of metamorphc rocks and granodiorite. They are fissure filling of shear fracture. There are no record of the distribution regulation type 3(NE system).

The results of the present survey shows that No.5 of type 3 (NE system) is considered to be the fissure filling of shear fracture since it is the alteration vein of sheared rock, however there are no sign that other systems(type 1, 2) suffered a shear. The widely distributed field of stress prior to the formation of quartz vein groups has its main stress running in E-W direction and the shear fracture occurred in NE system runs at an angle of 45 degrees to the main stress is represented by No. 5. The main vein and E-W system quartz veins of type 2 (NNW system) seems to be created along the tension fracture that occurred in the right angle to and nearly parallel to the main stress.

It may further be assumed that the dyke rocks such as quartz porphyry and lamprophyre had been regulated by the same field of stress although subject to the period of intrusion.

The distribution of quartz veins at the contact part of heterogenic rock may show that, even in the same field of stress, fractures are likely to occur.

#### 1-3 Analysis and test

Analysis and test in a semi-detailed survey include the thin sections of rocks, X-ray diffraction, chemical analysis of rocks and ores, polished sections of ores and the radioisotopic dating.

On the other hand, the following detailed survey include the thin sections of rocks, X-ray diffraction, chemical analysis of ores, the polished sections of orcs, radioisotopic dating of altered minerals and measurement of the homogenization temperature of fluid inclusions.

Some of the above are associated to both of the semi-detailed and detailed geological survey in describing. They are classified as follows and part of them is globally described. The results of the analysis and test of the drilling survey will be discussed in each chapter with exception of the measurement of the homogenization temperature of fluid inclusions.

# Items described in this chapter are,

Items described in the chapter of detailed survey

Thin sections of rocks (global description) X-ray diffraction (global description) Chemical analysis of rocks Radioisotopic dating

X-ray diffraction of specimens sampled across quartz vein Chemical analysis of ores (global description) Polished sections of ores (global description) Measurement of the homogenization temperature of fluid inclusions.

#### 1. Thin sections of rocks

Number of specimens: 32 in semi-detailed survey area

22 in detailed survey area

The results of the identification of the both areas are shown globally in Table II-1-2. The major findings are as discussed in the geological description (1-1 through 1-4).

#### 2. X-ray diffraction of powder

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Number of X-ray diffraction: Semi-detailed survey area: 22

#### Detailed survey area:

The thin sections of rocks are in general subjected to the x-ray diffraction test. The conditions of the X-ray diffraction are shown in Table II-1-3 and the results of mineral identification for the both area are shown together in Table II-1-4. For the quantitative description of mineral identified, quartz index from the calculation formula by Masao Kobayashi is used.

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Item	Specification
X-ray diffraction analyzer	Geiger flex (RIGAKU DENKI Ltd.)
Counter cathode	Cu
Filter	Ni
Tube voltage	30 KV
Tube current	15 mA
Full scales of count	2 0 0 0 cps
Time constant	1 sec
Scanning speed	2°/min
Recorder speed	2 cm/min
Divergence slit	1°
Scatter slit	1°
Receiving slit	0.3mm
Scanning range	$2 \theta = 2 \sim 40^{\circ}$

Table II-1-3 Conditions of X-ray diffraction

Identified alteration minerals are sericite, chlorite, smectite and calcite. Kaoline, pyrophillite and alunite that are characteristic to gold veins were not detected. Sericite are sometimes hard to distinguish from muscovite of rock-forming minerals. Rocks distributed in the region were subjected to the X-ray diffraction for several pieces of each kind of rock, the results of which were already described in the description of geology (1-1through 1-4). The summary for the entire region is as follows.

Sericite and chlorite were mostly found in gneiss excluding crystalline limestone. Chlorite, sericite (including rock form muscovite) were detected in two-mica granodiorite and biotite adamellite with a few exception.

Diorite has hardly shown a sign of being altered with the naked eye but some specimens were

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microscopic observation	Cordinate	. I 🗠	9.04	9.51	7.40	7.37	6, 10	6.04	9.42	9.50	11.96	12.40	12.40	12.40	12.30	12.41	10.88	1. /8	11. 91	10.25	11.44	13.16	17.77	12.04	11.04	11.96	11.0/	10.12	10.01	19 60	00.71	10 01	10.04	10.00	10.03	CQ . 21	12.02	11 47	19 60	12. 40
	CPS COT		8.25	8.45	9.12	9.05	8.63	8.43	8.11	7, 91	7.56	1.20	1.20	(.20	0.40	0.00	1.43	8. (5	1.32	8. 6/	6.10	0. 34	0.01	1. W	1.21	1.33	0.10	0.4J	0.00	2. 04 00	0. 33	0.41	0 42	1.41	7 41		0 00 00 0	6 01 5 07	0 10	0.15
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Table I-1-2 Result of microscopic observation of thin section (Semidetailed-detailed area) (1)

Longit.: longitude Latit.: latitude

GPS:global positioning system

+: strong +: medium -: weak

\* ◎ : abundance ○ : common △ : minor • : rare thbreviation of rock and mineral name

\* O : abundance

Adamil:adamellite Gr-Drigramodiorite Hobdt:hornblendite Ampbit:amphibolite Pegmatipegmatite Psamt:psammitic Pelit:pelitic Bi:biotite Hs.muscovite Q2:quartz Gt.garnet Au:augite Hy:hypersthene P1:plagioclase Kf:potassium feidspar Ho:hornblende O1:o1ivine Cc:calcite Ser:sericite Ch1:chlorite Ep:epidote Sph:sphene Apt:apatite Opq:opaque mineral metam.:metamorphism

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Table I-1-2 Result of the microscopic observation of thin section (Semidetailed-detailed area) (2)

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Rock name	(determined)	B1-MS-GR-DF	81-HS-GR-Dr	Adame I i te	Marble.	Basalt	Amphibolite	Peit Gneiss	Marble	Ri-Gneiss	Drumbini to	POPUINT LE	Psamt-Gheiss	8 i-HS-GR-DP	81-02-Dr	Psamt-Gneiss	Ms-Gneiss	
linate	Latit. 100E	9.35	9.65	9.16	9.14	6.51	6.43	6.38	7.02	7 13	26.4	1.32	1.61	11.75	11.75	11.75	11.75	
GPS Cordinate	Longit. 46N	8.25	8.04	7.63	7.61	8.59	8.51	8.44	7.93	7 93		1.65	1.85	8.03	8,03	8.03	8.03	
	Sample # [	960921073	960921074	960921075	960921076	960923001	960923004	960923006	960923018	960923019	00000000	200323020	960923024	960924012	960924013	960924014	960924015	
	NO.	40	41	42	43	44	45	97	17	48		27	20	5	52	53	54	

\*◎: abundance ○: common △: minor •: Fare +: strong ±: medium -: weak @PS:global positioning system tongit.:longitude Latit.:latitude Abbreviation of rock and mineral name tongit..longitude Latit.:latitude Abbreviation of rock and mineral name tondt:hornblendite Ambbit:amphibolite Pegnat:pegnatite Psammitic Pelit:pelitic Bi:biotite Hs:muscovite 02:quartz Gt:garnet Au:augite Hy:hypersthene P1:plagioclase Kf:potassium feldspar Ho:hornblende 01:olivine Cc:calcite Serisericite Ch1:chlorite Ep:epidote Sph:sphene Abt:apatite Opq:opaque mineral metam.:metamorphism

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An Ÿ Pl:plagioclase To:tournaline 8 AUG ⊽ НΥ ~ 21 呈 é -0 5 4 5 20 В QZ:Quartz 2 5 2 ŝ Ċ V 5 ÷ 20 AD 2  $\nabla$ 20 3 20 ຂ I 2013 22 20 ç 2 말 Ca:calcite An:Anorthite 2 2 G ¢ Σ 2 l assemblages Ca 07 P <1 41 Chicklorite Serisericite SmiSmectite Kikaolin-Hineral Hb:hornblende Hy:hypersthene Aug;augite Do:Dolomite 82  $\simeq$ 2 ھ ം ò 8 8 13 100 8 5 3 3 ŝ ŝ 52 <u>г</u> 3 V ▽ V  $\nabla$ 5 5 5 Hineral ωN ~  $\nabla$ n,  $\overline{\nabla}$ ⊽  $\nabla$ ⊽ ν 5 Ser ~  $\nabla \nabla$ ∿¦∽  $\nabla$ -ഗിപ n 5⊽ ;=  $\nabla$ ∣⊽ ∾\⊽  $\nabla$  $\overline{\nabla}$  $\overline{\nabla}$ 7 nane 02-vein 81-Gr-Dr 81-02-Hornblendite Gt-Amphibolite Bi-Adamellite Au-Hy-Porphyrite HS-Adamellite Psammitic Gneiss Psammitic Gneiss Psammitic Gneiss Au-Hy-Porphyry Hornblendite Bi-Adamelite Bi-Adamelite Bi-Adamelite Bi-Adamelite Bi-Adamelite Hs-Gneis Au-Hv-Porphyrite Bi-Adamellite Bi Alaamellite Bi Alaamellite Basalf-dyle Basalf-dyle Amphibolite Amphibolite BI-HS-Adamellite Pegnatite Bi-MS-Gr-Ur alt-vein alt-vein Au-Hy-Porphyrite Bi-Porphyriteor) <u> Fr-Porphyry</u> Psammitic Gneiss Porphyrite dyke Bi-HS-Gr-Dr Bi-HS-Gr-Dr Gt:garnet Bi:biotite 31-Gneiss alt-vein 02-vein Marble 100E Gr-Dr.granodiorite Hs.muscovite Ab:Albite Kf:potassium feidsper Cordinate 2 10.88 <u>81</u> 22 23 1 96 2.60 2.84 2.65 5 2. 56 20 6 <u>လ</u>လ <u>ന</u> 5 т, <u>[6PS\_Cu.</u> 100011.46N\_1 9 68 54 83 88 <u>888</u> 1005 1006 0.34918043 004 თ 918041 8 9 e Samp] 6000 50 je g 96 ŝ 96 96 96 ğ 96 ğ NO. <u>ر</u>، 2 þ 5 ଜୁକୁ 갑압

Table I-1-4 Result of X-ray diffraction (semidetailed-detailed area) (1)

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		i-Ms-Gr-D	Bí-MS-Gr-D	8i-Ms-Gr-D	BI-HS-Gr-Dr	8i-Hs-Gr-Dr	81-HS-Gr-D1	Bi-Hs-Gr-Dr	Bi-HS-Gr-Dr	Bi-Hs-Gr-Dr	8 i - HS-Gr-Dr	Bi-GZ-Dr	sammiti	Ms-Gnetss	81-MS-Gr-Dr	Bi-Hs-Gr-Dr	-HS-GI	-HS-61	8i-Ms-Gr-Dr	8i-Hs-Gr-Dr	-MS-6	8i-Ms-Gr-D	Bi-HS-Gr-D	Bi-MS-Gr-Dr	-HS-GI	<u>Bi-Ms-Gr-Dr</u>	81-HS-Gr-DF	81-MS-Gr-Dr	8 i - Ms-Gr-D	8 i - Hs-Gr-Dr	81-Hs-Gr-Dr
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nate	Lat i t. 100E	11.43	11.47	11.49	11.51	11.55	1.59	11.63	1.67	11.71	11.75	11.75	11.75	1.75	8.87	8.91	8.95	8.99	9.03	9.07	9.11	9.15	9.19	9.22	9.26	9.29	9, 33	9.37	9.39	9.41	9.45
GPS Cordinate										1		ļ	ļ																		
GPS	Longit 46N	7.99	7.99	7.99	8.00	8.00	8.01	8.01	8.02	8.02	8.03	8.03	8.03	8.03	8.13	8.14	8.14	8.15	8.15	8.16	8.16	8.17	8.17	8.18	8.18	8.19	8.19	8.20	8.20		8.21
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Table I-1-4 Result of X-ray diffraction (semidetailed-detailed area) (2)

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Oz:Quartz Pi:plagioclase To:tourmaline Ch:chlorite Ser:sericite Sm:Smectite K:kaolin-Mineral Ca:calcite Hb:hornblende Hy:hypersthene Aug:augite Do:Dolomite An:Anorthite Gt:garnet Bi:biotite Gr-Dr:granodiorite Ms:muscovite Ab:Albite Kf:potassium feldsper

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strongly altered as mentioned earlier and produces chlorite and sericite.

