12	60	1.46	<10	182	1	10	0.21	30	34	6.48	0.02	0.10	0.03	0.010	<10	15	<10
LA-S-016	704910	6897978	-28.02760	-54.91568	São Gabriel	<0.2	<0.5	168	1434	2	37	5	84	2.70	<10	267	2	<10	0.47	41	52	9.23	0.05	0.25	0.05	0.028	<10	24	<10
LA-S-017	707321	6898486	-28.02264	-54.89125	São Gabriel	<0.2	<0.5	145	1739	3	34	15	84	2.76	<10	245	2	15	0.31	50	45	9.41	0.04	0.14	0.03	0.028	<10	28	<10
LA-S-018	710181	6886946	-28.12629	-54.86012	São Gabriel	<0.2	<0.5	183	1489	4	42	6	105	2.69	<10	249	2	<10	0.28	43	57	11.26	0.04	0.12	0.03	0.025	<10	30	<10
LA-S-019	698640	6893994	-28.06449	-54.97875	São Gabriel	<0.2	<0.5	110	2050	<2	25	10	65	2.93	<10	252	2	<10	0.27	38	21	7.84	0.03	0.13	0.03	0.013	<10	36	<10
LA-S-020	717389	6893424	-28.06670	-54.78628	São Gabriel	<0.2	<0.5	125	1542	5	30	9	92	2.22	<10	221	2	<10	0.26	41	43	8.92	0.03	0.09	0.03	0.025	<10	23	<10
LA-S-021	720613	6893945	-28.06236	-54.75527	São Gabriel	<0.2	<0.5	186	1537	<2	46	9	90	2.61	<10	227	2	<10	0.47	46	63	9.90	0.03	0.25	0.05	0.020	<10	29	<10
LA-S-022	732180	6895270	-28.04754	-54.63814	São Gabriel	<0.2	<0.5	249	1626	4	45	7	105	1.91	<10	202	2	<10	0.35	49	56	12.01	0.03	0.18	0.04	0.016	<10	26	<10
LA-S-023	738718	6895599	-28.04341	-54.57154	São Gabriel	<0.2	<0.5	203	1814	<2	43	13	83	2.91	<10	219	2	29	0.31	53	57	12.35	0.04	0.09	0.03	0.024	15	41	<10
LA-S-024	741162	6889130	-28.10131	-54.54537	São Gabriel	<0.2	<0.5	227	1939	<2	44	12	100	2.94	<10	231	2	<10	0.24	57	74	13.78	0.03	0.07	0.03	0.025	16	35	<10
LA-S-025	726450	6872786	-28.25133	-54.69184	São Gabriel	<0.2	<0.5	97	1503	<2	32	6	60	2.44	<10	129	2	<10	0.08	36	32	8.67	0.02	0.06	0.03	0.024	<10	23	<10
LA-S-026	713768	6871981	-28.26071	-54.82088	São Gabriel	<0.2	<0.5	82	1086	<2	35	20	71	2.25	<10	223	2	<10	0.19	37	20	8.91	0.02	0.04	0.03	0.029	<10	25	<10
LA-S-027	727067	6855844	-28.40403	-54.68224	São Gabriel	<0.2	<0.5	122	1709	<2	35	11	75	2.53	<10	226	2	<10	0.31	52	27	9.63	0.02	0.09	0.03	0.032	<10	35	<10
LA-S-028	734611	6830270	-28.63334	-54.60007	São Gabriel	<0.2	<0.5	137	1613	<2	41	12	89	2.46	<10	271	2	<10	0.45	45	54	8.90	0.03	0.18	0.03	0.033	20	31	<10
LA-S-029	724368	6836477	-28.57944	-54.70601	São Gabriel	<0.2	<0.5	95	1244	<2	27	11	61	1.04	<10	196	1	<10	0.20	36	27	6.27	0.02	0.08	0.03	0.010	<10	15	<10
LA-S-030	714120	6839077	-28.56647	-54.81101	São Gabriel	<0.2	<0.5	102	1716	<2	32	12	70	1.61	<10	182	2	<10	0.25	37	33	8.80	0.02	0.10	0.03	0.014	<10	23	<10
LA-S-031	729830	6881020	-28.17647	-54.65904	São Gabriel	<0.2	<0.5	102	701	<2	23	15	74	0.75	<10	109	1	<10	0.17	31	21	7.56	0.01	0.05	0.03	0.014	15	11	<10
LA-S-032	742237	6821786	-28.70844	-54.52032	São Gabriel	<0.2	<0.5	226	1219	<2	46	21	121	1.38	<10	161	2	<10	0.17	42	74	10.28	0.01	0.05	0.02	0.025	<10	23	<10
LA-S-033	741963	6803978	-28.86908	-54.51932	São Gabriel	<0.2	<0.5	151	1101	<2	27	16	116	0.62	<10	98	2	29	0.12	40	17	10.64	<0.01	0.02	0.02	0.010	<10	15	<10
LA-S-034	752853	6818595	-28.73518	-54.41103	São Gabriel	<0.2	<0.5	145	1864	<2	35	12	85	1.59	<10	289	2	25	0.37	55	57	9.09	0.03	0.11	0.03	0.027	<10	22	<10
LA-S-035	799587	6812089	-28.79250	-54.34065	São Gabriel	<0.2	<0.5	121	1627	<2	34	10	80	2.20	<10	286	2	<10	0.53	48	40	8.95	0.04	0.22	0.03	0.022	<10	28	<10
LA-S-036	756246	6823024	-28.69458	-54.37732	São Gabriel	<0.2	<0.5	140	1452	<2	36	12	73	2.35	<10	230	2	16	0.31	38	52	8.94	0.03	0.12	0.03	0.025	<10	28	<10
LA-S-037	758759	6825844	-28.66865	-54.35226	São Gabriel	<0.2	<0.5	104	1002	<2	25	7	60	1.29	<10	135	1	<10	0.19	27	36	6.31	0.02	0.08	0.02	0.014	<10	16	<10
LA-S-038	766999	6844085	-28.50251	-54.27227	São Gabriel	<0.2	<0.5	135	1878	<2	34	16	82	2.08	<10	251	2	<10	0.39	55	59	9.72	0.03	0.11	0.02	0.036	<10	29	<10
LA-S-039	751743	6831213	-28.82162	-54.42517	São Gabriel	<0.2	<0.5	182	1796	<2	40	18	89	2.39	<10	287	2	<10	0.35	50	66	10.29	0.03	0.16	0.03	0.025	16	32	<10
LA-S-040	740823	6861566	-28.34996	-54.54312	São Gabriel	<0.2	<0.5	239	1772	<2	52	20	137	1.25	<10	207	2	19	0.25	51	71	11.41	0.02	0.11	0.03	0.014	19	19	<10
LA-S-041	741185	6855620	-28.40352	-54.53820	São Gabriel	<0.2	<0.5	262	1756	<2	50	15	129	1.83	<10	234	2	<10	0.32	56	76	12.14	0.02	0.14	0.03	0.021	31	27	<10
LA-S-042	749813	6856306	-28.39571	-54.45035	São Gabriel	<0.2	<0.5	152	1649	<2	35	20	82	2.09	<10	244	2	12	0.28	48	59	10.31	0.03	0.09	0.02	0.030	<10	31	<10
LA-S-043	756998	6856119	-28.39601	-54.37704	São Gabriel	<0.2	<0.5	168	1637	<2	41	13	85	2.10	<10	249	2	<10	0.31	50	69	11.20	0.03	0.10	0.02	0.030	14	32	<10
LA-S-044	764067	6854714	-28.40727	-54.30463	São Gabriel	<0.2	<0.5	116	1253	<2	29	11	73	1.52	<10	233	2	<10	0.32	39	42	7.74	0.03	0.11	0.02	0.037	<10	24	<10
LA-S-045	780758	6827138	-28.65240	-54.12770	São Gabriel	<0.2	<0.5	274	1969	<2	41	24	132	1.57	<10	264	2	33	0.26	45	74	11.49	0.02	0.15	0.02	0.037	<10	27	<10
LA-S-046	779947	6825929	-28.66347	-54.13569	São Gabriel	<0.2	<0.5	151	1625	<2	37	39	105	1.94	<10	323	2	<10	0.34	55	53	8.59	0.03	0.17	0.02	0.038	<10	25	<10
LA-S-047	764150	6834740	-28.65312	-54.09492	São Gabriel	<0.2	<0.5	193	1077	2	37	23	122	1.37	<10	182	2	22	0.20	37	57	8.79	0.02	0.09	0.02	0.026	<10	19	<10
LA-S-048	775831	6858111	-28.37426	-54.18752	São Gabriel	<0.2	<0.5	236	1828	<2	50	20	130	2.61	<10	236	2	24	0.24	51	94	11.81	0.03	0.07	0.02	0.039	11	34	<10
LA-S-049	763240	6858348	-28.37055	-54.11002	São Gabriel	<0.2	<0.5	243	2799	<2	57	30	130	2.63	<10	288	2	<10	0.26	50	68	11.91	0.04	0.11	0.02	0.029	<10	27	<10
LA-S-050	790826	6856768	-28.36308	-54.09230	São Gabriel	<0.2	<0.5	267	1473	<2	48	17	153	2.00	<10	214	2	14	0.26	50	60	11.72	0.03	0.11	0.02	0.029	<10	27	<10
LA-S-051	774040	6870330	-28.26442	-54.20662	São Gabriel	<0.2	<0.5	240	1876	<2	41	27	120	2.84	<10	245	2	18	0.26	40	52	11.05	0.03	0.11	0.02	0.042	<10	35	<10
LA-S-052	782936	6872300	-28.24478	-54.11651	São Gabriel	<0.2	<0.5	175	2206	<2	40	24	121	2.67	<10	382	2	<10	0.43	62	70	10.93	0.04	0.15	0.02	0.041	<10	33	<10
LA-S-053	769730	6869957	-28.27769	-54.25019	São Gabriel	<0.2	<0.5	221	3032	2	33	28	127	1.58	<10	571	2	27	0.29	65	68	12.02	0.02	0.06	0.02	0.046	11	23	<10
LA-S-054	772445	6897766	-28.01738	-54.22924	São Gabriel	<0.2	<0.5	112	1244	<2	34	20	87	2.01	<10	204	2	14	0.26	35	63	10.69	0.01	0.02	0.01	0.046	<10	36	<10
LA-S-055	778410	6890530	-27.99122	-54.16929	São Gabriel	<0.2	<0.5	264	1626	<2	42	12	130	1.54	<10	196	2	25	0.18	52	55	12.37	0.01	0.03	0.01	0.025	<10	28	<10
LA-S-056	762874	6882315	-28.15869	-54.32609	São Gabriel	<0.2	<0.5	211	1653	<2	40	26	133	1.87	<10	231	2	13	0.30	50	60	11.66	0.02	0.10	0.02	0.030	<10	31	<10
LA-S-057	751767	6868237	-28																										

Geochemical grade assay for stream sediments collected in the São Gabriel district

Sample No.	Sr ppm	Tl %	V ppm	W ppm	Y ppm	Zr ppm	S %
LA-S-013	29	0.15	296	<10	37	70	0.025
LA-S-014	18	0.16	256	<10	16	62	0.008
LA-S-015	19	0.36	366	<10	13	72	0.006
LA-S-016	39	0.29	400	<10	25	82	0.014
LA-S-017	26	0.28	416	<10	25	84	0.012
LA-S-018	22	0.39	623	<10	24	101	0.011
LA-S-019	24	0.22	305	<10	55	85	0.011
LA-S-020	23	0.37	476	<10	24	82	0.013
LA-S-021	40	0.41	487	<10	27	103	0.011
LA-S-022	29	0.69	830	<10	19	138	0.009
LA-S-023	24	0.45	603	<10	27	114	0.012
LA-S-024	18	0.50	730	<10	24	118	0.010
LA-S-025	11	0.28	343	<10	20	79	0.005
LA-S-026	18	0.21	335	<10	23	75	0.012
LA-S-027	22	0.22	431	<10	36	80	0.024
LA-S-028	42	0.28	374	<10	36	73	0.013
LA-S-029	20	0.38	350	<10	16	68	0.004
LA-S-030	19	0.38	478	<10	23	84	0.008
LA-S-031	12	0.44	590	<10	12	80	0.006
LA-S-032	15	0.81	696	<10	17	101	0.010
LA-S-033	11	0.66	952	<10	12	114	0.005
LA-S-034	48	0.44	499	<10	24	97	0.009
LA-S-035	72	0.37	465	<10	33	104	0.011
LA-S-036	43	0.37	426	<10	27	100	0.010
LA-S-037	16	0.35	365	<10	15	73	0.008
LA-S-038	29	0.28	418	<10	29	88	0.014
LA-S-039	44	0.38	513	<10	43	106	0.012
LA-S-040	29	1.08	859	<10	17	130	0.007
LA-S-041	30	0.88	791	<10	24	138	0.009
LA-S-042	25	0.36	492	<10	27	99	0.014
LA-S-043	32	0.44	532	<10	28	109	0.012
LA-S-044	26	0.30	350	<10	24	98	0.015
LA-S-045	17	0.48	765	<10	29	149	0.006
LA-S-046	30	0.39	486	<10	30	112	0.012
LA-S-047	16	0.66	626	<10	15	128	0.009
LA-S-048	19	0.68	703	<10	21	143	0.015
LA-S-049	25	0.60	754	<10	21	135	0.014
LA-S-050	19	0.91	805	<10	20	155	0.011
LA-S-051	15	0.66	639	<10	22	126	0.013
LA-S-052	31	0.40	571	<10	35	124	0.013
LA-S-053	20	0.49	680	15	23	118	0.009
LA-S-054	17	0.24	547	<10	30	132	0.013
LA-S-055	13	0.57	855	<10	19	146	0.013
LA-S-056	21	0.57	728	<10	25	139	0.014
LA-S-057	25	0.81	947	<10	17	158	0.009
LA-S-058	38	0.61	672	<10	24	141	0.011
LA-S-059	31	0.51	669	<10	15	111	0.012
LA-S-060	40	0.48	755	<10	19	132	0.012
LA-S-061	48	0.46	530	<10	28	124	0.013
LA-S-062	24	0.53	657	<10	27	145	0.012
LA-S-063	17	0.93	934	<10	21	155	0.010
LA-S-064	26	0.34	510	<10	32	106	0.013
LA-S-065	36	0.29	398	<10	25	102	0.029
LA-S-066	24	0.35	678	<10	20	105	0.014

Geochemical grade assay for stream sediments collected in the São Gabriel district

(3/7)

Sample No.	UTM_E	UTM_N	Latitude	Longitude	Location	Ag	Cd	Cu	Mn	Mo	Ni	Pb	Zn	Al	As	Ba	Be	Bi	Ca	Co	Cr	Fe	K	Mg	Na	P	Sb	Sc	Sn
LA-S-067	711278	6933431	-27.70678	-54.85724	São Gabriel	<0.2	<0.5	203	1686	<2	42	19	108	2.32	<10	279	2	17	0.49	53	56	10.39	0.04	0.25	0.04	0.040	<10	27	<10
LA-S-068	705163	6921590	-27.81455	-54.91719	São Gabriel	<0.2	<0.5	219	2737	<2	51	18	120	2.36	<10	416	2	<10	0.53	77	67	12.64	0.05	0.29	0.04	0.037	<10	25	<10
LA-S-069	698840	6911486	-27.90666	-54.97966	São Gabriel	<0.2	<0.5	237	4116	<2	78	20	97	4.15	17	745	3	<10	0.51	127	117	12.65	0.07	0.21	0.03	0.055	12	38	<10
LA-S-070	760478	6924632	-27.77751	-54.35675	São Gabriel	<0.2	<0.5	247	1460	<2	57	29	120	3.45	<10	226	2	<10	0.25	49	85	12.95	0.04	0.09	0.02	0.032	<10	38	<10
LA-S-071	759730	6924498	-27.77886	-54.36429	São Gabriel	<0.2	<0.5	255	1705	<2	61	22	147	3.16	<10	296	2	<10	0.32	62	87	14.07	0.04	0.15	0.03	0.032	<10	36	<10
LA-S-072	760607	6913856	-27.87466	-54.35307	São Gabriel	<0.2	<0.5	300	1684	<2	54	23	139	3.31	<10	272	2	<10	0.28	57	75	13.96	0.05	0.13	0.03	0.029	<10	34	<10
LA-S-073	766473	6912307	-27.88747	-54.29319	São Gabriel	<0.2	<0.5	367	1726	<2	75	21	210	1.55	<10	160	3	33	0.16	63	83	18.49	0.02	0.03	0.02	0.014	<10	24	<10
LA-S-074	781780	6918283	-27.89045	-54.13928	São Gabriel	<0.2	<0.5	215	1594	<2	57	14	139	2.88	<10	307	2	11	0.40	50	81	10.93	0.04	0.23	0.04	0.029	<10	33	<10
LA-S-075	781549	6929161	-27.79242	-54.14419	São Gabriel	<0.2	<0.5	254	1751	<2	65	11	140	3.14	<10	332	2	<10	0.48	57	77	12.78	0.05	0.28	0.04	0.040	<10	34	<10
LA-S-076	781636	6939138	-27.64244	-54.14565	São Gabriel	<0.2	<0.5	221	1848	<2	56	19	133	2.70	<10	340	2	11	0.41	56	83	11.45	0.05	0.22	0.03	0.039	<10	31	<10
LA-S-077	776860	6943227	-27.60656	-54.19494	São Gabriel	<0.2	<0.5	323	1883	<2	76	24	211	1.97	<10	266	2	21	0.35	61	101	14.42	0.03	0.19	0.03	0.030	<10	24	<10
LA-S-078	786963	6944986	-27.58860	-54.09311	São Gabriel	<0.2	<0.5	346	1767	<2	66	14	181	2.21	<10	234	2	30	0.28	55	68	15.94	0.04	0.12	0.03	0.028	<10	29	<10
LA-S-079	773423	6952723	-27.52163	-54.23188	São Gabriel	<0.2	<0.5	306	1963	<2	56	29	142	3.13	<10	273	3	21	0.29	70	125	19.20	0.03	0.12	0.05	0.022	<10	21	<10
LA-S-080	757496	6943709	-27.60604	-54.39106	São Gabriel	<0.2	<0.5	306	1963	<2	56	29	142	3.13	<10	273	3	21	0.29	70	125	19.20	0.03	0.12	0.05	0.022	<10	21	<10
LA-S-081	729619	6940494	-27.84006	-54.67270	São Gabriel	<0.2	<0.5	231	2130	<2	53	22	131	3.29	<10	342	2	13	0.45	64	69	12.18	0.05	0.20	0.05	0.044	<10	36	<10
LA-S-082	786279	6976058	-27.30858	-54.10735	São Gabriel	<0.2	<0.5	305	1892	<2	65	24	163	3.65	<10	367	2	27	0.47	67	80	13.48	0.06	0.26	0.05	0.042	<10	29	<10
LA-S-083	780134	6966659	-27.39461	-54.16722	São Gabriel	<0.2	<0.5	289	1852	<2	60	10	147	3.65	<10	331	2	27	0.47	57	80	13.48	0.06	0.26	0.05	0.044	<10	36	<10
LA-S-084	779126	6963307	-27.42504	-54.17664	São Gabriel	<0.2	<0.5	267	2170	<2	67	11	194	2.84	<10	304	2	18	0.43	68	93	14.28	0.05	0.25	0.04	0.044	<10	30	<10
LA-S-085	764680	6956491	-27.48933	-54.31811	São Gabriel	<0.2	<0.5	293	1656	<2	67	16	145	1.71	<10	174	3	20	0.21	59	88	16.40	0.03	0.12	0.03	0.019	<10	21	<10
LA-S-086	752217	6957031	-27.48689	-54.44729	São Gabriel	<0.2	<0.5	342	1559	<2	78	12	137	2.60	<10	206	2	17	0.49	57	93	16.30	0.03	0.09	0.03	0.026	<10	31	<10
LA-S-087	749799	6954311	-27.51187	-54.47118	São Gabriel	<0.2	<0.5	342	1559	<2	78	12	137	2.60	<10	206	2	17	0.49	57	93	16.30	0.03	0.09	0.03	0.026	<10	31	<10
LA-S-088	216075	6764591	-29.21552	-53.92045	São Gabriel	<0.2	<0.5	8	808	<2	4	20	32	0.74	<10	223	2	<10	0.10	7	4	2.48	0.02	0.04	0.03	0.036	<10	5	<10
LA-S-089	213641	6776632	-29.10644	-53.94236	São Gabriel	<0.2	<0.5	10	145	<2	5	12	12	0.39	<10	53	<1	<10	0.05	3	5	0.78	0.01	0.02	0.03	0.006	<10	2	<10
LA-S-090	231349	6782942	-29.05432	-53.75909	São Gabriel	<0.2	<0.5	7	321	<2	4	10	17	0.63	<10	97	<1	<10	0.06	5	6	0.92	0.02	0.04	0.03	0.011	<10	3	<10
LA-S-091	255640	6785915	-29.03151	-53.50916	São Gabriel	<0.2	<0.5	28	372	<2	9	12	38	0.98	<10	86	<1	<10	0.06	10	10	2.77	0.02	0.05	0.03	0.012	<10	6	<10
LA-S-092	275942	6782264	-29.01264	-53.30033	São Gabriel	<0.2	<0.5	25	310	<2	7	10	32	0.86	<10	65	<1	<10	0.04	8	8	2.34	0.01	0.03	0.02	0.010	<10	5	<10
LA-S-093	242100	6776568	-29.11311	-53.65023	São Gabriel	<0.2	<0.5	17	437	<2	6	13	36	0.72	<10	135	2	<10	0.12	9	5	1.94	0.02	0.07	0.03	0.021	<10	5	<10
LA-S-094	215895	6755411	-29.29823	-53.92465	São Gabriel	<0.2	<0.5	9	817	<2	4	24	32	0.76	<10	241	2	<10	0.18	5	5	2.22	0.03	0.06	0.02	0.024	<10	5	<10
LA-S-095	221698	6746389	-29.38085	-53.86725	São Gabriel	<0.2	<0.5	13	544	<2	6	25	34	0.81	<10	213	2	<10	0.17	5	5	2.22	0.03	0.06	0.02	0.024	<10	5	<10
LA-S-096	244393	6753935	-29.31853	-53.63233	São Gabriel	<0.2	<0.5	8	229	<2	4	7	17	0.36	<10	42	<1	<10	0.06	5	3	0.96	0.01	0.03	0.02	0.011	<10	3	<10
LA-S-097	260324	6762303	-29.24532	-53.46618	São Gabriel	<0.2	<0.5	5	91	<2	3	5	8	0.25	<10	40	<1	<10	0.04	2	2	0.42	<0.01	0.02	0.02	0.006	<10	1	<10
LA-S-098	277190	6759866	-29.28129	-53.29349	São Gabriel	<0.2	<0.5	25	471	<2	6	11	42	0.50	<10	96	<1	<10	0.08	10	4	3.70	0.02	0.05	0.02	0.010	<10	5	<10
LA-S-099	275095	6739850	-29.45051	-53.31888	São Gabriel	<0.2	<0.5	43	984	<2	9	22	88	1.54	<10	265	2	<10	0.32	17	7	5.81	0.07	0.18	0.03	0.029	<10	11	<10
LA-S-100	251473	6735511	-29.48517	-53.56322	São Gabriel	<0.2	<0.5	49	702	<2	13	14	56	1.08	<10	213	2	<10	0.35	16	9	4.09	0.05	0.18	0.04	0.024	<10	8	<10
LA-S-101	222594	6725312	-29.57104	-53.86337	São Gabriel	<0.2	<0.5	43	935	<2	13	14	56	1.08	<10	213	2	<10	0.48	17	11	4.37	0.05	0.22	0.06	0.032	<10	12	<10
LA-S-102	242605	6710415	-29.70963	-53.66054	São Gabriel	<0.2	<0.5	18	637	<2	8	13	39	0.83	<10	155	1	<10	0.23	10	4	2.13	0.05	0.10	0.02	0.026	<10	6	<10
LA-S-103	267330	6709300	-29.72457	-53.40541	São Gabriel	<0.2	<0.5	18	637	<2	8	13	39	0.83	<10	155	1	<10	0.11	3	3	0.95	0.04	0.06	0.02	0.010	<10	2	<10
LA-S-104	292103	6698853	-29.81163	-53.77186	São Gabriel	<0.2	<0.5	4	66	<2	4	4	9	0.29	<10	50	<1	<10	0.07	2	3	0.31	0.02	0.03	0.02	0.006	<10	<1	<10
LA-S-105	233879	6684973	-29.93713	-53.75694	São Gabriel	<0.2	<0.5	4	151	<2	3	5	8	0.16	<10	51	<1	<10	0.05	2	3	0.26	0.01	0.02	0.02	0.006	<10	<1	<10
LA-S-106	217670	6678987	-29.98746	-53.92627	São Gabriel	<0.2	<0.5	8	58	<2	5	12	13	0.19	<10	49	<1	<10	0.07	<1	4	0.21	0.02	0.02	0.02	0.007	<10	<1	<10
LA-S-107	259186	6725412	-29.57774	-53.48596	São Gabriel	<0.2	<0.5	16	225	<2	6	9	26	0.41	<10	80	<1	<10	0.13	5	3	1.60	0.04	0.07	0.02	0.011	<10	3	<10
LA-S-108	278621	6727272	-29.56457	-53.26509	São Gabriel	<0.2	<0.5	75	738	<2	19	12	90	0.90	<10	137	1	<10	0.29	22	14	5.66	0.05	0.17	0.03	0.020	<10	8	<10
LA-S-109	290746	6708611	-29.73496	-53.16361	São Gabriel	<0.2	<0.5	15	229	<2	6	6	23	0.37	<10	63	<1	<10	0.07	6	5	1.48	0.04	0.06	0.02	0.007	<10	3	<10
LA-S-110	301036	6707126	-29.75005	-53.05755	São Gabriel	<0.2	<0.5	14	224	<2	7	6	23	0.37	<10	63	<1	<10	0.19	6	5	1.24	0.05	0.09	0.02	0.008	<10	3	<10
LA-S-111	311020	6713632	-29.69113	-52.95318	São Gabriel	<0.2	<0.5	26	295	<2	12	15																	

Geochemical grade assay for stream sediments collected in the São Gabriel district

