

THE STATE OF SERGIPE GEOMORPHOLOGICAL MAP

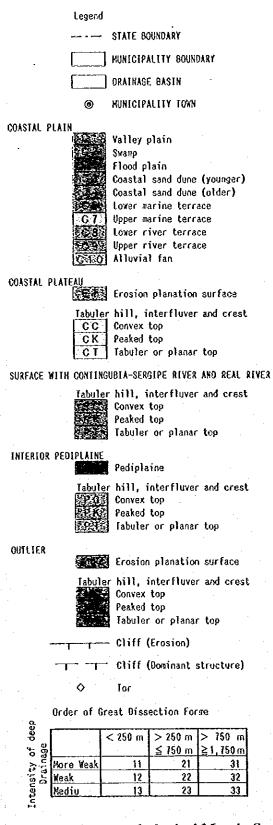
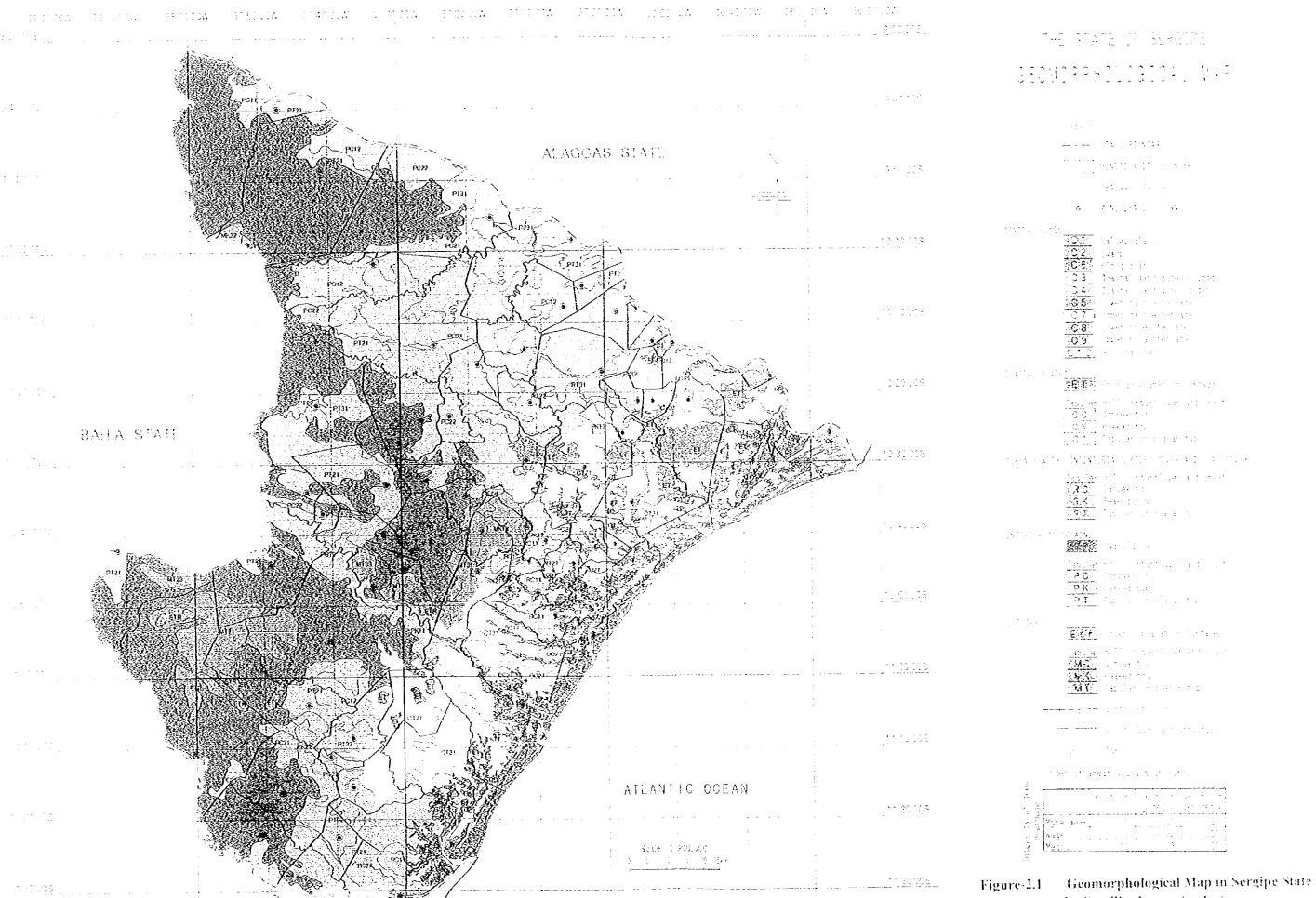
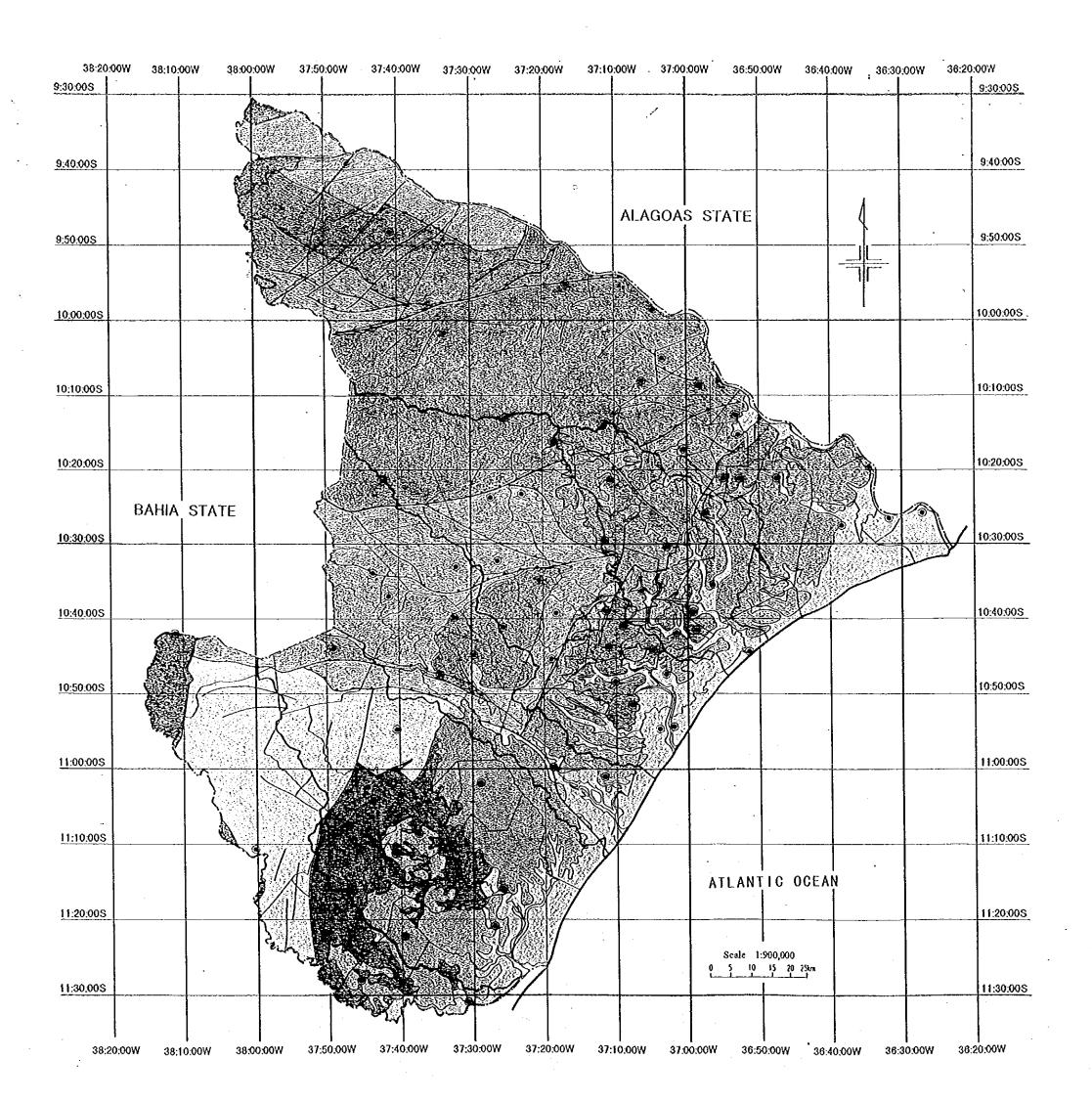


Figure-2.1 Geomorphological Map in Sergipe State by Satellite Image Analysis



16 01 01 A 38 1 2 DOW 38 03 DOW 31 50 DOW 37 40 DOW 37 30 DOW 37 20 DOW 37 13 DOW 37 60 DOW 38 50 DOW 36 40 DOW 36 30 DOW 36 20 DOW