Amphibolite produces chlorite and quartz porphyry produces sericite and lamprophyre produces chlorite respectively.

From the above, it is assumed that this region has been suffered widely by the hydrothermal alteration after the forming of quartz veins and the intrusion of dyke rocks.

The possibility of the alteration and the degree of alteration due to the forming of quartz vein is discussed in the next chapter from clay mineral survey by POSAM of the detailed survey areas I, II and the results of X-ray diffraction of specimens taken from sections across quartz veins.

## 3. Chemical analysis of rocks

A thorough chemical analysis was conducted on every kind of rocks distributed in this area. The results are shown in Table II-1-5. The result of calculation of seven normative minerals of the following hypabyssal-abyssal seven rocks is shown in Table II-1-6 and Harker diagram in Fig. II-1-6, alkali/SiO<sub>2</sub> diagram in Fig. II-1-7 and ACF diagram in Fig. II-1-8 respectively.

No.1	Two-mica granodiorite	(abbreviation in diagram: TW-mica Gr-Dr)
No.2	Diorite	(abbreviation in diagram: Diorite)
N0.3	Quartz porphyry	(abbreviation in diagram: Qz Porphyry)
No.4	Bi Adamellite	(abbreviation in diagram: Bi Adamellite)
No.5	Granite porphyry	(abbreviation in diagram: Gr-Dr dyke)
No.10	Pegmatite	(abbreviation in diagram: Pegmatite)
No.11	Tow-mica granodiorite	(abbreviation in diagram: Tw-mica Gr-Dr)

# Table II-1-6 C I PW Classification (Norm. calculation)

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No		- Q -	С	- OF.	ab .	an	ac	ns	mt	il
1	Tw-mica Gr-Dr	40.53	0	19.34	33.80	0	1	0.06	0	0.232
2	Diorite	8.495	0	18.21	37.50	13.64	0	0	5.657	2.793
3	Qz Porphyry	50.91	0	23.29	21.21	0	1.26	1.77	0	0.116
4	Bi Adamellite	27.96	0.461	17.36	39.39	7.961	0	0	0.512	0.767
5	Gr-Dr dyke	40.62	0.078	19.47	31.25	4.554	0	0	1.223	0.488
10	Pegmatite	34.92	0	33.85	20.23	9.133	0	0	0.044	0.097
11	Tw-mica Gr-Dr	51.34	0	22.90	17.97	0 :	0.64	4.41	0	0.115

		:				and the state of the	<u> </u>
No.	ap	wo-di	en-di	fs-di	cn-hy	fs-hy	total
1	0.141	2.374	0.506	2.025	0	0	100.008
2	1.762	3.088	2.438	0.302	5.436	0.674	99.995
3	0.118	0.282	0.079	0.216	0.199	0.544	99.914
4	0.374	0	0	0	1.861	3.357	100.003
5	0.357	0	0	0	1.102	0.757	99.899
10	0.141	1.197	0.431	0.790	0	0	100.833
11	0.211	0.115	0.328	0.906	0	0	98.935

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Table1-1-5 Result of chemical analysis of whole rock (Semidetailed area )

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		GPS C	GPS Cordinate	Rock name			Ch	Chemical	compositions	itions							·
No.	o. Sample#	46N	100E		Si0 <sub>2</sub>	A12(%)03	Na <sub>2</sub> 0 (%)	K <sub>2</sub> 0 (%)	CaO MgO	SHO NHO	Ti0_2 (%) <sup>2</sup>	Fez 0 3	Eeo Eeo	P <sub>2</sub> 0 ₅ (%) ₅	H 200	Ig. loss	·
	960827013	8.11	9.42	Tw mica Gr-Dr	78. 25	9.95	4.09	3. 22	1.43 0.	20 0.44	0.12	0.34	0. 75	0.06	0.12	0.85	
5	960902034	6.93	12.30	Diorite	58.69	15.62	4.43	3.08	5.24 3.	16 0.10	1.47	3.90	3.49	0.76	0.18	0.10	
က	3 960918038	7.27	11.04	Qz Pophyry	81.48	8.26	3.52	3.88	0.20 0.	11 0.02	0.06	0.43	0.44	0.05	0.17	1.14	· . · · · · ·
4	960921011	8.99	12.60	Bi Adamellite	71.67	14.08	4.61	2.91	1.80 0.	74 0.05	0.40	0.35	2.27	0.16	0.17	1.58	-
ഹ	960921067	8.12	12.56	Gr-Dr dyke	75. 59	11.09	3. 59	3. 22	1. 09 0. /	43 0. 03	0.25	0.82	0.96	0.15	0.18	1.48	
G	960923001	8. 59	6.51	Lamprophyre	45.62	14.13	3.73	1.77	8. 38 9. ]	11 0.17	2.22	3.57	8.28	0.77	0.38	1.50	
-	960923004	8.51	6.43	Amphibolite	50.16	12.79	2.16	0. 29	9.20 5.8	87 0.21	2.90	2.54	13. 26	0.31	0.12	0.12	
~	960923006	8.44	6.38	Pelit gneiss	65.40	17.04	1.06	3. 77	0.36 2.1	17 0.04	1.26	1.01	5.17	0.05	0.22	3.12	
6	960923018	7. 93	7.02	Marble	·1.60	2.46	0.44	0.10 5	52. 53 0. 4	40 0.03	0.01	0.04	0.12	0.03	0.09	42.49	
9	960923019	7.93	7.13	Pegmatite	73.98	13. 25	2.35	5. 63	2. 61 0. 1	17 0.01	0.05	0.03	0.47	0.06	0.35	0.90	
Ħ	960924012	8.03	11. 75	Tw mica Gr-Dr 81.	81.41	7.60	4.40	3. 83 (	0.78 0.1	13 0.04	0.06	0.22	0.50	0.09	0.09	0.40	
12	960924014	8. 03	11.75	Psamt Gneiss	70.15	11.51	2.35	1.84	5.41 2.2	23 0.07	0.81	1.21	2.30	0.16	0. 24	1.95	
13	960924015	8.03	11.75	Ms Gneiss	79.08	10.59	2.59	4.74	1. 23 0. 1	18 0.01	0.08	0.20	0.50 (	0.02	0. 07	0.69	· .
*	* Tw: Two Gr-	-Dr:Gra	no-Diori	Gr-Dr:Grano-Diorite Qz:Quartz H	Bi:Biotite		Pelt:Pelitic	Psant	Psamt:Psammitic	1	Ms:Muscovite	e					

IL: L'SAUDITIC MS: MUSCOVITE

-53-

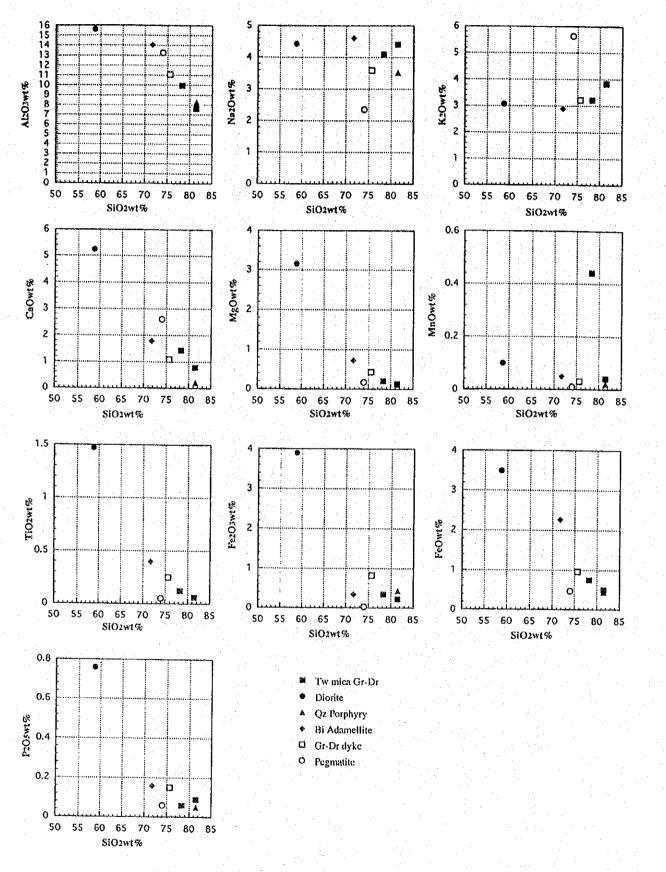
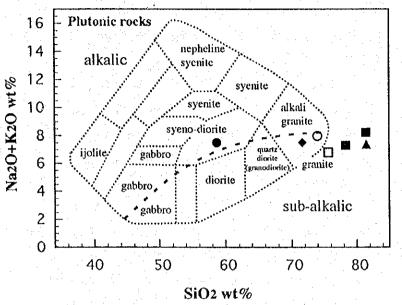


Fig. II -1-6 Harker variation diagram

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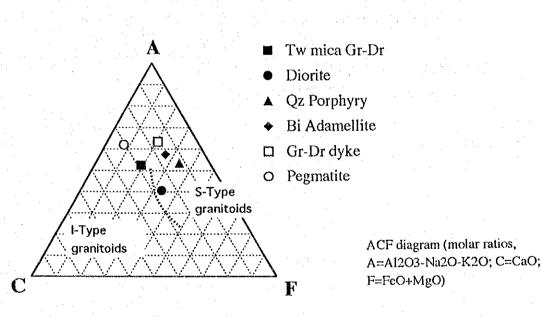
-54-



- Tw mica Gr-Dr
- Diorite
- ▲ Qz Porphyry
- ♦ Bi Adamellite
- □ Gr-Dr dyke
- O Pegmatite

Nomenclature of normal (i.e. non-potassic igneous rocks(after Cox et al. 1979). The divided line between alcalic and subalcalic magma series is from Miyashiro(1978).

# Fig. II -1-7 Alkali (Na<sub>2</sub>O+ K<sub>2</sub>O)/SiO<sub>2</sub> diagram





The results of analysis indicate that the specimens excepted diorite are granitoids containing relatively rich amount of SiO<sub>2</sub> exceeding 80wt%.

The results of calculations of normative minerals indicate the probability that they belong to ilmenite-granitoids since they have a relatively high level of normative magnetite with a exception of diorite.