Sample No.	Sr ppm	Ti %	V ppm	W ppm	Y ppm	Zr ppm	S %
LA-S-067	31	0.40	604	<10	25	122	0.014
LA-S-068	42	0.51	611	<10	26	132	0.012
LA-S-069	42	0.37	596	<10	32	136	0.023
LA-S-070	17	0.58	756	<10	23	147	0.011
LA-S-071	23	0.65	871	<10	25	156	0.013
LA-S-072	19	0.69	908	<10	21	160	0.009
LA-S-073	12	1.03	1624	<10	15	187	0.009
LA-S-074	30	0.59	643	<10	30	126	0.013
LA-S-075	34	0.59	775	<10	25	146	0.015
LA-S-076	30	0.54	704	<10	26	124	0.014
LA-S-077	25	0.89	1208	<10	18	160	0.009
LA-S-078	20	0.82	1194	<10	19	169	0.011
LA-S-079	21	1.15	1660	<10	15	184	0.010
LA-S-080	22	0.61	869	<10	26	149	0.012
LA-S-081	28	0.44	636	<10	30	117	0.016
LA-S-082	38	0.58	823	<10	19	133	0.017
LA-S-083	33	0.54	701	<10	25	137	0.015
LA-S-084	28	0.51	860	<10	18	97	0.019
LA-S-085	16	0.73	1216	<10	12	100	0.005
LA-S-086	17	0.64	968	<10	14	95	0.010
LA-S-087	34	0.59	968	<10	16	92	0.006
LA-S-088	17	0.07	34	<10	16	46	0.010
LA-S-089	5	0.06	27	<10	4	18	0.008
LA-S-090	9	0.05	24	<10	7	22	0.006
LA-S-091	9	0.21	114	<10	10	52	0.007
LA-S-092	6	0.17	94	<10	7	43	0.006
LA-S-093	18	0.08	59	<10	10	40	0.007
LA-S-094	27	0.06	95	<10	25	74	0.014
LA-S-095	25	0.06	30	<10	20	72	0.020
LA-S-096	9	0.05	28	<10	7	22	0.005
LA-S-097	5	0.03	13	<10	3	10	0.004
LA-S-098	13	0.24	142	<10	8	77	0.003
LA-S-099	49	0.40	184	<10	20	149	0.011
LA-S-100	62	0.30	200	<10	15	91	0.014
LA-S-101	54	0.17	162	<10	23	78	0.015
LA-S-102	26	0.09	81	<10	17	45	0.013
LA-S-103	11	0.04	31	<10	6	24	0.005
LA-S-104	5	0.01	10	<10	3	4	0.005
LA-S-105	5	<0.01	7	<10	3	4	0.009
LA-S-106	6	<0.01	6	<10	3	2	0.009
LA-S-107	19	0.09	68	<10	8	52	0.006
LA-S-108	32	0.50	434	<10	10	125	0.004
LA-S-109	9	0.13	84	<10	6	41	0.003
LA-S-110	14	0.08	63	<10	6	29	0.004
LA-S-111	17	0.10	79	<10	6	36	0.010
LA-S-112	33	0.24	206	<10	9	85	0.003
LA-S-113	36	0.12	133	<10	10	46	0.008
LA-S-114	31	0.22	238	<10	6	70	0.002
LA-S-115	29	0.31	309	<10	7	83	0.008
LA-S-116	73	0.22	196	<10	26	195	0.021
LA-S-117	62	0.38	389	<10	23	144	0.012
LA-S-118	63	0.37	429	<10	19	138	0.012
LA-S-119	28	0.36	370	<10	16	145	0.011
LA-S-120	28	0.46	481	<10	15	173	0.004

Geochemical grade assay for stream sediments collected in the São Gabriel district

Sample No.	UTM_E	UTM_N	Latitude	Longitude	Location	Ag	Cd	Cu	Mn	Mo	Ni	Pb	Zn	Al	As	Ba	Be	Bi	Ca	Co	Cr	Fe	K	Mg	Na	P	Sb	Sc	Sn
						ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	%	ppm	ppm	ppm	ppm	ppm	%	ppm	ppm	%	%	%	%	ppm	ppm	ppm
LA-S-121	340280	6765127	-29.23248	-52.64340	São Gabriel	<0.2	<0.5	32	1122	3	8	33	85	2.14	<10	384	3	<10	0.32	18	8	5.35	0.10	0.25	0.02	0.041	<10	12	<10
LA-S-122	341012	6762249	-29.25854	-52.63629	São Gabriel	<0.2	<0.5	32	993	<2	10	33	72	1.83	<10	316	3	<10	0.36	12	9	4.02	0.08	0.25	0.02	0.034	<10	11	<10
LA-S-123	338737	6773400	-29.15765	-52.65807	São Gabriel	<0.2	<0.5	27	947	3	7	32	92	1.80	<10	533	3	<10	0.31	12	8	4.98	0.08	0.19	0.02	0.032	<10	13	<10
LA-S-124	338291	6773566	-29.15609	-52.66263	São Gabriel	<0.2	<0.5	20	965	3	6	35	81	1.87	<10	980	3	<10	0.27	10	9	4.51	0.06	0.16	0.02	0.028	<10	11	<10
LA-S-125	339716	6785653	-29.04723	-52.64625	São Gabriel	<0.2	<0.5	16	927	<2	6	31	59	1.57	<10	260	3	<10	0.23	7	5	3.15	0.04	0.10	0.02	0.031	<10	9	<10
LA-S-126	297150	6764921	-29.26639	-53.08988	São Gabriel	<0.2	<0.5	38	952	2	6	21	76	1.21	<10	220	3	<10	0.25	17	3	5.12	0.04	0.14	0.02	0.034	<10	10	<10
LA-S-127	298020	6760697	-29.26639	-53.07882	São Gabriel	<0.2	<0.5	51	1144	3	8	34	83	1.49	<10	270	3	<10	0.30	20	4	5.79	0.05	0.16	0.02	0.036	<10	12	<10
LA-S-128	298625	6764784	-29.22978	-53.06157	São Gabriel	<0.2	<0.5	59	1383	3	9	29	97	1.74	<10	407	7	<10	0.41	26	7	6.25	0.08	0.21	0.02	0.036	<10	15	<10
LA-S-129	310790	6774069	-29.14775	-52.94518	São Gabriel	<0.2	<0.5	32	1115	<2	8	27	91	1.58	<10	322	2	<10	0.20	19	11	5.69	0.06	0.12	0.02	0.022	<10	12	<10
LA-S-130	308072	6779075	-29.10219	-52.97224	São Gabriel	<0.2	<0.5	46	894	<2	9	23	77	1.74	<10	279	3	<10	0.19	15	13	5.52	0.05	0.10	0.02	0.026	<10	13	<10
LA-S-131	302174	6766242	-29.03663	-53.03155	São Gabriel	<0.2	<0.5	26	994	2	6	29	58	1.68	<10	344	3	<10	0.22	11	8	3.89	0.06	0.12	0.02	0.024	<10	11	<10
LA-S-132	311509	6815761	-28.77176	-52.93080	São Gabriel	<0.2	<0.5	63	860	2	16	27	72	2.25	<10	241	2	<10	0.16	21	23	6.36	0.04	0.08	0.02	0.028	<10	17	<10
LA-S-133	312680	6816568	-28.76465	-52.91867	São Gabriel	<0.2	<0.5	89	865	3	21	27	86	1.67	<10	143	2	<10	0.14	23	25	6.51	0.03	0.06	0.02	0.023	<10	14	<10
LA-S-134	319789	6820299	-28.73200	-52.84550	São Gabriel	<0.2	<0.5	143	896	2	28	14	93	1.16	<10	86	1	<10	0.09	30	35	7.97	0.02	0.05	0.02	0.015	<10	14	<10
LA-S-135	324751	6821593	-28.72101	-52.79451	São Gabriel	<0.2	<0.5	104	896	2	27	18	79	2.08	<10	163	2	<10	0.15	26	43	7.32	0.03	0.09	0.02	0.027	<10	20	<10
LA-S-136	335275	6825509	-28.68707	-52.68603	São Gabriel	<0.2	<0.5	67	870	2	14	21	78	1.69	<10	220	2	<10	0.19	18	19	5.84	0.04	0.10	0.02	0.025	<10	15	<10
LA-S-137	342208	6831244	-28.63619	-52.61430	São Gabriel	<0.2	<0.5	22	769	2	7	37	65	2.41	<10	294	3	<10	0.21	10	8	4.93	0.05	0.09	0.02	0.027	<10	14	<10
LA-S-138	340868	6843431	-28.52607	-52.62631	São Gabriel	<0.2	<0.5	121	1869	4	35	25	63	2.62	<10	207	1	<10	0.20	42	64	8.27	0.05	0.09	0.03	0.039	<10	28	<10
LA-S-139	337288	6851779	-28.45030	-52.66171	São Gabriel	<0.2	<0.5	185	1952	5	56	18	89	3.65	<10	906	2	<10	0.26	66	109	12.07	0.04	0.12	0.03	0.043	<10	41	<10
LA-S-140	333102	6853877	-28.43084	-52.70414	São Gabriel	<0.2	<0.5	176	1022	3	45	12	91	2.07	<10	123	2	<10	0.13	36	126	10.62	0.02	0.08	0.03	0.030	<10	26	<10
LA-S-141	315312	6845861	-28.50077	-52.86699	São Gabriel	<0.2	<0.5	182	1117	4	45	18	82	3.46	<10	133	2	<10	0.19	38	66	11.66	0.03	0.10	0.04	0.028	<10	37	<10
LA-S-142	312190	6845399	-28.50449	-52.91895	São Gabriel	<0.2	<0.5	157	1720	3	46	19	86	3.55	<10	190	2	<10	0.30	55	59	11.36	0.04	0.15	0.04	0.032	<10	37	<10
LA-S-143	309529	6846814	-28.49043	-52.94588	São Gabriel	<0.2	<0.5	187	1369	3	43	14	96	2.88	<10	160	2	<10	0.23	47	58	12.00	0.03	0.12	0.04	0.028	<10	33	<10
LA-S-144	298777	6843430	-28.52024	-53.05627	São Gabriel	<0.2	<0.5	164	1339	4	35	23	95	2.78	<10	127	2	<10	0.19	34	47	11.70	0.03	0.10	0.04	0.038	<10	32	<10
LA-S-145	295689	6829322	-28.61996	-53.08979	São Gabriel	<0.2	<0.5	145	1170	5	29	17	110	1.79	<10	127	2	<10	0.11	30	34	10.85	0.02	0.06	0.03	0.026	<10	21	<10
LA-S-146	296029	6825033	-28.72628	-55.15397	São Gabriel	<0.2	<0.5	98	2279	3	20	21	60	2.12	<10	402	2	<10	0.49	59	11	7.86	0.06	0.22	0.07	0.017	<10	24	<10
LA-S-147	681893	6821986	-28.72196	-55.13766	São Gabriel	<0.2	<0.5	97	1333	4	21	21	53	2.17	<10	332	2	<10	0.49	46	20	7.74	0.06	0.27	0.04	0.016	<10	23	<10
LA-S-148	684440	6821704	-28.71873	-55.11165	São Gabriel	<0.2	<0.5	81	1542	4	22	22	61	1.43	<10	294	1	<10	0.46	32	13	5.32	0.04	0.19	0.06	0.015	<10	20	<10
LA-S-149	684561	6821666	-28.71906	-55.11040	São Gabriel	<0.2	<0.5	90	1476	3	20	17	67	1.81	<10	319	1	<10	0.54	41	14	6.54	0.04	0.24	0.05	0.015	<10	20	<10
LA-S-150	691568	6820433	-28.72913	-55.03849	São Gabriel	<0.2	<0.5	111	2271	3	23	20	75	2.69	<10	506	3	<10	0.60	53	14	8.58	0.07	0.26	0.05	0.023	<10	26	<10
LA-S-151	692468	6820412	-28.72922	-55.02927	São Gabriel	<0.2	<0.5	97	2489	4	23	21	65	2.71	<10	483	3	<10	0.48	53	16	9.04	0.05	0.18	0.05	0.031	<10	27	<10
LA-S-152	696261	6817332	-28.75643	-54.98992	São Gabriel	<0.2	<0.5	110	1749	3	28	16	73	2.85	<10	347	3	<10	0.46	50	16	8.77	0.06	0.21	0.04	0.029	<10	32	<10
LA-S-153	699064	6815897	-28.77082	-54.95275	São Gabriel	<0.2	<0.5	110	1347	4	19	23	69	2.64	<10	267	3	<10	0.31	38	17	9.64	0.04	0.11	0.03	0.030	<10	25	<10
LA-S-154	702134	6815047	-28.77613	-54.92940	São Gabriel	<0.2	<0.5	125	1428	3	21	32	82	2.19	<10	227	3	<10	0.31	38	17	9.64	0.04	0.11	0.03	0.030	<10	25	<10
LA-S-155	707186	6815930	-28.76737	-54.87804	São Gabriel	<0.2	<0.5	121	1722	3	22	22	73	2.37	<10	255	2	<10	0.32	45	26	10.10	0.03	0.12	0.03	0.031	<10	27	<10
LA-S-156	707904	6811035	-28.81140	-54.86959	São Gabriel	<0.2	<0.5	150	1403	4	21	19	83	2.07	<10	293	3	<10	0.24	43	19	10.11	0.03	0.10	0.03	0.023	<10	28	<10
LA-S-157	685577	6813177	-28.79550	-55.09862	São Gabriel	<0.2	<0.5	83	2948	3	16	20	62	2.12	<10	371	3	<10	0.43	47	15	5.96	0.03	0.15	0.03	0.023	<10	25	<10
LA-S-158	692002	6808479	-28.83693	-55.03202	São Gabriel	<0.2	<0.5	75	1219	<2	14	12	59	1.46	<10	295	2	<10	0.54	30	8	5.97	0.05	0.23	0.03	0.025	<10	18	<10
LA-S-159	698749	6801997	-28.89438	-54.96176	São Gabriel	<0.2	<0.5	81	1132	3	14	17	62	1.77	<10	342	3	<10	0.48	23	9	5.41	0.05	0.18	0.03	0.015	<10	20	<10
LA-S-160	701745	6797919	-28.93069	-54.93032	São Gabriel	<0.2	<0.5	27	1191	2	10	42	58	1.48	<10	607	4	<10	0.31	13	10	3.96	0.09	0.10	0.03	0.028	<10	11	<10
LA-S-161	703028	6796237	-28.94566	-54.91686	São Gabriel	<0.2	<0.5	15	656	<2	5	25	46	1.26	<10	391	3	<10	0.19	7	5	3.13	0.07	0.09	0.03	0.021	<10	8	<10
LA-S-162	707300	6801928	-28.89384	-54.87410	São Gabriel	<0.2	<0.5	118	1507	3	17	12	73	1.70	<10	330	2	<10	0.54	41	12	7.26	0.06	0.25	0.06	0.023	<10	21	<10
LA-S-163	705800	6798750	-28.92255	-54.88690	São Gabriel	<0.2	<0.5	83	1764	4	18	20	60	2.02	<10	378	2	<10	0.28	46	16	5.94	0.05	0.17	0.04	0.021	<10	22	<10
LA-S-164	711182	6795126	-28.95356	-54.77151	São Gabriel	<0.2	<0.5	66	803	<2	11	16	45	1.09	<10	279	2	<10	0.24	16	13								

Geochemical grade assay for stream sediments collected in the São Gabriel district

(4continued/7)

Sample No.	Sr ppm	Ti %	V ppm	W ppm	Y ppm	Zr ppm	S %
LA-S-121	51	0.30	107	<10	26	198	0.010
LA-S-122	41	0.25	78	<10	28	157	0.022
LA-S-123	48	0.37	105	<10	32	222	0.015
LA-S-124	35	0.28	67	<10	26	248	0.016
LA-S-125	32	0.12	43	<10	27	137	0.023
LA-S-126	24	0.25	139	<10	22	159	0.008
LA-S-127	32	0.28	150	<10	25	176	0.018
LA-S-128	53	0.41	229	<10	27	224	0.017
LA-S-129	34	0.45	155	<10	24	236	0.014
LA-S-130	29	0.35	181	<10	23	175	0.011
LA-S-131	39	0.24	81	<10	33	197	0.013
LA-S-132	25	0.36	241	<10	26	168	0.013
LA-S-133	14	0.51	308	<10	13	124	0.013
LA-S-134	7	0.87	554	<10	8	114	0.006
LA-S-135	16	0.48	353	<10	19	153	0.009
LA-S-136	25	0.42	212	<10	24	201	0.013
LA-S-137	32	0.24	57	<10	34	319	0.014
LA-S-138	18	0.31	333	<10	19	96	0.021
LA-S-139	23	0.47	535	<10	21	120	0.016
LA-S-140	13	0.56	657	<10	11	100	0.010
LA-S-141	17	0.52	575	<10	18	123	0.008
LA-S-142	33	0.46	510	<10	21	105	0.012
LA-S-143	24	0.70	616	<10	17	120	0.009
LA-S-144	22	0.49	543	<10	19	122	0.013
LA-S-145	13	0.58	622	<10	12	106	0.007
LA-S-146	61	0.29	378	<10	44	100	0.011
LA-S-147	64	0.22	407	<10	40	114	0.005
LA-S-148	68	0.29	297	<10	24	97	0.013
LA-S-149	77	0.34	343	<10	30	109	0.007
LA-S-150	104	0.24	365	<10	50	101	0.009
LA-S-151	70	0.19	370	<10	49	88	0.016
LA-S-152	59	0.24	404	<10	48	96	0.019
LA-S-153	36	0.27	500	<10	37	103	0.016
LA-S-154	34	0.33	510	<10	33	101	0.018
LA-S-155	32	0.31	493	<10	32	104	0.010
LA-S-156	23	0.46	603	<10	34	127	0.009
LA-S-157	47	0.11	266	<10	59	46	0.017
LA-S-158	86	0.21	318	<10	36	97	0.012
LA-S-159	72	0.22	251	<10	61	155	0.008
LA-S-160	123	0.24	70	<10	46	188	0.030
LA-S-161	69	0.20	54	<10	28	148	0.018
LA-S-162	58	0.30	400	<10	32	115	0.007
LA-S-163	66	0.26	260	<10	47	119	0.013
LA-S-164	46	0.26	166	<10	22	85	0.009
LA-S-165	67	0.28	238	<10	28	124	0.017
LA-S-166	9	0.21	201	<10	5	47	0.002
LA-S-167	48	0.20	269	<10	30	94	0.018
LA-S-168	26	0.43	532	<10	12	109	0.009
LA-S-169	30	0.43	681	<10	14	114	0.009
LA-S-170	66	0.73	664	<10	25	177	0.015
LA-S-171	49	0.75	742	<10	23	145	0.017
LA-S-172	100	0.63	587	<10	23	184	0.021
LA-S-173	43	1.02	1020	<10	17	171	0.016
LA-S-174	52	0.52	555	<10	26	147	0.008

Geochemical grade assay for stream sediments collected in the São Gabriel district

Table with columns: Sample No., UTM_E, UTM_N, Latitude, Longitude, Location, Ag, Cd, Cu, Mn, Mo, Ni, Pb, Zn, Al, As, Ba, Be, Bi, Ca, Co, Cr, Fe, K, Mg, Na, P, Sb, Sc, Sn. The table contains 100 rows of geochemical data for various sediment samples.

Geochemical grade assay for stream sediments collected in the São Gabriel district

Sample No.	Sr	Ti	V	W	Y	Zr	S
	ppm	%	ppm	ppm	ppm	ppm	%
LA-S-175	51	0.41	352	<10	28	131	0.017
LA-S-176	57	0.51	499	<10	20	147	0.015
OC-S-001	4	0.04	10	<10	3	5	0.004
OC-S-002	3	0.03	23	<10	4	6	0.003
OC-S-003	19	0.32	355	<10	10	32	0.006
OC-S-004	44	0.39	450	<10	25	49	0.010
OC-S-005	3	0.06	19	<10	3	9	0.002
OC-S-006	12	0.32	342	<10	10	34	0.005
OC-S-007	26	0.20	201	<10	10	34	0.005
OC-S-008	6	0.07	44	<10	5	18	0.003
OC-S-009	12	0.13	116	<10	7	25	0.003
OC-S-010	8	0.12	175	<10	7	20	0.004
OC-S-011	38	0.25	203	<10	15	60	0.005
OC-S-012	10	0.02	33	<10	8	6	0.005
OC-S-013	25	0.19	169	<10	14	42	0.006
OC-S-014	11	0.10	72	<10	6	23	0.003
OC-S-015	123	0.49	489	<10	33	105	0.011
OC-S-016	114	0.26	56	<10	37	133	0.016
OC-S-017	69	0.64	447	<10	37	116	0.014
OC-S-018	99	0.18	53	<10	32	119	0.011
OC-S-019	52	0.10	40	<10	37	58	0.013
OC-S-020	68	0.21	293	<10	25	70	0.012
OC-S-021	63	0.16	81	<10	20	67	0.007
OC-S-022	15	0.09	41	<10	5	29	0.003
OC-S-023	7	0.08	22	<10	3	18	0.003
OC-S-024	7	0.20	54	<10	4	21	0.003
OC-S-025	23	0.11	55	<10	6	32	0.003
OC-S-026	9	0.32	303	<10	7	49	0.003
OC-S-027	43	0.55	532	<10	23	99	0.009
OC-S-028	30	0.26	232	<10	19	63	0.006
OC-S-029	40	0.61	739	<10	27	88	0.009
OC-S-030	44	0.10	41	<10	38	62	0.013
OC-S-031	57	0.24	318	<10	27	76	0.009
OC-S-032	41	0.20	184	<10	18	57	0.007
OC-S-033	2	0.04	19	<10	4	4	0.005
OC-S-034	31	0.15	186	<10	12	40	0.004
OC-S-035	<1	0.03	9	<10	1	2	0.002
OC-S-036	10	0.11	97	<10	6	22	0.003
OC-S-037	140	0.25	122	<10	32	105	0.020
OC-S-038	3	0.08	47	<10	4	11	0.003
OC-S-039	7	0.25	241	<10	4	29	0.003
OC-S-040	16	0.08	53	<10	7	20	0.004
OC-S-041	33	0.26	779	<10	26	71	0.006
OC-S-042	34	0.12	234	<10	28	41	0.009
OC-S-043	58	0.37	897	<10	42	91	0.004
OC-S-044	58	0.40	424	<10	34	70	0.008
OC-S-045	42	0.23	670	<10	38	78	0.009
OC-S-046	82	0.23	174	<10	47	83	0.012
OC-S-047	24	0.22	89	<10	20	81	0.012
OC-S-048	4	0.03	24	<10	3	13	0.003
OC-S-049	48	0.42	483	<10	29	68	0.005
OC-S-050	101	0.25	649	<10	39	51	0.005
OC-S-051	86	0.27	438	<10	38	85	0.006
OC-S-052	50	0.36	433	<10	34	110	0.008

Geochemical grade assay for stream sediments collected in the São Gabriel district

Sample No.	UTM_E	UTM_N	Latitude	Longitude	Location	Ag	Cd	Cu	Mn	Mo	Ni	Pb	Zn	Al	As	Ba	Be	Bi	Ca	Co	Cr	Fe	K	Mg	Na	P	Sb	Sc	Sn
OC-S-053	667712	6772539	-29.16460	-55.27550	São Gabriel	-0.2	<0.5	6	694	<2	<1	3	5	0.12	<10	51	<1	<10	0.02	11	2	0.54	<0.01	<0.01	0.004	<10	<1	<10	
OC-S-054	653950	6771219	-29.17820	-55.41680	São Gabriel	-0.2	<0.5	5	40	<2	<1	2	3	0.03	<10	10	<1	<10	<0.01	1	1	0.36	<0.01	<0.01	0.002	<10	<1	<10	
OC-S-055	651211	6776603	-29.22530	-55.44420	São Gabriel	-0.2	<0.5	14	250	<2	<1	4	15	0.45	<10	40	<1	<10	0.07	8	9	1.46	0.02	0.06	0.003	<10	3	<10	
OC-S-056	646632	6776400	-29.13240	-55.49270	São Gabriel	-0.2	<0.5	4	98	<2	<1	4	4	0.18	<10	16	<1	<10	0.01	1	3	0.23	<0.01	0.01	0.002	<10	<1	<10	
OC-S-057	650688	6809979	-28.82890	-55.45550	São Gabriel	-0.2	<0.5	83	1344	<2	21	14	63	2.43	<10	240	<1	<10	0.90	30	37	5.78	0.04	0.44	0.06	0.016	<10	20	<10
OC-S-058	652061	6810714	-28.82210	-55.44160	São Gabriel	-0.2	<0.5	27	829	<2	6	10	25	0.68	<10	99	<1	<10	0.17	17	7	2.30	0.02	0.11	0.04	0.011	<10	6	<10
OC-S-059	651728	6810749	-28.94010	-55.03290	São Gabriel	-0.2	<0.5	81	1044	<2	3	30	76	2.07	<10	799	3	<10	0.35	13	8	4.97	0.13	0.24	0.04	0.028	<10	13	<10
OC-S-060	680011	6800416	-28.87760	-55.15420	São Gabriel	-0.2	<0.5	91	1821	<2	16	22	77	2.00	<10	499	3	<10	0.89	41	12	7.26	0.09	0.36	0.07	0.050	<10	15	<10
OC-S-061	679613	6804012	-28.87900	-55.15820	São Gabriel	-0.2	<0.5	82	1547	<2	15	17	66	2.82	<10	336	2	<10	0.60	38	13	6.49	0.07	0.36	0.05	0.028	<10	20	<10
OC-S-062	688088	6819383	-28.73910	-55.07390	São Gabriel	-0.2	<0.5	174	4485	<2	29	25	83	2.61	<10	1133	5	<10	0.63	120	22	17.60	0.08	0.39	0.06	0.052	<10	21	<10
OC-S-063	676238	6821283	-28.72370	-55.19550	São Gabriel	-0.2	<0.5	88	1845	<2	9	15	30	0.89	<10	153	<1	<10	0.28	25	13	2.99	0.02	0.13	0.02	0.023	<10	9	<10
OC-S-064	645966	6818924	-28.74870	-55.50490	São Gabriel	-0.2	<0.5	38	1079	<2	9	15	30	0.89	<10	153	<1	<10	0.28	25	13	2.99	0.02	0.13	0.02	0.023	<10	9	<10
OC-S-065	652937	6821710	-28.72280	-55.43410	São Gabriel	-0.2	<0.5	100	1638	<2	14	28	87	1.33	<10	288	4	<10	0.37	60	27	10.99	0.05	0.22	0.04	0.075	<10	16	<10
OC-S-066	660655	6828961	-28.65640	-55.35610	São Gabriel	-0.2	<0.5	56	1390	<2	12	18	54	1.38	<10	242	2	<10	0.37	96	15	5.17	0.03	0.21	0.04	0.032	<10	14	<10
OC-S-067	661702	6830885	-28.63920	-55.34570	São Gabriel	-0.2	<0.5	79	1831	<2	22	15	63	1.84	<10	376	2	<10	0.68	59	20	7.41	0.03	0.36	0.06	0.035	<10	17	<10
OC-S-068	670807	6831543	-28.63180	-55.25260	São Gabriel	-0.2	<0.5	77	1467	<2	16	13	58	1.48	<10	280	1	<10	0.50	36	16	5.86	0.03	0.24	0.05	0.037	<10	13	<10
OC-S-069	614562	6831489	-28.63870	-55.82790	São Gabriel	-0.2	<0.5	187	4556	<2	44	40	65	2.91	<10	1209	3	<10	0.51	91	68	13.59	0.05	0.29	0.03	0.054	<10	26	<10
OC-S-070	648250	6730760	-29.54380	-55.46990	São Gabriel	-0.2	<0.5	19	641	<2	5	6	26	0.56	<10	75	<1	<10	0.05	18	23	2.69	0.01	0.03	0.02	0.015	<10	4	<10
OC-S-071	673348	6713331	-29.69790	-55.20920	São Gabriel	-0.2	<0.5	5	54	<2	<1	6	2	0.20	<10	14	<1	<10	<0.01	<1	3	0.18	<0.01	<0.01	0.002	<10	<1	<10	
OC-S-072	663042	6716303	-29.67250	-55.31520	São Gabriel	-0.2	<0.5	2	62	<2	<1	4	1	0.23	<10	19	<1	<10	0.01	2	4	0.31	<0.01	<0.01	0.002	<10	<1	<10	
OC-S-073	653766	6717686	-29.66140	-55.41120	São Gabriel	-0.2	<0.5	1	52	<2	1	5	1	0.17	<10	17	<1	<10	<0.01	1	3	0.21	<0.01	<0.01	0.002	<10	<1	<10	
OC-S-074	641336	6711507	-29.71850	-55.53980	São Gabriel	-0.2	<0.5	3	44	<2	<1	4	2	0.12	<10	14	<1	<10	0.01	1	3	0.21	<0.01	<0.01	0.002	<10	<1	<10	
OC-S-075	638711	6723270	-29.61280	-55.58810	São Gabriel	-0.2	<0.5	24	295	<2	5	9	36	0.48	<10	58	<1	<10	0.08	13	7	3.02	0.01	0.08	0.01	0.009	<10	6	<10
OC-S-076	641769	6720638	-29.63600	-55.53560	São Gabriel	-0.2	<0.5	9	419	<2	2	3	10	0.12	<10	39	<1	<10	0.02	12	6	1.02	<0.01	0.01	0.02	0.007	<10	1	<10
OC-S-077	651729	6725989	-29.59200	-55.43330	São Gabriel	-0.2	<0.5	5	183	<2	<1	3	3	0.07	<10	22	<1	<10	0.01	2	2	0.20	<0.01	<0.01	0.002	<10	<1	<10	
OC-S-078	625607	6757150	-29.30630	-55.70660	São Gabriel	-0.2	<0.5	7	174	<2	<1	5	6	0.16	<10	23	<1	<10	0.03	3	3	0.49	<0.01	0.01	0.004	<10	1	<10	
OC-S-079	625780	6741563	-29.44890	-55.70300	São Gabriel	-0.2	<0.5	5	97	<2	2	6	6	0.24	<10	28	<1	<10	0.01	4	6	0.56	<0.01	0.01	0.004	<10	2	<10	
OC-S-080	6217048	6744863	-29.41960	-55.75220	São Gabriel	-0.2	<0.5	88	1697	<2	16	18	52	1.79	<10	305	2	<10	0.61	41	14	6.58	0.04	0.39	0.03	0.025	<10	19	<10
OC-S-081	615071	6748132	-29.39070	-55.61410	São Gabriel	-0.2	<0.5	9	136	<2	1	5	8	0.25	<10	29	<1	<10	0.04	3	4	0.86	<0.01	0.03	0.01	0.005	<10	2	<10
OC-S-082	628005	6790056	-29.01120	-55.68570	São Gabriel	-0.2	<0.5	11	489	<2	2	16	13	0.21	<10	61	<1	<10	0.03	10	8	1.29	<0.01	0.03	0.01	0.004	<10	4	<10
OC-S-083	697509	6757846	-29.29280	-54.96670	São Gabriel	-0.2	<0.5	35	894	<2	7	26	25	0.73	<10	133	<1	<10	0.27	25	15	3.34	0.02	0.14	0.02	0.022	<10	7	<10
OC-S-084	671231	6725320	-29.59010	-55.23200	São Gabriel	-0.2	<0.5	5	84	<2	<1	2	2	0.08	<10	14	<1	<10	0.01	1	1	0.14	<0.01	<0.01	0.002	<10	<1	<10	
OC-S-085	676452	6733361	-29.51680	-55.17940	São Gabriel	-0.2	<0.5	4	576	<2	<1	<2	1	0.07	<10	40	<1	<10	0.01	4	1	0.22	<0.01	<0.01	0.002	<10	<1	<10	
OC-S-086	708016	6722050	-29.61430	-54.87240	São Gabriel	-0.2	<0.5	6	265	<2	<1	3	4	0.10	<10	39	<1	<10	0.05	4	2	0.43	<0.01	0.02	0.01	0.006	<10	<1	<10
OC-S-087	748111	6706649	-29.74550	-54.43450	São Gabriel	-0.2	<0.5	5	24	<2	<1	3	2	0.05	<10	19	<1	<10	0.02	<1	<1	0.12	<0.01	<0.01	0.002	<10	<1	<10	
OC-S-088	746927	6710745	-29.70880	-54.44760	São Gabriel	-0.2	<0.5	3	77	<2	1	4	4	0.12	<10	28	<1	<10	0.04	2	2	0.35	<0.01	0.02	0.01	0.002	<10	<1	<10
OC-S-089	6703170	6703170	-29.79880	-54.64500	São Gabriel	-0.2	<0.5	6	32	<2	<1	2	2	0.08	<10	23	<1	<10	0.02	<1	1	0.11	<0.01	<0.01	0.003	<10	<1	<10	
OC-S-090	614851	6841533	-28.54790	-55.82600	São Gabriel	-0.2	<0.5	138	855	<2	30	7	55	1.93	<10	180	<1	<10	0.79	25	18	4.85	0.06	0.55	0.13	0.042	<10	9	10
OC-S-091	622000	6833087	-28.62340	-55.75200	São Gabriel	-0.2	<0.5	106	4709	<2	23	31	55	1.76	<10	530	2	<10	0.61	72	32	9.58	0.03	0.28	0.03	0.045	<10	18	<10
OC-S-092	623391	6833752	-28.61730	-55.73780	São Gabriel	-0.2	<0.5	40	960	<2	7	19	32	0.45	<10	129	<1	<10	0.16	18	13	2.77	0.01	0.09	0.02	0.014	<10	7	<10
OC-S-093	645793	6845261	-28.51110	-55.51020	São Gabriel	-0.2	<0.5	23	229	<2	6	13	22																