by Satellite Image Analysis



THE STATE OF SERGIPE GEOLOGICAL MAP

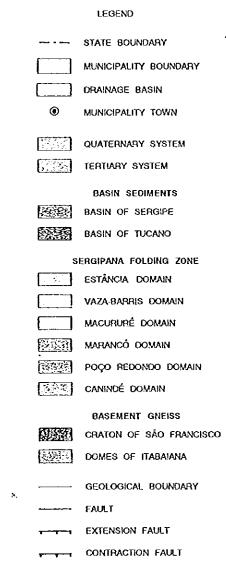
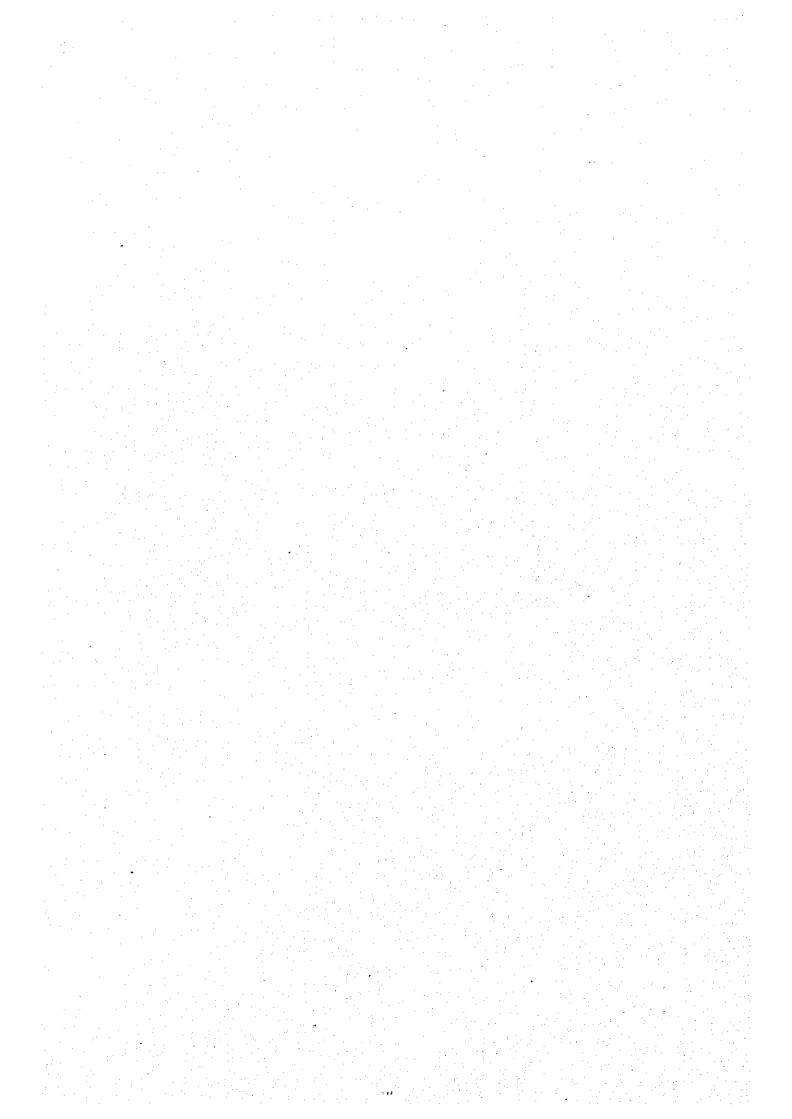


Figure-2.2 Geological Map in Sergipe State by Satellite Image Analysis



2.3 Topography and Geology by River Basin

2.3.1 Topography by River Basin

Sergipe State is divided into 6 main river basins. Every basin has the same geomorphological component from upstream to downstream, namely mountain area - flat plateau area - coastal plain. However, each geomorphological unit has different weight by basin. Basins in the northern part of Sergipe State have less mountain area, on the other hand, basins in southern part have more mountain area. This fact is closely related to size of basin. Sao Francisco River Basin, located in the northern-most part of Sergipe State, has less mountain area and more flat plateau. It results in milder slope than the other basins. On the other hand, Piaui River Basin and Real River Basin, located in the southern-most part of Sergipe State, have more mountain area. The other basins, Japaratuba, Sergipe, Vaza-Barris River Basin, located in the middle part of Sergipe State, have intermediate characteristics between the formers. In these Basin, however, Itabaiana Dome is notable with steep mountain dividing Sergipe Basin and Vaza-Barris Basin.

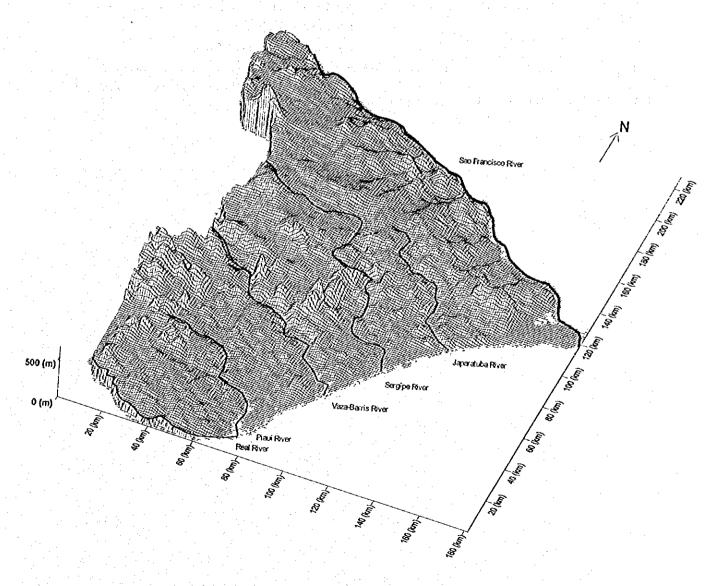


Figure-2.3 Bird's-eyes View of Sergipe State

2.3.2 Geology by River Basin

Each river basin has different geological unit. Representative geological unit by river basin is shown in Table-2.3.

Table-2.3 Geology of Basin

Basin	Upper stream	Middle stream	Lower stream
Sao Francisco	Maranco Domain (3%) Poco Redondo Domain (6%) Caninde Domain (3%)	Macurure Domain (40%)	Quaternary (5%) Barreiras Formation (11%)
Innovetula		Barreiras Formation (48%)	
Japaratuba	Macurure Domain(28%)	Sergipe Basin (12%)	Quaternary (12%)
Sergipe	Macurure Domain (26%)	Vaza-Barris Domain (24%) Itabaiana Dome (7%)	Quaternary (11%) Barreiras Formation (16%) Sergipe Basin (11%)
Vaza-Barris	Vaza-Barris Domain (51%)	Itabaiana Dome (11%)	Barreiras Formation(17%)
Piaui	Estancia Domain (20%)	Barreiras Formation(3 Sao Francisco Craton	
Real	Estancia Do	main (59%)	Barreiras Formation (9%) Sao Francisco Craton (18%)

Note: (%): Area of each geological unit represented by % against total basin area

CHAPTER 3 HYDROGEOLOGY

3.1 Hydrogeological Classification

Hydrogeology of the study area is dominated by geological condition. Hydrogeological classification should follow the geological classification. Table-3.1 shows hydrogeological classification and Figure-3.1 shows its distribution.

Table-3.1 Hydrogeological Unit the Study Area

Age		Stratigraphy	Rock Faces	Hydraulic Characteristics	
Cenozoic Quaternary		Alluvium	Clay, silt, sand, gravel	Unconfined stratum water	
	Tertiary	Barreiras Claystone, siltstone, sandstone, Formation conglomerate		Unconfined / confined stratum water	
Mesozoic 	Cretaceous	Tucano Basin	Limestone, sandstone, shale	Unconfined / confined stratum water	
Palaeozoic	Silurian	Sergipe Basin	Limestone, sandstone, shale	Unconfined / confined stratum water	
Late Proterozoic		Caninde Domain	Gabbro, amphibolite, metavolcanic rock, ultramafic rock	Unconfined fissure water	
		Poco Redondo Domain	Granites, migmatite, gneiss	Unconfined fissure water	
		Maranco Domain	Granites, meta conglomerate, phyllite	Unconfined fissure water	
	er egyéket	Macurure Domain .	Micaschist, quartzite, gabbro	Unconfined fissure water	
Middle Prote Proterozoic	rozoic - late	Vaza-Barris Domain	Carbonate, phyllite, argillaceous rock	Unconfined fissure water	
		Estancia Domain	Sandstone, argillaceous rock Conglomerate.	Unconfined fissure water	
Archaean - early Proterozoic		Sao Francisco Craton	Gneiss, migmatite, granodiorite.	Unconfined fissure water	
		Itabaiana Dome Craton	Migmatite.	Unconfined fissure water	

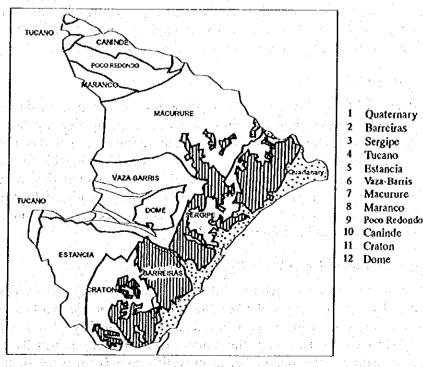


Figure-3.1 Hydrogeological Unit

3.2 Current Groundwater Use and Hydrogeological Information

Groundwater development has been carried out mainly by drilling deep wells in Sergipe State. More than 4,000 deep wells and great number of shallow wells were drilled until now. Other than wells, groundwater is used from springs. Characteristics of deep wells are described below based on the result of existing data analysis.

(1) Deep Well Data-base

SRH has deep well data-base storing more than 4,000 well data. This data base contains items listed below:

Table-3.2	Contents of	Deep Well	Data-base by	SRH

Items	Content
Organization	1) Name of organization which funded drilling
Well location	1) Municipality, 2) Village name
Well dimension	1) Diameter, 2) Length,
Well capacity	1) Yield, 2) Static water level, 3) Dynamic water level 4) Draw down
Water quality	1) pH, 2) Cl, 3) No, 4) No ₃ , 5) CaCO ₃
Well construction record	1) Year, month and data of commencement of drilling,
	2) Year, month, data of completion of drilling

Groundwater potential of this Study were analyzed using information derived from the Data-base which was expected to provide hydrogeological characteristics such as well capacity, hydraulic parameters and water quality and so on. In order to get hydrogeological characteristics mentioned above, lithology and hydrogeological unit of each deep well must be identified from the data-base. As shown in Table-3.2, however, the data-base does not have such items. In order to resolve this problem, the following method was taken; one representative hydrogeological unit was selected for each municipality in accordance with its hydrogeological location. Then, each deep well in the data-base was given one hydrogeological unit according to the municipality to which this deep well belongs. It means that all the deep wells belonging to the same municipality have the same hydrogeological unit. This method does not lead to precise result, however it seems that this method will be permitted in order to assess rough groundwater potential for the Master Plan under the condition of the current data shortage of the data-base.

(2) Aquifer of Deep Well

Deep wells were drilled in the sites belonging to almost all the hydrogeological units which are listed in Table-3.1, including Barreiras formation. It seems that boreholes having been drilled in Barreiras formation, in fact, aimed at underlying hydrogeological unit as a target because Barreiras formation is generally too thin and too impermeable for boreholes to get sufficient groundwater. Consequently, hydrogeological characteristics indicated by deep wells which are located in Barreiras formation area are considered to indicate those of underlying unit. Barreiras formation should be considered as the aquifer suitable for shallow well (dug well) and not for deep wells.

(3) Distribution of Deep Wells by Municipality

The number of existing deep wells is different by municipality (see Appendix-1). Some municipalities have many deep wells but others not. It is notable that tremendous number of deep wells were drilled in the past in Itabaiana municipality (758 wells) and Lagarto municipality (289 wells). However, the number of deep wells is usually less than 50 in most of municipalities. The number of deep wells depends on water demand and water quality by municipality.