Especially those with high SiO<sub>2</sub> (around 80wt%) as shown in Harker variation diagram (Fig.II-1-6) that has low content of  $Al_2O_3$ , which may suggest silicification alteration but a microscopic observation of a thin section does not support the suggestion as discussed earlier.

The figure of alkali/SiO<sub>2</sub> shows that these rocks are the diorite-granitoids of the subalkalics considered from the relationship between alkali (Na<sub>2</sub>O + K<sub>2</sub>O) and SiO<sub>2</sub>.

The ACF diagram shows the plot of these rocks. Plots are seen in four points in the generated region of granitoids originated from sedimentary rocks (S-type) (diorite, biotite adamellite,quartz porphyry, granodiorite), one point in the granitoids originated from igneous rocks region(I type) (pegmatite) and one point in the boundary composition region (two-mica granodiorite). (Note that No.11 Tw-mica Gr-Dr with low level of  $Al_2O_3$  is not plotted here.)

#### 4. Chemical analysis of ore minerals

This analysis are described globally in the chapter 2 as mentioned earlier.

#### 5. Polished sections of ore minerals

This will be described globally in the chapter 2.

#### 6. Radioisotopic dating

The classifications and quantities used in this measurement are as follows:

1		
Classification	Rocks	Quantity
K-Ar	Whole rocks	2
Rb-Sr	Whole rocks	2 sets
K-Ar	Altered minerals	2
Rb-Sr	Altered minerals	2 sets
	K-Ar Rb-Sr K-Ar	K-ArWhole rocksRb-SrWhole rocksK-ArAltered minerals

For Rb-Sr, in the case of rock body, three rock specimens were taken from a location several hundred meters apart and used as a set for the analysis of whole rock. For altered minerals, specimens taken from 4 different locations are used as a set. These specimens especially the later were assumed to have attained isotopic equilibrium.

For K-Ar, in the case of rock body, the most fresh rock specimens were taken in the area for the analysis of whole rock and for altered minerals, the same specimens as part of Rb-Sr altered minerals were used.

Among them the sericite was detected through X-ray diffraction conducted earlier from the specimens of altered mineral for the altered clay of the quartz veins Nos. 1, 5, 7 and clay vein of drilling MJMT-2.

Table II-1-7 shows the list of the specimens used in this measurement.

No.	Method	Area	GPS coord	inate	Rock name	Remark
		and the Area	46N	100E	and a state of the state	
RD1-1	Rb-Sr	Semi-det.	06.56	12.16	Diorite	whole rock
RD1-2	Rb-Sr	Semi-det.	06.63	12.31	Diorite	whole rock
RD1-3	Rb-Sr	Semi-det.	06.73	12.56	Diorite	whole rock
RD2-1	Rb-Sr	Semi-det.	08.86	11.26	Lamprophyre	whole rock
RD2-2	Rb-Sr	Semi-det.	08.92	11.36	Lamprophyre	whole rock
RD2-3	Rb-Sr	Semi-det.	09.02	11.60	Lamprophyre	whole rock
RD3-1	Rb-Sr	Detailed	08.25	09.04	Ser+Ab clay	alter. By No.1 Q v
RD3-2	Rb-Sr	Detailed	08.45	09.51	Ser+Ab clay	alter. By No.5 Q v
RD3-3	Rb-Sr	Detailed	08.14	09.42	Ser+Ab clay	alter. By No.1 Q v
RD3-4	Rb-Sr	Detailed	07.43	10.88	Ser+Ab clay	alter. By No.7 Q v
RD4-1	Rb-Sr	Detailed	MJMT-2	192.40m	Ser+Ab clay	argil.part of Gr-Dr
RD4-2	Rb-Sr	Detailed	MJMT-2	196.60m	Ser+Ab clay	clay vein
RD4-3	Rb-Sr	Detailed	MJMT-2	232.45m	Ser+Ab clay	clay vein
KD1	K-Ar	Semi-det.	08.10	09.42	Two-M- Gr-Dr	whole rock
KD2	K-Ar	Semi-det.	08.99	12.60	Bi-Adamellite	whole rock
KD3	K-Ar	Detailed	08.45	09.51	Ser+Ab clay	same sample with RD3-2
KD4	K-Ar	Detailed	07.43	10.88	Ser+Ab clay	same sample with RD3-4

Table II-1-7 Sample list of radioisotopic age determination

Table II-1-8 shows the result of isotopic composition of Rb/Sr.

	· · · ·	1 able 11 - 1-8	isotopic cor	nposition of Ki	D/Sr and a second
	No.	Rb(ppm)	Sr(ppm)	87Rb/86Sr	87Sr/86Sr(2 σ)
÷	RD1-1	101	972	0.3030	0.70640(1)
	RD1-2	102	971	0.3025	0.70638(1)
	RD1-3	- 107	942	0.3281	0.70648(1)
	RD2-1	126	897	0.4075	0.70755(1)
	RD2-2	133	904	0.4268	0.70763(1)
•	RD2-3	116	820	0.4089	0.70752(1)
:	RD3-1	28.7	38.6	2.1490	0.71709(1)
	RD3-2	21.5	59.7	1.0450	0.71426(1)
	RD3-3	36.5	19.2	5.5260	0.72583(1)
	RD3-4	58.5	38.7	4.3830	0.72421(1)
	RD4-1	75.6	219	1.0010	0.71400(1)
	RD4-2	113	133	1.9720	0.71695(1)
	RD4-3	98.5	208	1.3710	0.71596(1)

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Fig II-1-9 shows the isochron diagram of RD1 series (diorite) and Fig. II-1-10 shows RD2 series (lamprophyre), Fig II-1-11 shows RD3series (altered minerals) and Fig. II-1-12 shows RD 4 series (core altered minerals) respectively.

The test results for 4 specimens of K-Ar are shown in Table II-1-9.

Among the results of Rb/Sr isotopic, the diorite and lamprophyre do not show good isochron because of the large margin of error due to the narrow range of Rb-Sr of the specimens. It is necessary to review the selection of specimens.

As for the altered clay in the field and the cores, there were a wide dispersion in contents of Rb, Sr and those isotopic ratios of individual specimen. It is therefore doubtful whether the isotopic equilibrium among the specimens has been attained.

On the other hand, the results of four K-Ar specimens are reliable because of the low content of air.

As described above, the result of Rb-Sr leaves some problems. The measured periods are as follows:

1) Diorite (Rb-Sr)	250.4 $\pm$ 85.7 Ma (Permian period)
2) Lamprophyre (Rb-Sr)	351.8 ± 236.7 Ma
	(Devonian to carboniferous period)
3) Quartz vein altered clay (Rb-Sr)	193.0 ± 27.3 Ma
	(Jurassic period)
4) Core clay vein (Rb-Sr)	208.3 $\pm$ 136.5 Ma (Permian period)
5) Two-mica granodiorite (K-Ar)	268.1 $\pm$ 5.6 Ma (Permian period)
6) Biotite adamellite (K-Ar)	384.7 $\pm$ 7.8 Ma (Devonian period)
7) Altered clay of quartz vein No.5 (K-Ar)	234.4 ± 4.9 Ma
	(Permian period)
8) Altered clay of quartz vein No. 7 (K-Ar)	$225.8 \pm 4.8$
	(Permian period)

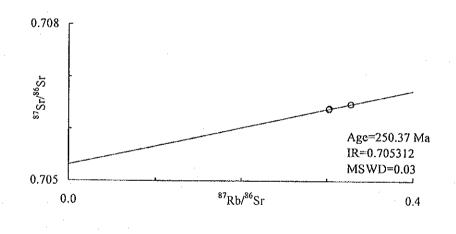
According to the documents of former East Germany, K-Ar ages of two-mica granodiorite and diorite are 5.2-5.5 Ma and 3.1 Ma respectively. In the case of holocrystalline rocks, the regular method is to separate composite minerals into the same sizes for an isotopic analysis. Measurement values would differ whether or not minerals are separated or not. We do not know which method the former East Germany used.

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Sample No.	<sup>87</sup> Rb/ <sup>86</sup> Sr	ERR%	<sup>87</sup> Sr/ <sup>86</sup> Sr	ERR%
RD1-1	0.303000	0.500	0.706400	0.010
RD1-2	0.302500	0.500	0.706380	0.010
RD1-3	0.328100	0.500	0.706480	0.010

()

CALCULATED RESULT! Sample Numbers = 3 Best Slope = 3.56158984217964D-03MSWD = 0.03Age :  $250.37\pm85.65$  Ma Initial Ratio :  $0.705312\pm0.000380$ 





-59-

Sample No.	<sup>87</sup> Rb/ <sup>86</sup> Sr	ERR%	<sup>87</sup> Sr/ <sup>86</sup> Sr	ERR%
RD2-1	0.407500	0.500	0.707550	0.010
RD2-2	0.426800	0.500	0.707630	0.010
RD2-3	0.408900	0.500	0.707520	0.010

# CALCULATED RESULT!

Sample Numbers = 3

Best Slope = 5.00734014438708D-03MSWD = 0.14Age :  $351.75 \pm 236.69$  Ma Initial Ratio :  $0.705492 \pm 0.001400$ 

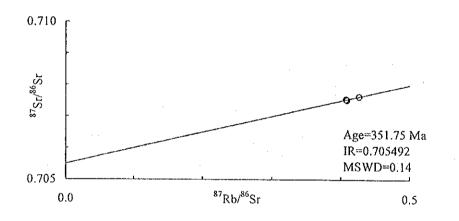


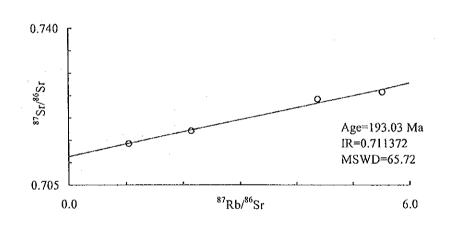
Fig. II-1-10 Isotopic age determination by Rb-Sr (Lamprophyre)

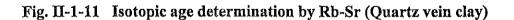
-60-

Sample No.	<sup>87</sup> Rb/ <sup>86</sup> Sr	ERR%	<sup>87</sup> Sr/ <sup>86</sup> Sr	ERR%
RD3-1	2.149000	0.500	0.717090	0.010
RD3-2	1.045000	0.500	0.714260	0.010
RD3-3	5.526000	0.500	0.725830	0.010
RD3-4	4.383000	0.500	0.724210	0.010

# CALCULATED RESULT!

Sample Numbers = 4 Best Slope = 2.74484375741825D-03MSWD = 65.72Age :  $193.033 \pm 27.32$ Ma Initial Ratio :  $0.711372 \pm 0.001280$ 

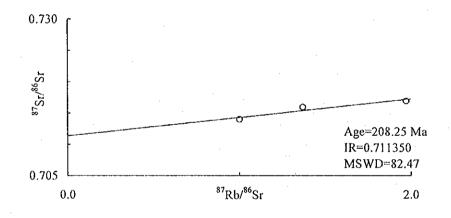




Sample No.	<sup>87</sup> Rb/ <sup>86</sup> Sr	ERR%	<sup>87</sup> Sr/ <sup>86</sup> Sr	ERR%
RD4-1	1.001000	0.500	0.714000	0.010
RD4-2	1.972000	0.500	0.716950	0.010
RD4-3	1.371000	0.500	0.715960	0.010