Geochemical grade assay for stream sediments collected in the São Gabriel district

Sample No.	Sr ppm	Ti %	V ppm	W ppm	Y ppm	Zr ppm	S %
OC-S-053	2	0.01	22	<10	2	3	0.002
OC-S-054	<1	0.01	24	<10	<1	3	0.002
OC-S-055	9	0.06	72	<10	4	13	0.003
OC-S-056	2	0.04	10	<10	2	3	0.002
OC-S-057	54	0.27	206	<10	29	46	0.005
OC-S-058	14	0.29	107	<10	9	29	0.004
OC-S-059	108	0.27	113	<10	39	66	0.014
OC-S-060	88	0.25	314	<10	39	37	0.007
OC-S-061	58	0.21	239	<10	41	56	0.010
OC-S-062	89	0.22	870	<10	53	81	0.005
OC-S-063	56	0.22	229	<10	46	58	0.009
OC-S-064	21	0.08	140	<10	17	20	0.008
OC-S-065	37	0.24	505	<10	37	54	0.009
OC-S-066	39	0.31	251	<10	28	45	0.006
OC-S-067	57	0.28	387	<10	27	47	0.005
OC-S-068	39	0.30	281	<10	22	40	0.006
OC-S-069	51	0.15	527	<10	40	53	0.005
OC-S-070	5	0.05	109	<10	5	16	0.004
OC-S-071	<1	0.04	9	<10	4	3	0.003
OC-S-072	1	0.03	12	<10	2	2	0.003
OC-S-073	<1	0.02	7	<10	2	2	0.003
OC-S-074	1	0.02	11	<10	1	2	0.002
OC-S-075	6	0.12	183	<10	9	30	0.005
OC-S-076	3	0.01	35	<10	2	4	0.002
OC-S-077	1	0.01	9	<10	1	3	0.002
OC-S-078	3	0.03	23	<10	2	7	0.002
OC-S-079	2	0.03	24	<10	3	4	0.003
OC-S-080	39	0.13	228	<10	32	82	0.005
OC-S-081	4	0.07	43	<10	3	13	0.002
OC-S-082	3	0.19	58	<10	4	25	0.002
OC-S-083	16	0.13	185	<10	17	37	0.004
OC-S-084	1	<0.01	6	<10	1	2	0.002
OC-S-085	2	<0.01	11	<10	<1	2	0.001
OC-S-086	4	<0.01	15	<10	3	4	0.002
OC-S-087	2	<0.01	4	<10	1	1	0.002
OC-S-088	3	<0.01	11	<10	2	3	0.003
OC-S-089	2	<0.01	3	<10	1	1	0.003
OC-S-090	49	0.16	222	<10	21	51	0.003
OC-S-091	47	0.14	337	<10	32	60	0.005
OC-S-092	15	0.33	164	<10	12	48	0.003
OC-S-093	7	0.68	157	<10	4	44	0.004
OC-S-094	20	0.32	306	<10	21	61	0.003
OC-S-095	43	0.25	407	<10	44	54	0.013
OC-S-097	53	0.03	475	<10	44	35	0.008
OC-S-098	19	0.10	137	<10	15	25	0.005
OC-S-099	36	0.07	638	<10	40	43	0.008
OC-S-101	51	0.12	232	<10	28	48	0.006
OC-S-102	30	0.26	195	<10	18	48	0.003
OC-S-103	42	0.20	172	<10	22	52	0.008
OC-S-104	31	0.11	409	<10	36	57	0.004
OC-S-105	36	0.08	502	<10	23	60	0.004
OC-S-106	58	0.11	699	<10	59	61	0.007
OC-S-107	17	0.11	110	<10	11	30	0.003
OC-S-109	46	0.14	441	<10	47	44	0.003

Geochemical grade assay for stream sediments collected in the São Gabriel district

(7/7)

Sample No.	UTM_E	UTM_N	Latitude	Longitude	Location	Ag	Cd	Cu	Mn	Mo	Ni	Pb	Zn	Al	As	Ba	Be	Bi	Ca	Co	Cr	Fe	K	Mg	Na	P	Sb	Sc	Sn
						ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	%	ppm	ppm	ppm	ppm	%	%	ppm	ppm	%	%	%	ppm	ppm	ppm	ppm
OC-S-110	601429	6750601	-29.36960	-55.95490	São Gabriel	<0.2	<0.5	10	232	<2	2	5	10	0.24	<10	51	<1	<10	0.15	6	5	0.93	0.01	0.07	0.02	0.007	<10	3	<10
OC-S-111	583293	6759756	-29.28830	-56.14250	São Gabriel	<0.2	<0.5	26	383	<2	9	22	38	0.51	<10	122	<1	<10	0.34	18	12	2.42	0.02	0.16	0.02	0.026	<10	4	<10
OC-S-112	589562	6807613	-28.85600	-56.08180	São Gabriel	<0.2	<0.5	42	3375	<2	14	33	23	0.57	<10	329	1	<10	0.30	68	14	3.24	0.02	0.14	0.02	0.041	<10	7	<10
OC-S-113	589598	6800499	-28.91970	-56.01580	São Gabriel	<0.2	<0.5	26	1290	<2	6	15	20	0.65	<10	192	<1	<10	0.19	22	11	3.41	0.02	0.09	0.02	0.028	<10	6	<10
OC-S-114	567077	6799146	-28.93380	-56.31180	São Gabriel	<0.2	1.6	162	4090	<2	30	33	101	2.68	<10	616	6	<10	0.51	114	68	19.89	0.04	0.25	0.03	0.154	<10	28	<10
OC-S-115	558484	6781967	-29.08920	-56.39920	São Gabriel	<0.2	0.9	177	2125	<2	29	16	98	3.05	<10	517	3	<10	0.87	71	24	10.40	0.08	0.72	0.06	0.039	<10	39	11
OC-S-116	679090	6695750	-29.84700	-55.14630	São Gabriel	<0.2	<0.5	4	305	<2	1	4	5	0.21	<10	38	<1	<10	0.04	4	7	0.58	0.01	0.02	0.02	0.008	<10	<1	<10
OC-S-117	698335	6696393	-29.84700	-54.94700	São Gabriel	<0.2	<0.5	2	66	<2	1	3	2	0.21	<10	26	<1	<10	0.04	1	3	0.23	0.01	0.03	0.02	0.004	<10	<1	<10
OC-S-118	720018	6694685	-29.85870	-54.72230	São Gabriel	<0.2	<0.5	6	38	<2	1	3	1	0.06	<10	17	<1	<10	0.01	1	2	0.23	<0.01	<0.01	0.02	0.003	<10	<1	<10
OC-S-119	732056	6696733	-29.83800	-54.59820	São Gabriel	<0.2	<0.5	4	24	<2	1	3	<1	0.04	<10	16	<1	<10	<0.01	<1	2	0.20	<0.01	<0.01	0.02	0.002	<10	<1	<10
OC-S-120	730391	6687313	-30.10360	-54.60910	São Gabriel	<0.2	<0.5	2	26	<2	1	4	2	0.19	<10	28	<1	<10	0.03	<1	3	0.21	0.02	0.02	0.02	0.003	<10	<1	<10
OC-S-121	716817	6681180	-29.98100	-54.75270	São Gabriel	<0.2	<0.5	7	26	<2	1	4	<1	0.07	<10	16	<1	<10	0.01	<1	4	0.18	<0.01	<0.01	0.02	0.001	<10	<1	<10
OC-S-122	729068	6674922	-30.09630	-54.68660	São Gabriel	<0.2	<0.5	3	30	<2	1	3	1	0.04	<10	16	<1	<10	0.01	<1	1	0.12	<0.01	<0.01	0.02	0.001	<10	<1	<10
OC-S-123	734021	6706247	-29.75180	-54.58000	São Gabriel	<0.2	<0.5	3	22	<2	1	5	1	0.13	<10	30	<1	<10	0.03	<1	3	0.12	<0.01	0.01	0.02	0.004	<10	<1	<10
OC-S-124	702880	6712366	-29.68930	-54.20690	São Gabriel	<0.2	<0.5	8	174	<2	2	7	10	0.42	<10	52	<1	<10	0.08	4	3	0.88	0.02	0.04	0.02	0.010	<10	3	<10
OC-S-125	760917	6706194	-29.74690	-54.30210	São Gabriel	<0.2	<0.5	7	31	<2	1	3	2	0.10	<10	23	<1	<10	0.02	<1	2	0.15	0.01	0.01	0.02	0.003	<10	<1	<10
OC-S-126	766001	6715174	-29.66490	-54.25180	São Gabriel	<0.2	<0.5	12	324	<2	1	4	4	0.14	<10	46	<1	<10	0.03	6	2	0.60	0.01	0.01	0.03	0.006	<10	<1	<10
OC-S-127	717867	6700412	-29.80740	-54.74570	São Gabriel	<0.2	<0.5	6	45	<2	1	2	<1	0.07	<10	19	<1	<10	0.01	<1	2	0.11	0.01	<0.01	0.03	0.003	<10	<1	<10
OC-S-128	778046	6689295	-29.89550	-54.12080	São Gabriel	<0.2	<0.5	6	284	<2	2	9	7	0.51	<10	152	<1	<10	0.11	4	5	0.86	0.03	0.05	0.03	0.011	<10	1	<10
OC-S-129	770896	6686082	-29.92610	-54.19390	São Gabriel	<0.2	<0.5	13	1831	<2	7	25	23	1.27	18	243	3	<10	0.29	20	16	4.69	0.08	0.17	0.03	0.059	<10	3	<10
OC-S-130	779023	6671786	-30.05310	-54.10610	São Gabriel	<0.2	<0.5	4	88	<2	2	5	5	0.36	<10	72	<1	<10	0.09	2	3	0.41	0.05	0.08	0.03	0.006	<10	<1	<10
OC-S-131	771336	6688828	-30.17160	-54.18240	São Gabriel	<0.2	<0.5	11	169	<2	5	6	14	0.56	<10	187	<1	<10	0.22	4	7	0.89	0.06	0.13	0.04	0.014	<10	2	<10
OC-S-132	739211	6663886	-30.13280	-54.51680	São Gabriel	<0.2	<0.5	2	108	<2	1	5	2	0.22	<10	37	<1	<10	0.05	1	3	0.25	0.02	0.03	0.02	0.004	<10	<1	<10
OC-S-133	734452	6662994	-30.14270	-54.56600	São Gabriel	<0.2	<0.5	9	290	<2	3	14	14	0.67	<10	114	<1	<10	0.16	5	5	1.11	0.05	0.09	0.03	0.013	<10	2	<10
OC-S-133A	735648	6663170	-30.13990	-54.55370	São Gabriel	<0.2	<0.5	2	149	<2	1	4	<1	0.24	<10	47	<1	<10	25.80	1	3	0.29	0.03	0.04	0.02	0.005	<10	<1	<10
OC-S-134	755461	6706181	-29.74820	-54.35840	São Gabriel	<0.2	<0.5	7	55	<2	1	3	3	0.12	<10	25	<1	<10	0.02	2	2	0.50	0.01	0.01	0.02	0.007	<10	<1	<10
OC-S-135	772094	6740928	-29.43150	-54.19590	São Gabriel	<0.2	0.8	103	1341	<2	13	17	86	2.43	<10	429	3	<10	0.52	32	6	8.76	0.11	0.38	0.05	0.051	<10	21	<10

Geochemical grade assay for stream sediments collected in the São Gabriel district

Sample No.	Sr ppm	Ti %	V ppm	W ppm	Y ppm	Zr ppm	S %
OC-S-110	22	0.04	49	<10	5	13	0.003
OC-S-111	26	0.14	216	<10	13	30	0.010
OC-S-112	26	0.06	168	<10	16	26	0.002
OC-S-113	14	0.06	134	<10	11	24	0.003
OC-S-114	49	0.20	889	<10	56	72	0.009
OC-S-115	90	0.22	352	<10	85	114	0.008
OC-S-116	4	<0.01	13	<10	3	3	0.003
OC-S-117	3	0.02	7	<10	2	2	0.003
OC-S-118	2	<0.01	5	<10	1	2	0.003
OC-S-119	1	<0.01	6	<10	1	2	0.002
OC-S-120	3	0.01	5	<10	2	2	0.003
OC-S-121	2	0.02	5	<10	1	3	0.002
OC-S-122	1	<0.01	3	<10	<1	2	0.002
OC-S-123	2	0.02	5	<10	6	1	0.003
OC-S-124	14	0.04	28	<10	7	17	0.003
OC-S-125	2	0.01	4	<10	2	2	0.003
OC-S-126	2	<0.01	17	<10	2	4	0.002
OC-S-127	2	<0.01	3	<10	1	2	0.002
OC-S-128	10	0.03	19	<10	7	5	0.004
OC-S-129	26	<0.01	66	<10	14	13	0.006
OC-S-130	13	0.01	7	<10	4	7	0.004
OC-S-131	25	<0.01	14	<10	5	7	0.022
OC-S-132	5	0.02	5	<10	3	3	0.003
OC-S-133	20	<0.01	22	<10	6	8	0.006
OC-S-133A	6	<0.01	7	<10	3	4	<0.001
OC-S-134	3	<0.01	11	<10	2	3	0.002
OC-S-135	66	0.33	313	<10	32	76	0.007

Geochemical grade assay for stream sediments collected in the Ponta Grossa district

Sample No.	UTM_E	UTM_N	Latitude	Longitude	Location	Ag	Cd	Cu	Mn	Mo	Ni	Pb	Zn	Al	As	Ba	Be	Bi	Ca	Co	Cr	Fe	K	Mg	Na	P	Sb	Sc	Sn
ES-S-49	518065	7289248	-24.51035	-50.82168	Ponta Grossa	<0.2	<0.5	71	269	<2	14	19	54	1.85	<10	86	2	15	0.07	10	44	7.51	<0.01	<0.01	0.01	0.076	<10	24	<10
ES-S-50	517121	7286093	-24.53885	-50.83096	Ponta Grossa	<0.2	<0.5	103	1211	4	35	3	76	1.91	<10	174	1	17	0.85	36	21	7.40	0.02	0.35	0.03	0.108	<10	25	<10
ES-S-51	519727	7282463	-24.57161	-50.80518	Ponta Grossa	<0.2	<0.5	22	389	<2	13	18	54	1.43	<10	150	1	<10	0.16	13	12	3.26	0.04	0.13	0.02	0.030	<10	7	<10
ES-S-52	519951	7282870	-24.56793	-50.80297	Ponta Grossa	<0.2	<0.5	24	632	3	17	14	84	1.07	<10	134	1	<10	0.20	20	13	4.28	0.08	0.16	0.02	0.038	<10	7	<10
ES-S-53	519871	7287894	-24.52260	-50.84332	Ponta Grossa	<0.2	<0.5	29	331	3	8	12	41	1.81	<10	84	<1	<10	0.10	7	11	4.37	0.03	0.06	0.02	0.036	<10	12	<10
ES-S-54	518900	7276105	-24.62904	-50.81326	Ponta Grossa	<0.2	<0.5	133	759	<2	19	6	74	2.88	<10	213	2	17	0.28	27	13	8.43	0.04	0.15	0.03	0.089	<10	27	<10
ES-S-55	519211	7276841	-24.62239	-50.81020	Ponta Grossa	<0.2	<0.5	46	803	<2	14	9	70	1.56	<10	144	1	<10	0.18	20	15	5.08	0.04	0.09	0.02	0.041	<10	11	<10
ES-S-56	515438	7275863	-24.63126	-50.84746	Ponta Grossa	<0.2	<0.5	62	643	<2	10	14	54	2.43	<10	146	1	13	0.11	13	9	6.43	0.03	0.05	0.02	0.058	<10	20	<10
ES-S-57	519160	7274059	-24.64752	-50.81066	Ponta Grossa	<0.2	<0.5	75	772	<2	32	19	93	1.74	<10	111	1	<10	0.16	30	66	6.92	0.03	0.07	0.02	0.046	<10	19	<10
ES-S-58	519082	7271757	-24.66831	-50.81160	Ponta Grossa	<0.2	<0.5	135	1310	2	48	5	111	1.78	<10	176	2	21	0.44	35	59	7.78	0.04	0.29	0.03	0.048	<10	23	<10
ES-S-59	516496	7268915	-24.69400	-50.83693	Ponta Grossa	<0.2	<0.5	55	783	<2	19	7	78	1.07	<10	110	1	<10	0.10	20	24	5.88	0.02	0.03	0.02	0.024	<10	13	<10
ES-S-60	516637	7267877	-24.70338	-50.83552	Ponta Grossa	<0.2	<0.5	11	438	<2	11	10	31	0.81	<10	124	<1	<10	0.10	9	10	1.81	0.03	0.07	0.01	0.020	<10	4	<10
ES-S-61	516769	7264833	-24.72986	-50.83418	Ponta Grossa	<0.2	<0.5	22	454	<2	12	15	57	1.38	<10	174	1	<10	0.27	15	12	3.57	0.05	0.15	0.02	0.046	<10	7	<10
ES-S-62	519637	7271135	-24.67392	-50.80591	Ponta Grossa	<0.2	<0.5	25	342	2	13	18	43	0.58	<10	85	<1	<10	0.09	10	18	2.12	0.02	0.04	0.02	0.014	<10	4	<10
ES-S-63	518218	7261590	-24.76014	-50.81981	Ponta Grossa	<0.2	<0.5	11	181	2	8	9	30	0.79	<10	87	<1	<10	0.11	6	10	1.75	0.07	0.09	0.02	0.015	<10	3	<10
ES-S-64	522088	7268979	-24.79346	-50.82045	Ponta Grossa	<0.2	<0.5	5	75	<2	4	5	15	0.24	<10	37	<1	<10	0.03	2	3	0.57	0.01	0.02	0.02	0.006	<10	1	<10
ES-S-65	518149	7257901	-24.80173	-50.82045	Ponta Grossa	<0.2	<0.5	19	492	<2	10	12	73	1.07	<10	100	<1	<10	0.16	12	12	3.60	0.03	0.10	0.02	0.027	<10	7	<10
ES-S-66	519708	7251282	-24.84498	-50.80493	Ponta Grossa	<0.2	<0.5	13	287	<2	9	8	35	1.08	<10	112	<1	<10	0.12	7	8	1.65	0.07	0.11	0.02	0.011	<10	3	<10
ES-S-67	526520	7252182	-24.85261	-50.73752	Ponta Grossa	<0.2	<0.5	11	98	<2	3	<2	18	0.25	<10	42	<1	<10	0.03	3	4	0.93	0.01	0.02	0.02	0.006	<10	1	<10
ES-S-68	505014	7248283	-24.88042	-50.95036	Ponta Grossa	<0.2	<0.5	39	852	2	21	21	113	1.09	<10	184	2	<10	0.37	24	19	5.59	0.06	0.23	0.03	0.040	<10	8	<10
ES-S-69	503925	7244764	-24.91220	-50.96113	Ponta Grossa	<0.2	<0.5	15	445	<2	12	14	55	0.52	<10	164	1	<10	0.21	20	19	5.08	0.04	0.16	0.02	0.039	<10	7	<10
ES-S-70	500941	7238934	-24.96486	-50.90668	Ponta Grossa	<0.2	<0.5	15	445	<2	12	14	55	0.52	<10	164	1	<10	0.21	20	19	5.08	0.04	0.16	0.02	0.039	<10	7	<10
ES-S-71	513098	7236573	-24.98612	-50.87022	Ponta Grossa	<0.2	<0.5	25	673	4	27	21	78	3.02	21	247	2	<10	0.21	25	34	6.78	0.09	0.14	0.03	0.039	<10	16	<10
ES-S-72	509655	7244348	-24.91593	-50.90439	Ponta Grossa	<0.2	<0.5	22	364	<2	13	15	61	0.98	<10	207	1	<10	0.24	13	14	2.37	0.11	0.20	0.03	0.019	<10	4	<10
ES-S-73	512655	7242867	-24.92929	-50.87466	Ponta Grossa	<0.2	<0.5	23	624	5	15	15	56	1.32	<10	199	1	<10	0.22	11	16	2.66	0.13	0.21	0.03	0.017	<10	5	<10
ES-S-74	512681	7242689	-24.93080	-50.87440	Ponta Grossa	<0.2	<0.5	23	624	5	15	15	56	1.32	<10	199	1	<10	0.22	11	16	2.66	0.13	0.21	0.03	0.017	<10	5	<10
ES-S-75	509875	7246620	-24.89541	-50.89145	Ponta Grossa	<0.2	<0.5	21	492	<2	12	10	51	0.96	<10	173	<1	<10	0.20	10	11	2.07	0.10	0.19	0.03	0.015	<10	4	<10
ES-S-76	507787	7245836	-24.90251	-50.92289	Ponta Grossa	<0.2	<0.5	16	654	5	15	19	78	0.93	<10	144	1	<10	0.27	17	17	3.50	0.06	0.20	0.03	0.039	<10	5	<10
ES-S-77	511163	7242725	-24.93058	-50.88944	Ponta Grossa	<0.2	<0.5	12	90	3	8	10	21	0.50	<10	169	<1	<10	0.12	10	10	1.80	0.07	0.11	0.03	0.013	<10	3	<10
ES-S-78	522315	7240572	-24.94950	-50.77895	Ponta Grossa	<0.2	<0.5	12	90	3	8	10	21	0.50	<10	169	<1	<10	0.12	10	10	1.80	0.07	0.11	0.03	0.013	<10	3	<10
ES-S-79	520233	7242829	-24.93406	-50.79960	Ponta Grossa	<0.2	<0.5	12	90	3	8	10	21	0.50	<10	169	<1	<10	0.12	10	10	1.80	0.07	0.11	0.03	0.013	<10	3	<10
ES-S-80	519952	7243194	-24.92626	-50.80240	Ponta Grossa	<0.2	<0.5	16	367	<2	15	20	48	1.62	<10	235	1	<10	0.18	12	20	4.47	0.12	0.15	0.02	0.023	<10	11	<10
ES-S-81	507733	7243008	-24.92776	-50.69562	Ponta Grossa	<0.2	<0.5	14	122	<2	5	4	17	0.40	<10	183	1	<10	0.13	12	11	2.94	0.07	0.10	0.02	0.025	<10	4	<10
ES-S-82	530742	7242176	-24.92776	-50.69562	Ponta Grossa	<0.2	<0.5	6	35	<2	3	2	8	0.19	<10	22	<1	<10	0.01	1	2	0.34	<0.01	<0.01	0.02	0.003	<10	<1	<10
ES-S-83	532808	7242176	-24.93523	-50.67505	Ponta Grossa	<0.2	<0.5	6	35	<2	3	2	8	0.19	<10	22	<1	<10	0.01	1	2	0.34	<0.01	<0.01	0.02	0.003	<10	<1	<10
ES-S-84	534277	7241304	-24.94307	-50.66048	Ponta Grossa	<0.2	<0.5	16	353	<2	12	8	27	1.52	<10	96	<1	<10	0.04	9	14	2.19	0.03	0.03	0.02	0.023	<10	4	<10
ES-S-85	465141	7284184	-24.55560	-51.34422	Ponta Grossa	<0.2	<0.5	13	391	<2	15	3	29	1.82	<10	106	<1	<10	0.06	11	18	2.25	0.06	0.05	0.02	0.033	<10	4	<10
ES-S-86	459837	7281766	-24.57751	-51.39666	Ponta Grossa	<0.2	<0.5	135	2178	2	25	7	108	1.68	<10	217	2	<10	0.28	52	11	10.28	0.04	0.12	0.03	0.035	<10	23	<10
ES-S-87	459099	7286107	-24.53828	-51.40382	Ponta Grossa	<0.2	<0.5	46	692	<2	10	7	61	0.75	<10	81	<1	<10	0.11	15	12	3.97	0.03	0.05	0.02	0.016	<10	6	<10
ES-S-88	466908	7280306	-24.59066	-51.32696	Ponta Grossa	<0.2	<0.5	30	410	<2	10	3	45	0.71	<10	59	<1	<10	0.09	7	7	3.09	0.01	0.04	0.02	0.016	<10	6	<10
ES-S-89	473241	7276588	-24.62456	-51.26438	Ponta Grossa	<0.2	<0.5	29	363	<2	10	4	36	0.44	<10	65	<1	<10	0.10	9	6	2.15	0.02	0.06	0.01	0.010	<10	4	<10
ES-S-90	476887	7274711	-24.64157	-51.23035	Ponta Grossa	<0.2	<0.5	29	363	<2	10	4	36	0.44	<10	65	<1	<10	0.10	9	6	2.15	0.02	0.06	0.01	0.010	<10	4	<10
ES-S-91	478377	7264601	-24.73290	-51.21382	Ponta Grossa	<0.2	<0.5	18	271	<2	7	4	36	0.59	<10	62	<1	<10	0.11	7	4	1.24	0.03	0.06	0.01	0.011	<10	2	<10
ES-S-92	479068	7266894	-24.71220	-51.20695	Ponta Grossa	<0.2	<0.5	18	271	<2	7	4	36	0.59	<10	62	<1	<10	0.11	7	4	1.24	0.03	0.06	0.01	0.011	<10	2	<10
ES-S-93	483807	7260747	-24.76778	-51.16017	Ponta Grossa	<0.2	<0.5	11	258	<2	6	7	28	0.48	<10	86	<1	<10	0.14	4	6	2.22	0.04	0.12	0.02	0.012	<10	2	<10
ES-S-94																													