(4) Basic Capacity of Deep Well

Yield, specific capacity, success rate and water quality are most important parameters of deep well (see Appendix-1). Representative values of these parameters are shown in Table-3.3. It is clear that well capacity is different by each geological unit, and also groundwater development potential seems to be different by each geological unit. Especially difference in water quality is dominant. Generally in terms of water quality, deep wells in sedimentary rock area (Cretaceous and Quaternary) is more excellent than deep wells in crystalline rock area in quality and quantity. Barreiras formation which distributes in wide area of the Study area, is out of Table-3.3 because of its poor capacity for deep wells.

Specific Rate of fresh Aquifer Yield Success rate Capacity water (m³/day) (m³/day/m) (%)(%) Alluvium covering Sergipe Basin 600 140 95 100 Alluvium covering Craton Tucano Basin 100 60 60 Sergipe Basin covering Barreiras 140 17 80 85 Sergipe Basin outcropping 140 13 70 60 Caninde Domain 40 2 45 10 Poco Redondo Domain 2 40 45 10 Maranco Domain 40 2 45 10 Macurure Domain 40 2 60 15 Vaza-Barris Domain 80 4 75 40 Estancia Domain 50 3 70 50 Sao Francisco Craton covered by Barreiras 70 4 85 90 Sao Francisco Craton outcropping 40 75 30 Itabaiana Dome Craton 70 75 35

Table-3.3 Basic Capacity of Deep Well

(5) Groundwater Quality

It is notable that groundwater in Precambrian rock usually contains high salinity in the study area (see Appendix-1). Especially chlorine (Cl) density is high, usually more than 250 ppm. Groundwater in Maranco Domain and Macurure Domain, which belong to Precambrian rock, show especially high Cl with more than 250 ppm in most of deep wells. On the other hand, groundwater in Itabaiana Dome, which also belongs to Precambrian rock, shows less Cl concentration, and some wells show Cl of less than 250 ppm. Compared with Precambrian rock, sedimentary rock (Cretaceous, Tertiary and Quaternary) in the Study area has lower Cl concentration of usually less than 250 ppm.

(6) Depth and Diameter of Deep Wells

Most of wells have depth of 40 m~80 m, and the average is 60m (see Appendix-1). Diameter of deep wells is 6 inch (15 cm).

(7) Groundwater Static Level

Groundwater static level is usually less than 15 m from ground surface, and average is 10 m (see Appendix-1).

(8) Number of Deep Wells by Aquifer Drilled in the Past

The number of deep wells by aquifer is shown in Table-3.4. In the past, many deep wells

were drilled in the area covered by Barreiras Formation, and in the area of Itabaiana Dome, Estancia Belt and Vaza-Barris Belt.

Hydrogeological unit Number Hydrogeological unit Number Quaternary 13 Vaza-Barris Domain 599 Area covered by 247 1,078 Macurure Domain **Barreiras Formation** Sergipe Basin 272 Maranco Domain 108 Tucano Basin 97 Sao Francisco Craton 25 Estancia Dómain 603 Itabaiana Dome Craton 950

Table-3.4 Number of Deep Wells Drilled by Aquifer

3.3 Groundwater Field Survey

Groundwater level and water quality were observed for all Sergipe State by the Study Team. The observation was carried out twice, the first observation was in September 1998, and the second one was in November 1998. Observed items were below:

- Groundwater level
- Water quality: pH, temperature, electric conductivity, dissolved oxygen

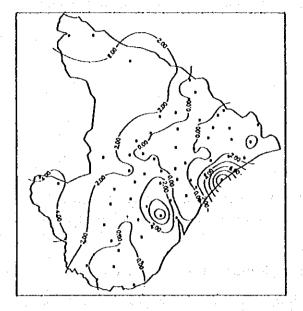
All the results of the groundwater survey are attached to Appendix-2.

(1) Groundwater Level

The observation was done twice in different season in order to obtain groundwater level fluctuation. Groundwater level was measured for 70 deep wells and 30 shallow wells. The result is shown in Figure-3.2 (a)-(b). As shown in this figure, groundwater level fluctuation is about 1m to 2 m in the whole Sergipe State during the observation period. The observation wells sometimes showed locally unusual large fluctuation which seemed to be caused by pumping. These unusual groundwater fluctuations were easily identified by comparing them with the other normal fluctuations.

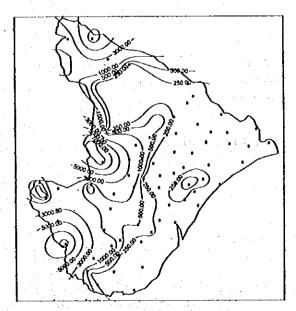
(2) Water Quality

Measured electric conductivity is shown in Figure-3.2 (c)-(d). Electric conductivity has strong relationship with salinity. The relation is approximated as Conductivity (μ S/cm) = 0.5 x Cl (ppm). As shown in Figure-3.2 (c)-(d), electric conductivity, namely salinity, becomes gradually higher from the coastal area to the inland area.

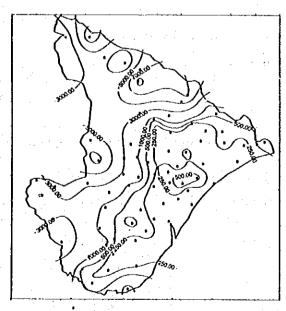


a) Groundwater Level Draw Down of Deep b) Groundwater Level Draw Down of Shallow Wells during Sep. and Nov. 1998 (m)

Wells during Sep. and Nov. 1998 (m)



c) Electric Conductivity of Groundwater in Sep. 1998 (µS/cm)



d) Electric Conductivity of Groundwater in Sep. 1998 (μS/cm)

Figure-3.2 Result of Field Groundwater Survey

CHAPTER 4 GROUNDWATER POTENTIAL

Groundwater exists in whole Sergipe State, though its quantity and quality is different in each site. Groundwater development is possible in any place in Sergipe State depending on groundwater development potential of each place. Groundwater development potential is dominated by three factors shown below;

- Groundwater recharge
- -- Well capacity
- Groundwater quality

Items above are examined below.

4.1 Modified Aquifer Classification

This Study used deep well data-base established by SRH. Hydrogeological classification for the Master Plan formulation was made as shown in Table-4.1 taking account of effective utilization of the data-base.

Table-4.1 Hydrogeological Classification for the Master Plan

Hydrogeological Unit					
Alluvium covering Sergipe Basin	Maranço Domain				
Alluvium covering Sao Francisco Craton	Macurure Domain				
Tucano Basin	Vaza-Barris Domain				
Sergipe Basin covered by Barreiras	Estancia Domain				
Sergipe Basin outcropping	Sao Francisco Craton covered by Barreiras				
Caninde Domain	Sao Francisco Craton outcropping				
Poco Redondo Domain	Itabaiana Dome Craton				

4.2 Groundwater Recharge

Groundwater recharge is analyzed by two methods below.

Method - (I) :

Analysis of groundwater level fluctuation

Method - (II)

Numerical Simulation

4.2.1 Method-(I): Analysis of Groundwater Level Fluctuation

(1) Principle of Method-(I)

Principle of Method-(I) is shown in Figure-4.1.

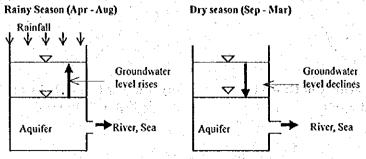


Figure-4.1 Principle of Method-(I)

Annual groundwater level fluctuation suggests the change in annual amount of groundwater stored in aquifers which is provided by rainfalls. Consequently, annual groundwater recharge can be calculated by formula below;

$$R = u \times dh \times F$$

Where,

R: Annual groundwater recharge (m³/year) u: Specific yield of aquifer dh: Annual groundwater level fluctuation (m) F: Area of aquifer (m²)

(2) Specific Yield

There are no existing data showing representative values of specific yield in the Study area. Therefore, the values are assumed using general values with wide range by each geological unit, and the result is shown in Table-4.2.

' i			
Aquifer	Specific yield	Aquifer	Specific yield
Alluvium covering Sergipe	0.15 - 0.20	Maranco Domain	0.005 - 0.01
Alluvium covering Craton	0.10 - 0.20	Macurure Domain	0.005 - 0.01
Tucano	0.05 - 0.10	Vaza-Barris Domain	0.03 - 0.05
Sergipe covered by Barreiras	0.05 - 0.15	Estancia Domain	0.01 - 0.02
Sergipe outcropping	0.10 - 0.15	Craton covered by Barreiras	0.05 - 0.15
Caninde Domain	0.005 - 0.01	Sao Francisco Craton outcropping	0.03 - 0.05
Poco Redondo Domain	0.005 - 0.01	Itabaiana Dome	0.03 - 0.05

Table-4.2 Specific Yield of Aquifer

(3) Annual Groundwater Level Fluctuation (dh)

Therefore annual groundwater level fluctuation (dh) was assumed using the result of groundwater level survey carried out by the Study Team. However, groundwater level fluctuation was observed for only 3 month between Sep. and Nov. 1998 in this survey. Therefore, annual groundwater level fluctuation must be assumed from groundwater level fluctuation observed for 3 month. However, there is no such a theoretical methods as foreseen the annual groundwater fluctuation from short time fluctuation without long term data. Consequently only an empirical method was applicable for this purpose with assumed pattern of annual groundwater fluctuation shown in Figure-4.2 which was assumed from rainfall pattern of the Study area. After careful examination by which abnormal data were excluded, annual groundwater level fluctuation was assumed as shown in Table-4.3.

dh = dh' × (between 2 and 3)

where dh: Annual groundwater level fluctuation,

dh': Groundwater level fluctuation during Sep. and Nov. 1998

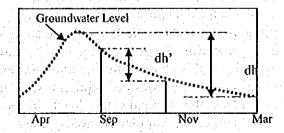


Figure-4.2 Groundwater Fluctuation Pattern Empirically Assumed

Table-4.3 Presumed Annual Groundwater Level Fluctuation

Hydrogeological Unit	Observed groundwater level fluctuation during Sep. to Nov.	Assumed annual groundwater level fluctuation	
Sedimentary Rock Area	0.7m	1.4m - 2.1m	
Crystalline Rock Area	0.8m	1.6m – 2.4m	

(4) Result of Analysis by Method-(I)

Annual groundwater recharge was analyzed, and the result is shown in Table-4.4.