#### CALCULATED RESULT!

Sample Numbers = 3 Best Slope = 2.961532429298864D-03MSWD = 82.47Age :  $208.47 \pm 136.51$  Ma Initial Ratio :  $0.711350 \pm 0.002880$ 





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Table II-1-9 Result of isotopic age determination by K-Ar

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Sample	K	Weight	<sup>[36</sup> Ar]	<sup>40</sup> Ar/ <sup>36</sup> Ar	f <sup>40</sup> Arlrad	K-Ar age	Air-fract.
	(wt%)	(g)	$(10^{-10} \text{cm}^3 \text{STP/g})$		$(10^{-8} \text{cm}^3 \text{STP/g})$	(Ma)	(%)
	2.69	0.0647	$19.06 \pm .51$	$15790. \pm 366$	$3080.7 \pm 31.0$	$273.3 \pm 5.7$	1.8
KD-1	(2.0%)	0.0486	18.93 ± .63	15766. ± 462	$3003.9 \pm 30.3$	$267.0 \pm 5.6$	1.8
		0.0521	$25.66 \pm .61$	$11663. \pm 235$	$2967.2 \pm 30.1$	$263.9 \pm 5.5$	2.5
	2.47	0.0560	$30.87 \pm .67$	$13094. \pm 237$	$4031.3 \pm 40.6$	$378.6 \pm 7.6$	2.2
KD-2	(2.0%)	0.0575	27.48 ± .60	$14951. \pm 279$	$4110.0 \pm 41.4$	$385.2 \pm 7.8$	1.9
		0.0580	24.56 ± .59	$16801. \pm 352$	$4167.5 \pm 42.0$	$390.1 \pm 7.9$	1.7
	0.560	0.0685	67.68 ± .81	$1104.9 \pm 5.1$	$548.19 \pm 5.57$	$235.9 \pm 5.0$	26.8
KD-3	(2.0%)	0.0497	64.93 ± .89	$1129.7 \pm 7.4$	$541.92 \pm 5.52$	233.4 土 4.9	26.2
		0.0531	64.19 土 .85	$1140.5 \pm 7.3$	$543.13 \pm 5.53$	233.9 ± 4.9	25.9
KD-4	1.27	0.0542	$92.57 \pm 1.09$	$1570.9 \pm 7.4$	$1180.5 \pm 12.0$	225.6 ± 4.8	18.8
	(2.0%)	0.0607	$91.66 \pm 1.05$	1576.4 土 6.7	$1174.6 \pm 12.0$	224.5 ± 4.7	18.8

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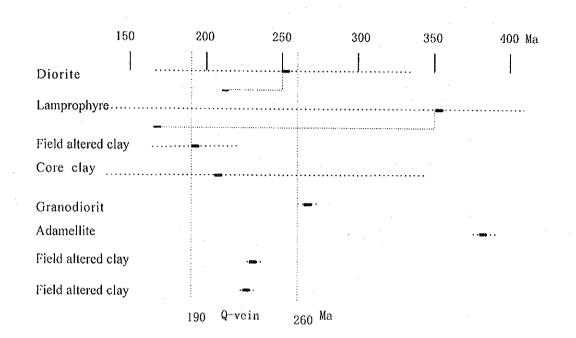
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# 7. Estimation of the age of quartz vein formation

The result of the radioisotopic dating described in the preceding clause is outlined as follows:



From the above figure, the age of the quartz vein formation based on Rb-Sr, is assumed to be 190 to 260 Ma and it is between 220 and 260 Ma (Permian period) if limited to K-Ar.

#### Chapter 2 Geological detailed survey

The areas for the geological detailed survey consist of the following two areas:

	Survey areas
Area I	7.75 km <sup>2</sup>
Area II	2.25 km <sup>2</sup>
Total	10.00 km <sup>2</sup>

#### 2-1 Survey method

Survey was performed by using a map obtained by enlarging a 1/10000 topographic map to a 1/2000 map. Particularly, the quartz vein survey method is described later.

#### 2-2 Geology

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The geological map of the 2 areas of the detailed survey is shown in Fig. II-2-1 (1) and (2).

The geology of the two detailed survey areas are composed of metamorphic rocks, twomica granodiorite, biotite adamellite, diorite, felsitic rock, quartz porphyry and lamprophyre.

As to the generation facies, lithofacies and microscopic observations of these rocks are already discussed in the chapter 1, this chapter only complements what had previously been explained.

#### 1. Metamorphic rocks

Metamorphic rocks distribute mainly in the eastern peripheral of the Area I and also distribute small amount in the south-west part of the Area I and II. They are composed of the psammitic gneiss and, as described earlier, are penetrated by small scale irregular dike-sheet of biotite adamcllite and pegmatite. The structure of psammitic gneiss is disturbed by the intrusion of two-mica granodiorite. In the north region of the Area 1, psammitic gneiss with a long diameter of 100m are caught in two-mica granodiorite.

#### 2. Biotite adamellite

Biotite adamellite distribute in the north-east end of the area II. Relation with two-mica granodiorite do not clear but their color is whitish and contains little amount of muscovite. The geological boundary of these rocks is assumed to run in the north extension trend of the boundary of two-mica granodiorite and psammitic gneiss.

A microscopic observation shows that the biotite adamellite is granitic with larger amount of potassium feldspar but there are little difference between the both rocks by whole rock analysis.

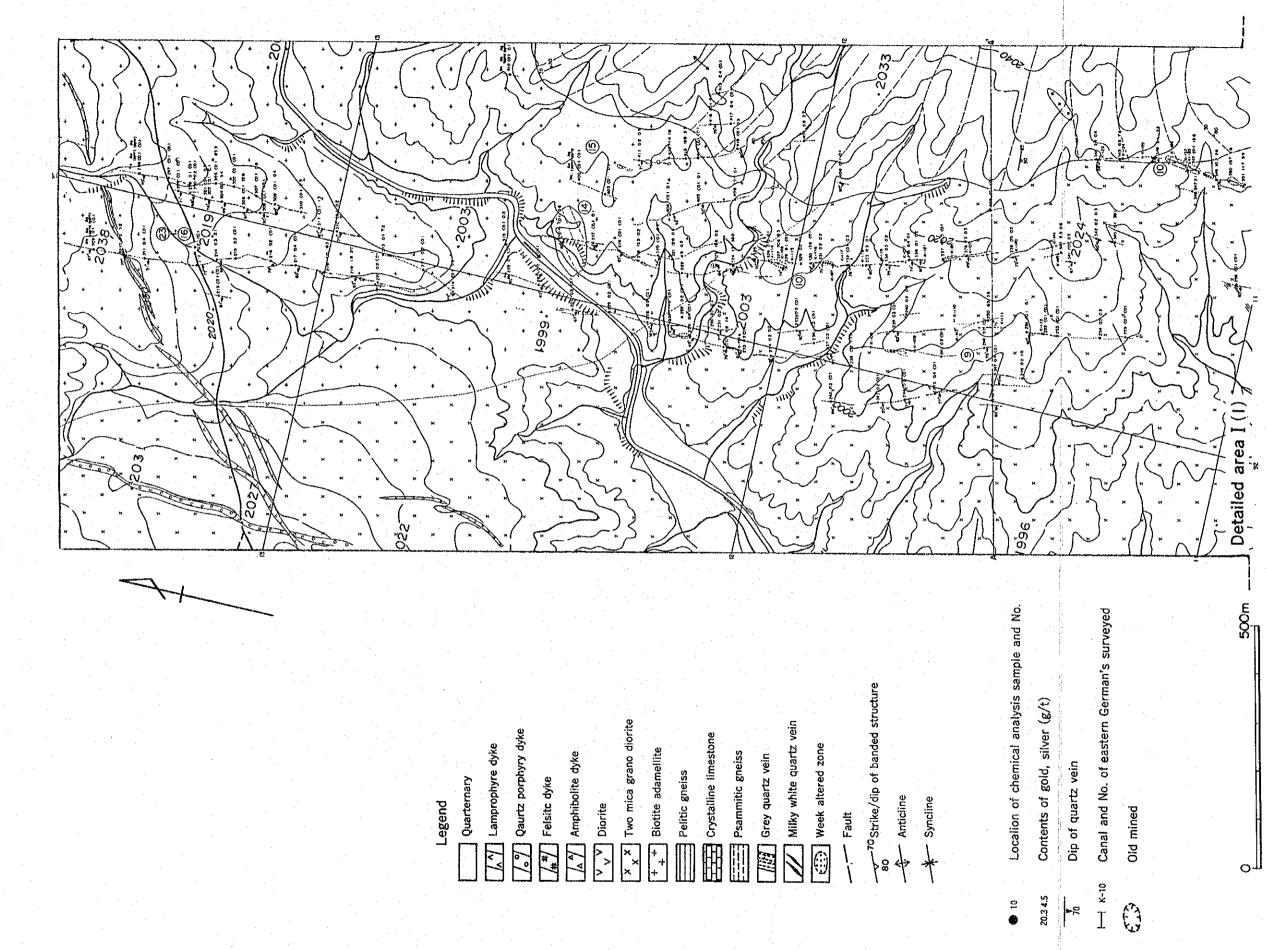
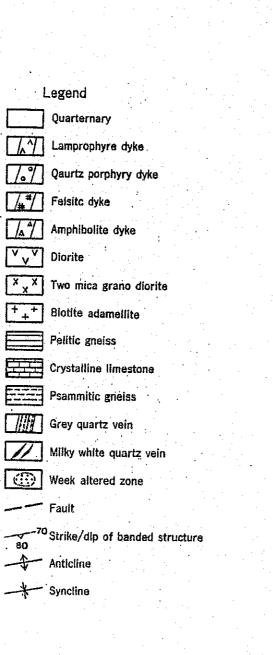


Fig. I-2-1 Map of geological and quartz veins surveyed (Area I ) (I)

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Locallon of chemical analysis sample and No. Contents of gold, silver (g/t)

Dip of quartz vein

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20.3 4.5

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Canal and No. of eastern German's surveyed