Geochemical grade assay for stream sediments collected in the Ponta Grossa district

Sample No.	Sr ppm	Ti %	V ppm	W ppm	Y ppm	Zr ppm	S %
ES-S-49	7	0.19	327	<10	6	77	0.022
ES-S-50	34	0.25	343	<10	31	60	0.008
ES-S-51	16	0.14	152	<10	8	32	0.011
ES-S-52	17	0.35	369	<10	7	42	0.007
ES-S-53	11	0.28	146	<10	7	51	0.010
ES-S-54	23	0.24	291	<10	24	74	0.013
ES-S-55	19	0.30	251	11	10	51	0.010
ES-S-56	17	0.24	239	<10	12	61	0.017
ES-S-57	12	0.41	469	<10	11	61	0.019
ES-S-58	25	0.45	572	<10	21	68	0.008
ES-S-59	10	0.37	508	<10	8	55	0.011
ES-S-60	10	0.04	57	<10	7	15	0.008
ES-S-61	20	0.05	135	<10	11	28	0.007
ES-S-62	8	0.20	142	<10	6	22	0.011
ES-S-63	10	0.10	81	<10	7	19	0.004
ES-S-64	2	0.12	30	<10	4	13	0.003
ES-S-65	13	0.30	266	<10	6	38	0.007
ES-S-66	11	0.07	36	<10	6	22	0.004
ES-S-67	3	0.05	25	<10	4	6	0.002
ES-S-68	25	0.32	526	<10	11	45	0.011
ES-S-69	16	0.31	404	<10	8	44	0.009
ES-S-70	14	0.13	113	<10	6	21	0.008
ES-S-71	23	0.22	232	<10	19	61	0.051
ES-S-72	23	0.11	95	<10	7	24	0.016
ES-S-73	17	0.06	46	<10	6	23	0.007
ES-S-74	21	0.08	68	<10	8	27	0.013
ES-S-75	19	0.08	72	<10	6	22	0.013
ES-S-76	22	0.24	231	<10	8	33	0.013
ES-S-77	14	0.03	44	<10	6	14	0.011
ES-S-78	4	0.01	12	<10	5	7	0.009
ES-S-79	25	0.06	106	<10	13	54	0.010
ES-S-80	20	<0.01	43	<10	10	27	0.010
ES-S-81	2	<0.01	12	<10	3	6	0.005
ES-S-82	1	<0.01	8	<10	2	3	0.002
ES-S-83	4	0.02	44	<10	6	15	0.009
ES-S-84	7	<0.01	41	<10	5	16	0.010
ES-S-85	12	0.04	47	<10	5	15	0.010
ES-S-86	21	0.41	617	<10	18	106	0.011
ES-S-87	10	0.31	316	<10	6	45	0.009
ES-S-88	7	0.23	218	<10	5	37	0.006
ES-S-89	8	0.16	142	<10	3	24	0.009
ES-S-90	9	0.15	151	<10	4	25	0.006
ES-S-91	10	0.09	85	<10	4	15	0.003
ES-S-92	15	0.07	101	<10	8	29	0.003
ES-S-93	13	0.05	26	<10	4	7	0.006
ES-S-94	16	0.03	19	<10	4	11	0.006
ES-S-95	17	0.17	88	<10	5	17	0.009
ES-S-96	19	0.04	26	<10	5	14	0.006
ES-S-97	19	0.27	194	<10	7	29	0.006
ES-S-98	12	0.15	100	<10	5	17	0.005
ES-S-99	24	0.27	134	<10	7	32	0.007
ES-S-100	18	0.26	203	<10	7	27	0.010
ES-S-101	26	0.15	131	<10	10	29	0.008
ES-S-102	27	0.16	128	<10	8	37	0.013

Geochemical grade assay for stream sediments collected in the Ponta Grossa district

Sample No.	UTM_E	UTM_N	Latitude	Longitude	Location	Ag	Cd	Cu	Mn	Mo	Ni	Pb	Zn	Al	As	Ba	Be	Bi	Ca	Co	Cr	Fe	K	Mg	Na	P	Sb	Sc	Sn
						ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	%	ppm	ppm	ppm	ppm	%	%	ppm	ppm	%	%	%	%	ppm	ppm	ppm
ES-S-103	495310	7245126	-24.90893	-51.04644	Ponta Grossa	<0.2	<0.5	19	385	3	13	16	66	0.95	<10	172	<1	<10	0.29	7	13	1.88	0.08	0.25	0.02	0.022	<10	3	<10
ES-S-104	632699	7337603	-24.06813	-49.69276	Ponta Grossa	<0.2	<0.5	13	77	<2	9	10	29	1.21	<10	48	<1	<10	0.05	3	16	1.68	0.02	0.02	0.02	0.016	<10	4	<10
ES-S-105	631097	7339799	-24.10263	-49.71014	Ponta Grossa	<0.2	<0.5	27	582	<2	12	15	34	1.20	<10	110	<1	<10	0.09	9	15	2.95	0.03	0.02	0.02	0.018	<10	6	<10
ES-S-106	633006	7335930	-24.08323	-49.69155	Ponta Grossa	<0.2	<0.5	30	691	<2	11	5	36	0.76	<10	83	<1	<10	0.06	10	16	3.65	0.01	<0.01	0.02	0.014	<10	8	<10
ES-S-107	635733	7336561	-24.07730	-49.66479	Ponta Grossa	<0.2	<0.5	67	556	3	13	5	70	0.84	<10	89	1	<10	0.29	16	8	5.43	0.04	0.07	0.02	0.055	<10	10	<10
ES-S-108	636146	7337025	-24.07312	-49.66646	Ponta Grossa	<0.2	<0.5	67	567	<2	11	9	72	1.07	<10	116	<1	<10	0.27	17	8	4.59	0.04	0.07	0.02	0.061	<10	9	<10
ES-S-109	636146	7336999	-24.10311	-49.66046	Ponta Grossa	<0.2	<0.5	12	114	<2	2	2	12	0.27	<10	30	<1	<10	0.03	2	3	0.96	<0.01	<0.01	0.02	0.011	<10	2	<10
ES-S-110	637923	7336405	-24.07852	-49.64324	Ponta Grossa	<0.2	<0.5	73	560	3	15	5	89	0.77	<10	81	1	<10	0.32	18	8	6.03	0.04	0.08	0.03	0.053	<10	9	<10
ES-S-111	636675	7337458	-24.06912	-49.65561	Ponta Grossa	<0.2	<0.5	42	615	2	13	5	65	0.56	<10	69	<1	<10	0.11	15	22	4.87	0.02	0.05	0.02	0.023	<10	7	<10
ES-S-112	633900	7333800	-24.10239	-49.68256	Ponta Grossa	<0.2	<0.5	59	369	<2	8	6	53	0.74	<10	87	1	12	0.13	13	5	5.20	0.02	0.02	0.01	0.043	<10	11	<10

Geochemical grade assay for stream sediments collected in the Ponta Grossa district

Sample No.	Sr	Ti	V	W	Y	Zr	S
	ppm	%	ppm	ppm	ppm	ppm	%
ES-S-103	23	0.12	70	<10	6	20	0.009
ES-S-104	6	0.04	55	<10	5	16	0.018
ES-S-105	10	0.11	92	<10	6	30	0.010
ES-S-106	7	0.23	293	<10	6	34	0.008
ES-S-107	16	0.22	408	<10	14	44	0.005
ES-S-108	16	0.16	246	<10	18	32	0.016
ES-S-109	4	0.12	42	<10	2	10	0.006
ES-S-110	17	0.26	493	<10	12	44	0.006
ES-S-111	8	0.19	473	<10	8	40	0.005
ES-S-112	9	0.12	306	<10	15	48	0.008

List of stream water collected in the Lomba Grande district

Sample No.	E-W	N-S	Sample No.	E-W	N-S
CK-A- 1	22J 495450	UTM 6698500	CK-A- 69	22J 505500	UTM 6701250
CK-A- 2	22J 494900	UTM 6699000	CK-A- 70	22J 505100	UTM 6700550
CK-A- 3	22J 492450	UTM 6698150	CK-A- 71	22J 504900	UTM 6699650
CK-A- 4	22J 493000	UTM 6699100	CK-A- 72	22J 504900	UTM 6699550
CK-A- 5	22J 494800	UTM 6699550	CK-A- 73	22J 504800	UTM 6698950
CK-A- 6	22J 495000	UTM 6699650	CK-A- 74	22J 504850	UTM 6699250
CK-A- 7	22J 495700	UTM 6700200	CK-A- 75	22J 502050	UTM 6700200
CK-A- 8	22J 495850	UTM 6700400	CK-A- 76	22J 502550	UTM 6699550
CK-A- 9	22J 494250	UTM 6700350	CK-A- 77	22J 502600	UTM 6701150
CK-A- 10	22J 494150	UTM 6701550	CK-A- 78	22J 503100	UTM 6700350
CK-A- 11	22J 493200	UTM 6700600	CK-A- 79	22J 503000	UTM 6700200
CK-A- 12	22J 493100	UTM 6701400	CK-A- 80	22J 503400	UTM 6700100
CK-A- 13	22J 492200	UTM 6702000	CK-A- 81	22J 503400	UTM 6698900
CK-A- 14	22J 498150	UTM 6698500	CK-A- 82	22J 503900	UTM 6700150
CK-A- 15	22J 497700	UTM 6698500	CK-A- 83	22J 503850	UTM 6700050
CK-A- 16	22J 499300	UTM 6699550	CK-A- 84	22J 503950	UTM 6699800
CK-A- 17	22J 498800	UTM 6699350	CK-A- 85	22J 502700	UTM 6698100
CK-A- 18	22J 499100	UTM 6699650	CK-A- 86	22J 501750	UTM 6698900
CK-A- 19	22J 499550	UTM 6701100	CK-A- 87	22J 500700	UTM 6698400
CK-A- 20	22J 498200	UTM 6703000	CK-A- 88	22J 501100	UTM 6697550
CK-A- 21	22J 499050	UTM 6703100	CK-A- 89	22J 501050	UTM 6700200
CK-A- 22	22J 498500	UTM 6704000	CK-A- 90	22J 501450	UTM 6701400
CK-A- 23	22J 498700	UTM 6704000	CK-A- 91	22J 501200	UTM 6701600
CK-A- 24	22J 498650	UTM 6704650	CK-A- 92	22J 501100	UTM 6701100
CK-A- 25	22J 498650	UTM 6705000	CK-A- 93	22J 500050	UTM 6700950
CK-A- 26	22J 498300	UTM 6705200	CK-A- 94	22J 502250	UTM 6694300
CK-A- 27	22J 499150	UTM 6705800	CK-A- 95	22J 500200	UTM 6693250
CK-A- 28	22J 496450	UTM 6706100	CK-A- 96	22J 501700	UTM 6691650
CK-A- 29	22J 494200	UTM 6705600	CK-A- 97	22J 513500	UTM 6704750
CK-A- 30	22J 496900	UTM 6703400	CK-A- 98	22J 513400	UTM 6704200
CK-A- 31	22J 496050	UTM 6704250	CK-A- 99	22J 511950	UTM 6703750
CK-A- 32	22J 496100	UTM 6704400	CK-A- 100	22J 511500	UTM 6702450
CK-A- 33	22J 495250	UTM 6704300	CK-A- 101	22J 512850	UTM 6701750
CK-A- 34	22J 494100	UTM 6702900	CK-A- 102	22J 513950	UTM 6705950
CK-A- 35	22J 493800	UTM 6704000	CK-A- 103	22J 515100	UTM 6705200
CK-A- 36	22J 494750	UTM 6703700	CK-A- 104	22J 515700	UTM 6706700
CK-A- 37	22J 494900	UTM 6703700	CK-A- 105	22J 515750	UTM 6703400
CK-A- 38	22J 495300	UTM 6702750	CK-A- 106	22J 509600	UTM 6697650
CK-A- 39	22J 493350	UTM 6704900	CK-A- 107	22J 510200	UTM 6698000
CK-A- 40	22J 493200	UTM 6704900	CK-A- 108	22J 511100	UTM 6698200
CK-A- 41	22J 492300	UTM 6703450	CK-A- 109	22J 511600	UTM 6698700
CK-A- 42	22J 503750	UTM 6705700	CK-A- 110	22J 512300	UTM 6698250
CK-A- 43	22J 502850	UTM 6705300	CK-A- 111	22J 512250	UTM 6696900
CK-A- 44	22J 502900	UTM 6705450	CK-A- 112	22J 511850	UTM 6697650
CK-A- 45	22J 502400	UTM 6705350	CK-A- 113	22J 511050	UTM 6697450
CK-A- 46	22J 502700	UTM 6703000	CK-A- 114	22J 510850	UTM 6696700
CK-A- 47	22J 502900	UTM 6703050	CK-A- 115	22J 510750	UTM 6696450
CK-A- 48	22J 501750	UTM 6703600	CK-A- 116	22J 510150	UTM 6697000
CK-A- 49	22J 507400	UTM 6706700	CK-A- 117	22J 509600	UTM 6696150
CK-A- 50	22J 509500	UTM 6705950	CK-A- 118	22J 508850	UTM 6696550
CK-A- 51	22J 509400	UTM 6703850	CK-A- 119	22J 511400	UTM 6695950
CK-A- 52	22J 510000	UTM 6703900	CK-A- 120	22J 510950	UTM 6694250
CK-A- 53	22J 509000	UTM 6703500	CK-A- 121	22J 510100	UTM 6695550
CK-A- 54	22J 507950	UTM 6703000	CK-A- 122	22J 509550	UTM 6695000
CK-A- 55	22J 504750	UTM 6703850	CK-A- 123	22J 514950	UTM 6699950
CK-A- 56	22J 506300	UTM 6704050	CK-A- 124	22J 514250	UTM 6699600
CK-A- 57	22J 505700	UTM 6703800	CK-A- 125	22J 514900	UTM 6699050
CK-A- 58	22J 506700	UTM 6703200	CK-A- 126	22J 513350	UTM 6699000
CK-A- 59	22J 506450	UTM 6703150	CK-A- 127	22J 515700	UTM 6699400
CK-A- 60	22J 506600	UTM 6702900	CK-A- 128	22J 515600	UTM 6699200
CK-A- 61	22J 507150	UTM 6702000	CK-A- 129	22J 516300	UTM 6698250
CK-A- 62	22J 506650	UTM 6701700	CK-A- 130	22J 507350	UTM 6701050
CK-A- 63	22J 506700	UTM 6701450	CK-A- 131	22J 507450	UTM 6601100
CK-A- 64	22J 505800	UTM 6702150	CK-A- 132	22J 507700	UTM 6700300
CK-A- 65	22J 504850	UTM 6702500	CK-A- 133	22J 508450	UTM 6701250
CK-A- 66	22J 503900	UTM 6702950	CK-A- 134	22J 517550	UTM 6700450
CK-A- 67	22J 503900	UTM 6702100	CK-A- 135	22J 516550	UTM 6700350
CK-A- 68	22J 503900	UTM 6701100	CK-A- 136	22J 516750	UTM 6699250

List of stream water collected in the Lomba Grande district

Sample No.	E-W		N-S	
CK-A- 137	22J	516650	UTM	6698900
CK-A- 138	22J	517500	UTM	6698250
CK-A- 139	22J	517400	UTM	6698050
CK-A- 140	22J	517900	UTM	6699500
CK-A- 141	22J	518900	UTM	6697800
CK-A- 142	22J	517350	UTM	6696050
CK-A- 143	22J	515500	UTM	6698400
CK-A- 144	22J	515600	UTM	6698450
CK-A- 145	22J	516550	UTM	6696650
CK-A- 146	22J	513800	UTM	6695850
CK-A- 147	22J	514050	UTM	6695850
CK-A- 148	22J	512700	UTM	6693450
CK-A- 149	22J	515530	UTM	6694200
CK-A- 150	22J	515550	UTM	6694200
CK-A- 151	22J	515400	UTM	6694000
CK-A- 152	22J	514450	UTM	6692450
CK-A- 153	22J	514050	UTM	6692150
CK-A- 154	22J	512850	UTM	6690900
CK-A- 155	22J	512950	UTM	6691100
CK-A- 156	22J	513650	UTM	6690400
CK-A- 157	22J	515500	UTM	6693000
CK-A- 158	22J	514700	UTM	6691250
CK-A- 159	22J	511800	UTM	6690900
CK-A- 160	22J	512950	UTM	6687850
CK-A- 161	22J	513050	UTM	6689200
CK-A- 162	22J	513350	UTM	6688800
CK-A- 163	22J	514350	UTM	6690150
CK-A- 164	22J	515000	UTM	6689600
CK-A- 165	22J	512450	UTM	6689550
CK-A- 166	22J	509700	UTM	6692700
CK-A- 167	22J	509600	UTM	6692850
CK-A- 168	22J	508300	UTM	6689550
CK-A- 169	22J	508250	UTM	6689800
CK-A- 170	22J	506550	UTM	6690350
CK-A- 171	22J	505800	UTM	6689800
CK-A- 172	22J	503150	UTM	6690450
CK-A- 173	22J	502950	UTM	6690250
CK-A- 174	22J	503200	UTM	6691950
CK-A- 175	22J	505400	UTM	6692600
CK-A- 176	22J	505950	UTM	6690700
CK-A- 177	22J	508150	UTM	6692950
CK-A- 178	22J	504650	UTM	6695800
CK-A- 179	22J	507900	UTM	6694450
CK-A- 180	22J	506100	UTM	6695350
CK-A- 181	22J	504550	UTM	6696400
CK-A- 182	22J	491600	UTM	6698850

Geochemical grade assay of stream water collected in the Lomba Grande district

Sample No.	Li	Be	Na	Mg	Al	Si	K	Ca	Sc	Ti	V	Cr	Mn	Fe	Co	Ni	Cu	Zn	Ga	Ge	As	Se	Br	Rb	Sr	Y	Zr	Nb	Mo	Ru	Pd	Ag	Cd	In
CK-A-01	<1	<0.1	2470	813	170	4440	10400	982	4	4.3	0.42	1.4	82.2	704	0.533	0.6	10.4	23.4	0.05	<0.01	0.43	<0.2	54	3.093	11.64	0.900	0.16	0.013	<0.1	<0.01	<0.01	<0.01	<0.01	<0.01
CK-A-02	<1	<0.1	3370	1910	103	4220	599	2490	4	2.9	0.44	0.8	224.9	1310	1.099	0.6	1.4	5.1	0.04	<0.01	0.59	<0.2	107	2.356	23.63	0.251	0.08	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	<0.01
CK-A-03	2	<0.1	3140	1150	85	6070	1380	1540	5	3.6	0.64	0.7	78.9	706	0.650	0.6	12.5	8.6	0.03	<0.01	0.76	<0.2	38	4.457	13.92	0.362	0.10	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	<0.01
CK-A-04	2	<0.1	3080	975	54	6590	1700	1450	5	3.4	0.49	0.9	24.7	337	0.211	0.4	0.8	5.0	0.02	<0.01	0.29	<0.2	40	5.252	11.93	0.268	0.08	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	<0.01
CK-A-05	<1	<0.1	3380	1290	84	4550	1860	2730	4	3.1	0.43	1.1	426.2	1160	1.334	0.9	1.7	6.6	0.05	<0.01	0.53	<0.2	48	6.141	19.07	0.473	0.07	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	<0.01
CK-A-06	<1	<0.1	7300	7050	161	10400	1840	6230	8	5.1	0.83	1.7	323.8	991	1.678	2.0	5.4	10.1	0.05	<0.01	0.29	<0.2	59	5.247	32.96	0.731	0.04	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	<0.01
CK-A-07	<1	<0.1	4590	1610	38	6820	2790	1810	5	3.9	0.48	0.8	88.4	611	0.653	1.1	5.9	5.0	0.03	<0.01	0.41	<0.2	48	7.862	17.87	0.211	0.06	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	<0.01
CK-A-08	<1	<0.1	16500	9210	136	12100	1370	11800	9	6.7	1.63	1.5	184.5	470	0.716	1.3	1.4	1.8	0.06	0.02	0.16	<0.4	55	3.222	47.51	0.174	0.05	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	<0.01
CK-A-09	<1	<0.1	3850	752	82	6290	1850	959	4	3.3	0.61	0.7	30.2	184	0.181	0.5	1.1	4.3	0.03	<0.01	1.22	<0.2	41	6.069	11.51	0.183	0.06	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	<0.01
CK-A-10	<1	<0.1	4410	1530	81	6230	4960	1400	4	7.8	1.32	0.9	90.0	474	0.905	0.7	1.7	11.1	0.03	0.02	0.59	<0.2	48	11.456	17.36	0.946	0.36	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	<0.01
CK-A-11	<1	<0.1	2950	538	61	4690	2090	<50	4	2.7	0.52	1.0	22.3	190	0.099	<0.3	0.6	3.3	0.03	<0.01	0.17	<0.2	35	7.338	8.63	0.413	0.05	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	<0.01
CK-A-12	<1	<0.1	2420	804	61	4240	954	203	3	2.7	0.40	0.9	31.8	266	0.247	<0.3	3.6	6.9	0.03	<0.01	0.35	<0.2	37	3.103	8.94	0.169	0.08	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	<0.01
CK-A-13	<1	<0.1	2590	727	41	3860	968	134	3	2.6	0.34	0.8	35.3	195	0.228	<0.3	0.5	6.0	0.02	0.01	0.34	<0.2	34	3.238	8.32	0.141	0.10	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	<0.01
CK-A-14	<1	<0.1	5040	2590	65	9840	1570	2900	7	4.9	0.56	1.4	199.3	888	0.995	1.0	0.9	3.4	0.05	<0.01	0.31	<0.2	46	4.459	23.05	0.213	0.07	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	<0.01
CK-A-15	1	<0.1	3540	1100	56	6540	2410	1690	5	3.6	0.36	1.6	265.5	721	1.191	0.5	0.7	3.6	0.05	<0.01	0.33	<0.2	42	6.465	15.80	0.212	0.04	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	<0.01
CK-A-16	<1	<0.1	6890	5650	36	11000	884	5540	9	5.6	0.60	1.5	208.3	960	0.954	1.0	0.7	2.3	0.05	<0.01	0.30	<0.2	46	2.304	29.15	0.158	0.05	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	<0.01
CK-A-17	<1	<0.1	5040	2450	52	9540	2900	2450	7	5.0	0.58	1.1	218.1	903	0.753	0.7	0.8	1.4	0.06	<0.01	0.30	<0.2	40	7.958	20.52	0.158	0.03	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	<0.01
CK-A-18	<1	<0.1	4980	1850	71	9420	1580	2310	7	5.0	0.59	1.3	103.3	602	0.644	0.7	2.4	2.8	0.05	<0.01	0.31	<0.2	42	4.706	19.83	0.243	0.16	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	<0.01
CK-A-19	<1	<0.1	4280	916	68	9500	1730	1080	7	4.3	0.57	1.0	65.8	334	0.611	0.4	0.4	1.4	0.04	<0.01	0.26	<0.2	42	4.885	13.51	0.272	0.06	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	<0.01
CK-A-20	2	<0.1	8230	3250	35	13600	2700	5690	10	6.4	1.20	1.1	50.2	369	0.404	0.4	0.7	0.9	0.03	<0.01	0.33	<0.2	48	6.584	30.11	0.161	0.05	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	<0.01
CK-A-21	1	<0.1	5820	1080	61	11500	1650	1710	8	5.3	0.77	1.0	37.9	266	0.441	0.8	2.2	2.8	0.03	<0.01	0.28	<0.2	33	4.240	16.13	0.190	0.12	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	<0.01
CK-A-22	1	<0.1	6240	2310	108	9850	2650	3750	7	5.1	1.30	1.1	104.3	563	0.980	0.8	1.8	3.7	0.05	<0.01	0.30	<0.2	40	6.305	22.82	0.375	0.11	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	<0.01
CK-A-23	1	0.1	4260	1170	207	7370	1860	1780	5	4.0	1.17	1.2	139.9	993	0.968	0.9	2.9	6.2	0.08	<0.01	0.50	<0.2	41	4.990	17.24	0.551	0.22	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	<0.01
CK-A-24	<1	<0.1	4760	1260	87	7520	2430	2010	5	3.7	0.71	0.9	164.2	693	0.853	0.5	0.9	3.0	0.06	<0.01	0.44	<0.2	39	6.136	17.71	0.235	0.12	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	<0.01
CK-A-25	<1	<0.1	4470	675	146	7320	1150	1070	5	4.1	0.57	0.9	54.4	253	0.569	0.5	0.6	2.9	0.05	<0.01	0.31	<0.2	32	3.315	11.33	0.261	0.15	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	<0.01
CK-A-26	<1	<0.1	3290	1360	169	6240	2010	1030	5	3.3	1.11	1.6	102.4	1010	1.017	0.7	4.0	3.3	0.07	<0.01	0.53	<0.2	38	6.292	17.31	0.601	0.15	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	<0.01
CK-A-27	<1	<0.1	4390	836	127	7330	2250	928	5	4.1	0.89	0.9	17.2	143	0.211	0.3	1.2	2.7	0.04	<0.01	0.26	<0.2	33	6.106	13.03	0.445	0.14	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	<0.01
CK-A-28	<1	<0.1	3710	1810	288	6320	3380	2120	4	3.5	3.59	1.4	32.3	480	0.402	0.7	1.2	2.7	0.08	<0.01	0.52	<0.2	49	9.206	21.10	0.672	0.20	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	<0.01
CK-A-29	<1	<0.1	3480	1330	214	6420	2060	1380	4	3.3	1.69	1.3	58.1	912	0.602	0.7	4.5	4.7	0.07	<0.01	0.51	<0.2	42	6.355	15.11	0.594	0.17	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	<0.01
CK-A-30	<1	<0.1	4350	2780	279	8250	2680	2940	6	7.9	1.15	1.9	22.4	373	0.299	0.9	1.3	2.0	0.08	<0.01	0.21	<0.2	45	7.580	20.67	0.358	0.17	0.014	<0.1	<0.01	<0.01	<0.01	<0.01	<0.01
CK-A-31	<1	<0.1	3790	1890	112	7530	2420	2110	5	4.1	0.87	1.4	97.7	576	0.838	0.9	1.4	3.0	0.05	<0.01	0.34	<0.2	41	7.299	19.35	0.288	0.09	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	<0.01
CK-A-32	1	0.1	3070	918	59	7450	2430	573	5	3.5	0.89	1.3	57.4	236	0.384	0.4	0.5	2.3	0.03	<0.01	0.26	<0.2	34	7.661	13.97	0.253	0.06	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	<0.01
CK-A-33	<1	<0.1	2840	1660	130	4730	1090	1890	4	2.4	0.97	1.1	34.8	1230	0.522	0.6	2.8	4.3	0.05	<0.01	0.80	<0.2	39	4.015	21.00	0.341	0.13	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	<0.01
CK-A-34	<1	<0.1	5200	1500	75	7630	2380	1490	5	3.7	0.76	1.3	26.7	230	0.261	0.4	0.8	2.3	0.03	<0.01	0.24	<0.2	45	6.081	16.58	0.211	0.11	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	<0.01
CK-A-35	<1	<0.1	3410	1410	101	2960	1810	988	2	1.8	1.12	1.1	44.5	688	0.251	<0.3	0.7	1.9	0.04	<0.01	0.60	<0.2	44	5.591	17.80	0.324	0.14	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	<0.01
CK-A-36	2	<0.1	5230	1810	109	10100	2440	2540	7	4.6	1.03	1.7	67.7	720	0.643	0.6	1.1	3.1	0.05	<0.01	0.49	<0.2	55	6.463	21.23	0.314	0.13	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	<0.01
CK-A-37	<1	<0.1	3280	1130	54	5840	1580	1280	4	2.7	0.54	1.1	62.7	474	0.326	0.4	0.4																	