Table-4.4 Annual Groundwater Recharge by Method-(I)

Hydrogeological Unit	Area	Annual rainfall	Annual Recharge			
	Km²	mm/y	mm/y	lit/s/km²	m³/s	% of annual rainfall
Alluvium covering Sergipe	1,061	1,398	210 - 420	6.7 - 13.3	7.1 - 14.1	15.0 - 30.1
Alluvium covering Craton	434	1,672	140 - 420	4.4 - 13.3	1.9 - 5.8	8.4 – 25.1
Tucano Basin	310	613	70 - 210	2.2 - 6.7	0.7 - 2.1	11.4 - 34.3
Sergipe covered by Barreiras	2,688	1,271	70 - 315	2.2 - 9.9	6.0 - 26.8	5.5 – 24.8
Sergipe outcropping	962	1,160	140 - 315	4.4 - 9.9	4.3 - 9.6	12.1 - 27.2
Caninde Domain	854	521	8 - 24	0.25 - 0.76	0.2 - 0.6	1.5 - 4.6
Poco Redondo Domain	1,050	570	8 - 24	0.25 - 0.76	0.3 - 0.8	1.4 - 4.2
Maranco Domain	569	639	8 - 24	0.25 - 0.76	0.1 - 0.4	1.3 - 3.8
Macurure Domain	4,909	785	8 - 24	0.25 - 0.76	1.2 - 3.7	1.0 - 3.1
Vaza-Barris Domain	2,656	972	48 - 120	1.52 - 3.81	4.0 - 10.1	4.9 – 12.3
Estancia Domain	2,391	921	16 - 48	0.51 - 1.52	1.2 - 3.6	1.7 - 5.2
Craton covered by Barreiras	2,092	1,425	70 - 315	2.2 - 9.9	4.6 - 20.9	4.9 – 22.1
Craton outcropping	1,435	1,205	48 - 120	1.52 - 3.81	2.2 - 5.5	4.0 – 10.0
Itabaiana Dome	639	1,082	48 - 120	1.52 - 3.81	1.0 - 2.4	4.4 – 11.1
Total	22,050	1,015	7 F g		34.8 - 107	9 – 15.0

4.2.2 Method-(II): Numerical Simulation

(1) Principal of Method-(II)

This method is the same as the method-(II) in principle. The groundwater level near the ground surface fluctuates in accordance with the amount of the recharge from rainfall into the ground. It is reasonable to consider the aquifers of crystalline rock as unconfined aquifers, hence the assumption described above will be allowed. On the other hand, in sedimentary rock area such as Sergipe basin, where many layers with great thickness constitute complicated multiple aquifers with confined condition, it is considered that the groundwater recharge from the rainfall may infiltrate into the ground through long and complicated passes to finally reach each confined aquifer. Even in such multiple aquifers, however, the amount of groundwater recharge to the multiple aquifers can be estimated by groundwater recharge to the unconfined aquifer which are located on the top of the multiple aquifers. Consequently in this study, only the aquifers near the ground surface, which are allowed to be considered as unconfined aquifers, were modeled for numerical simulation which can calculate the groundwater level in accordance with given groundwater recharges. The numerical simulation were repeated until the calculated groundwater level fitted to the actual observed groundwater level with enough accuracy,

then the given groundwater recharge at the moment was adopted as the solution of this simulation.

Procedure of the simulation is as follows;

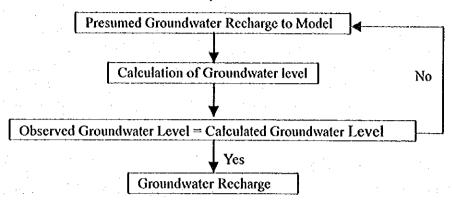


Figure-4.3 Procedure of Simulation

(2) 😘 Aguifer Model

(a) Simulation Program

A program for the numerical simulation which was used for this study was MODFLOW, the three dimensional finite difference model which was developed by USGS(United Nation Geological Survey). This program can solve the basic differential equations, constituted by the combination of the Darcy formula low and fundamental continuity equation, under the given boundary and initial conditions, like many conventional simulation models.

(b) Mesh of Modeling

All the Sergipe state was covered by a great number of square meshes. The total number of the meshes was 69 x 73=5,371 and the size of the meshes was 3.15 km x 3.15 km. Moreover, the meshes were vertically divided to represent superposition of layers. An element divided by the meshes is called as a 'Cell' which is individually given hydraulic conductivity and layer thickness. All the calculations are performed at the center of each cell and calculated results are also given to the center of the cell.

(c) Layer of Modeling

Simplified simulation model was developed with the hydrogeological unit following the conventional geological classification (Geologia e Recursos Minerars do Estado de Sergipe, 1988). The ground surface with topographical variety was roughly reproduced on the simulation model referring topographical maps. The model consists of four layers in the every parts (thickness; 1st~3rd layer~50m, 4th layer~250m). There is no theoretical background that each layer should has thickness above. However as explained later, hydraulic conductivity of the first layer were obtained from specific capacity of the existing deep well with the average depth of about 50m. Consequently the hydraulic conductivity obtained by method above should be applied for the 1st layer with 50m thickness. On the other hand, almost no data were available for 2nd~4th layer compared with the 1st layer. The 2nd~4th layer with 50m and 250m thickness were applied for necessity requested from hydrogeological judgment and numerical simulation technique. As the

result, the model can analyze groundwater flow within the aquifers of 400m thickness from the ground surface. In this simulation, aquifer were modeled as unconfined aquifers, and groundwater flow were analyzed within the aquifers of relatively shallow parts of the ground which is subject to rainfall. On the contrary, Sergipe basin consists of multiple layers which constitute confined aquifers with the total thickness of more than 1,000m. It is clear that the method described above can not cover groundwater flow throughout the Sergipe basin which is of more than 1000m thickness. This study, however, were interested in the groundwater movement which is driven in repose to rainfall, so only the uppermost part of the Sergipe basin, which may be possibly regarded as unconfined aquifers, were analyzed as aquifer which was 400m in total thickness and was divided into four layers of 50m~250m thickness each.

(d) Hydraulic Conductivity

As permeability of the aquifers varies wide in place by place, numerical simulation for large area needs huge number of data with high accuracy. The purpose of the study, however, is to get rough assessment concerning groundwater recharge for all the Sergipe state within the limited existing data. Consequently, the simulation model should be simplified in accordance with accuracy of the existing data. There are few detailed hydraulic parameters obtained by pumping tests (e.g. Avaliacao dos aquiferos da Bacia Sergipe/Alagoas Entre Aracaju e Capela, 1998 etc.), however these data are too limited to reflect the hydrogeological characteristics of large area. On the contrary, specific capacity of deep well is tested for almost all the deep wells all over the Sergipe State. It is generally admitted that specific capacity has strong relationship with transmissivity of wells, that is = specific capacity = transmissivity. Consequently representative transmissivities are assumed for each hydrogeological unit from the representative specific capacities which were obtained statistically from deep well data-base established by SRII. Subsequently hydraulic conductivity was calculated from the transmissivity for each hydrogeological unit. The aquifers in the simulation model were subdivided into 4 aquifers.

- 1) Sedimentary rock area
 - In Sedimentary rock area (Sergipe basin), hydraulic conductivities were calculated from specific capacities which were obtained from the well data-base, then hydraulic conductivities were given to each cell of the simulation model which constitutes each hydrogeological unit. The subdivided four layers have the same hydraulic conductivity.
- 2) Crystalline rock area
 Hydraulic conductivities have different values among the 1st layer to 4th layer as explained below;
 - 1st layer (from ground surface to GL-50m)

Representative specific capacities were obtained from the well data-base, then hydraulic conductivity were calculated and given to each cell which constitutes each hydrogeological unit of the model.

2nd layer (GL-50~GL-100m)

The 2nd layers were given the hydraulic conductivities of 1/10 of the 1st layers except the areas covered by Barreiras formation, where the 2nd layers were given the same hydraulic conductivities as the 1st layers.

3rd layer and 4th layer (GL-100m~GL-400m)

The 3rd and 4th layers are regarded as almost impermeable, then hydraulic conductivity of 0.0001m/day were given to all the 3rd and 4th layers.

Discussion above is concerning horizontal hydraulic conductivity. On the other hand, vertical hydraulic conductivity were set as 1/10 of horizontal hydraulic conductivity for sedimentary rock area and 1.0 of horizontal hydraulic conductivity for crystalline rock area in experimental aspect.

(e) Boundary Condition

Boundary condition must be set along all the border of the model and sometimes within the model. Two types of boundary condition are usually applied in accordance with hydrogeological situation, these are 'groundwater head is constant' and 'groundwater flux is constant' along the border. These boundary conditions are indispensable for performance of numerical simulation. Boundary conditions were set as described below:

Border with the Atlantic Ocean

The boundary condition that groundwater head is constant was given along the border with the Atlantic ocean (see Appendix-(III)).

River

Groundwater head is constant. The 6 major rivers and the other main rivers in Sergipe state were set boundary condition in this analysis, of which the Sao Francisco river and the Real river constitute the border with this boundary condition and the others are located inside the model with boundary conditions (see Appendix-(III)).

Water Shed which lies between Sergipe State and Bahia State

Water shed which lies between Sergipe state and Bahia state were given the boundary condition that groundwater flow is zero(0). This boundary with this condition means the groundwater-shed through which groundwater can not flow.

All the border of the model were fixed with one of above three conditions. In addition some rivers which exist inside the model were also given internal boundary condition that groundwater head are constant along the river course.

(f) Groundwater Recharge from Rainfall

The purpose of this simulation is to get assessment of groundwater recharge from rainfall. Many rainfall data sets were assumed for unknown actual groundwater recharge and given to the simulation model, and groundwater recharge with high reality was finally chosen from input data set which gave the highest correspondence between the calculated groundwater level and the actual observed groundwater level.