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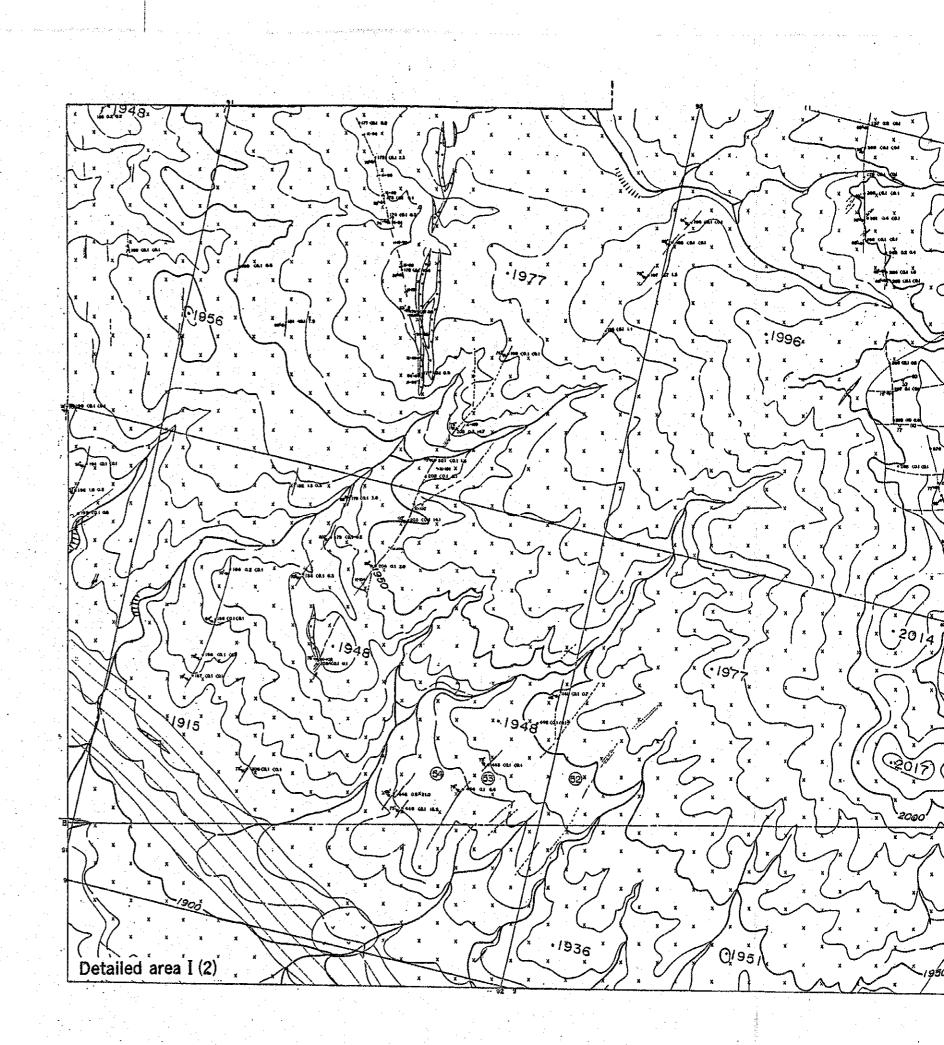
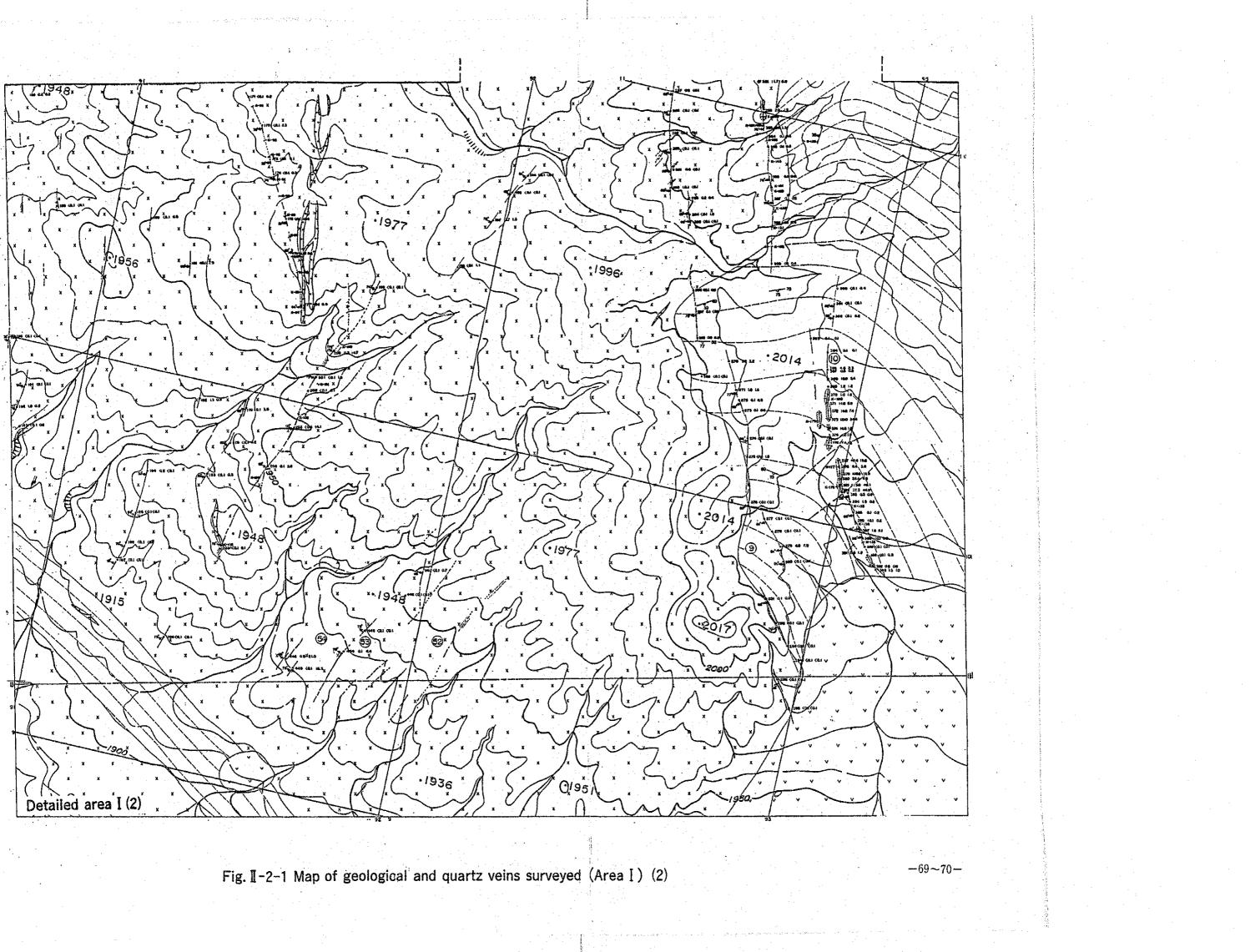
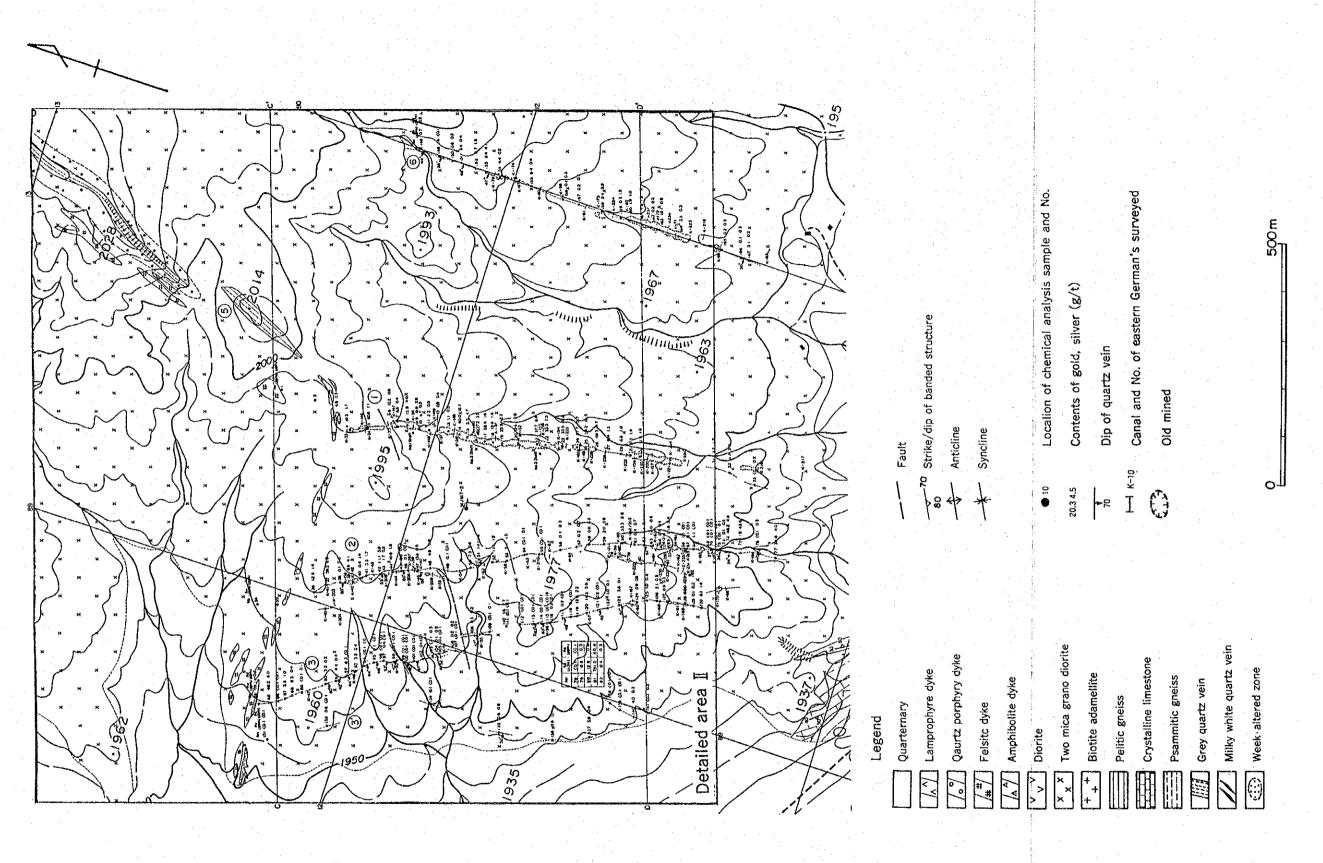


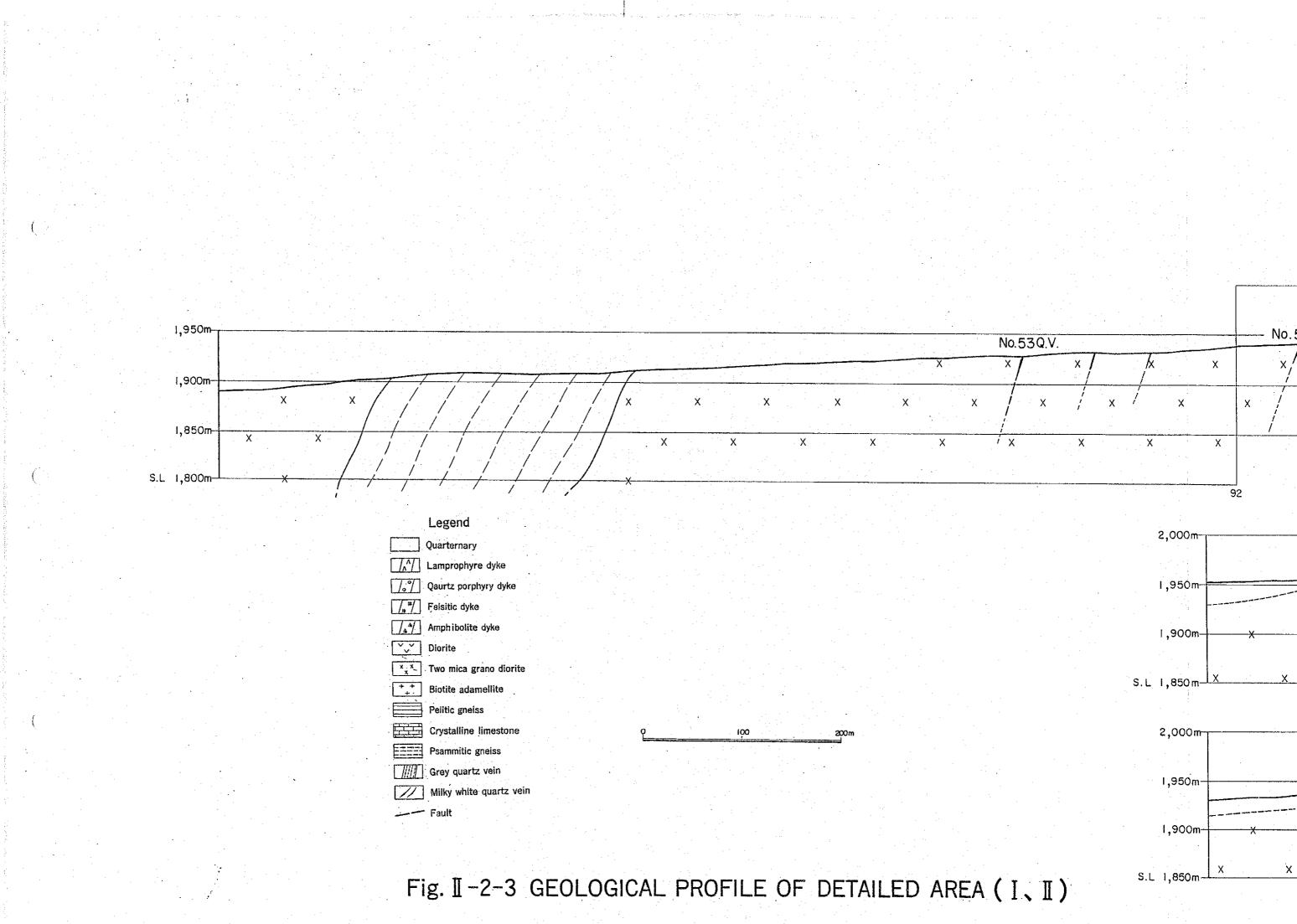
Fig. II-2-1 Map of geological and quartz veins surveyed (Area I) (2)