Geochemical grade assay of stream water collected in the Lomba Grande district

Sample No.	Sn	Sb	Te	I	Cs	Ba	La	Ce	Pr	Nd	Sm	Eu	Gd	Tb	Dy	Ho	Er	Tm	Yb	Lu	Hf	Ta	W	Re	Os	Pt	Au	Hg	Tl	Pb	Bi	Th
CK-A-01	33.9	0.39	-0.01	8	0.028	57.4	9.335	1.113	0.248	0.937	3.539	0.033	0.283	0.039	0.142	0.262	0.068	0.007	0.045	0.007	0.068	<0.001	0.08	<0.001	<0.002	<0.01	<0.002	<0.01	0.029	44.5	0.01	0.057
CK-A-02	<0.1	0.07	<0.01	13	0.038	43.0	0.278	0.718	0.833	0.992	0.024	0.084	0.084	0.009	0.043	0.008	0.021	0.003	0.022	0.005	0.003	<0.001	<0.02	<0.001	<0.002	<0.01	<0.002	<0.01	0.015	0.8	0.01	0.046
CK-A-03	<0.1	0.10	<0.01	7	0.032	44.9	0.279	0.660	0.078	0.330	1.105	0.023	0.099	0.012	0.052	0.111	0.030	0.004	0.032	0.005	0.004	<0.001	<0.02	<0.001	<0.002	<0.01	<0.002	<0.02	0.020	1.0	0.01	0.039
CK-A-04	<0.1	0.12	<0.01	10	0.035	47.8	0.133	0.314	0.039	0.173	0.052	0.023	0.050	0.007	0.033	0.011	0.039	0.004	0.012	0.005	0.005	<0.001	<0.02	<0.001	<0.002	<0.01	<0.002	<0.02	0.020	0.3	0.01	0.029
CK-A-05	<0.1	0.03	<0.02	12	0.034	55.8	0.162	0.422	0.056	0.219	0.068	0.026	0.074	0.008	0.039	0.007	0.022	0.003	0.023	0.003	<0.002	<0.001	<0.02	<0.001	<0.002	<0.01	<0.002	<0.02	0.042	3.6	<0.01	0.033
CK-A-06	<0.1	0.04	<0.02	11	0.048	53.8	0.073	0.192	0.023	0.109	0.038	0.044	0.039	0.005	0.031	0.006	0.021	0.003	0.022	0.003	0.002	<0.001	<0.02	<0.001	<0.002	<0.01	<0.002	<0.02	0.036	0.7	0.01	0.029
CK-A-07	<0.1	0.07	<0.01	19	0.026	26.9	0.131	0.283	0.039	0.153	0.041	0.016	0.042	0.007	0.029	0.005	0.016	0.003	0.017	0.003	0.002	<0.001	<0.02	<0.001	<0.002	<0.01	<0.002	<0.02	0.014	0.3	<0.01	0.018
CK-A-08	<0.1	0.05	<0.01	7	0.055	40.3	0.108	0.264	0.030	0.121	0.047	0.016	0.042	0.007	0.026	0.006	0.017	0.002	0.017	0.003	<0.002	<0.001	<0.02	<0.001	<0.002	<0.01	<0.002	<0.02	0.019	0.5	<0.01	0.017
CK-A-09	<0.1	0.05	<0.01	7	0.131	51.9	0.203	0.484	0.060	0.263	0.075	0.028	0.088	0.015	0.103	0.027	0.075	0.009	0.078	0.010	0.014	<0.001	<0.02	<0.001	<0.002	<0.01	<0.002	<0.02	0.061	0.4	<0.01	0.014
CK-A-10	<0.1	0.04	<0.01	5	0.114	36.6	0.080	0.178	0.025	0.106	0.044	0.017	0.055	0.010	0.052	0.100	0.030	0.004	0.032	0.005	0.004	<0.001	<0.02	<0.001	<0.002	<0.01	<0.002	<0.02	0.032	0.2	<0.01	0.009
CK-A-11	<0.1	0.05	<0.01	3	0.050	25.5	0.131	0.045	0.182	0.050	0.051	0.047	0.066	0.025	0.060	0.010	0.030	0.003	0.018	0.003	0.003	<0.001	<0.02	<0.001	<0.002	<0.01	<0.002	<0.02	0.021	0.3	<0.01	0.015
CK-A-12	<0.1	0.05	<0.01	3	0.041	28.2	0.084	0.212	0.026	0.104	0.031	0.011	0.032	0.005	0.017	0.004	0.013	0.002	0.013	0.002	0.003	<0.001	<0.02	<0.001	<0.002	<0.01	<0.002	<0.02	0.019	0.2	<0.01	0.011
CK-A-13	<0.1	0.07	<0.02	8	0.043	41.7	0.164	0.398	0.048	0.186	0.056	0.023	0.050	0.007	0.036	0.007	0.021	0.002	0.016	0.003	0.004	<0.001	<0.02	<0.001	<0.002	<0.01	<0.002	<0.02	0.017	0.3	<0.01	0.026
CK-A-14	<0.1	0.08	<0.02	7	0.032	76.6	0.075	0.202	0.022	0.094	0.033	0.022	0.041	0.005	0.025	0.005	0.017	0.002	0.012	0.002	<0.002	<0.001	<0.02	<0.001	<0.002	<0.01	<0.002	<0.02	0.028	0.3	<0.01	0.021
CK-A-15	<0.1	0.07	<0.01	10	0.027	39.1	0.106	0.275	0.029	0.112	0.032	0.014	0.034	0.004	0.018	0.004	0.011	0.002	0.010	0.002	<0.002	<0.001	<0.02	<0.001	<0.002	<0.01	<0.002	<0.02	0.008	0.2	<0.01	0.017
CK-A-16	<0.1	0.07	<0.01	7	0.028	50.3	0.120	0.299	0.032	0.137	0.043	0.017	0.038	0.006	0.025	0.005	0.012	0.002	0.016	0.002	<0.002	<0.001	<0.02	<0.001	<0.002	<0.01	<0.002	<0.02	0.016	0.2	<0.01	0.020
CK-A-17	<0.1	0.09	<0.01	8	0.025	41.0	0.171	0.385	0.047	0.200	0.064	0.021	0.062	0.009	0.035	0.008	0.024	0.003	0.020	0.003	0.003	<0.001	<0.02	<0.001	<0.002	<0.01	<0.002	<0.02	0.017	0.4	<0.01	0.027
CK-A-18	<0.1	0.08	<0.01	6	0.034	42.1	0.136	0.352	0.039	0.172	0.060	0.023	0.063	0.009	0.040	0.009	0.022	0.003	0.025	0.003	0.004	<0.001	<0.02	<0.001	<0.002	<0.01	<0.002	<0.02	0.028	0.3	<0.01	0.022
CK-A-19	<0.1	0.09	<0.01	5	0.019	59.6	0.084	0.220	0.029	0.108	0.038	0.021	0.038	0.004	0.024	0.005	0.015	0.002	0.015	0.003	<0.002	<0.001	<0.02	<0.001	<0.002	<0.01	<0.002	<0.02	0.016	0.2	<0.01	0.020
CK-A-20	<0.1	0.07	<0.01	3	0.020	43.0	0.121	0.294	0.037	0.170	0.062	0.019	0.053	0.006	0.031	0.006	0.017	0.002	0.019	0.003	0.005	<0.001	<0.02	<0.001	<0.002	<0.01	<0.002	<0.02	0.016	0.4	<0.01	0.019
CK-A-21	<0.1	0.07	<0.01	5	0.042	61.2	0.215	0.561	0.067	0.289	0.100	0.037	0.098	0.012	0.059	0.012	0.031	0.004	0.032	0.005	0.005	<0.001	<0.02	<0.001	<0.002	<0.01	<0.002	<0.02	0.026	0.4	<0.01	0.039
CK-A-22	<0.1	0.09	<0.03	8	0.043	49.7	0.454	1.140	0.134	0.549	0.178	0.052	0.177	0.019	0.092	0.018	0.053	0.006	0.047	0.006	0.012	<0.001	<0.02	<0.001	<0.002	<0.01	<0.002	<0.02	0.023	1.1	<0.01	0.072
CK-A-23	<0.1	0.07	<0.01	6	0.048	52.1	0.141	0.373	0.040	0.171	0.058	0.021	0.049	0.008	0.043	0.008	0.021	0.003	0.023	0.003	0.005	<0.001	<0.02	<0.001	<0.002	<0.01	<0.002	<0.02	0.029	0.4	<0.01	0.038
CK-A-24	<0.1	0.06	<0.01	4	0.032	30.6	0.099	0.260	0.029	0.121	0.047	0.021	0.049	0.008	0.043	0.008	0.024	0.003	0.022	0.004	0.009	<0.001	<0.02	<0.001	<0.002	<0.01	<0.002	<0.02	0.017	0.3	<0.01	0.027
CK-A-25	<0.1	0.07	<0.01	4	0.054	47.1	0.410	1.005	0.120	0.512	0.166	0.048	0.157	0.021	0.108	0.019	0.054	0.007	0.051	0.007	0.007	<0.001	<0.02	<0.001	<0.002	<0.01	<0.002	<0.02	0.030	0.8	<0.01	0.047
CK-A-26	<0.1	0.10	<0.01	4	0.062	35.3	0.121	0.305	0.042	0.175	0.068	0.028	0.077	0.012	0.065	0.012	0.040	0.005	0.037	0.006	0.007	<0.001	<0.02	<0.001	<0.002	<0.01	<0.002	<0.02	0.027	0.3	<0.01	0.012
CK-A-27	<0.1	0.06	<0.02	4	0.116	47.4	0.974	2.241	0.261	0.956	0.267	0.068	0.237	0.029	0.123	0.023	0.063	0.007	0.055	0.008	0.011	<0.001	<0.02	<0.001	<0.002	<0.01	<0.002	<0.02	0.039	1.0	<0.01	0.041
CK-A-28	<0.1	0.06	<0.01	6	0.063	43.2	0.533	1.348	0.151	0.585	0.189	0.050	0.181	0.022	0.101	0.020	0.052	0.006	0.044	0.008	0.007	<0.001	<0.02	<0.001	<0.002	<0.01	<0.002	<0.02	0.033	1.0	<0.01	0.058
CK-A-29	<0.1	0.07	<0.01	6	0.072	51.4	0.199	0.508	0.059	0.265	0.094	0.034	0.090	0.012	0.061	0.012	0.031	0.004	0.036	0.003	0.008	<0.001	<0.02	<0.001	<0.002	<0.01	<0.002	<0.02	0.028	0.4	<0.01	0.033
CK-A-30	<0.1	0.06	<0.01	5	0.078	51.7	0.221	0.558	0.062	0.250	0.079	0.030	0.083	0.010	0.045	0.010	0.027	0.003	0.027	0.004	0.004	<0.001	<0.02	<0.001	<0.002	<0.01	<0.002	<0.02	0.036	0.5	<0.01	0.026
CK-A-31	<0.1	0.05	<0.01	3	0.070	58.7	0.120	0.350	0.039	0.162	0.056	0.026	0.062	0.008	0.042	0.008	0.023	0.003	0.020	0.003	0.003	<0.001	<0.02	<0.001	<0.002	<0.01	<0.002	<0.02	0.044	0.3	<0.01	0.014
CK-A-32	<0.1	0.05	<0.01	7	0.060	31.2	0.455	1.048	0.130	0.517	0.143	0.037	0.130	0.015	0.059	0.012	0.030	0.002	0.031	0.005	0.007	<0.001	<0.02	<0.001	<0.002	<0.01	<0.002	<0.02	0.017	0.6	<0.01	0.059
CK-A-33	<0.1	0.06	<0.01	4	0.038	41.2	0.130	0.371	0.040	0.160	0.060	0.020	0.056	0.007	0.033	0.006	0.022	0.002	0.023	0.004	0.003	<0.001	<0.02	<0.001	<0.002	<0.01	<0.002	<0.02	0.024	0.3	<0.01	0.024
CK-A-34	<0.1	0.09	<0.01	4	0.069	31.4	0.472	1.008	0.127	0.503	0.146	0.039	0.126	0.015	0.063	0.011	0.034	0.004	0.028	0.004	0.005	<0.001	<0.02	<0.001	<0.002	<0.01	<0.002	<0.02	0.017	0.6	0.01	0.044
CK-A-35	<0.1	0.07	<0.01	5	0.063	56.2	0.254	0.701	0.080	0.330	0.096	0.034	0.094	0.011	0.053	0.010	0.033	0.004	0.034	0.005	0.005	<0.001	<0.02	<0.001	<0.002	<0.01	<0.002	<0.02	0.028	0.4	<0.01	0.042
CK-A-36	<0.1	0.04	<0.01	7	0.030	37.4	0.124	0.313	0.037	0.156	0.040																					

Geochemical grade assay of stream water collected in the Lomba Grande district

Sample No.	U	SO ₄
CK-A-01	0.040	14
CK-A-02	0.023	9
CK-A-03	0.024	16
CK-A-04	0.014	15
CK-A-05	0.013	15
CK-A-06	0.025	26
CK-A-07	0.012	18
CK-A-08	0.022	52
CK-A-09	0.013	18
CK-A-10	0.020	11
CK-A-11	0.010	12
CK-A-12	0.013	14
CK-A-13	0.010	15
CK-A-14	0.016	15
CK-A-15	0.010	17
CK-A-16	0.011	11
CK-A-17	0.012	18
CK-A-18	0.015	16
CK-A-19	0.014	18
CK-A-20	0.018	18
CK-A-21	0.014	16
CK-A-22	0.027	15
CK-A-23	0.043	14
CK-A-24	0.017	16
CK-A-25	0.021	19
CK-A-26	0.034	12
CK-A-27	0.020	14
CK-A-28	0.081	12
CK-A-29	0.046	13
CK-A-30	0.020	17
CK-A-31	0.018	13
CK-A-32	0.017	14
CK-A-33	0.039	9
CK-A-34	0.014	20
CK-A-35	0.039	9
CK-A-36	0.024	19
CK-A-37	0.013	16
CK-A-38	0.015	17
CK-A-39	0.026	14
CK-A-40	0.041	15
CK-A-41	0.009	15
CK-A-42	0.010	16
CK-A-43	0.017	17
CK-A-44	0.010	17
CK-A-45	0.010	14
CK-A-46	0.006	15
CK-A-47	0.011	14
CK-A-48	0.005	17
CK-A-49	0.008	10
CK-A-50	0.007	12
CK-A-51	0.007	11
CK-A-52	0.005	11
CK-A-53	0.012	10
CK-A-54	0.012	9
CK-A-55	0.010	14
CK-A-56	0.020	11
CK-A-57	0.020	13
CK-A-58	0.015	10
CK-A-59	0.013	11
CK-A-60	0.015	9
CK-A-61	0.017	10

Geochemical grade assay of stream water collected in the Lomba Grande district

Sample No.	Li	Be	Na	Mg	Al	Si	K	Ca	Sc	Ti	V	Cr	Mn	Fe	Co	Ni	Cu	Zn	Ga	Ge	As	Se	Br	Rb	Sr	Y	Zr	Nb	Mo	Ru	Pd	Ag	Cd	In
CK-A-62	<1	<0.1	2050	669	79	3570	1200	563	3	2.8	0.39	0.8	27.4	294	0.184	<0.3	0.4	1.6	0.03	0.01	0.27	<0.2	32	4.428	11.49	0.163	0.08	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	
CK-A-63	<1	<0.1	2030	769	98	4470	1080	352	3	3.6	0.58	0.9	10.4	170	0.101	<0.3	0.5	1.1	0.03	<0.01	0.20	<0.2	32	4.382	12.20	0.203	0.10	0.005	<0.1	<0.01	<0.01	<0.01	<0.01	
CK-A-64	<1	<0.1	1970	648	66	3660	1220	401	3	2.4	0.40	0.9	30.6	253	0.169	<0.3	0.3	0.3	0.02	<0.01	0.20	<0.2	31	4.542	10.27	0.220	0.07	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	
CK-A-65	<1	<0.1	2020	630	50	3190	1030	226	3	2.3	0.23	0.8	9.1	53	0.072	0.6	0.9	3.3	0.02	<0.01	0.08	<0.2	29	3.047	9.88	0.363	0.04	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	
CK-A-66	<1	<0.1	1850	537	43	3710	1090	106	3	2.2	0.23	0.8	10.5	95	0.079	<0.3	0.4	0.4	0.01	<0.01	0.12	<0.2	29	3.972	8.65	0.376	0.09	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	
CK-A-67	<1	<0.1	2280	661	57	3760	1330	464	3	2.4	0.40	0.9	37.1	269	0.210	<0.3	0.6	0.6	0.02	<0.01	0.26	<0.2	37	4.539	10.35	0.171	0.08	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	
CK-A-68	<1	<0.1	2300	658	62	4830	1560	486	3	2.5	0.41	0.8	23.1	261	0.160	<0.3	0.5	1.6	0.02	<0.01	0.28	<0.2	37	5.379	10.76	0.211	0.07	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	
CK-A-69	<1	<0.1	2300	665	87	6110	1980	1170	4	3.3	0.58	1.1	7.8	67	0.095	<0.3	0.5	1.7	0.03	<0.01	0.19	<0.2	32	6.707	9.65	0.319	0.04	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	
CK-A-70	<1	<0.1	3460	765	81	6840	1580	1170	5	3.5	0.61	0.8	8.3	112	0.090	<0.3	0.4	1.1	0.02	<0.01	0.18	<0.2	32	5.045	12.09	0.178	0.06	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	
CK-A-71	<1	<0.1	3880	1160	28	6790	1280	1920	4	3.3	0.17	0.8	122.0	394	0.452	<0.3	1.2	2.6	0.04	<0.01	0.31	<0.2	37	3.778	18.23	0.122	0.05	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	
CK-A-72	<1	<0.1	4830	1090	66	10300	2380	2300	7	5.0	0.58	1.3	101.9	404	0.467	<0.4	0.3	2.4	0.05	<0.02	0.35	<0.2	34	6.153	16.03	0.192	0.05	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	
CK-A-73	<1	<0.1	5120	2020	63	7230	1700	3360	5	3.4	0.44	1.0	87.1	878	0.439	0.5	0.5	1.7	0.05	<0.01	0.52	<0.2	59	3.855	30.59	0.195	0.09	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	
CK-A-74	<1	<0.1	4960	1570	89	9060	1870	3910	6	4.4	0.58	1.3	177.6	812	0.385	0.4	0.5	1.6	0.07	0.01	0.45	<0.2	51	5.543	25.29	0.201	0.09	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	
CK-A-75	<1	<0.1	5390	1050	65	10200	2270	1200	6	4.5	0.89	0.9	23.7	237	0.290	0.5	0.8	2.4	0.03	<0.01	0.21	<0.2	39	5.896	15.07	0.164	0.04	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	
CK-A-76	<1	<0.1	6290	2020	69	9260	2680	6200	6	4.9	0.64	0.9	127.2	539	0.454	0.3	1.3	1.3	0.05	<0.01	0.38	<0.2	44	5.293	29.99	0.191	0.06	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	
CK-A-77	<1	<0.1	5290	1070	28	9170	1330	1750	6	3.9	1.21	1.0	11.9	613	0.098	0.3	0.5	1.3	0.01	0.01	0.22	<0.2	43	3.440	14.91	0.164	0.06	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	
CK-A-78	<1	<0.1	5420	1430	39	9030	1560	3770	6	3.8	0.56	0.9	191.8	316	0.478	0.4	0.4	1.2	0.06	<0.01	0.38	<0.2	45	4.079	23.55	0.134	0.06	0.020	<0.1	<0.01	<0.01	<0.01	<0.01	
CK-A-79	<1	<0.1	5810	1380	80	10200	2220	3340	7	4.9	0.63	1.1	116.3	696	0.444	0.4	1.6	1.7	0.06	0.02	0.40	<0.2	45	5.530	22.06	0.197	0.36	0.066	<0.1	<0.01	<0.01	<0.01	<0.01	
CK-A-80	<1	<0.1	6270	2080	47	9180	1810	5640	6	4.8	0.62	1.3	110.0	619	0.318	0.4	0.4	1.1	0.05	0.01	0.43	<0.2	46	3.815	29.61	0.158	0.07	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	
CK-A-81	<1	<0.1	5670	1700	32	8660	2060	3080	6	4.0	0.67	0.9	60.1	188	0.338	0.4	1.7	2.2	0.03	<0.01	0.25	<0.2	39	5.778	24.09	0.112	0.03	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	
CK-A-82	<1	<0.1	2470	718	106	5150	1590	733	3	2.8	0.63	1.0	92.3	564	0.431	0.4	1.3	2.7	0.05	<0.01	0.32	<0.2	39	5.331	24.67	0.277	0.09	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	
CK-A-83	<1	<0.1	5110	1480	58	9240	1870	4480	7	5.3	0.56	1.2	99.0	603	0.320	0.4	1.1	2.0	0.04	0.01	0.42	<0.2	42	4.805	24.65	0.165	0.07	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	
CK-A-84	<1	<0.1	5740	1930	46	9220	1840	4160	7	4.5	0.60	1.0	131.9	572	0.415	0.4	0.6	1.5	0.04	<0.01	0.42	<0.2	44	5.039	27.15	0.155	0.05	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	
CK-A-85	<1	<0.1	3680	1020	48	7240	2440	1670	5	3.4	0.68	1.0	12.6	41	0.100	0.4	1.0	4.0	0.01	<0.01	0.15	<0.2	36	8.614	16.23	0.176	0.03	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	
CK-A-86	<1	<0.1	3600	1260	55	6250	1510	3730	5	3.4	0.75	1.0	15.8	119	0.107	<0.3	0.3	1.0	0.02	<0.01	0.25	<0.2	43	4.006	17.97	0.165	0.07	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	
CK-A-87	<1	<0.1	4270	1730	78	9760	1740	2280	7	5.3	0.75	2.2	26.2	294	0.213	0.8	0.4	1.1	0.03	<0.01	0.20	<0.2	39	5.047	16.36	0.217	0.08	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	
CK-A-88	<1	<0.1	3040	726	51	6950	2270	579	5	3.8	0.57	0.9	20.5	180	0.261	0.4	0.4	1.1	0.02	<0.01	0.20	<0.2	36	6.998	10.21	0.221	0.06	<0.006	<0.1	<0.01	<0.01	<0.01	<0.01	
CK-A-89	<1	<0.1	4650	1050	153	10100	2430	1600	7	6.2	1.00	1.0	38.5	529	0.453	0.5	1.0	2.3	0.05	0.01	0.44	<0.2	42	7.517	17.89	0.451	0.15	0.007	<0.1	<0.01	<0.01	<0.01	<0.01	
CK-A-90	<1	<0.1	30200	1080	126	9160	12400	1850	6	12.5	0.75	1.3	44.8	414	0.404	<0.3	2.6	6.7	0.17	<0.01	0.40	<0.2	45	5.276	18.53	0.227	0.12	0.018	0.1	<0.01	<0.01	<0.01	<0.01	
CK-A-92	<1	<0.1	5180	1580	89	8470	2010	3350	5	7.8	0.66	0.9	93.4	910	0.602	0.5	0.6	2.4	0.05	<0.01	0.65	<0.2	52	6.330	23.40	0.283	0.16	0.007	<0.1	<0.01	<0.01	<0.01	<0.01	
CK-A-93	<1	<0.1	3740	1030	86	7430	1350	1420	5	4.8	0.54	0.9	62.8	493	0.446	0.4	0.6	2.4	0.04	<0.01	0.38	<0.2	39	4.320	15.76	0.234	0.10	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	
CK-A-94	<1	<0.1	6950	2230	142	2800	2370	4050	3	5.0	1.33	0.9	69.5	1250	0.491	0.6	1.2	2.7	0.06	0.01	1.07	<0.2	56	5.672	32.51	0.392	0.31	0.008	<0.1	<0.01	<0.01	<0.01	<0.01	
CK-A-95	<1	<0.1	5120	2150	186	4700	2140	3900	4	8.0	1.50	1.1	38.3	1040	0.424	0.7	1.3	2.3	0.07	<0.01	0.90	<0.2	52	4.281	30.78	0.453	0.40	0.014	<0.1	<0.01	<0.01	<0.01	<0.01	
CK-A-96	<1	<0.1	4140	1540	88	2870	696	2040	3	3.3	0.64	0.9	104.9	1370	0.463	<0.3	0.5	11.9	0.05	<0.01	0.59	<0.2	71	1.946	24.07	0.272	0.19	0.005	<0.1	<0.01	<0.01	<0.01	<0.01	
CK-A-100	<1	<0.1	2160	493	134	2940	883	468	3	2.7	0.51	0.9	41.3	359	0.212	0.4	0.7	2.3	0.04	<0.01	0.28	<0.2	23	2.905	7.83	0.237	0.22	0.009	<0.1	<0.01	<0.01	<0.01	<0.01	
CK-A-101	<1	<0.1	1930	384	111	2550	476	<50	2	3.7	0.51	0.7	32.3	201	0.227	0.4	0.7	2.8	0.03	<0.01	0.21	<0.2	25	1.985	5.99	0.273	0.23	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	
CK-A-102	<1	<0.1	3290	1770	339	7460	960	2870	6	12.2	2.24	1.3	34.0	579	0.412	0.6	1.1	4.3	0.10	0.01	0.16	<0.2	28	3.453	17.98	0.507	0.34	0.018	<0.1	<0.01	<0.01	<0.01	<0.01	
CK-A-103	<1	<0.1	3400	1620	345	7660	833	2480	6	11.1	2.23	1.5	23.4	453	0.381	0.5	1.5	6.2	0.09	0.01	0.14	<0.2	27	2.821	16.74	0.527	0.32	0.018	&					