(g) Criteria for Completion of Numerical Simulation

The groundwater level were calculated at each cell. As mentioned above, after the comparison between the calculated groundwater level and the observed groundwater level, the simulation was completed when both of groundwater levels gave high correspondence. For the criteria of the completion of numerical simulation, it is not proper to compare individually calculated groundwater levels with the actual observed ones which are distributed in wide area and include seasonal fluctuation and artificial fluctuation. It is

better to employ statistical method, that is to compare the histograms of groundwater levels which are made of the calculated ones and the actual observed ones. The histogram, which was made of 4,000 drilling results, shows the smooth frequency curve and indicates that the majority of deep wells have groundwater levels of shallower than GL-15m. This histogram may suggest that groundwater near the ground surface may be approached as usual unconfined groundwater, and discontinuities of groundwater levels may not be taking place. Hence the method mentioned above may be applicable.

The aquifer model is outlined as shown in Table-4.5 and Appendix-(III).

Table-4.5 Simplified Aquifer Model

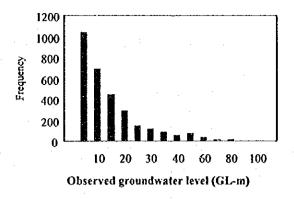
ltem	Content						
Aquifer model	4 layer model: Permeability of layers gradually reduces from the 1 st layer to the 4 th layer.						
Boundary condition	There are three types of boundary conditions. - Sea : Constant groundwater level boundary - Water shed : No groundwater flow boundary - River : Groundwater discharge boundary						
Conductivity	Conductivity is examined for each aquifer from specific capacities of existing boreholes recorded in the data-base which stores 4,000 borehole data						

(3) Result of Method-(II)

The simulation was completed when simulated groundwater level fitted with observed one, and groundwater recharge at this time was considered to be the most likely groundwater recharge. The groundwater recharge at this time is shown in Table-4.6. Histogram of observed groundwater levels and simulated ones is shown in Figure-4.4. Both of the histograms look to correspond enough to each other. Histogram of observed groundwater level is obtained from borehole data-base which shows actual groundwater level during drilling.

Table-4.6 Groundwater Recharge by Method-(II)

	Area	Annual rainfall	i I Annual aroundwater recharge			
Hydrogeological Unit	km²	mm/y	mm/y	lit/s/km²	m³/s	% of Annual Rainfall
Alluvium covering Sergipe	1,061	1,398	360	11.42	12.11	25.8
Alluvium covering Craton	434	1,672	170	5.39	2.34	10.2
Tucano Basin	310	613	110	3.49	1.08	18.0
Sergipe covered by Barreiras	2,688	1,271	250	7.93	21.31	19.7
Sergipe outcropping	962	1,160	280	8.88	8.54	24.1
Caninde Domain	854	521	10	0.32	0.27	1.9
Poco Redondo Domain	1,050	570	10	0.32	0.33	1.8
Maranco Domain	569	639	10	0.32	0.18	1.6:
Macurure Domain	4,909	785	15	0.48	2.34	1.9
Vaza-Barris Domain	2,656	972	60	1.90	5.05	6.2
Estancia Domain	2,391	921	30	0.95	2.27	3.3
Craton covered by Barreiras	2,092	1,425	90	2.85	5.97	6.3
Craton outcropping	1,435	1,205	60	1.90	2.73	5.0
Itabaiana Dome Craton	639	1,082	80	2.54	1.62	7.4
Total	22,050	1,015	95	3.00	66.15	9.3



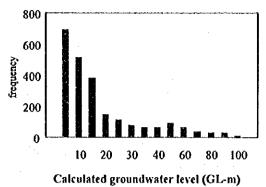


Figure-4.4 Histogram of Groundwater Level

4.2.3 Examination of Groundwater Recharge

Groundwater recharge was estimated by two methods as explained before. The result of the method-(I) shows a wide range of values, the annual groundwater recharge of 34.8m³/s – 107.0 m³/s with the average of 70.9m³/s, as shown in Table-4.4. On the other hand, the result of the method-(II) has a single value, the annual groundwater recharge of 66.15 m³/s, as shown in Table-4.6. Comparing the two results, result of the method-(I) is, if its average is taken, almost the same as the result of method-(II) with negligible difference between them. Consequently, in this report, the result of the method-(II) is finally applied as value of groundwater recharge, though the actual groundwater recharge is considered to be between 34.8m³/s – 107.0 m³/s. The result is shown in Table-4.7. The numerical simulation model employed for this Study is a simplified model to obtain a rough estimation of groundwater recharge for the Master Plan. More detailed simulation model constructed by more detailed data is necessary for farther study of groundwater recharge of individual groundwater development.

Table-4.7 Annual Groundwater Recharge

	Area	Annual rainfall	Annual ground water recharge				
Hydrogeological Unit	Km²	mm/y	mm/y	lit/s/km²	m³/s	% of Annual Rainfall	
Alluvium covering Sergipe	1,061	1,398	360	11.42	12.11	25.8	
Alluvium covering Craton	434	1,672	170	5.39	2.34	10.2	
Tucano Basin	310	613	110	3.49	1.08	18.0	
Sergipe covered by Barreiras	2,688	1,271	250	7.93	21.31	19.7	
Sergipe outcropping	962	1,160	280	8.88	8.54	24.1	
Caninde Domain	854	521	10	0.32	0.27	1.9	
Poco Redondo Domain	1,050	570	10	0.32	0.33	1.8	
Maranco Domain	569	639	10	0.32	0.18	1.6	
Macurure Domain	4,909	785	15	0.48	2.34	1.9	
Vaza-Barris Domain	2,656	972	60	1.90	5.05	6.2	
Estancia Domain	2,391	921	30	0.95	2.27	3.3	
Craton covered by Barreiras	2,092	1,425	90	2.85	5.97	6.3	
Craton Outcropping	1,435	1,205	60	1.90	2.73	5.0	
Itabaiana Dome Craton	639	1,082	80	2.54	1.62	7.4	
Total	22,050	1,015	95	3.00	66.15	9.3	

4.3 Well Capacity and Water Quality

Well capacity and water quality were analyzed using the existing data-base. The result is shown in Table-4.8.

Table-4.8 Well Capacity and Water Quality by Aquifer

Hydrogeological Unit	Expected yield (m³/day)	Specific capacity (m³/day/m)	Success rate (%)	Rate of fresh water (%)
Alluvium covering Sergipe Basin Alluvium covering Craton Basin	600	140	95	100
Tucano Basin	100	4	60	60
Sergipe Basin covered by Barreiras	s 140 see	17	80	85
Sergipe Basin outcropping	140	13	70	60
Caninde Domain	40	2	45	10
Poco Redondo Domain	40	2	45	10
Maranco Domain	40	2	45	10
Macurure Domain	40	2	60	15
Vaza-Barris Domain	80	4	75	40
Estancia Domain	50	3	70	50
Craton covered by Barreiras	70	4	85	90
Sao Francisco Craton outcropping	40	2	75	30
Itabaiana Dome	70	4.4	75	35

Note:

'Fresh water' means chlorine (Cl) is less than 250 ppm.

4.4 Groundwater Development Potential

Groundwater development potential is shown in brief by basin in Table-4.9, and in detail by municipality in Table-4.10.

Table-4.9 Groundwater Potential by River Basin

Basin	Area	Total amount with consideration	out water quality	Total amount wi consideration (C	
	(km²)	mm/year	mil m³/year	mm/year	mil m³/year
Sao Francisco	7,276.3	84	611	61	444
Japaratuba	1,722.0	152	262	113	195
Scrgipe	3,673.0	131	481	91	334
Vaza Barris	2,559.0	99	253	64	164
Piaui	4,262.0	80	341	56	239
Real	2,558.0	54	138	30	77
Total	22,050.3	95	2,086	66	1,453

Using Table-4.10 for groundwater development plan, items below should be noted.

- Some municipalities locate not only in a single geological unit, but also in several geological units. Such a situation is already took into consideration by weighted mean method in all the parameters shown in Table-4.10,
- Success rate shown in Table-4.10 is the success rate of well with yield of more than 8 m³/day. This success rate is, according to data-base analysis, almost the same as the success rate of well with yield of more than 0. Therefore, the success rate can be used for rural groundwater development plan, and at same time for urban groundwater development plan with the expected yield in Table-4.10.