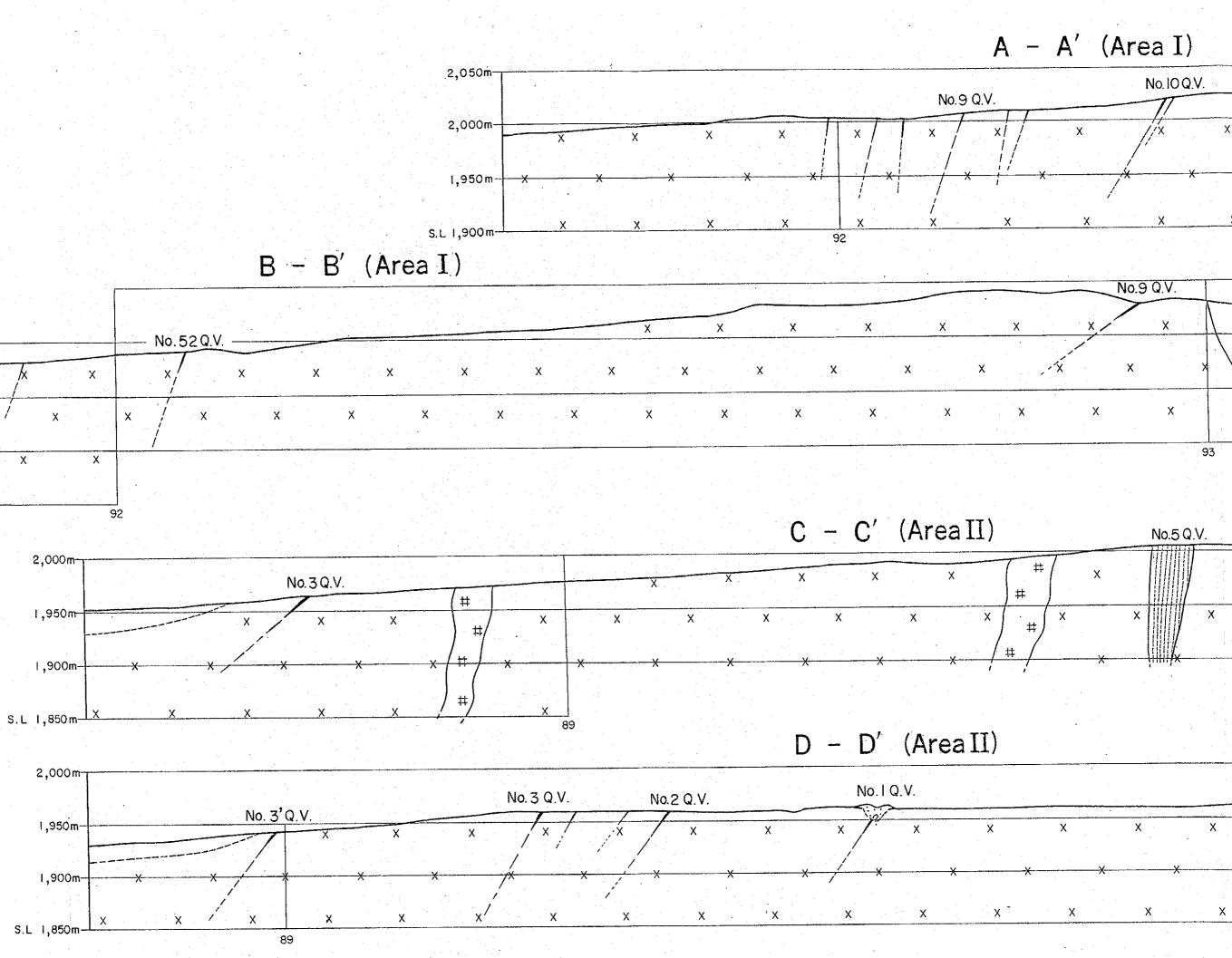


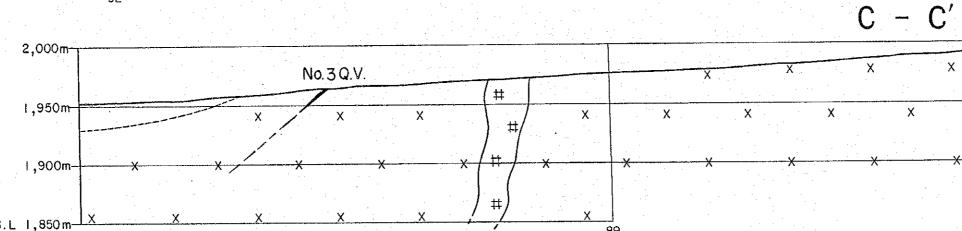
Map of geological and quartz veins surveyed (Area  ${\rm I\!I}$  ) Fig. I -2-2

 $-71 \sim 72 -$ 



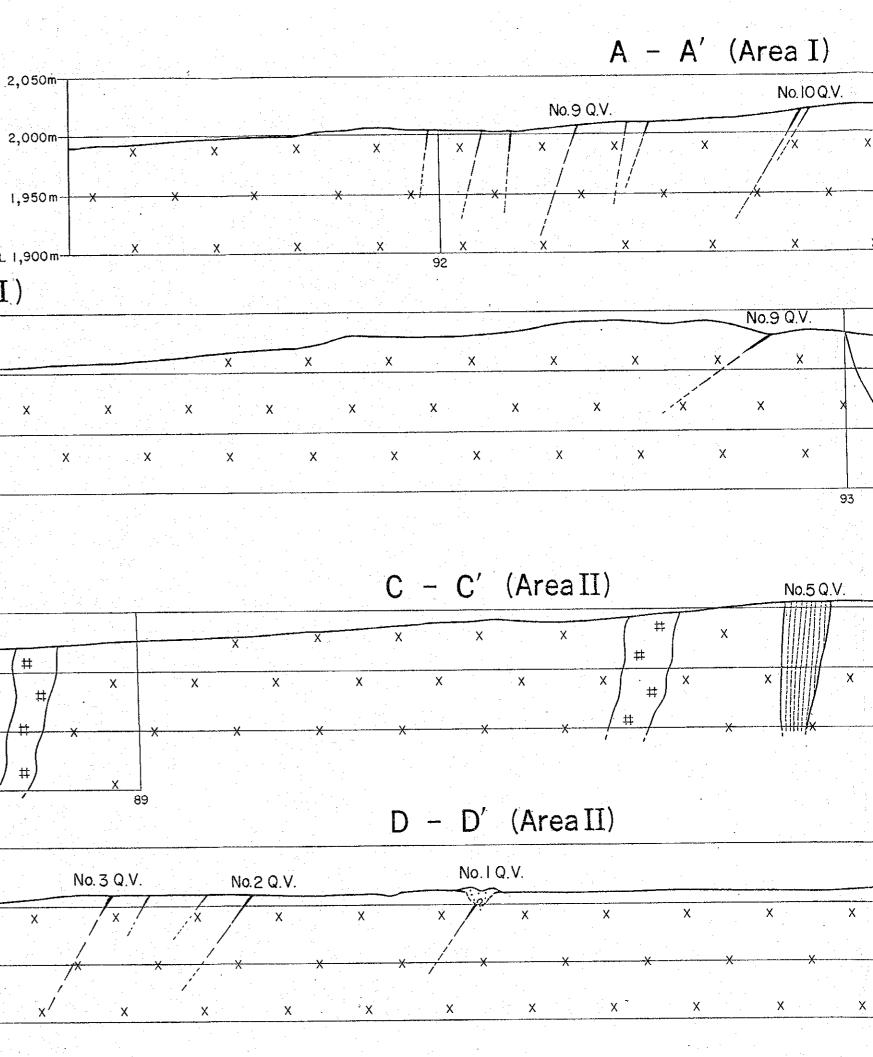
# AREA(I,I)