Geochemical grade assay of stream water collected in the Lomba Grande district

Sample No.	Sn	Sb	Te	I	Cs	Ba	La	Ce	Pr	Nd	Sm	Eu	Gd	Tb	Dy	Ho	Er	Tm	Yb	Lu	Hf	Ta	W	Re	Os	Pt	Au	Hg	Tl	Pb	Bi	Th
CK-A-62	<-0.1	0.01	<-0.01	3	0.062	28.5	1.104	2.63	0.029	0.129	0.033	0.04	0.041	0.005	0.030	0.006	0.017	0.002	0.021	0.003	0.004	<-0.001	<-0.02	<-0.001	<-0.002	<-0.01	<-0.002	<-0.2	0.019	0.3	<-0.1	0.018
CK-A-63	<-0.1	0.01	<-0.01	2	0.082	30.5	1.146	3.76	0.057	0.189	0.057	0.022	0.061	0.008	0.033	0.006	0.021	0.003	0.018	0.003	0.004	<-0.001	<-0.02	<-0.001	<-0.002	<-0.01	<-0.002	<-0.2	0.023	0.3	<-0.1	0.013
CK-A-64	<-0.1	<-0.01	<-0.01	3	0.065	29.1	0.992	2.53	0.030	0.122	0.045	0.015	0.049	0.006	0.034	0.006	0.021	0.003	0.024	0.003	0.004	<-0.001	<-0.02	<-0.001	<-0.002	<-0.01	<-0.002	<-0.2	0.022	0.2	<-0.1	0.013
CK-A-65	<-0.1	<-0.01	<-0.01	3	0.061	26.1	0.994	2.06	0.032	0.141	0.054	0.021	0.073	0.010	0.053	0.013	0.038	0.004	0.042	0.006	0.003	<-0.001	<-0.02	<-0.001	<-0.002	<-0.01	<-0.002	<-0.2	0.024	0.3	<-0.1	0.006
CK-A-66	<-0.1	0.02	<-0.01	7	0.057	28.9	0.966	1.73	0.022	0.094	0.029	0.04	0.045	0.007	0.031	0.005	0.022	0.003	0.021	0.003	0.004	<-0.001	<-0.02	<-0.001	<-0.002	<-0.01	<-0.002	<-0.2	0.019	0.3	<-0.1	0.007
CK-A-67	<-0.1	0.02	<-0.01	4	0.058	24.5	0.998	2.76	0.030	0.121	0.035	0.015	0.040	0.006	0.029	0.005	0.017	0.002	0.019	0.003	0.003	<-0.001	<-0.02	<-0.001	<-0.002	<-0.01	<-0.002	<-0.2	0.023	0.2	<-0.1	0.016
CK-A-68	<-0.1	0.01	<-0.01	4	0.046	26.7	1.106	2.84	0.030	0.138	0.038	0.016	0.041	0.006	0.034	0.006	0.021	0.002	0.019	0.003	0.002	<-0.001	<-0.02	<-0.001	<-0.002	<-0.01	<-0.002	<-0.2	0.019	0.3	<-0.1	0.016
CK-A-69	<-0.1	<-0.01	<-0.01	3	0.088	32.8	1.131	2.93	0.038	0.185	0.054	0.020	0.066	0.011	0.053	0.011	0.032	0.002	0.036	0.006	0.002	<-0.001	<-0.02	<-0.001	<-0.002	<-0.01	<-0.002	<-0.2	0.028	0.3	<-0.1	0.007
CK-A-70	<-0.1	0.05	<-0.01	3	0.034	26.6	0.996	2.16	0.028	0.112	0.035	0.014	0.042	0.006	0.032	0.006	0.018	0.002	0.020	0.003	0.002	<-0.001	<-0.02	<-0.001	<-0.002	<-0.01	<-0.002	<-0.2	0.014	0.2	<-0.1	0.010
CK-A-71	<-0.1	0.01	<-0.01	5	0.014	34.6	0.949	1.16	0.015	0.062	0.019	0.010	0.023	0.003	0.012	0.003	0.009	<-0.001	0.008	0.002	0.003	<-0.001	<-0.02	<-0.001	<-0.002	<-0.01	<-0.002	<-0.2	0.009	0.2	<-0.1	0.012
CK-A-72	<-0.1	0.02	<-0.01	7	0.035	43.0	0.995	2.52	0.026	0.112	0.032	0.014	0.035	0.004	0.022	0.004	0.012	0.001	0.012	0.003	0.003	<-0.001	<-0.02	<-0.001	<-0.002	<-0.01	<-0.002	<-0.2	0.016	0.2	<-0.1	0.019
CK-A-73	0.4	0.02	<-0.01	10	0.020	66.3	1.161	3.97	0.049	0.203	0.054	0.023	0.057	0.007	0.043	0.006	0.016	0.002	0.019	0.003	0.005	<-0.001	<-0.02	<-0.001	<-0.002	<-0.01	<-0.002	<-0.2	0.013	0.8	<-0.1	0.035
CK-A-74	<-0.1	0.02	<-0.01	12	0.042	42.5	1.136	3.58	0.040	0.185	0.046	0.019	0.049	0.007	0.030	0.007	0.017	0.003	0.021	0.003	0.004	<-0.001	<-0.02	<-0.001	<-0.002	<-0.01	<-0.002	<-0.2	0.016	0.3	<-0.1	0.036
CK-A-75	<-0.1	0.01	<-0.01	11	0.025	45.0	1.107	2.86	0.033	0.144	0.049	0.019	0.046	0.006	0.028	0.005	0.014	0.003	0.016	0.002	0.002	<-0.001	<-0.02	<-0.001	<-0.002	<-0.01	<-0.002	<-0.2	0.023	0.4	<-0.1	0.018
CK-A-76	<-0.1	0.02	<-0.01	9	0.021	66.1	1.106	2.85	0.034	0.135	0.043	0.022	0.043	0.006	0.032	0.006	0.015	0.002	0.013	0.002	<-0.002	<-0.001	<-0.02	<-0.001	<-0.002	<-0.01	<-0.002	<-0.2	0.012	0.3	<-0.1	0.027
CK-A-77	<-0.1	<-0.01	<-0.01	5	0.022	33.2	0.941	0.996	0.014	0.077	0.027	0.012	0.027	0.004	0.027	0.004	0.016	0.003	0.019	0.003	0.004	<-0.001	<-0.02	<-0.001	<-0.002	<-0.01	<-0.002	<-0.2	0.011	0.2	<-0.1	0.009
CK-A-78	<-0.1	<-0.01	<-0.01	10	0.013	47.0	0.954	1.31	0.017	0.076	0.023	0.012	0.029	0.004	0.018	0.004	0.010	0.001	0.014	0.002	0.002	<-0.001	<-0.02	<-0.001	<-0.002	<-0.01	<-0.002	<-0.2	0.013	0.2	<-0.1	0.019
CK-A-79	<-0.1	0.04	<-0.01	9	0.032	51.8	1.149	3.75	0.044	0.169	0.055	0.020	0.052	0.007	0.029	0.005	0.018	0.002	0.017	0.003	0.010	<-0.001	<-0.02	<-0.001	<-0.002	<-0.01	<-0.002	<-0.2	0.015	0.3	<-0.1	0.031
CK-A-80	<-0.1	0.01	<-0.01	9	0.022	56.9	0.996	2.42	0.028	0.129	0.045	0.019	0.045	0.005	0.027	0.005	0.012	0.002	0.015	0.002	0.005	<-0.001	<-0.02	<-0.001	<-0.002	<-0.01	<-0.002	<-0.2	0.008	0.2	<-0.1	0.025
CK-A-81	<-0.1	<-0.01	<-0.01	5	0.020	49.5	0.960	1.44	0.018	0.076	0.023	0.013	0.027	0.003	0.017	0.004	0.010	<-0.001	0.009	0.001	<-0.002	<-0.001	<-0.02	<-0.001	<-0.002	<-0.01	<-0.002	<-0.2	0.016	0.2	<-0.1	0.012
CK-A-82	<-0.1	<-0.01	<-0.01	5	0.055	28.9	0.971	1.483	0.047	0.196	0.059	0.020	0.070	0.009	0.047	0.010	0.027	0.003	0.026	0.004	0.006	<-0.001	<-0.02	<-0.001	<-0.002	<-0.01	<-0.002	<-0.2	0.022	0.4	<-0.1	0.027
CK-A-83	<-0.1	0.03	<-0.01	9	0.033	43.4	1.119	2.96	0.034	0.155	0.048	0.016	0.046	0.007	0.030	0.006	0.016	0.002	0.014	0.003	0.004	<-0.001	<-0.02	<-0.001	<-0.002	<-0.01	<-0.002	<-0.2	0.016	0.3	<-0.1	0.034
CK-A-84	<-0.1	0.03	<-0.01	6	0.018	51.6	1.102	2.54	0.029	0.119	0.037	0.015	0.033	0.005	0.029	0.006	0.012	0.002	0.016	0.002	0.002	<-0.001	<-0.02	<-0.001	<-0.002	<-0.01	<-0.002	<-0.2	0.012	0.2	0.04	0.024
CK-A-85	<-0.1	0.02	<-0.01	2	0.024	43.6	0.997	2.24	0.025	0.097	0.029	0.015	0.033	0.005	0.027	0.005	0.017	0.002	0.016	0.002	<-0.002	<-0.001	<-0.02	<-0.001	<-0.002	<-0.01	<-0.002	<-0.2	0.023	0.5	<-0.1	0.008
CK-A-86	<-0.1	0.07	<-0.01	5	0.024	37.9	0.983	1.82	0.027	0.126	0.038	0.015	0.039	0.005	0.032	0.006	0.018	0.002	0.015	0.003	0.006	<-0.001	<-0.02	<-0.001	<-0.002	<-0.01	<-0.002	<-0.2	0.012	0.2	<-0.1	0.015
CK-A-87	<-0.1	0.02	<-0.01	6	0.031	40.1	1.137	3.01	0.039	0.163	0.050	0.020	0.051	0.007	0.034	0.006	0.014	0.002	0.020	0.003	0.003	<-0.001	<-0.02	<-0.001	<-0.002	<-0.01	<-0.002	<-0.2	0.018	0.2	<-0.1	0.018
CK-A-88	<-0.1	0.03	<-0.01	3	0.041	42.7	1.102	2.58	0.031	0.121	0.042	0.017	0.052	0.006	0.034	0.006	0.019	0.003	0.020	0.003	0.003	<-0.001	<-0.02	<-0.001	<-0.002	<-0.01	<-0.002	<-0.2	0.031	0.4	<-0.1	0.012
CK-A-89	<-0.1	0.03	<-0.01	5	0.046	50.6	1.038	2.86	0.092	0.399	0.117	0.042	0.116	0.017	0.075	0.014	0.043	0.006	0.040	0.005	0.011	<-0.001	<-0.02	<-0.001	<-0.002	<-0.01	<-0.002	<-0.2	0.033	0.3	<-0.1	0.043
CK-A-90	<-0.1	0.06	<-0.01	7	0.036	37.0	1.151	3.49	0.030	0.211	0.064	0.021	0.061	0.008	0.039	0.007	0.023	0.002	0.022	0.004	0.005	<-0.001	<-0.02	<-0.001	<-0.002	<-0.01	<-0.002	<-0.2	0.022	1.2	<-0.1	0.028
CK-A-91	<-0.1	0.03	<-0.01	10	0.027	37.3	0.987	2.58	0.064	0.144	0.024	0.024	0.083	0.010	0.044	0.010	0.029	0.003	0.021	0.004	0.008	<-0.001	<-0.02	<-0.001	<-0.002	<-0.01	<-0.002	<-0.2	0.014	0.3	<-0.1	0.043
CK-A-92	<-0.1	0.03	<-0.01	7	0.028	30.7	1.132	3.42	0.039	0.176	0.056	0.018	0.058	0.007	0.033	0.007	0.021	0.003	0.021	0.003	0.003	<-0.001	<-0.02	<-0.001	<-0.002	<-0.01	<-0.002	<-0.2	0.016	0.2	<-0.1	0.026
CK-A-93	<-0.1	0.05	<-0.01	14	0.076	50.4	1.181	3.61	0.143	0.597	0.154	0.043	0.139	0.016	0.073	0.014	0.042	0.005	0.033	0.005	0.012	<-0.001	<-0.02	<-0.001	<-0.002	<-0.01	<-0.002	<-0.2	0.022	0.6	<-0.1	0.073
CK-A-94	<-0.1	0.06	<-0.01	11	0.048	49.1																										

Geochemical grade assay of stream water collected in the Lomba Grande district

Sample No.	U	SO ₄
CK-A-62	0.011	12
CK-A-63	0.017	10
CK-A-64	0.011	12
CK-A-65	0.020	13
CK-A-66	0.007	15
CK-A-67	0.010	11
CK-A-68	0.013	15
CK-A-69	0.013	16
CK-A-70	0.010	16
CK-A-71	0.005	9
CK-A-72	0.010	16
CK-A-73	0.014	14
CK-A-74	0.013	14
CK-A-75	0.016	18
CK-A-76	0.016	16
CK-A-77	0.014	18
CK-A-78	0.008	17
CK-A-79	0.015	17
CK-A-80	0.014	17
CK-A-81	0.010	12
CK-A-82	0.020	13
CK-A-83	0.016	15
CK-A-84	0.013	15
CK-A-85	0.008	16
CK-A-86	0.013	16
CK-A-87	0.016	17
CK-A-88	0.010	18
CK-A-89	0.027	14
CK-A-90	0.021	3
CK-A-92	0.017	9
CK-A-93	0.011	13
CK-A-94	0.050	14
CK-A-95	0.055	13
CK-A-96	0.024	15
CK-A-100	0.017	12
CK-A-101	0.018	9
CK-A-102	0.023	11
CK-A-103	0.027	10
CK-A-104	0.028	12
CK-A-105	0.013	10
CK-A-106	0.033	15
CK-A-107	0.041	11
CK-A-108	0.024	15
CK-A-109	0.030	16
CK-A-110	0.035	19
CK-A-111	0.022	16
CK-A-112	0.020	17
CK-A-113	0.052	11
CK-A-114	0.022	17
CK-A-115	0.029	17
CK-A-116	0.028	17
CK-A-117	0.041	13
CK-A-118	0.033	13
CK-A-119	0.028	21
CK-A-120	0.042	14
CK-A-121	0.037	18
CK-A-122	0.044	17
CK-A-123	0.012	16
CK-A-124	0.017	18
CK-A-125	0.041	15
CK-A-126	0.025	15

Geochemical grade assay of stream water collected in the Lombha Grande district

Sample No.	Li	Be	Na	Mg	Al	Si	K	Ca	Sc	Ti	V	Cr	Mn	Fe	Co	Ni	Cu	Zn	Ga	Ge	As	Se	Br	Rb	Sr	Y	Zr	Nb	Mo	Ru	Pd	Ag	Cd	In
CK-A-127	2	0.1	2790	621	97	6780	845	418	6	3.6	0.86	1.1	21.1	217	0.231	0.4	0.5	9.0	0.03	<0.01	0.23	<0.2	29	2.881	7.82	0.326	0.56	<0.005	<0.1	<0.01	<0.01	<0.01	<0.01	
CK-A-128	1	<0.1	3020	983	121	6210	892	945	5	3.3	1.18	0.9	19.5	563	0.293	0.6	1.0	5.7	0.03	<0.01	0.46	<0.2	36	3.481	12.76	0.461	0.19	<0.005	<0.1	<0.01	<0.01	<0.01	<0.001	
CK-A-129	1	<0.1	2840	776	182	5250	978	1120	4	3.3	1.18	1.0	26.9	563	0.325	0.5	1.1	8.0	0.05	<0.01	0.37	<0.2	33	3.275	11.20	0.601	0.22	<0.005	<0.1	<0.01	<0.01	<0.01	<0.001	
CK-A-130	<1	<0.1	2010	772	119	3200	1130	787	3	2.7	0.77	0.9	20.6	402	0.179	<0.3	1.6	7.0	0.03	<0.01	0.32	<0.2	30	4.237	12.46	0.313	0.17	<0.005	<0.1	<0.01	<0.01	<0.01	<0.001	
CK-A-131	<1	<0.1	2080	868	119	2560	1200	878	3	3.8	0.87	0.9	22.1	691	0.199	0.3	1.6	15.7	0.04	<0.01	0.44	<0.2	35	3.615	12.45	0.232	0.19	<0.006	<0.1	<0.01	<0.01	<0.01	<0.001	
CK-A-132	<1	<0.1	2330	997	119	4410	1140	1330	4	2.9	0.93	0.9	46.6	704	0.307	0.3	0.8	17.4	0.04	<0.01	0.39	<0.2	36	3.761	14.15	0.280	0.16	<0.005	<0.1	<0.01	<0.01	<0.01	<0.001	
CK-A-133	<1	<0.1	1590	463	111	3010	659	129	3	2.4	0.50	1.0	8.3	314	0.090	<0.3	0.6	21.7	0.03	<0.01	0.16	<0.2	27	2.775	7.42	0.279	0.19	<0.005	<0.1	<0.01	<0.01	<0.01	<0.001	
CK-A-134	1	<0.1	3080	741	48	5660	654	1000	4	3.1	0.37	1.0	58.7	477	0.263	0.6	2.3	4.0	0.02	<0.01	0.31	<0.2	31	2.184	10.37	0.142	0.14	<0.005	<0.1	<0.01	<0.01	<0.01	<0.001	
CK-A-135	2	<0.1	3020	766	37	5840	687	632	5	2.9	0.35	0.8	40.1	352	0.232	0.6	1.2	4.0	0.02	<0.01	0.33	<0.2	36	2.871	10.69	0.111	0.10	<0.005	<0.1	<0.01	<0.01	<0.01	<0.001	
CK-A-136	<1	<0.1	2260	966	141	2360	326	800	3	3.2	0.51	0.8	12.7	688	0.150	<0.3	0.6	3.5	0.04	<0.01	0.45	<0.2	48	1.512	14.31	0.311	0.21	<0.007	<0.1	<0.01	<0.01	<0.01	<0.001	
CK-A-137	2	<0.1	2970	742	77	6560	707	583	5	3.3	0.79	0.8	17.9	355	0.197	0.3	0.7	3.3	0.02	<0.01	0.34	<0.2	35	2.590	10.00	0.291	0.13	<0.005	<0.1	<0.01	<0.01	<0.01	<0.001	
CK-A-138	1	<0.1	2980	1020	104	5610	937	1410	4	3.2	0.80	0.9	31.3	511	0.280	0.4	0.8	2.6	0.03	<0.01	0.62	<0.2	38	3.367	15.90	0.387	0.18	<0.005	<0.1	<0.01	<0.01	<0.01	<0.001	
CK-A-139	1	<0.1	3190	967	108	5980	889	1240	5	3.5	0.84	0.9	32.9	540	0.297	0.4	1.1	1.7	0.03	<0.01	0.35	<0.2	36	2.883	13.89	0.383	0.16	<0.005	<0.1	<0.01	<0.01	<0.01	<0.001	
CK-A-140	1	<0.1	3210	799	54	5740	656	1090	4	3.0	0.53	0.9	61.0	621	0.273	<0.3	0.7	2.7	0.03	<0.01	0.47	<0.2	35	2.417	12.82	0.184	0.14	<0.005	<0.1	<0.01	<0.01	<0.01	<0.001	
CK-A-141	<1	<0.1	3180	862	63	5310	633	1120	4	3.0	0.57	0.9	58.8	736	0.292	<0.3	0.5	2.8	0.03	<0.01	0.48	<0.2	36	2.447	13.72	0.215	0.18	<0.005	<0.1	<0.01	<0.01	<0.01	<0.001	
CK-A-142	<1	<0.1	3630	1230	179	1370	789	1350	2	5.3	1.41	1.0	29.5	790	0.374	0.5	4.1	7.9	0.05	<0.01	1.04	<0.2	52	2.656	14.57	0.417	0.55	<0.009	<0.1	<0.01	<0.01	<0.01	<0.001	
CK-A-143	1	<0.1	3480	1010	101	7140	1170	1610	6	3.8	0.83	1.0	31.1	454	0.273	0.6	1.5	16.1	0.03	<0.01	0.32	<0.2	35	3.689	13.39	0.324	0.14	<0.005	<0.1	<0.01	<0.01	<0.01	<0.001	
CK-A-144	1	<0.1	3360	1120	145	6840	731	1870	5	4.0	1.07	1.0	24.5	645	0.273	0.7	0.9	19.9	0.05	<0.01	0.50	<0.2	41	2.694	16.43	0.483	0.17	<0.005	<0.1	<0.01	<0.01	<0.01	<0.001	
CK-A-145	<1	<0.1	4250	1290	142	849	944	1320	1	5.0	0.86	1.0	39.6	905	0.412	1.1	2.3	33.3	0.04	<0.01	0.86	<0.2	62	2.780	15.89	0.370	0.51	0.014	<0.1	<0.01	<0.01	<0.01	<0.001	
CK-A-146	4	0.2	4480	1200	196	9720	1270	1900	8	6.2	1.31	1.2	62.0	562	0.678	1.1	1.7	34.3	0.06	0.01	0.36	<0.2	41	3.405	16.02	0.774	0.19	<0.006	<0.1	<0.01	<0.01	<0.01	<0.001	
CK-A-147	4	0.1	2630	705	87	6210	955	582	5	3.3	0.85	0.9	30.3	646	0.356	0.4	0.7	2.9	0.03	<0.01	0.39	<0.2	30	3.741	10.49	0.457	0.14	<0.005	<0.1	<0.01	<0.01	<0.01	<0.001	
CK-A-148	4	0.1	3700	1030	107	8380	1090	1370	6	3.9	1.01	0.9	32.1	555	0.369	0.5	0.6	5.6	0.03	<0.01	0.43	<0.2	35	3.155	14.40	0.479	0.21	<0.005	<0.1	<0.01	<0.01	<0.01	<0.001	
CK-A-149	3	0.1	3460	830	73	8940	1450	520	7	4.2	1.00	1.1	27.8	253	0.560	0.9	0.8	5.5	0.03	<0.01	0.25	<0.2	33	4.778	10.49	0.465	0.08	<0.005	<0.1	<0.01	<0.01	<0.01	<0.001	
CK-A-150	<1	<0.1	3300	1800	67	2830	399	2150	3	2.1	0.88	0.9	84.2	1490	0.307	0.6	4.9	11.6	0.03	<0.01	0.85	<0.2	53	1.784	23.98	0.261	0.21	<0.005	<0.1	<0.01	<0.01	<0.01	<0.001	
CK-A-151	3	<0.1	3440	1120	60	7870	1010	1340	6	3.9	0.79	0.9	31.7	831	0.353	0.6	0.8	13.3	0.03	<0.01	0.52	<0.2	35	3.498	16.44	0.377	0.13	<0.005	<0.1	<0.01	<0.01	<0.01	<0.001	
CK-A-152	<1	<0.1	3000	1550	128	2890	536	1630	3	2.2	1.68	0.9	69.8	1490	0.580	0.6	1.7	19.4	0.05	<0.01	1.35	<0.2	63	2.170	19.27	0.421	0.25	<0.005	<0.1	<0.01	<0.01	<0.01	<0.001	
CK-A-153	<1	<0.1	7030	2060	176	1610	2160	2720	2	9.8	2.87	1.0	71.4	2010	0.714	0.5	11.0	0.07	<0.01	1.66	<0.2	82	5.796	24.75	0.474	0.47	0.018	<0.1	<0.01	<0.01	<0.01	<0.001		
CK-A-154	<1	<0.1	3370	1780	162	3020	1130	2670	3	5.3	1.64	1.1	58.3	914	0.607	0.7	3.5	7.3	0.06	<0.01	0.73	<0.2	60	3.986	19.14	0.441	0.72	0.012	<0.1	<0.01	<0.01	<0.01	<0.001	
CK-A-155	<1	<0.1	4970	2690	215	2450	7430	4580	3	6.7	4.79	1.1	75.3	1170	0.661	0.5	2.6	4.4	0.08	0.01	1.23	<0.2	66	12.786	24.76	0.542	0.83	0.011	<0.1	<0.01	<0.01	<0.01	<0.001	
CK-A-156	<1	<0.1	2910	1220	124	1320	1040	1660	2	3.4	2.91	1.1	34.0	1010	0.300	0.4	1.1	3.7	0.06	<0.01	1.01	<0.2	48	2.816	13.57	0.289	0.51	<0.005	<0.1	<0.01	<0.01	<0.01	<0.001	
CK-A-157	2	<0.1	3110	1250	115	5360	1000	1560	4	3.2	1.15	1.0	60.6	1200	0.556	0.5	0.8	2.7	0.04	<0.01	0.65	<0.2	42	3.515	16.69	0.496	0.23	<0.005	<0.1	<0.01	<0.01	<0.01	<0.001	
CK-A-158	<1	<0.1	3800	1530	103	3700	1320	1950	4	3.0	1.64	1.0	62.6	1690	0.506	0.5	0.9	3.2	0.04	<0.01	0.97	<0.2	57	4.433	20.30	0.487	0.39	<0.006	<0.1	<0.01	<0.01	<0.01	<0.001	
CK-A-159	<1	<0.1	4520	1630	115	4310	798	2020	4	3.7	1.20	1.2	90.1	847	0.739	0.5	1.2	17.8	0.06	<0.01	0.56	<0.2	73	2.778	19.37	0.247	0.37	0.007	<0.1	<0.01	<0.01	<0.01	<0.001	
CK-A-160	<1	<0.1	2820	1260	137	4160	540	1420	4	3.4	1.21	1.5	51.1	1380	0.546	0.9	1.7	48.8	0.05	<0.01	0.55	<0.2	70	1.838	13.88	0.530	0.24	<0.006	<0.1	<0.01	<0.01	<0.01	<0.001	
CK-A-161	<1	<0.1	2810	1230	116	2150	445	1450	2	3.8	1.40	1.2	34.0	1270	0.334	0.4	2.4	55.6	0.04	<0.01	0.67	<0.2	53	1.698	15.37	0.324	1.00	0.008	<0.1	<0.01	<0.01	<0.01	<0.001	
CK-A-162	<1	<0.1	2760	1350	89	2420	560	1780	3	3.3	1.41	1.1	44.5	1220	0.441	0.4	1.2	13.7	0.03	<0.01	0.64	<0.2	54	1.987	17.38	0.403	0.33	<0.005	<0.1	<0.01	<0.01	<0.01	<0.001	
CK-A-163	<1	<0.1	2830	1220	99	1500	826	1590	2	2.2	2.35	1.0	23.0	1110	0.228	1.1	77.4	8.6	0.03	<0.01	0.94	<0.2	50	2.505	14.40	0.288	0.32	<0.005	<0.1	<0.01	<0.01	<0.01	<0.001	
CK-A-164	<1	<0.1	2620	1130	105	2370	697	1350	3	2.2	1.41	0.9	42.1	1170	0.407	0.5	3.2	8.3	0.03	<0.01	0.74	<0.2	46	2.562										

Geochemical grade assay of stream water collected in the Lomba Grande district

Table with 32 columns (Sample No., Sn, Sb, Te, I, Cs, Ba, La, Ce, Pr, Nd, Sm, Eu, Gd, Tb, Dy, Ho, Er, Tm, Yb, Lu, Hf, Ta, W, Re, Os, Pt, Au, Hg, Tl, Pb, Bi, Th) and 42 rows of data for samples CK-A-127 to CK-A-182.