Table-4.10 Groundwater Development Potential by Municipality

	Table-4.10 Grou	Huwater 1	JUYUI	opine	111 1 01	ichthal i	Jy IVAUII	пстран	ıy	
					nual	Expected	Specific	Success	Rate of	
Cođe	Municipality	Main aquifer	Area	1	dwater harge	Yield	Capacity	Rate	Fresh	Rank
			(km²)	(mm/y)		(m³/day)	(m³/day/m)	(%)	Water (%)	ŀ
01-0120	CANINDE DO SAO FRANCISCO	Caninde	908	27	776	40	3	50	15	D
01-0220	FEIRA NOVA	Macurure	189	15	92	40	2	60	15	Ď
01-0240	GARARU	Macurure	640	15	313	35	2	60	15	Ď
01-0260	GRACHO CARDOSO	Macurure	236	15	115	40	2	60	15	D
01-0310		Macurure	202	15	99	40	2	60	15	D
	MONTE ALEGRE	Macurure	418	15	204	40	2	60	15	D
01-0450	NOSSA SENHORA DA GLORIA	Macurure	745	15	364	40	2 .	60	15	D
01-0540 01-0560	POCO REDONDO PORTO DA FOLHA	Poco Redondo Poco Redondo		14 20	533	40	2 3	45	10	D
02-0140		Macurure	895 634	32	555 642	40 50	4	50 60	$\frac{10}{20}$	D D
	FREI PAULO	Vaza-Barris	406	67	860	75	8	75	35	c
02-0145		Macurure	347	39	430	60	5	70	25	Ď
02-0500	PEDRA MOLE	Vaza-Barris	79	62	154	80	Ž	75	40	c.
	PINHAO	Vaza-Barris	152	62	298	80	7	75	40	С
02-0600		Vaza-Barris	263	61	507	75	7	75	40	С
03-0020	AQUIDABA	Macurure	370	15	181	40	2	60	15	D
	CUMBE MALHADA DOC DOLC	Macurure	131	15	64	40	2	60	15	D
03-0380	MALHADA DOS BOIS MURIBECA	Macurure Sergipe	59 82	70 150	132 389	60 90	9	65 65	25 40	D.
		Macurure	482	32	489	50	4	65	25	C D
	SAO MIGUEL DO ALEIXO	Vaza-Barris	143	45	206	65	5	70	30	b D
04-0050		Vaza-Barris	129	209	857	150	17	75	65	A-B
04-0100	CAMPO DO BRITO	Dome .	200	75	476	70	5	75	35	c
	ITABAIANA	Dome	338	80	854	70	5	75	35	С
	MACAMBIRA	Vaza-Barris	137	65	285	75	6	75	40	C
	MALHADOR	Vaza-Barris	102	95	308	85	9	75	45	C
01-0410 01-0680		Dome Vaza-Barris	95 102	74 75	223	70	5	75 80	35 50	C
	POCO VERDE	Tucano	380	81	243 982	75 60	5 8	65	55	C
05-0330	SIMAO DIAS	Vaza-Barris	560	55	978	70	6	75	40	D
05 0740		Estancia	1119	33	1,170	50	3	70	50	Ιŏ
06-0350	LAGARTO	Estancia	962	57	1,740	60	4	75	55	č
06-0580		Estancia	528	46	776	45	3	75	40	Ď
	AMPARO DE SAO FRANCISCO	Macurure	39	15	. 19	40	2	60	15	D
07-0070	BREJO GRANDE	Q/S	149	369	1,747	600	90	95	100	A
07-0110	CANHOBA	Macurure	165	15	81	40	2	60	15	D
07-0160	CEDRO DE SAO JOAO	Macurure	73	89	207	65	6	65	25	Ð
07-0270 07-0440	ILHA DAS FLORES NEOPOLIS	Q/S Sergipe	57 249	369 288	675 2,281	600 250	90 34	- 95 80	100 85	A A-B
07-0170	NOSSASENHORA DE LOURDES	Macurure	80	200 15	. 39	250 40	2	60	15	D D
07-0570	PROPRIA	Sergipe	95	262	794	300	42	75	65	A·B
07-0730	TELHA	Macurure	56	37	67	45	4	60	15	D
07-0999	SANTANA DO SAO FRANCISCO	Sergipe	47	273	406	180	23	80	85	A-B
08-0130	CAPELA	Macurure	431	45	615	50	4	60	20	Đ
		Sergipe	93	273	806	170	19	75	70	A-B
08-0650	SANTA ROSA DE LIMA	Vaza-Barris	66	122	256	95	10	75	50	C
08-0720	JAPARATUBA	Sergipe	167	238	1,262	175	23	80	80	A·B
	JAPOATA	Sergipe	374 397	278 269	3,299	215	29 19	80 80	85 80	A A-B
09-0490	PACATUBA	Sergipe Sergipe	407	323	3,391 4,167	160 405	59	90	90	A-B A
09-0530	PIRAMBU	Sergipe	199	317	2,000	385	56	85	90	Ã
09-0690	SAO FRANCISCO	Sergipe	86	275	758	140	15	75	70	A-B
10-0150	CARMOPOLIS	Sergipe	40	305	386	295	41	80	85	A
10-0250	GENERAL MAYNARD	Sergipe	18	289	166	180	21	75	70	A
10-0360	LARANJEIRAS	Sergipe	163	302	1,564	235	30	75	70	A
10-0100	MARUIM	Sergipe	95	292	882	195	23	75	70	A-B
10-0590	RIACHUELO ROSARIO DO CATETE	Sergipe Sergipe	78 103	293 302	731 993	180 260	20 34	75 80	65 75	A·B A·B
10-0610	SANTO AMARO DAS BROTAS	Sergipe Sergipe	237	323	2,434	395	31 57	80 85	90	A-B A
11-0030	ARACAJU	Sergipe	181	330	1,904	440	65	90	95	A
11-0060		Q/S	87	369	1,030	600	90	95	100	Â
11-0480	NOSSA SENHORA DO SOCORRO	Sergipe	157	305	1,518	285	39	80	80	A·B
11-0670		Sergipe	432	276	3,790	215	29	8Ô	85	Λ
12-0040	ARAUA	Craton	194	72	444	50	3	80	50	C
12-0067	BOQUIM	Craton	213	76	514	60	3.	80	60	C
12-0170 12-0300	CRISTINAPOLIS	Craton	251	71 65	564	50	3	80	60	C
12-0300	ITABAIANINHA PEDRINHAS	Craton Craton	480 39	65 82	991 103	45 60	2 3	80 80	40 70	C B
		Craton	255	88	716	65	4	80	85	В
12-0620	LOMINIADA				535	45	3	75	45	Ď
12-0620 12-0750	TOMAR DO GERU	Estancia	337	50	030.1	40			7.0	
		Estancia Craton	337 124	81	319	55	3	80	70	В
12-0750 12-0760 13-0210	TOMAR DO GERU									B
12-0750 12-0760 13-0210 13-0280	TOMAR DO GERU UMBAUBA ESTANCIA INDIAROBA	Craton Craton Craton	124 649 311	81 114 108	319 2,347 1,067	55 70 70	3 9 8	80 80 80	70 85 80	B B
12-0750 12-0760 13-0210	TOMAR DO GERU UMBAUBA ESTANCIA INDIAROBA ITAPORANGA D'AJUDA	Craton Craton	124 649	81 114	319 2,347	55 70	3 9	80 80	70 85	В

2) Main aquifer: 2) Rate of fresh water:

Q/S- Quaternary covering Sergipe Basin Fresh water means CI is less than 250 mg/L.

4.5 Total Evaluation of Groundwater Development Potential

Groundwater development potential was evaluated based on three factors, groundwater recharge, well capacity and water quality. Taking the three factors into the consideration, rank of groundwater potential by aquifer was evaluated, and the results is shown in Table-4.11 and Figure-4.5. The total evaluation gives important criteria on possibility of new groundwater development in the future.

Table-4.11		Evaluation	of	Groundwate	r	Potential !	by	Aq	uifer
IIIOIV IIII	- 1	***************************************	V.	Orounaniate		T Official	~,		*****

Hydrogeological Unit	Groundwater Recharge	Well Capacity	Water Quality	Total Evaluation
Alluvium covering Sergipe	Λ :	A - B	Α	A - B
Alluvium covering Craton	В	В	A	В
Tucano Basin	В	В	В	В
Sergipe covered by Barreiras	Λ	A - B	Λ	A - B
Sergipe outcropping	A	A - B	В	A - B
Caninde Domain	D	D	D	D
Poco Redondo Domain	D	D	D	D
Maranco Domain	D	D	D	14.1 D
Macurure Domain	D	D	D	D
Vaza-Barris Domain	С	C	C	С
Estancia Domain	D	D	С	D
Craton covered by Barreiras	С	С	Α	В
Sao Francisco Craton outcropping	С	D	C	C
Itabaiana Dome Craton	С	С	С	С

Note: 1) A - High, B - Medium, C - Low, D - Very low

²⁾ Barreiras Formation is excluded from Table above because of its poor capacity for deep wells

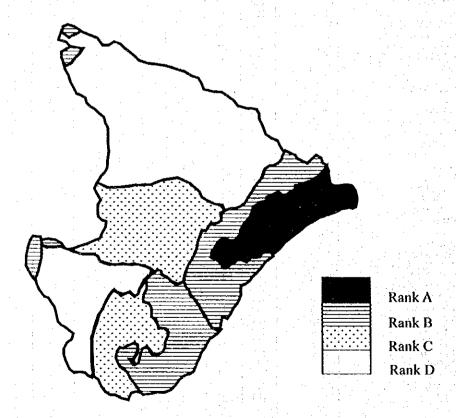


Figure-4.5 Rank of Groundwater Development Potential

CHAPTER 5 RECOMMENDATION

5.1 Promising Groundwater Development Site

Promising groundwater development sites are described below.

5.1.1 Alluvial Basin Aquifer

Alluvial aquifer is located in the coastal area. This aquifer has high permeability, and great deal of groundwater is possible to be pumped up from this aquifer. On the other hand, thickness of the Alluvium is not great, and there is a possibility that groundwater discharge by pumping exceeds groundwater recharge, and sea water intrudes into the aquifer. Therefore, proper well location / yield must be designed.

5.1.2 Sergipe Basin Aquifer

Sergipe Basin aquifer is the most promising aquifer in State. This aquifer expands in large area with high permeability and good water quality. On the other hand, this aquifer has not yet highly developed so far. Therefore, this aquifer is the most promising for new groundwater development. Figure-5.1 shows regional geological section including this aquifer. This aquifer sometimes locates deep in the ground, and drilling length must be 100 to 200m, which is longer than the existing wells. There is a possibility that groundwater discharge by pumping exceeds groundwater recharge, because the aquifer has high permeability. Sergipe Basin aquifer is divided into some formations, and promising ones are described below.

(1) Penedo Formation and Serraia Formation

Sandstone forms good aquifer in Penedo Formation and Serraia Formation. These aquifers locate in the northern part of Sergipe Basin. Muribeca, Malhada dos Bois and Japoata municipalities are located in this area.

(2) Sapucari Formation, Angico Formation, Maruim Formation and Agulhada Formation

Limestone of Sapucari Formation, Angico Formation, Maruim Formation and Agulhada Formation form good aquifers. These aquifers locate in the middle part of Sergipe Basin, in the north - east of Aracaju city, and now provides groundwater to Aracaju city as one of important water resources.

(3) Marituba Formation

Marituba Formation locates eastern-most part of Sergipe Basin along the coast. This Formation forms confined aquifer covered by Alluvium, and is composed of limestone and sandstone with total thickness of 500m (see Figure-5.2). This Formation forms the most excellent aquifer with the highest permeability of all the aquifers in Sergipe State. in this aquifer, wells need 100 to 200m depth for new groundwater development.

(4) Tucano – Jatoba Basin

Sao Sbrastiao Formation and Curituba Formation locate eastern-most part of Sergipe State. Sandstone of these formations forms good aquifer, and is important water resource because this formation is surrounded by Precambrian rock with low yield.