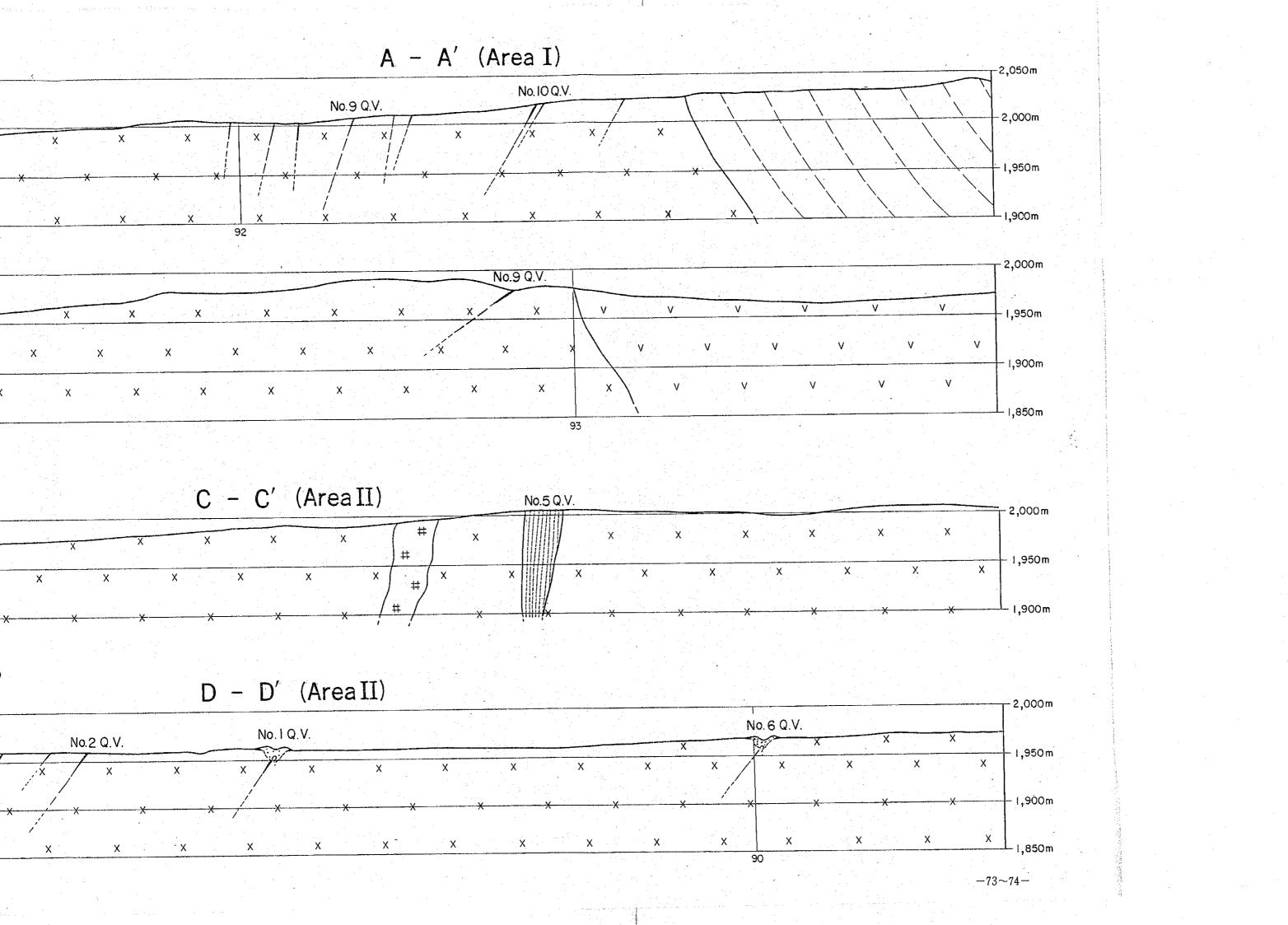




B - B (Area 1)	
No.52Q.V	<u> </u>
No.53Q.V.	
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	X	X	* / *	<del>X</del>	X	X		<del>- X</del>	X	<del>, ,</del>	<del> </del>		-X	X	—X
x	X	X	X	X	X	X	X		X	X	X	X	X	<b>X</b>	X

They belong to sub-alkalic granitoids as mentioned earlier.

If biotite adamellite that penetrated in psammitic gneiss is identical to this rock, then the rock is older than two-mica granodiorite as described earlier.

#### 3. Two-mica granodiorite

Two-mica granodiorite distributes in the center through the northern part of the Area I and in all part of the Area II. At the point where they contact with metamorphic rocks outer than the southwest area II, there is so called the contact quenching facies where coarse granule of holocrystalline changes transitionally to medium granule semi-holocrystalline or fine granule porphyritic texture.

The description of rocks is omitted.

#### 4. Diorite

Diorite distributes mainly along the south-east end of the area I forming the largest rocks among the present survey areas and there are also stock 100 meters in diameter in the south-western part of this area.

The description of rocks is omitted.

#### 5. Felsitic rocks

Felsitic rocks distributes in the northern part of the area II with many small rocks in undefined or lenticular shapes at a distance of several to several ten meters to the north side of No. 5 quartz vein in parallel to the vein intermittently. The sizes of independent rocks are several meters to several ten meters in width and 10 to 100 meters in length.

The tone of color is dark gray to gray as compared to the white gray of No.5 quartz vein. The composition is entirely non phenocryst and heterogeneous compared to quartz veins. They may be hard to distinguish each other. We will have to make thin section out of these rocks.

If any case those felsitic rock and No.5 quartz vein are assumed to be formed in the same period.

#### 6. Dykes

Quartz porphyry is distributed in the area I and in its peripherals. One to three dikes in S-N system run intermittently in echelons. There are two ladder shaped dikes in the southern part of the area.

Lamprophyre dike developed in NE-SW system trend in the northernpart of the area I and one dike outside the area branches out into several dikes in this area forming an echelon arrangement. As mentioned earlier this dike penetrates No.23 quartz vein (No.9 quartz vein system).

#### 2-3 Survey of quartz vein

#### 1. Survey and sampling method

The survey of quartz vein was carried out for both semi-detailed and detailed survey areas for all known quartz veins. The width (maximum, average), directions, inclinations and the extension of a single vein were recorded. For the width of veins, the outcrops and waste sample were distinguished, particularly in the case of goaves, were assumable in the waste have been recorded and those not assumable were excluded from the record.

At the same time, specimens for chemical analysis during the survey were collected. We used as reference the survey drawings (1/2000) along the vein made by the former East Germany. The locations of sampling were identified by GPS. The intervals of sampling in the precision survey area were determined as following for the average values of major veins which depend on the continuity of the outcrops. It is about 20 meters in the case of goaves.

Vein No.	Interval of sampling (m)
1	27
2	25
3	21
6	38
9	61
10	46

The sampling method of specimens is that the debris of 1 to 3 cm in diameters are taken out by striking the outcrops using hammers and chisels at several points in the direction right angle to the vein. One set of samples weight 800g to 1kg.

In the case of goaves, about 30 boulders remaining abundant in the waste in general were taken out at random from the area ranging two to three meters in width in the right angle to the vein running direction. They were then crushed and 1 kg of them was collected as one set of sample.

In the course of sampling, the survey for gold particles and sulfides was conducted as well. The gold particles were found, though very rarely, in almost all the goaves.

The results of this survey is shown in Table II-2-1 as the list of chemical analysis of the quartz veins. The sampling locations were showed already in Fig. II-1-4 together with other test specimens.

#### 2. Mineralization through field observation

The result of the field survey for gold particles and sulfide minerals in major quartz vein

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using a magnifying glass is as follows. The gold particles were found in most cases near chalcopyrite and galena. However, these sulfide minerals do no always indicate the existence of the gold particles.

The gold particles and sulfide minerals are sporadically found in veins but have not banding structure or zonal arrangement.

#### 1) No. 1 vein

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57 m m

The gold particles were often found in the goaves. Normally the gold particles occurred individually in the quartz veins. Their diameters are usually about 0.1mm. Sometimes the chalcopyrite and rarely the galena are also found. They rarely occurred with the gold particles in the same specimen but found in the specimen nearby. This was common to all veins. (They may be scen coexisting under a microscope.)

#### 2) No. 2 vein

The gold particles were observed in many places in the branched vein goaves of the southern part of this vein. Occasionally the chalcopyrite and the galena were also found.

The gold particles were also found in the outcrops in the northern part of this vein, together with the chalcopyrite and the galena occasionally. When the chalcopyrite and the galena are found in the same specimen they are likely to co-exist. This is common to the other veins.

#### 3) No. 3 vein

The gold particles are observed relatively frequent in southern part of the area but the mining is being conducted locally. The chalcopyrite are found generally.

In the northern part of this vein, the gold particles and the galena are found in a part of its northern most.

4) No. 3' vein

The gold particles are found in one point of the outcrop.

#### 5) No. 6 vein

In the goaves of this vein, the gold particles are found individually or the chalcopyrite and the galena are found nearby.

6) No. 7 vein

No gold particles are found. As the mineralization of the chalcopyrite and the galena

were found in a considerable scope of field. The results of the chemical analysis described later indicate that the content of gold is 0.1/g/t or less.

#### 7) No. 8 vein

The mineralization of the galena was found in a considerable scope of field of this vein and the content of gold is 0.1g/t or less according to the result of analysis.

#### 8) No. 9 vein

This vein extends 3 km long but contains no gold particles. Only the chalcopyrite and the galena are found in some places. The content of gold is low as the result of analysis shows.

#### 9) No. 10 vein

The gold particles are found mainly in the goaves, alone or together with the chalcopyrite and the galena.

#### 2-4 Analysis and test

#### 1. Preparation of thin section of rocks

Since the microscopic observation of thin sections were explained generally in Chapter I, they are omitted in this section .

#### 2. X-ray diffraction

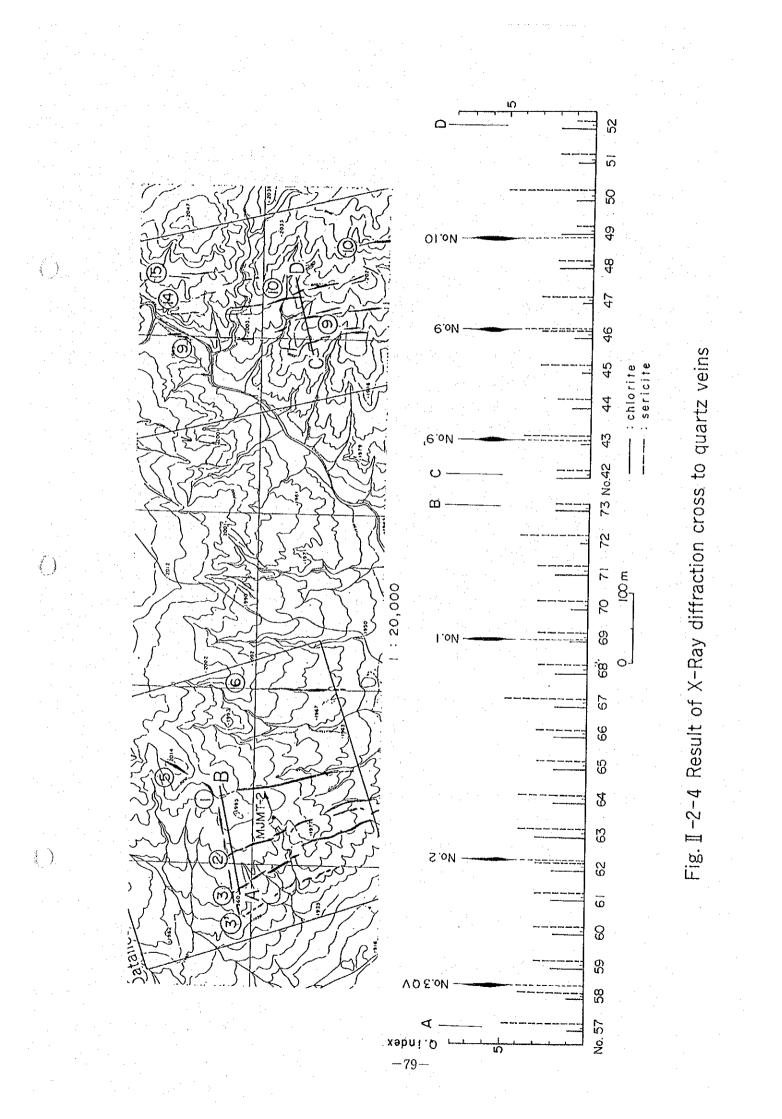
This is also discussed in the chapter 1, omitted in this section.

In this section, the result of the X-ray diffraction test on samples traversing the quartz veins in order to study if any alteration occurred and the degree of the alteration due to the forming of the quartz vein.

A survey line was made at the right angle to the quartz veins in areas I and II to collect samples along the survey line. Sampling intervals were set equally 50m to identify the alterations of the entire peripherals of quartz veins since no alteration was found visually on the host rocks (granodiorite, etc.) as described earlier.