Geochemical grade assay of stream water collected in the Lomba Grande district

Sample No.	U	SO ₄
CK-A-127	0.025	14
CK-A-128	0.039	14
CK-A-129	0.040	13
CK-A-130	0.022	11
CK-A-131	0.024	8
CK-A-132	0.022	10
CK-A-133	0.025	11
CK-A-134	0.010	13
CK-A-135	0.010	13
CK-A-136	0.034	8
CK-A-137	0.021	14
CK-A-138	0.030	13
CK-A-139	0.028	13
CK-A-140	0.015	13
CK-A-141	0.017	9
CK-A-142	0.052	6
CK-A-143	0.021	14
CK-A-144	0.035	13
CK-A-145	0.032	5
CK-A-146	0.039	18
CK-A-147	0.032	13
CK-A-148	0.028	16
CK-A-149	0.025	19
CK-A-150	0.034	9
CK-A-151	0.024	13
CK-A-152	0.056	8
CK-A-153	0.052	4
CK-A-154	0.059	6
CK-A-155	0.090	10
CK-A-156	0.050	6
CK-A-157	0.036	9
CK-A-158	0.047	8
CK-A-159	0.027	9
CK-A-160	0.053	12
CK-A-161	0.040	8
CK-A-162	0.041	7
CK-A-163	0.046	8
CK-A-164	0.050	7
CK-A-165	0.026	8
CK-A-166	0.050	9
CK-A-167	0.038	12
CK-A-168	0.022	13
CK-A-169	0.057	14
CK-A-170	0.022	9
CK-A-171	0.025	9
CK-A-172	0.091	11
CK-A-173	0.111	10
CK-A-174	0.073	14
CK-A-175	0.111	14
CK-A-176	0.118	13
CK-A-177	0.045	10
CK-A-178	0.034	11
CK-A-179	0.026	11
CK-A-180	0.021	12
CK-A-181	0.012	11
CK-A-182	0.014	13

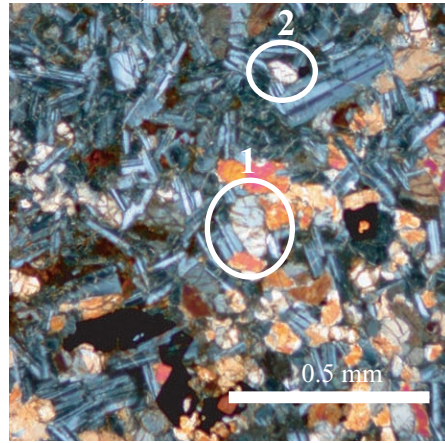
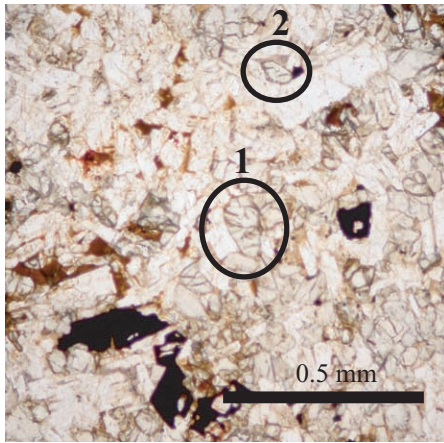
APPENDIX 9

Electron microprobe analysis for olivine in the Paraná basaltic lava and intrusion

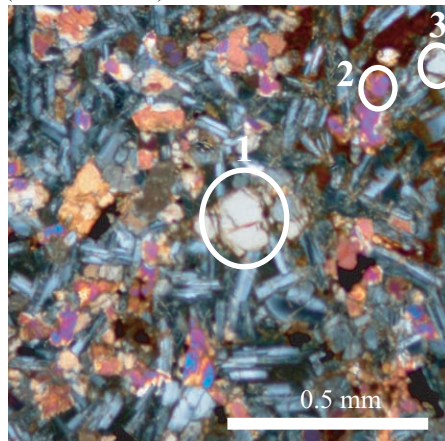
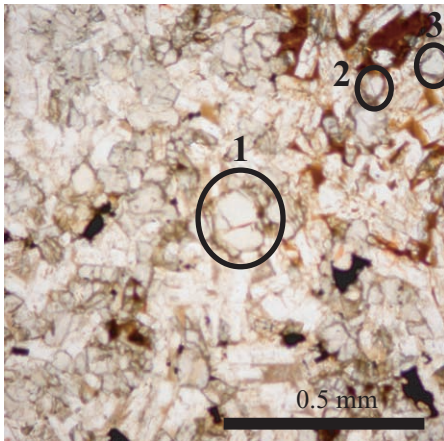
Sample No.	Ni	Mn	Mg	Fe	O	Total
	(wt%)					
AS004A Point9-1-core	0.015	0.389	12.131	18.169	29.334	60.038
AS004A Point9-2-core	0.064	0.431	10.451	19.875	30.979	61.800
AS004A Point10-1-core	0.043	0.379	13.097	16.436	29.911	59.866
AS004A Point10-2-core	—	0.435	12.176	17.497	31.521	61.629
AS004A Point10-3-core	0.035	0.212	10.668	7.993	23.800	42.708
AS021 Point7-1-core	0.020	0.303	13.317	14.772	33.057	61.469
AS021 Point7-2-core	0.035	0.297	14.300	14.185	31.566	60.383
AS021 Point8-1-core	—	0.231	9.953	10.547	22.245	42.976
KN003 Point5-1-core	0.195	0.155	29.989	9.177	38.155	77.671
KN003 Point5-2-core	0.120	0.204	28.709	12.287	36.573	77.893
KN003 Point5-3-core	0.183	0.168	28.117	11.979	38.350	78.797
KN003 Point6-1-core	0.214	0.118	28.472	10.027	41.271	80.102
KN003 Point6-2-core	0.243	0.127	29.746	9.340	39.757	79.213
TG62-226.3 Point1-1-core	0.184	0.211	27.577	13.517	37.950	79.439
TG62-226.3 Point1-1-rim	0.202	0.212	27.988	13.188	38.762	80.352
TG62-226.3 Point1-2-core	0.180	0.217	28.130	14.359	36.683	79.569
TG62-226.3 Point1-2-rim	0.166	0.209	27.015	13.913	39.886	81.189
TG62-226.3 Point2-1-core	0.136	0.172	28.494	11.824	36.054	76.680
TG62-226.3 Point2-1-rim	0.209	0.215	26.668	13.105	35.975	76.172
TG62-226.3 Point2-2-core	0.176	0.224	28.236	13.091	37.747	79.474
TG62-226.3 Point2-2-rim	0.204	0.179	28.291	11.310	35.888	75.872
TG114-289.9 Point3-1-core	0.160	0.231	26.470	14.638	36.858	78.357
TG114-289.9 Point3-1-rim	0.121	0.327	24.348	17.660	38.077	80.533
TG114-289.9 Point3-2-core	0.090	0.346	24.728	19.051	35.981	80.196
TG114-289.9 Point3-2-rim	0.193	0.307	23.273	18.155	39.814	81.742
TG114-289.9 Point4-1-core	0.176	0.224	26.244	14.149	37.686	78.479
TG114-289.9 Point4-1-rim	0.201	0.237	26.816	13.963	37.628	78.845
TG114-289.9 Point4-2-core	0.129	0.290	23.769	17.815	37.130	79.133
TG114-289.9 Point4-2-rim	0.114	0.328	24.446	18.794	36.823	80.505

Measuring Points of EPMA

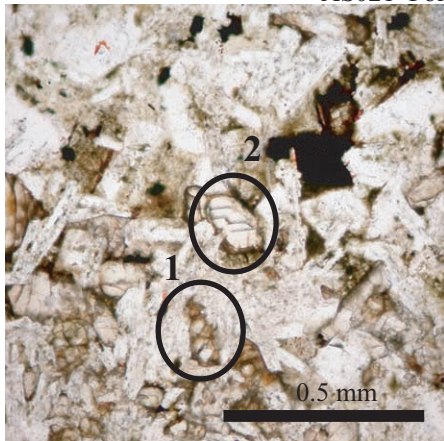
AS004 Point 9 (Lava: Ribeira)



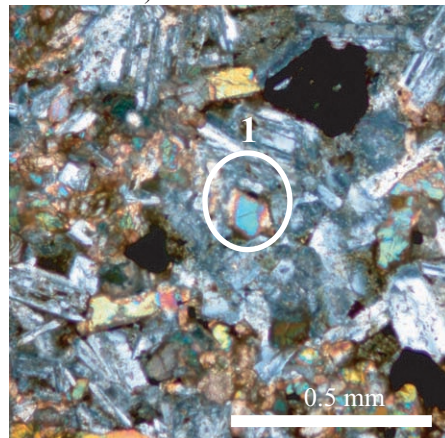
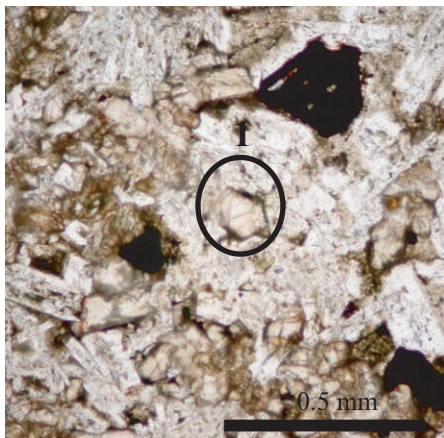
AS004A Point 10 (Lava: Ribeira)



AS021 Point 7 (Lava: Gramado)

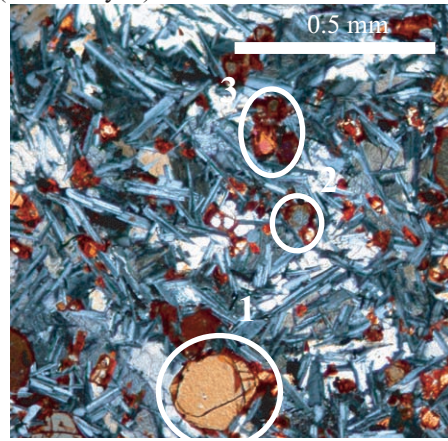
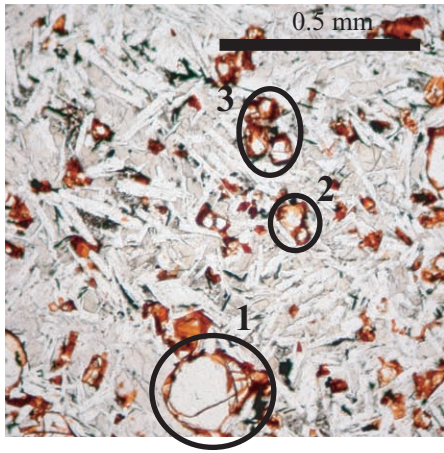


AS021 Point 8 (Lava: Gramado)

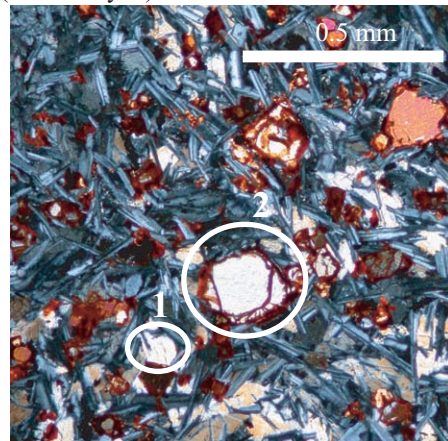
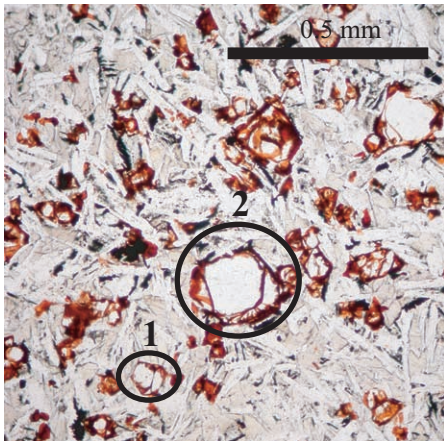


Measuring Points of EPMA

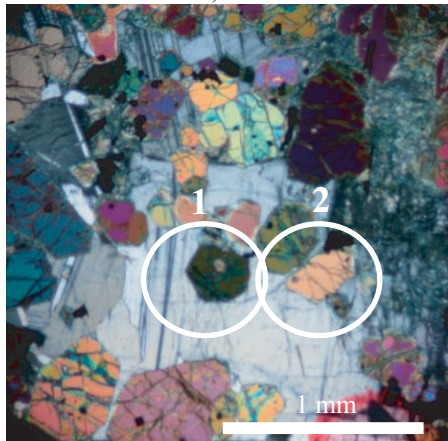
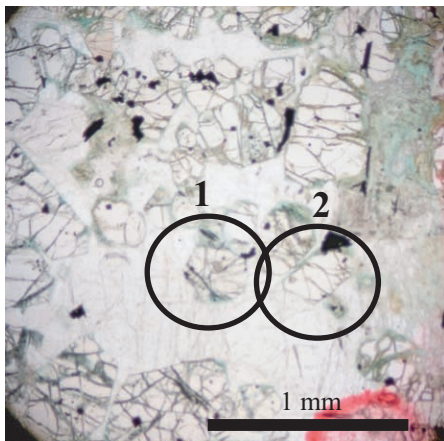
KN003 Point 5 (Feeder Dyke)



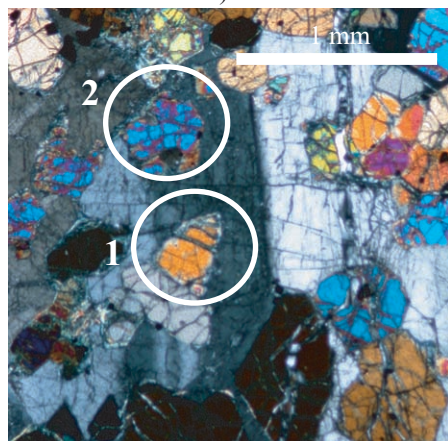
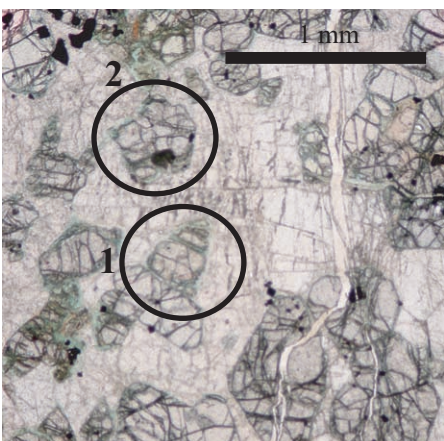
KN003 Point 6 (Feeder Dyke)



TG62-226.3 Point 1 (Intrusion: Drill Core)

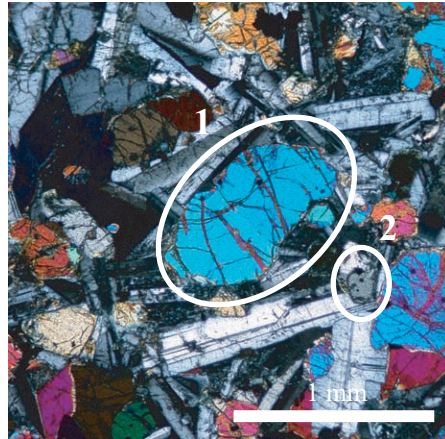
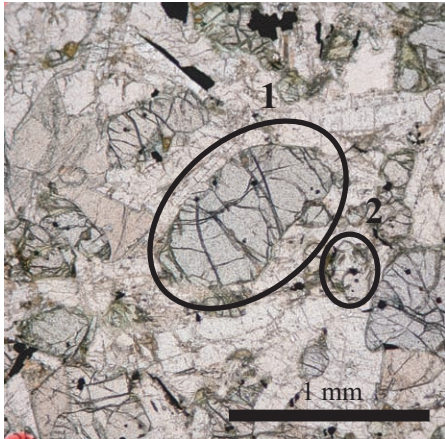


TG62-226.3 Point 2 (Intrusion: Drill Core)

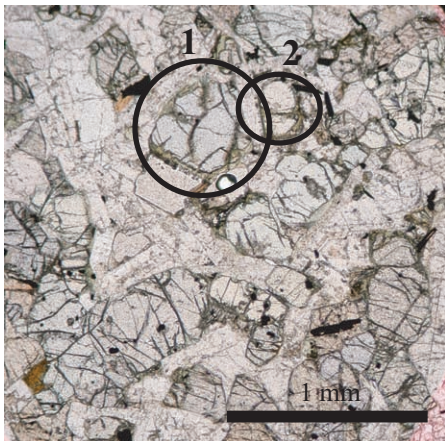


Measuring Points of EPMA

TG114-289.9 Point 3 (Intrusion: Drill Core)



TG114-289.9 Point 4 (Intrusion: Drill Core)



APPENDIX 10

$^{143}\text{Nd}/^{144}\text{Nd}$ and $^{87}\text{Sr}/^{86}\text{Sr}$ isotopic ratios for the Paraná basaltic lava and intrusion

Sample	$(^{143}\text{Nd}/^{144}\text{Nd})_{\text{p}}$	Err	$(^{87}\text{Sr}/^{86}\text{Sr})_{\text{p}}$	Err	Sm (ppm)	Nd (ppm)	Rb (ppm)	Sr (ppm)	^{147}Sm	^{144}Nd	^{87}Rb	^{86}Sr	$(^{147}\text{Sm}/^{144}\text{Nd})_{\text{p}}$	$(^{87}\text{Rb}/^{86}\text{Sr})_{\text{p}}$	λ (Sm)	λ (Rb)	$(^{143}\text{Nd}/^{144}\text{Nd})_{\text{initial}}$	$(^{87}\text{Sr}/^{86}\text{Sr})_{\text{initial}}$	ϵ Nd	ϵ Sr
AS001	0.512373	6	0.711215	15	3.8	17.0	40.9	210.0	0.5624785	4.0578098	11.3709620	20.7060000	0.1386163	0.5491627	6.53912E-12	1.42038E-11	0.512255	0.7102000	-7.47	80.91
AS003	0.512368	7	0.706376	13	3.9	17.6	21.6	316.0	0.5877084	4.1976001	6.0230663	31.1576000	0.1400106	0.1933097	6.53912E-12	1.42038E-11	0.512249	0.706019	-7.59	21.56
AS006	0.512407	6	0.706153	15	6.2	26.9	28.6	353.0	0.9339879	6.4019680	7.9600529	34.8080000	0.1458908	0.2289576	6.53912E-12	1.42038E-11	0.512283	0.705730	-6.93	17.46
AS011	0.512394	4	0.706363	14	6.5	29.4	31.8	345.0	0.9804617	6.9916089	8.8410105	34.0170000	0.1402341	0.2598998	6.53912E-12	1.42038E-11	0.512275	0.705883	-7.09	19.63
AS014	0.512379	5	0.706311	17	13.2	62.3	53.3	471.0	1.9779968	14.8291710	14.8217148	46.4406000	0.1333855	0.3191542	6.53912E-12	1.42038E-11	0.512266	0.705721	-7.27	17.33
AS023	0.512336	5	0.710155	16	6.3	26.2	46.4	216.0	0.9388733	6.2432498	12.9085414	21.2976000	0.1503821	0.6061031	6.53912E-12	1.42038E-11	0.512208	0.709035	-8.39	64.37
AS024A	0.512370	5	0.706881	16	3.7	18.5	16.2	162.0	0.5480981	2.9836813	4.6683581	15.9732000	0.1836986	0.2922619	6.53912E-12	1.42038E-11	0.512414	0.706341	-4.37	26.13
KM106	0.512397	8	0.705923	17	9.5	43.1	25.2	462.5	1.4190925	10.2587142	7.0256880	45.6013375	0.1383304	0.2940676	6.53912E-12	1.42038E-11	0.512279	0.705638	-7.00	16.16
KM120	0.512413	4	0.70618	14	6.4	26.7	22.8	218.3	0.9660897	6.3479746	6.3463187	21.5204135	0.1521887	0.1948976	6.53912E-12	1.42038E-11	0.512284	0.705635	-6.91	16.11
KM142	0.512418	5	0.706235	15	9.5	43.0	29.5	564.7	1.4259549	10.2416203	8.2107531	55.6836932	0.1392314	0.1474535	6.53912E-12	1.42038E-11	0.512300	0.705962	-6.60	20.76
KM153	0.512399	4	0.706208	14	8.6	37.1	27.0	512.4	1.2912470	8.8277634	7.5108976	50.5188503	0.1462711	0.1486751	6.53912E-12	1.42038E-11	0.512275	0.705933	-7.09	20.34
KN002	0.512502	6	0.706743	13	1.7	6.6	5.9	174.0	0.2534079	1.5825298	1.6463975	17.1564000	0.1601284	0.0959640	6.53912E-12	1.42038E-11	0.512366	0.706566	-5.31	29.32
KN003	0.512567	5	0.705988	16	3.8	15.6	8.5	318.0	0.5695040	3.7033471	2.3753072	31.3548000	0.1537809	0.0757558	6.53912E-12	1.42038E-11	0.512436	0.705848	-3.94	19.13
KN005	0.512334	5	0.709502	18	5.8	23.3	48.8	188.0	0.8649026	5.5336090	13.5696982	18.5368000	0.1562999	0.7320410	6.53912E-12	1.42038E-11	0.512201	0.708149	-8.52	51.80
KN006	0.512421	6	0.706829	18	4.0	15.1	14.7	170.0	0.6047099	3.5974490	4.0937012	16.7620000	0.1680941	0.2442251	6.53912E-12	1.42038E-11	0.512278	0.706378	-7.02	26.65
KN009	0.512510	7	0.707613	16	4.1	16.3	58.3	213.0	0.6181273	3.8704729	16.2196563	21.0018000	0.1597037	0.7722984	6.53912E-12	1.42038E-11	0.512374	0.706186	-5.15	23.93
KN012	0.512400	5	0.706922	19	4.0	15.3	17.4	185.0	0.6053614	3.6502805	4.8375080	18.2410000	0.1658393	0.2651997	6.53912E-12	1.42038E-11	0.512259	0.706482	-7.39	27.42
KN014	0.512384	6	0.705917	16	7.7	35.2	33.4	399.0	1.1509931	8.3706240	9.2931987	39.3414000	0.1375039	0.2362193	6.53912E-12	1.42038E-11	0.512267	0.705430	-7.24	13.92
KN016	0.512456	5	0.706240	19	6.3	27.2	28.1	229.0	0.9498608	6.4816711	7.8119840	22.5794000	0.1465457	0.3459784	6.53912E-12	1.42038E-11	0.512331	0.705601	-5.98	15.62
KN019	0.512391	4	0.705804	17	4.9	21.7	10.3	355.0	0.7292740	5.1555955	2.8579493	35.0030000	0.1414529	0.0816487	6.53912E-12	1.42038E-11	0.512271	0.705653	-7.16	16.37
KN020	0.512370	6	0.706225	14	5.6	25.1	23.0	360.0	0.8331737	5.9840361	6.4119399	35.4960000	0.1392327	0.1806384	6.53912E-12	1.42038E-11	0.512252	0.705891	-7.54	19.75
KN022A	0.512387	5	0.706141	17	5.3	24.0	21.3	370.0	0.7946826	5.7210771	5.9292316	36.4820000	0.1389044	0.1625249	6.53912E-12	1.42038E-11	0.512269	0.705841	-7.20	19.03
KN024A	0.512376	7	0.706238	15	5.3	24.7	23.3	375.0	0.7999073	5.8743148	6.4853952	36.9750000	0.1361703	0.1755995	6.53912E-12	1.42038E-11	0.512260	0.705914	-7.37	20.07
KN027	0.512399	6	0.706497	18	6.2	26.1	31.9	264.0	0.9299938	6.2144668	8.8805558	26.0304000	0.1496498	0.3411609	6.53912E-12	1.42038E-11	0.512272	0.705866	-7.14	19.40
KN031	0.512376	4	0.706416	15	6.4	27.6	29.3	289.0	0.9564408	6.5738732	8.1561515	28.4954000	0.1454912	0.2862270	6.53912E-12	1.42038E-11	0.512252	0.705887	-7.52	19.69
KN032	0.512413	8	0.706190	20	4.3	17.8	19.3	285.0	0.6464084	4.2362829	5.3577339	28.1010000	0.1525886	0.1906599	6.53912E-12	1.42038E-11	0.512283	0.705838	-6.92	18.99
KN040A	0.512373	5	0.706457	17	6.1	28.0	32.2	336.0	0.9096085	6.6689402	8.9655759	33.1296000	0.1363948	0.2706213	6.53912E-12	1.42038E-11	0.512257	0.705957	-7.43	20.68
KN050	0.512384	6	0.705936	17	5.5	24.0	12.0	355.0	0.8237702	5.7196655	3.3351901	35.0030000	0.1440242	0.0952830	6.53912E-12	1.42038E-11	0.512262	0.705760	-7.34	17.88
KN052	0.512382	4	0.706136	19	4.6	20.1	17.3	335.0	0.6905617	4.7850638	4.8133479	33.0310000	0.1443161	0.1457221	6.53912E-12	1.42038E-11	0.512259	0.705867	-7.39	19.40
KN104	0.512397	5	0.705907	13	7.8	36.5	27.1	444.8	1.1695365	8.6985927	7.5455805	43.8586560	0.1344512	0.1720431	6.53912E-12	1.42038E-11	0.512283	0.705589	-6.93	15.46
KN110A	0.512439	4	0.705882	15	6.2	27.0	20.4	404.2	0.9230642	6.4203630	5.6839369	39.8552619	0.1437713	0.1426145	6.53912E-12	1.42038E-11	0.512177	0.705618	-6.27	15.88
KN112	0.512393	5	0.70715	12	5.6	24.9	21.8	314.6	0.8462054	5.9173233	6.0741563	31.0181589	0.1430048	0.1958258	6.53912E-12	1.42038E-11	0.512271	0.706788	-7.15	32.48
KN128	0.512425	7	0.705976	17	11.8	54.3	37.4	635.9	1.7637871	12.9330554	10.4093462	62.7002360	0.1363782	0.1660177	6.53912E-12	1.42038E-11	0.512309	0.705669	-6.42	16.60
KN150	0.512424	5	0.706188	18	11.9	56.0	51.0	497.3	1.7810230	13.3169042	14.2054548	49.0350653	0.1337415	0.2896094	6.53912E-12	1.42038E-11	0.512306	0.705653	-6.47	16.36
KN163	0.512453	4	0.706134	14	9.2	41.5	31.7	470.2	1.3848660	9.8778985	8.8132834	46.3604733	0.1401984	0.1901039	6.53912E-12	1.42038E-11	0.512334	0.705785	-5.93	18.21
KN172	0.512394	6	0.706064	11	10.1	46.1	34.4	462.0	1.5114620	10.9664400	9.5715577	45.5484128	0.1378261	0.2101403	6.53912E-12	1.42038E-11	0.512267	0.705676	-7.24	16.69
KN173	0.512419	6	0.70612	13	6.6	27.3	23.1	256.4	0.9935032	6.4934437	6.4248046	25.2761400	0.1530010	0.2541846	6.53912E-12	1.42038E-11	0.512289	0.705650	-6.81	16.33
TM107	0.512425	9	0.706454	17	6.0	25.0	22.8	257.3	0.8948684	5.9460892	6.3316590	25.3663490	0.1504970	0.2496086	6.53912E-12	1.42038E-11	0.512297	0.705993	-6.65	21.19
TM121	0.512386	8	0.705887	14	8.2	35.8	21.0	511.0	1.2231840	8.5204000	5.8443000	50.3846000	0.1435595	0.1159938	6.53912E-12	1.42038E-11	0.512264	0.705673	-7.30	16.64