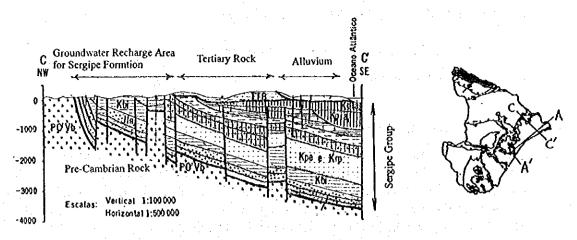


Figure-5.1 Regional Geological Section of Sergipe Basin

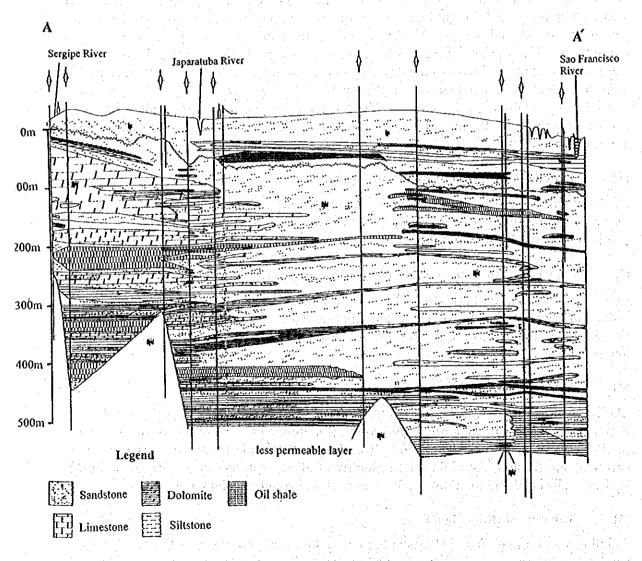


Figure-5.2 Geological Section of Marituba Formation

5.1.3 Crystalline Rock Aquifer

Precambrian aquifer is inferior to the other aquifers in terms of quantity and quality in Sergipe States, and groundwater development in Precambrian aquifer is possible on a small scale only for rural area. On the other hand, areas with higher development potential are locally identified as described below.

(1) Salgado, Lagarto, Estancia Area

Salgado, Lagarto and Estancia areas are located in Estancia Domain and Sao Francisco Craton. Lagarto Formation of Estancia Domain has boundary with gneiss of Sao Francisco Craton, and the boundary with many fissures sometimes forms good aquifer. However, Barreiras Formation covers this area wide, and the boundary is usually difficult to be found.

(2) Itabaiana Dome and Vaza-Barris Area

Gneiss of Itabaiana Dome forms good aquifers, and Frei Paulo Formation and Ribeiropolis Formation of Vaza-Barris Domain also form good aquifers.

5.2 Geophysical Method for Groundwater Survey

In Sergipe state, geophysical methods were not so far applied for groundwater survey to locate promising drilling sites. VLF Method, a kind of electromagnetic method which use special electromagnetic wave prevailing all over the world, is effective to locate new drilling sites for better success rate, especially in Precambrian rock. Fissure zones in rock can be easily and quickly detected by this method with low cost. This method, however, has not yet carried out in Sergipe State so far, therefore, the method may be recommended for new groundwater development. Other than VLF method, electric resistivity method is effective to make clear of geological situation, which lead to proper drilling plan in sedimentary rock area such as Sergipe basin.

5.3 More Detailed Groundwater Potential Assessment for Specified Area

In this study groundwater potential was roughly estimated for whole the Sergipe state in order to formulate the Master plan. However the data and hydrogeological model which were used for this estimation were not detailed ones. In the near future when groundwater development is planned on a large scale in specified area, the more detailed groundwater potential study is necessary including numerical simulations of groundwater movement in response to large scale pumping, which can assess the potential amount of groundwater to be developed without any environmental trouble.

5.4 Proper pumping Test for Aquifer Parameters

Pumping test, of course, is currently performed for every deep wells immediately after they are completed, however, this pumping test is aimed at getting rough estimation for its possible pumping rate without giving hydraulic parameters of its aquifer. It is necessary to obtain aquifer hydraulic parameters precisely for planning large scale groundwater development. For this purpose, execution of proper pumping test is necessary. Pumping test including step draw-down test should be performed with appropriate method.

5.5 Groundwater Level Monitoring

The groundwater monitoring including monitoring of groundwater level and groundwater quality is important in view points of both new groundwater development and environmental conservation. The Study Team performed groundwater field survey in 1998 which is considered a kind of groundwater monitoring. It is desirable that SRH or other adequate organization shall perform and continue groundwater monitoring. Groundwater level monitoring is indispensable for assessment of groundwater development potential and is also indispensable for effective counter-measures such as formulating regulations against environmental problems which are caused by inadequate groundwater use such as overpimping.

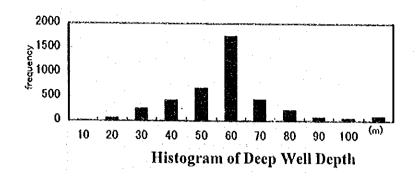
5.6 Groundwater Data-base

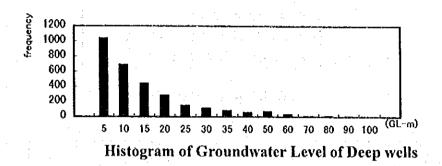
Data of more than 4,000 deep wells which were drilled in Sergipe state in the past are stored in a groundwater data-base established by SRII. This data-base was utilized for the Master Plan Study. The data stored in the data-base includes many items such as well owners, well locations, well dimension, well capacities and so on. This data-base can provide many kinds of useful information concerning hydrogeology and groundwater in Sergipe state. However it is considered that this data-base still need additional important items such as exact location and aquifer lithology/geological unit of the wells. These items are essential for hydrogeological study and may give the most important information for multiple purposes. SRH is now trying to add new items to the data-base including items above and in the near future this data-base may play significantly important roles in aspects of groundwater developments and groundwater environmental conservation, though the completion of the data-base needs great elaboration.

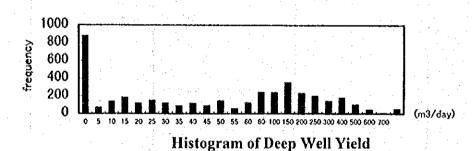
APPENDIX-1

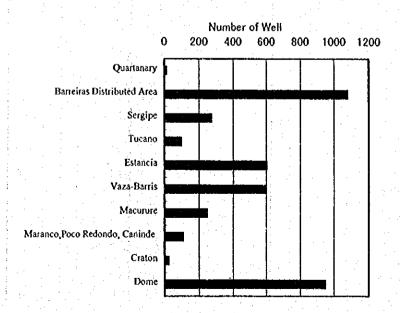
Current Groundwater Use and Hydrogeological Information

	e graja da karantar karantar
그리는 그리 어머니의 얼마를 하면 되었다면 하고 있다면 나는 것이 없는데 모든 것이다.	医乳球性皮肤 医苯甲基二酚
그는 그는 이 그들이 그는 전환에 하게 하면 하는 전 반속이다. 이 그리고 있는 것 같은 사이를 받았	
그들의 불자들은 그리고에 모르는 사람들은 보고 하는 것은 그리고 그리고 있다.	
그 집에 마음에 들면 내려가 들고 들고 들어가 그렇게 하고 이번 나면 가게 되어 있다.	
이 가는 이 가는데 걸어가 모든 눈이 하면 하는 이 사람들이 하는 것 같아.	
그 문의 문을 다듬어 하시다고 하는데 이 학생들이라고 살아 있다면 하시다는 사람은 밤	
그리다 살 살아 있는 일이 이 모두 있는 이 그리다면 하셨다는 것이라면 하면 하는 것이 하셨다면 나는	
그는 그 아이들은 사이에 가는 것 같습니다. 그들은 그 아이들이 가는 것은 것을 하는 것이다.	
그는 여러는 것들은 사고 이상되고 말아났는 말으로 사고 하다 젊을 사고를 받아 보는 말로	
그들은 노스에게 어느로 하나가는 그들은 얼룩하나 이렇지만 하는데 모든 목록을	
그는 이 나는 사람들이 다른 아는 이 불만하고 있다. 그리고 말했습니다. 그리고 그리고 그리고 있는 것으로	
一点,这一点,我们就是一个一点都没有,我们们的,我们们的一个一个一个一个一个一个一个一个一个一个一个一个,我们就是这个人,不是一个一个一个一个一个一个一个一个一	
그런데, 그 어디 아이를 살아보는 사람들이 사람들이 되었다. 그는	
그러지 그 사람은 등을 들었다면 회교를 받는 사람들은 가장 모양하는 사람들이 가능하는데 다	
	회사 등 경우는 기계 가는 그는
그리네 속이를 내려가 병하면 다른 한 시간을 하다면 살을 하다고 있다면 사람들이 하는데 되었다.	
그는 마이 마스 회사는 사람이는 그들은 살아 되다. 그 그는 사람들은 사람들이 모든 바람이 되었다.	
그리는 전 하는데 말로 할만하다는 하는데 하는 모든 모든 모든 하는데 하나 하고 모든데 들었어요?	과 등 경험을 되는 것을 보는 것이다. 역사 이 상실적은 것이 하는 것
그는 이 그는 얼마를 하다는 사람들은 사람들은 사람들이 있는 사람들이 가는 것 같아. 전 전환을 하다	
그들이라는 하다 나는 사람들이 되는 사람들이 하는 것이 없는 것이 없는 것이 없는 것이 없는 것이다.	
그 이렇게 한번을 보다 말이 하고 말을 하니까 그 살아 있다며 만든 맛이 되는 때문에 없는 말했다.	
그 사람이 하게 하면 모든 이 아이에게 하면 되었다. 그렇게 되었다. 하는데 나는데 나를 다 하다	
그들은 경화 이용하는 항상처럼 모양을 받아 아름답아 전략을 살려면 살고 있다.	
그 많은 하는데 이름은 살아보다 아들이 하는 것을 하는 것을 하는데 말했다고 말로 들었다.	
그가 얼마 집에 되지 않아 되었다고 한 회사가 하다면 하다는 모양 살아 하네요요 되었다.	
그리는 그림으로 하다는 모든 그는 하는 들으는 이는 살아서 이 자꾸 살아서 걸을 통해 들었습니다. 그	
그들은 아이, 아이는 말을 들었다. 아이는 말을 하는 사람들은 경험을 하는 것을 하는데 되었다.	등시대로를 걸는 그런 맛이
그리다 그리 일반에 되었다고 있다면 보았다는 하는 물 학생들 그릇 목욕에 들었다는데 되었다.	H [18] H [18] 프로젝트 [18] 프로
그 그 한 경험하는 사람들이 생각하다고 말하는 생각으로 모두 살아왔다. 사람은 생각을 다고 밝다.	
그 이번 그 일을 걸려 되고 있게 화막했다. 한 제 후 해 주는 사용 함께 함께 수 있다.	
그는 본문은 그의 이번 그들은 소리를 가지지고 하는 경기를 통해를 통해 가게 되었다. 그림은 사람들은 것은	
그림 그림으로 하는데 이번에 하는 그리고를 보면 하고 있었다. 그리고 있다면 하셨다고 하는데	
그는 나는 그 그는 선생님, 살을 내려왔다. 그는 그 때가 살린 사라 나를 맞지 않는 것을 받았다.	
그 그 그 그 그는 사람이 그 사이에 하루지만 되지만 말았다. 한국의 동시, 루시 글로, 본 경기 말했다.	
그리고 말이 있다. 그 하는 말리면이 보고 있다면 아니라 하고 있는 사람들은 사람들은 모든 살아왔다면 가장	
그리아, 사람들 마음이 하는 사람들은 것이 되었다면 하는 것 같은 이 중에 되었다면 얼굴을 하는 것이 되었다.	
그는 네이지는 그 사람들은 아이트를 된 방법을 하고 있었다면 방송은 회사들의 호텔을 다쳤다.	
그리는 그리는 항상 등을 되었다면 하는 것이 되었다. 그리는 그리는 그리는 것이 없었다. 그리는 것이 없는 것이 없는 것이 없었다.	
그 하는 경우를 들었다면 하는 것은 사람들이 가는 모양을 하고 있다면 가장 하는 것이 되었다.	
그는 사람이 되었다. 그는 사람들은 사람들은 사람들은 사람들은 사람들은 사람들은 사람들은 사람들은	경우리 하고 보고 살 있었다.
그러는 그는 마이 이 그를 맞는 것이 없는 것은 것이라면 그렇게 얼굴 바로 모든 것이 없었다.	
그 사람들은 사람들이 가는 사람들이 되었다. 그 사람들이 가장 사람들이 가장 사람들이 가장 하는 사람들이 되었다. 경우 사람들은 사람들이 되었다. 사람들은 사람들은 사람들은 사람들은 사람들은 사람들은 사람들은 사람들은	with white post in the development of the Control of Section (1997). We discuss the control of t

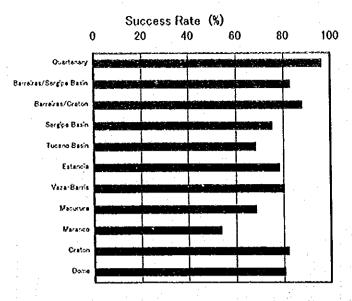




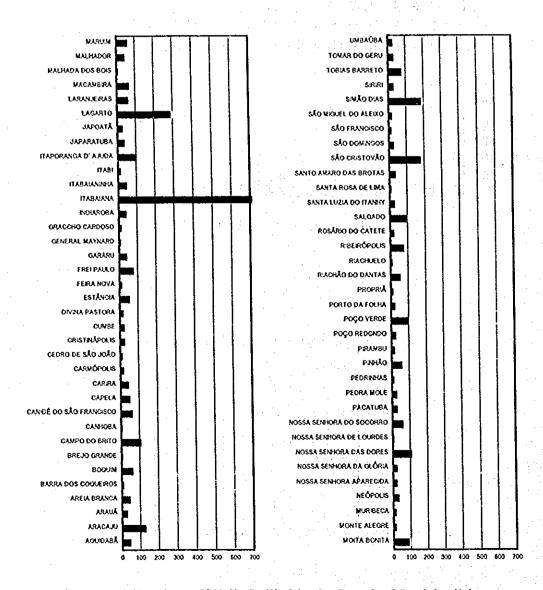




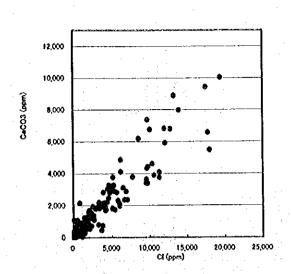
Number of Wells drilled in the Past by Aquifer

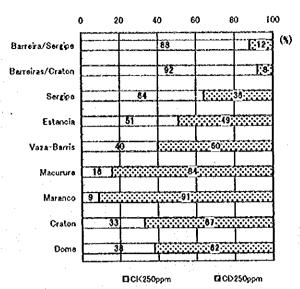


Well Success Rate By Aquifer



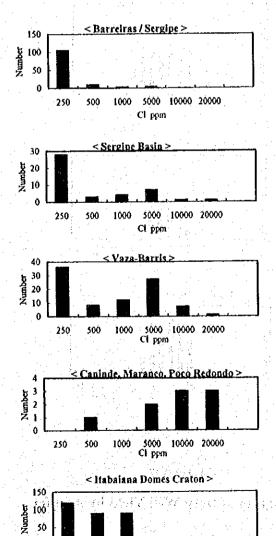
Number of Wells Drilled in the Past by Municipalities





Relationship between Cl and CaCO₃

Cl Concentration by Aquifer

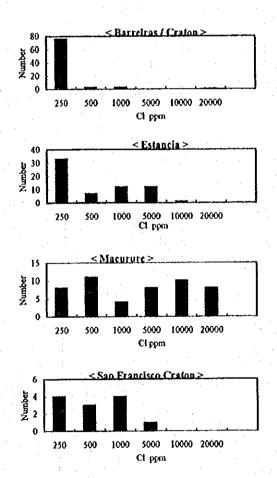


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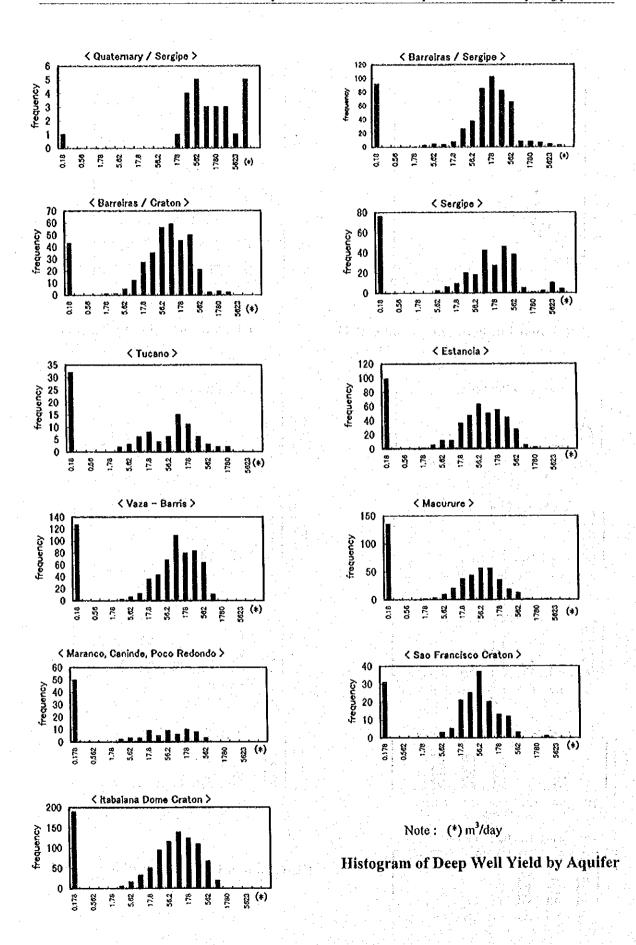
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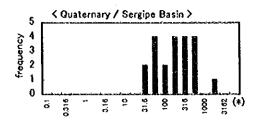
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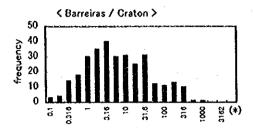
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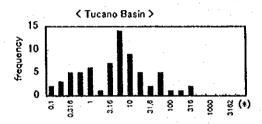


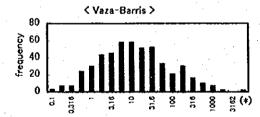
Histogram of Cl Concentration by Aquifer

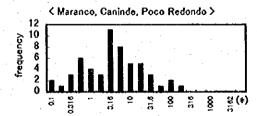


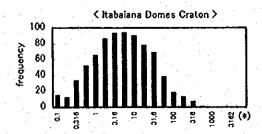


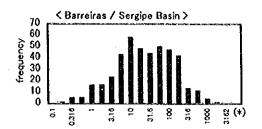


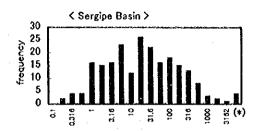


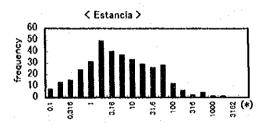


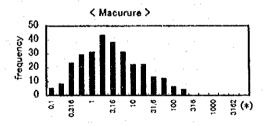


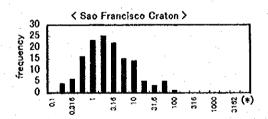










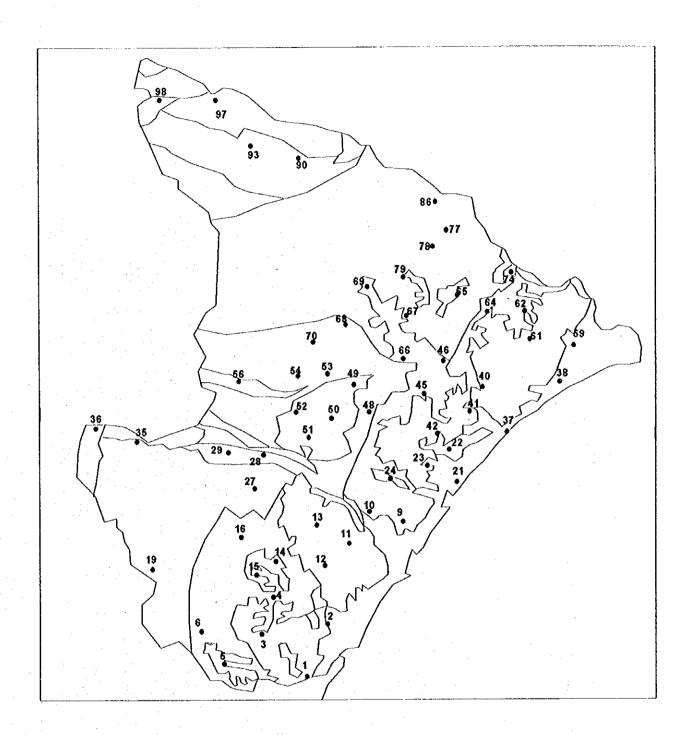


Note: (*) m³/day/m

Histogram of Deep Well Specific Capacity

APPENDIX-2

Result of Groundwater Field Survey



Water Quality Observation Point

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