$^{143}\text{Nd}/^{144}\text{Nd}$ and $^{87}\text{Sr}/^{86}\text{Sr}$ isotopic ratios for the Paraná basaltic lava and intrusion

Sample	$(^{143}\text{Nd}/^{144}\text{Nd})\text{p}$	Err	$(^{87}\text{Sr}/^{86}\text{Sr})\text{p}$	Err	Sm (ppm)	Nd (ppm)	Rb (ppm)	Sr (ppm)	^{147}Sm	^{144}Nd	^{87}Rb	^{86}Sr	$(^{147}\text{Sm}/^{144}\text{Nd})\text{p}$	$(^{87}\text{Rb}/^{86}\text{Sr})\text{p}$	λ (Sm)	λ (Rb)	$(^{143}\text{Nd}/^{144}\text{Nd})$ initial	$(^{87}\text{Sr}/^{86}\text{Sr})$ initial	ϵ Nd	ϵ Sr
TM126	0.512415	6	0.706844	14	11.1	50.2	33.0	572.0	1.6638900	11.9476000	9.1839000	56.3992000	0.1392656	0.1628374	6.53912E-12	1.42038E-11	0.512297	0.706543	-6.66	29.00
TM130	0.512393	6	0.705894	11	8.1	35.6	22.7	497.0	1.2107679	8.4628697	6.3099897	49.0010217	0.1430682	0.1287726	6.53912E-12	1.42038E-11	0.512271	0.705656	-7.15	16.41
TM153	0.512395	5	0.705907	18	7.4	32.1	23.0	442.0	1.1152560	7.6398000	6.4009000	43.5812000	0.1459797	0.1468730	6.53912E-12	1.42038E-11	0.512271	0.705636	-7.16	16.12
TM164	0.512404	4	0.705974	13	9.0	41.8	31.2	394.3	1.3447529	9.9394888	8.6948632	38.8730595	0.1353013	0.2236732	6.53912E-12	1.42038E-11	0.512289	0.705561	-6.81	15.05
WW024	0.512254	5	0.709798	17	5.4	24.2	43.5	232.0	0.8134024	5.7501593	12.0994777	22.8752000	0.1414574	0.5289343	6.53912E-12	1.42038E-11	0.512134	0.708820	-9.84	61.33
WW099	0.512289	4	0.711683	20	4.6	20.4	60.0	214.0	0.6860673	4.8621924	16.6955276	21.1004000	0.1411025	0.7912422	6.53912E-12	1.42038E-11	0.512169	0.710221	-9.15	81.20
WW122	0.512399	5	0.705282	22	11.7	55.0	25.4	801.0	1.7584783	13.0833829	7.0652077	78.9786000	0.1344055	0.0894572	6.53912E-12	1.42038E-11	0.512285	0.705117	-6.89	8.75
WW184	0.512348	8	0.71049	11	4.0	16.1	34.9	186.9	0.5992366	3.8282942	9.7060723	18.4319831	0.1565190	0.5265886	6.53912E-12	1.42038E-11	0.512215	0.709517	-8.25	71.21
YM141	0.512366	5	0.706006	12	8.5	37.8	25.9	412.7	1.2683321	8.9922736	7.2217116	40.6887299	0.1410469	0.1774868	6.53912E-12	1.42038E-11	0.512246	0.705678	-7.65	16.72
YM187	0.512337	7	0.706095	22	8.2	25.2	276.5	169	1.2249614	5.9872910	76.9457196	1.6637323	0.2045936	46.2488577	6.53912E-12	1.42038E-11	0.512196	0.980618	-8.62	3919.34
YM198	0.512529	7	0.706465	17	3.8	13.6	15.7	155.8	0.5683626	3.2389701	4.3590071	15.3652065	0.1754763	0.2836934	6.53912E-12	1.42038E-11	0.512443	0.705941	-3.81	20.45
YM223	0.512431	6	0.705551	16	11.7	57.1	44.6	803.8	1.7607918	13.5869352	12.4230192	79.2504875	0.1295945	0.1567564	6.53912E-12	1.42038E-11	0.512321	0.705261	-6.19	10.81
AT03-519.2	0.512331	4	0.711547	17	6.1	27.5	61.0	241.0	0.9128910	6.5450000	16.9763000	23.7626000	0.1394791	0.7144126	6.53912E-12	1.42038E-11	0.512212	0.710227	-8.30	81.29
AT08-845.1	0.512334	4	0.712020	16	6.5	29.4	64.0	210.0	0.9698530	6.9972000	17.8112000	20.7060000	0.1386059	0.8601951	6.53912E-12	1.42038E-11	0.512216	0.710430	-8.23	84.18
CA-86RS 496.25+496.75	0.512252	5	0.717142	13	7.8	35.5	100.2	180.3	1.1744862	8.4425260	27.8981657	17.7821379	0.1391200	1.5688870	6.53912E-12	1.42038E-11	0.512134	0.714242	-9.84	138.29
SA109 657.70+657.90	0.512405	7	0.710139	13	3.2	13.5	25.0	244.7	0.4730667	3.2051217	6.9706853	24.1274814	0.1475971	0.2889106	6.53912E-12	1.42038E-11	0.512279	0.709605	-6.99	72.46
PP-02-SP 1120-1130	0.51239	6	0.70596	12	8.4	37.6	28.5	429.6	1.2537206	8.9520861	7.9248533	42.3585684	0.1400479	0.1870897	6.53912E-12	1.42038E-11	0.512271	0.705614	-7.16	15.82
PP-02-SP 1280-1290	0.512403	4	0.706128	16	10.3	46.2	24.7	545.5	1.5504314	10.9958625	6.8845459	53.7819187	0.1410013	0.1280086	6.53912E-12	1.42038E-11	0.512283	0.705891	-6.92	19.75
PP-02-SP 1380-1390	0.512479	5	0.706405	13	9.2	38.4	34.6	337.5	1.3834765	9.1442556	9.6157492	33.2810382	0.1512946	0.2889258	6.53912E-12	1.42038E-11	0.512350	0.705871	-5.61	19.46
PP-02-SP 1460-1470	0.512925	6	0.706703	12	10.4	49.4	41.7	526.8	1.5940334	11.7660993	11.5959659	51.9424312	0.1325336	0.2232465	6.53912E-12	1.42038E-11	0.512812	0.706290	3.40	25.41
PP-02-SP 1540-1550	0.512411	5	0.706803	11	10.0	46.2	45.2	512.9	1.5033242	11.0066218	12.5897354	50.5710526	0.1365836	0.2489514	6.53912E-12	1.42038E-11	0.512295	0.706343	-6.69	26.16
PP-02-SP 1590-1600	0.512316	5	0.709391	11	6.0	25.7	54.6	348.6	0.8973651	6.1280904	15.1861419	34.3673049	0.1464347	0.4418776	6.53912E-12	1.42038E-11	0.512191	0.708574	-8.71	57.83
PP-02-SP 280-290	0.512412	4	0.706481	14	6.0	24.6	23.3	270.3	0.8934496	5.8567380	6.4850218	26.6509778	0.1525507	0.2433540	6.53912E-12	1.42038E-11	0.512282	0.706031	-6.94	21.74
PP-02-SP 440-450	0.51241	7	0.706261	13	5.0	20.6	21.9	255.6	0.7453952	4.8925446	6.0884737	25.2057334	0.1523533	0.2415511	6.53912E-12	1.42038E-11	0.512280	0.705815	-6.98	18.66
PP-02-SP 600-610	0.512362	6	0.706151	11	7.3	32.6	28.5	369.5	1.0897918	7.7584766	7.9358408	36.4333075	0.1404647	0.2178183	6.53912E-12	1.42038E-11	0.512243	0.705748	-7.71	17.72
PP-02-SP 760-770	0.512397	5	0.706355	14	9.2	41.2	35.7	358.2	1.3775181	9.8154215	9.9564482	35.3206227	0.1403422	0.2813214	6.53912E-12	1.42038E-11	0.512278	0.705835	-7.03	18.95
PP-02-SP 800-810	0.512371	6	0.706624	12	8.2	36.4	29.2	436.6	1.2621617	8.9849258	6.9372239	42.9316296	0.1404755	0.1615877	6.53912E-12	1.42038E-11	0.512261	0.705763	-7.36	17.93
PP-02-SP 960-970	0.512371	6	0.70624	12	8.2	36.4	29.2	436.6	1.2223740	8.6734076	8.1316249	43.0508834	0.1409335	0.1888840	6.53912E-12	1.42038E-11	0.512251	0.705891	-7.55	19.74
SW01 556.55+556.70	0.512383	4	0.717506	12	6.6	26.3	72.0	210.3	0.9939850	6.2523059	20.0299762	20.7378074	0.1589790	0.9658676	6.53912E-12	1.42038E-11	0.512248	0.715721	-7.61	159.27
TG114-289.9	0.512864	5	0.705162	15	2.5	9.7	9.8	188.0	0.3719789	2.3108101	2.7338333	18.5368000	0.1609734	0.1474814	6.53912E-12	1.42038E-11	0.512727	0.704889	1.74	5.53
TG228-817	0.512219	4	0.710876	17	5.2	23.6	37.0	243.0	0.7764820	5.6168000	10.2971000	23.9598000	0.1382428	0.4297657	6.53912E-12	1.42038E-11	0.512101	0.710082	-10.47	79.23
TG62-237.6	0.512851	7	0.705398	19	2.3	9.2	12.1	189.0	0.3472032	2.1880149	3.3702364	18.6354000	0.1586841	0.1808513	6.53912E-12	1.42038E-11	0.512716	0.705064	1.52	8.00
TG95-756.3	0.512210	5	0.710140	16	4.6	21.3	32.0	236.0	0.6925380	5.0694000	8.9056000	23.2696000	0.1366114	0.3827139	6.53912E-12	1.42038E-11	0.512094	0.709433	-10.62	70.02

$^{143}\text{Nd}/^{144}\text{Nd}$ CHUR: 0.512638 (Faure, 1986)

$^{87}\text{Sr}/^{86}\text{Sr}$ Nd UR: 0.7045 (White and Hofmann, 1982)

APPENDIX 11

AS001

Inc.	39/40	(39/40)err	36/40	(36/40)err
351	1.07E-02	3.10E-04	1.91E-03	3.00E-05
519	8.12E-03	1.70E-05	6.39E-04	3.20E-06
595	9.84E-03	9.00E-06	3.45E-04	2.10E-06
689	1.08E-02	9.90E-06	3.88E-04	3.00E-06
769	1.05E-02	1.00E-05	4.90E-04	5.30E-06
845	9.33E-03	1.30E-05	8.83E-04	7.90E-06
1102	1.14E-02	1.20E-05	7.86E-04	4.00E-06
1422	9.08E-03	5.20E-05	1.18E-03	1.00E-05

Inc.	% 39Ar	Age (Ma)	Age Err	%40Arrad
351	2.68	68.8	2.3	44
519	17.87	163.7	0.4	81
595	23.54	150.3	0.2	90
689	17.62	135.9	0.2	89
769	11.51	134.4	0.3	86
845	5.37	131.1	0.4	74
1102	17.53	111.6	0.2	77
1422	3.87	119.2	0.9	65

AS003

Inc.	39/40	(39/40)err	36/40	(36/40)err
421	1.79E-04	3.00E-06	3.29E-03	1.20E-05
512	4.54E-03	7.80E-06	1.68E-03	9.40E-06
590	9.41E-03	9.70E-06	5.87E-04	4.60E-06
688	1.01E-02	9.40E-06	5.28E-04	4.60E-06
769	1.05E-02	1.20E-05	4.61E-04	5.30E-06
841	9.57E-03	1.00E-05	8.12E-04	9.00E-06
1097	9.34E-03	1.00E-05	1.25E-03	8.00E-06
1422	6.02E-03	1.50E-05	1.84E-03	1.90E-05

Inc.	% 39Ar	Age (Ma)	Age Err	%40Arrad
421	2.08	239.4	29	2.6
512	6.32	181.1	1	50
590	21.05	145	0.3	82
688	25.17	137.7	0.3	84
769	16	135.9	0.3	86
841	10.4	131.5	0.5	76
1097	16.44	112.5	0.4	63
1422	2.55	125.6	1.6	46

AS006

Inc.	39/40	(39/40)err	36/40	(36/40)err
444	8.41E-03	8.10E-06	6.29E-04	9.20E-06
530	9.42E-03	1.10E-05	3.60E-04	8.00E-06
590	9.90E-03	1.10E-05	4.39E-04	6.10E-06
690	1.10E-02	1.10E-05	3.11E-04	3.90E-06
770	1.06E-02	1.10E-05	5.05E-04	4.70E-06
910	8.67E-03	2.50E-05	1.15E-03	2.80E-05
1100	8.34E-03	1.70E-05	1.55E-03	1.20E-05
1420	5.56E-03	7.10E-05	1.99E-03	3.00E-05

Inc.	% 39Ar	Age (Ma)	Age Err	%40Arrad
444	7.72	159.3	0.5	83
530	11.6	156.3	0.4	89
590	15.82	145.3	0.3	87
690	31	136.9	0.2	91
770	23.07	133.4	0.3	85
910	3.15	126.5	1.6	66
1100	6.18	108.6	0.7	54
1420	1.46	123.2	3	41

AS023

Inc.	39/40	(39/40)err	36/40	(36/40)err
432	1.08E-02	3.30E-05	1.39E-03	1.60E-05
514	1.03E-02	1.10E-05	9.64E-04	1.10E-05
588	9.22E-03	1.10E-05	9.40E-04	9.00E-06
771	8.84E-03	9.30E-06	1.02E-03	1.20E-05
845	7.69E-03	2.00E-05	1.75E-03	2.20E-05
909	6.39E-03	2.20E-05	2.20E-03	3.90E-05
1106	5.80E-03	8.20E-06	2.12E-03	1.60E-05
1420	2.79E-03	1.50E-05	2.28E-03	1.70E-05

Inc.	% 39Ar	Age (Ma)	Age Err	%40Arrad
432	27.89	91.2	0.8	59
514	24.19	116.2	0.5	71
588	22.98	130.3	0.5	72
771	12.18	131.5	0.7	70
845	4.6	105.3	1.4	48
909	1.98	92.1	3	35
1106	4.1	107.8	1.4	37
1420	2.09	191.2	2.9	33

AS024A

Inc.	39/40	(39/40)err	36/40	(36/40)err
438	3.62E-03	3.50E-05	2.29E-03	1.30E-05
523	5.23E-03	2.20E-05	1.72E-03	1.00E-05
590	7.01E-03	7.20E-06	1.30E-03	7.20E-06
690	7.15E-03	6.60E-06	1.37E-03	5.50E-06
767	6.20E-03	7.00E-06	1.55E-03	9.10E-06
843	5.74E-03	1.50E-05	1.79E-03	1.40E-05
908	5.54E-03	6.40E-06	1.94E-03	1.50E-05
1105	5.68E-03	8.50E-06	2.04E-03	9.00E-06
1424	3.25E-03	7.90E-06	2.39E-03	1.30E-05

Inc.	% 39Ar	Age (Ma)	Age Err	%40Arrad
438	5.18	148.3	2	32
523	12.68	155.7	1.1	49
590	16.97	145.5	0.5	61
690	30.18	138.5	0.4	59
767	10.64	144.9	0.7	54
843	5.05	136.6	1.2	47
908	4.77	128.8	1.3	43
1105	11.18	116.7	0.8	40
1424	3.35	150.2	2	29

KN002

Inc.	39/40	(39/40)err	36/40	(36/40)err
432	7.20E-03	1.90E-04	2.55E-03	4.50E-05
515	5.78E-03	1.80E-04	2.11E-03	5.80E-05
600	5.20E-03	1.70E-05	1.83E-03	1.70E-05
688	5.38E-03	7.10E-06	1.85E-03	8.10E-06
769	6.11E-03	9.70E-06	1.54E-03	1.00E-05
917	5.32E-03	1.50E-05	1.84E-03	2.00E-05
1110	5.06E-03	5.70E-06	2.06E-03	9.00E-06
1420	2.31E-03	7.60E-06	2.70E-03	1.70E-05

Inc.	% 39Ar	Age (Ma)	Age Err	%40Arrad
432	7.21	57.1	3.1	25
515	11.24	106.3	5.5	37
600	6.84	143	1.5	46
688	30.73	136.8	0.7	45
769	14.34	144.1	0.8	54
917	5.92	139.3	1.8	46
1110	19.15	125.4	0.8	39
1420	4.57	141.8	3.4	20

KN003

Inc.	39/40	(39/40)err	36/40	(36/40)err
450	3.23E-03	5.20E-06	2.11E-03	9.90E-06
550	4.52E-03	1.30E-05	1.84E-03	1.00E-05
600	4.51E-03	2.30E-05	1.92E-03	1.40E-05
690	4.12E-03	7.20E-06	2.09E-03	9.40E-06
770	5.82E-03	2.30E-05	1.47E-03	2.10E-05
910	5.34E-03	1.40E-05	1.70E-03	1.60E-05
1110	5.81E-03	8.10E-06	1.81E-03	1.20E-05
1410	3.40E-03	3.00E-05	2.26E-03	3.10E-05

Inc.	% 39Ar	Age (Ma)	Age Err	%40Arrad
450	15.15	185.8	1.4	37
550	14.24	163.2	1.1	46
600	12.48	155.4	1.5	43
690	19.54	150	1.1	38
770	7.12	156.9	1.8	56
910	8.61	151.1	1.4	50
1110	19.61	130.4	1	46
1410	3.26	157.1	4.4	33

KN005

Inc.	39/40	(39/40)err	36/40	(36/40)err
444	8.19E-03	4.90E-05	6.59E-04	1.70E-05
522	9.67E-03	1.10E-05	2.57E-04	4.90E-06
600	1.07E-02	1.10E-05	2.08E-04	3.30E-06
693	1.11E-02	1.10E-05	2.25E-04	4.00E-06
770	1.07E-02	1.70E-05	4.04E-04	1.40E-05
910	1.01E-02	2.60E-05	7.23E-04	1.40E-05
1110	1.04E-02	1.80E-05	1.13E-03	1.70E-05
1420	7.75E-03	2.30E-05	1.34E-03	2.90E-05

Inc.	% 39Ar	Age (Ma)	Age Err	%40Arrad
444	4.29	159	1.3	81
522	19.92	154.7	0.3	92
600	29.81	142.6	0.2	94
693	23.97	137	0.2	93
770	7.01	134.3	0.6	88
910	5.67	126.8	0.7	79
1110	6.98	105.2	0.8	67
1420	2.35	127.3	1.8	60

KN012

Inc.	39/40	(39/40)err	36/40	(36/40)err
370	7.35E-03	7.80E-06	1.07E-03	9.60E-06
530	8.51E-03	1.10E-05	4.70E-04	9.00E-06
595	9.60E-03	8.90E-06	3.88E-04	4.20E-06
688	9.86E-03	1.10E-05	4.86E-04	7.40E-06
769	9.38E-03	1.30E-05	6.44E-04	9.70E-06
910	7.75E-03	1.20E-05	1.21E-03	1.40E-05
1100	5.81E-03	2.20E-05	2.33E-03	1.80E-05
1420	3.23E-03	1.90E-05	2.57E-03	3.10E-05

Inc.	% 39Ar	Age (Ma)	Age Err	%40Arrad
370	10.97	151.3	0.6	68
530	16.01	163.6	0.5	86
595	29.49	149.8	0.2	88
688	20.41	141.4	0.4	86
769	11.85	140.5	0.5	81
910	5.65	135.4	0.9	64
1100	4.16	88.7	1.5	31
1420	1.45	122.1	4.5	24

KN016

Inc.	39/40	(39/40)err	36/40	(36/40)err
435	1.10E-04	1.80E-06	3.31E-03	1.20E-05
522	6.05E-03	5.50E-05	1.56E-03	1.50E-05
597	9.08E-03	1.10E-05	6.54E-04	6.40E-06
681	1.00E-02	9.40E-06	5.53E-04	2.60E-06
761	1.01E-02	1.00E-05	5.88E-04	4.70E-06
921	9.44E-03	1.10E-05	9.18E-04	6.00E-06
1110	9.26E-03	1.10E-05	1.28E-03	9.30E-06
1434	5.91E-03	1.10E-05	1.85E-03	1.20E-05

Inc.	% 39Ar	Age (Ma)	Age Err	%40Arrad
435	0.64	307.1	45	2
522	4.79	145	1.7	54
597	11.07	144.7	0.4	81
681	35.91	136	0.2	84
761	18.1	133.3	0.3	83
921	13.83	126.4	0.3	73
1110	12.01	110.6	0.5	62
1434	3.66	125.4	1	45

KN020

Inc.	39/40	(39/40)err	36/40	(36/40)err
438	7.51E-03	1.70E-05	6.52E-04	1.50E-05
525	9.29E-03	1.70E-05	3.50E-04	7.30E-06
593	9.91E-03	1.10E-05	3.26E-04	1.10E-05
686	1.07E-02	1.10E-05	3.33E-04	4.80E-06
776	1.08E-02	1.40E-05	3.64E-04	8.50E-06
840	9.62E-03	2.20E-05	7.18E-04	1.30E-05
1098	7.68E-03	1.60E-05	1.38E-03	2.10E-05
1420	5.26E-03	3.10E-05	2.03E-03	3.80E-05

Inc.	% 39Ar	Age (Ma)	Age Err	%40Arrad
438	5.24	173.8	1	55
525	28.52	156.8	0.5	69
593	12.84	148.5	0.5	77
686	22.33	137.9	0.3	80
776	17.41	135.1	0.4	85
840	7.25	134	0.7	87
1098	5.14	126.5	1.3	73
1420	1.25	124.4	3.5	55

KN024A

Inc.	39/40	(39/40)err	36/40	(36/40)err
433	6.70E-03	6.80E-05	6.31E-04	4.90E-05
533	9.44E-03	1.00E-05	3.35E-04	7.40E-06
602	1.04E-02	1.10E-05	2.36E-04	7.30E-06
693	1.11E-02	1.00E-05	2.91E-04	4.60E-06
771	9.48E-03	1.00E-05	7.62E-04	1.40E-05
912	8.03E-03	2.40E-05	1.25E-03	2.10E-05
1102	7.47E-03	2.90E-05	2.00E-03	2.60E-05
1418	4.77E-03	9.00E-05	1.99E-03	6.90E-05

Inc.	% 39Ar	Age (Ma)	Age Err	%40Arrad
433	1.44	195.7	3.8	81
533	34	155.5	0.4	90
602	17.75	145.8	0.4	93
693	31.38	135.3	0.2	91
771	7.86	133.9	0.7	77
912	3.86	128.9	1.3	63
1102	2.95	91	1.7	41
1418	0.77	141.6	7.2	41

KN027

Inc.	39/40	(39/40)err	36/40	(36/40)err
458	4.87E-03	2.20E-05	1.53E-03	1.50E-05
535	5.50E-03	2.40E-05	1.05E-03	1.50E-05
606	7.16E-03	1.10E-05	7.72E-04	8.90E-06
693	8.59E-03	1.30E-05	6.58E-04	6.20E-06
771	9.82E-03	1.20E-05	5.07E-04	7.80E-06
845	1.07E-02	9.90E-06	4.48E-04	2.20E-06
1100	1.11E-02	1.50E-05	8.97E-04	1.30E-05
1420	8.42E-03	2.60E-05	1.53E-03	2.90E-05

Inc.	% 39Ar	Age (Ma)	Age Err	%40Arrad
458	4.3	181.8	1.6	55
535	3.29	202.2	1.5	69
606	7.48	174.8	0.6	77
693	9.17	153.1	0.4	80
771	10.13	141.8	0.4	85
845	51.8	132.8	0.2	87
1100	11.33	109.5	0.6	73
1420	2.5	107.5	1.7	55

KN031

Inc.	39/40	(39/40)err	36/40	(36/40)err
437	6.65E-03	1.00E-05	1.09E-03	1.50E-05
520	8.37E-03	1.30E-05	6.10E-04	1.20E-05
601	9.58E-03	1.00E-05	4.93E-04	7.30E-06
693	1.05E-02	1.00E-05	3.60E-04	3.60E-06
772	1.05E-02	1.00E-05	5.26E-04	4.00E-06
916	1.00E-02	1.10E-05	7.08E-04	6.60E-06
1108	9.21E-03	1.30E-05	1.24E-03	6.40E-06
1430	7.74E-03	1.60E-05	1.50E-03	1.40E-05

Inc.	% 39Ar	Age (Ma)	Age Err	%40Arrad
437	3.9	166.3	1.1	68
520	4.54	159.8	0.7	81
601	23.91	146.2	0.4	85
693	21.61	139.7	0.2	89
772	20.18	131.8	0.2	84
916	11.73	129.9	0.3	79
1108	10.27	114	0.4	63
1430	3.85	119.1	0.9	56

KN050

Inc.	39/40	(39/40)err	36/40	(36/40)err
429	5.83E-03	1.90E-05	1.35E-03	1.90E-05
530	7.61E-03	8.80E-06	8.12E-04	5.40E-06
594	8.64E-03	9.00E-06	7.04E-04	5.20E-06
688	9.53E-03	8.80E-06	6.56E-04	4.70E-06
770	9.41E-03	1.20E-05	8.07E-04	8.20E-06
912	8.18E-03	1.30E-05	1.24E-03	1.00E-05
1151	6.02E-03	1.40E-05	2.01E-03	1.20E-05
1424	2.86E-03	2.60E-05	2.67E-03	2.30E-05

Inc.	% 39Ar	Age (Ma)	Age Err	%40Arrad
429	4.19	168.4	1.6	60
530	18.19	163.2	0.4	76
594	17.55	150.3	0.3	79
688	29.71	139.2	0.3	81
770	13.61	133.4	0.4	58
912	9.01	127.8	0.6	63
1151	6.47	112.1	1	41
1424	1.28	122.8	4	21

KN052

Inc.	39/40	(39/40)err	36/40	(36/40)err
264	4.99E-03	3.00E-04	3.05E-03	1.40E-05
419	4.32E-03	6.30E-06	2.01E-03	9.60E-06
517	7.28E-03	7.10E-06	1.02E-03	4.40E-06
589	7.68E-03	7.00E-06	1.12E-03	6.50E-06
687	8.52E-03	8.00E-06	9.78E-04	4.40E-06
770	9.42E-03	9.60E-06	8.25E-04	5.60E-06
841	8.11E-03	9.90E-06	1.20E-03	9.20E-06
1103	5.54E-03	9.40E-06	2.08E-03	9.70E-06
1418	2.85E-03	2.20E-05	2.50E-03	1.80E-05

Inc.	% 39Ar	Age (Ma)	Age Err	%40Arrad
264	8.5	33.6	2.4	9.8
419	5.76	154.1	1	40.6
517	25.27	157.7	0.3	70
589	13.96	143.3	0.4	67
687	19.85	137.7	0.3	71
770	12.38	132.6	0.3	76
841	6.39	131.4	0.6	65
1103	6.48	115.3	0.9	38
1418	1.41	151	3.3	26