Identified clay minerals were chlorite and (sericite) in samples from every specimens of the both areas. However, it is very likely to presume that the most part of this sericite may be muscovite of rock-forming minerals.

Fig. II-2-4 shows a figure of the result of identification. The figure clarify that these two areas are in the chlorite zone and no relations were found between the levels of quartz index and the locations of quartz veins. Therefore, this chlorite zone is not related to the forming of the quartz



veins. The zone is considered to be the alteration caused by regional hydrothermal alteration or by diagenesis.

#### 3. Chemical analysis of ore minerals

Semi-detailed survey area

Detailed survey area

Number of samples to be analyzed 46 samples 403 samples

Appendix A-3 shows the list of chemical analysis samples and

Appendix A-4 shows the results of the analysis. For the detailed survey area, the grades of Au and Ag are shown in the geological map (1/2,000) of Fig. II-2-1 and Fig. II-2-2. As for the Au, the grades were classified for the detailed survey area and shown in Fig. II-2-5 (1/10,000), attached to the end of the report.

#### 1) Abstract of the result of analysis

The five elements of Hg, Bi, Te, Se and Mo were mostly less than the detectable levels (Hg: 10ppb, Bi: 1ppm, Te, Se, Mo: 5ppm). Au:

The content is high at the old mining zone (waste sample: several 10ppm) and less generally in the quartz veins outcrops. No. 1, 2, and 3 veins of the major vein have higher contents and No. 9 and 10 are locally higher.

Ag:

The Ag is often found together with Au and the ratio of Au/Ag contents is approximately 5:1. The maximum level of Ag was 83ppm (No.10) and the same sample had the Au content of less than 0.1ppm or less.

As:

The As is found alone or together with Au or Ag and the maximum level was 94ppm.

Sb:

The Sb is relatively in a good association with Au. It also associated with Ag and As. The level is mostly less than order of several 10 ppm, the maximum level being 400ppm. Hg:

**A** 1

In No.1 quartz vein, the higher content is in a range of between 500 and 1,700ppb. Bi:

The Bi is found in extremely limited parts of veins, the content is between 20 and 80ppm nothing to do with the gold grade.

Te:

The content is occasionally between 20 and 40ppm and is associated with the high grade portion of the gold.

Se:

Not detectable in all samples.

Mo:

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The Mo is detectable in only 5 samples with the maximum level 44ppm. Gold content of those samples are less than 1 ppm.

#### 2) Average content of Au

The average width of the major quartz veins, the average content of Au and the total length are shown in Table 11-2-1.

Area	No. of vein	Thickness	Gold content	Total length
		(cm)	(g/t)	(m)
I	7	12.6	0.03	750
Ι	. 7'	8.5	0.19	±500
I .	8	15.3	0.27	1,500
I	9	26.6	0.71	3,000
<u> </u>	9'	17.8	0.21	400
I	9″	13.0	0.26	400
1	10	42.3	10.01	2,800
I	14	19.7	0.10	400
1	15	10.8	2.43	400
I	16~23	13.2	0.66	> 1,500
П	1	20.7	21.97	900
П	2	18.9	7.78	1,000
П	3	23.6	6.15	1,000
<u>П</u>	3'	22.8	0.92	800
Π	6	26.5	2.55	800

Table II -2-1 Gold content, thickness, length of main quartz veins

### 4. Polished sections of ore minerals

Quantity: Field samples 35 Drilling core 7

In the field, samples with the sign of mineralization (native gold, sulfide minerals) were collected. Cores had no sign of mineralization by naked eye excluding samples from the west vein of MJMT-1 No.10 vein (depth: 299.84m). Table II-2-2 shows the result of microscopic observation.

The ore minerals found through the microscopic observation were the primary minerals such as native gold, chalcopyrite, galena, the sphalerite, pyrite, tetrahedrite, and secondary minerals such as the covelline and the geothite. The altaite, tellurium sulfide, tellurium oxide minerals also may be found. The occurrence of the individual minerals are as follows:

#### Native gold

The average diameter of native gold is 0.01 to 0.05 mm and the maximum is 0.3 X 1.0mm. Most of the gold is found independently between quartz grain crystals. Some are found along the cracks of quartz or contained in chalcopyrite or with tetrahedrite.

#### Chalcopyrite

Co-exists with galena and pyrite or involves tetrahedrite. Chalcopyrite is secondarily replaced by covelline and sometimes by goethite.

#### Galena

Galena occurs between quartz grain crystals and involves chalcopyrite.

#### Sphalerite

Sphalerite is contained in tetrahedrite or associated with the chalcopyrite.

#### Pyrite

Pyrite is found in most of the samples and occurs between quartz grain crystals or together with the chalcopyrite or contained in it. In most cases, pyrite is replaced by goethite.

#### Tetrahedrite

Tetrahedrite occurs between quartz grain crystals or associated with chalcopyrite or sphalerite.

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-	Sample#	مر من م الس	ordinate	No. of	Average				emblage	S .			,	
No.	960-	46N	100E	Qz vein	width(m)			Primar	Y :			Secon	dary	Remar
					and and a second se	Au	Cp	Ру	Gn	Sp	Tet	CV	Go	[
1	827011	8, 12	9.42	1	0.20*	0	Ø.	•				Ø	0	*:
2	827019	8.04	9.44	· 1 .	0.20*	Ô					+			waste
3	830036	7.75	9.39	2	0.02	•		•						sa∎pi
4	830041	7.74	9.38	2	0.10*	-0		·		. <sup>1</sup> .				
5	830044	7.67	9.41	2	0.10*	Ø	ν.						- -	1 .
6	830046	7,88	9.30	2	?	+		•			+-		Ø	
7	826041	7.81	9.25	3	0.20	•	Ö	· · · ·		+	0	0		
8	826041	7.81	9.25	3	0.20	: . *	0	•	0	1. A.		1.11	0	· ·
9	919005	8.58	8.58	4	?	0		•						
10	901011	7.99	9.86	6	0.25*		0	•	0			0	0	
11	901012	7.97	9.86	6	0,25*	•	0					0	0	
12	901012	7.97	9,86	6	0.25*	+		•			·		Õ	
13	901016	7.92	9.85	6	0.30*	0		<b> </b>		   .			<u> </u>	1
14	907005	7.39	10.95	7	0.10			0	0			0	0	
15	907008	7.27	11.03	7	0,20		0	0	0	; '	1	Õ	0	1
16	909009	7.37	10.42	7'	0.20*			+					1 O	ł
17	906008	7.03	11.10	8	0.25				0					
18	912062	7.43	11.83	9	0.30	•	0	+				0	0	
19	912063	7.41	11.83	9	1.00	0		+			· · · · ·		0 0	1 - L
20	920027	6,38	12.38	none No.	?	Õ		+					0	· .
21	920028	6. 41	12.30	none No.	?		0	- <u>-</u>			· · · ·		0	
22	903044	8.09	11.60	10	0.50*	 •			   +				+	
23	903047	8.03	11.61	10	0.30		0	0	© i			+	$\overline{0}$	
24	903061	7.68	11.89	10	0.40*	0		- <u>·</u> -				·	<u> </u>	
25	903064	7.63	11.89	10	0.40*	+							·····	
26	903064	7.63	11.89	10	0.40*	•	©		+					
27	903092	7.11	12.23	10	2		0					+	+	- 1. - 1.
28	903107	6.98	12.32	10	0.30*		0	 +			+	1		
29	903107	6.98	12.32	10	0.30*		9			· ·		+	0	
30	920015	6.50	12. 32	none No.	<u>0.30</u> ≉ ?		0	L					0	
31	916003	8.28	12.40	15	0.20*	0		 -+-		-+-	<u> </u>	0		
32	919011	8.81	7.60	27	0.20* ?				0		1		+	
33	919011	8.81	7.60	27	?		+				+		0	
33 34	920011	6.55	12.95	48			-	+	0		0	- <del>1</del> -		
54 35	921007	9.04		····	?	+		- <b> </b> - 	0				+	, ,
55 36	921007 HJHT-1 20		11.81	silicif.	?					· · · · ·			0	
30 37		<u> </u>		10	0.067			0						
	KJHT-1 29			10(W)	0.04			0	+					
38	HJHT-1 29			10(W)	0.05			0	0	-1-				
39	HJHI-1 29			10(W)	0.05			Ô	0					
10	HJHT-2 15			1	0.10			0		•				
11	HJHI-2 15			1	0.14			0				·.		
12	HJHT-2 15	3.50m~1	53.55M	· 1	0.05			Ô		•				

Table I-2-2 Hicroscopic observation of polished sections (Semidetailed, Detailed & Drilling core)

Cv: covelline Go: goethite unknown tellurium sulphide (〇).

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te Py: pyrite Gn: galena Sp: sphalerite Tet: tetrahedrite • except table mark,occurs altaite (O),unknown tellurium oxide(©), • Relative content grade: @>O>+>•

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# Covelline

Covelline replaced secondarily to chalcopyrite and tetrahedrite.

Goethite

Goethite mostly replaced to pyrite.

## 5. Measuring homogenization temperature of fluid inclusions

Homogenization temperatures were measured on 27 samples of major quartz veins and drilling cores. Samples were collected from the middle of veins not having the flow structure for the outcrops. Other samples were taken from the waste sample in old mining where native gold or sulfide minerals were found.

The diameter of the fluid inclusion are generally less than 2.5 to 10 micrometers and most of them are in the shape of egg or polygon. Most of them are of vapor-liquid phase and some are of vapor phase. The content ratio of vapor-liquid phase is generally as low as 2 to 5 % and 10% in some of them.

Measurements of homogenization temperature were multiplied by 2,000 due to their small grain diameter to see primary inclusions on CRT.(Television monitor).

The microscope and heating device TH-600 manufactured by LINCOME was used for temperature measurement with the final rate of temperature increase of 1.0 to 0.1°C per minute. The inorganic reagents Benzanilide (163°C) and Sodium-nitrate (305°C) were used for the temperature correction.

The histogram of these measurement result of samples have almost normal distribution although some samples have two peaks.

The results of the measurements are shown in Table II-2-3 (1)  $\sim$  (5). The average values of homogenization temperatures of these samples are shown respectively in Table II-2-4.

Q.V.No.	No. of sample	Homogenization temperature(°C)	Remark
1	2	139.5, 139.8	
2	2	165.0, 151.0	
3	1	145.0	
4	1	177.4	
- 5	2	147.5, 153.9	
6	1	157.5	
- 7	1	128.1	barren quartz
7,	1	125.3	barren quartz
9	3	122.2, 116.4*, 114.1*	* barren quartz
10	5	115.5*,133.6*,154.1, 118.4, 147.1	* barren quartz
nonc No.	2	127.0,115.6	barren guartz
1	3	155.8,146.6, 156.2	MJMT-2 152.60m ~153.65m
10	2	156.2,197.1	MJMT-1 202.78m ~203.20m
10W	1	147.6	MJMT-1 298.84m ~299.04m

Table II-2-4 Average homogenization temperature of fluid inclusion of each veins

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Table I-2-3 Result of homogenization temperature mesured from fluid inclusion in quartz vein (detailed area, dril.core) (1)

											· · · · · · · · · · · · · · · · · · ·
	Hi ctodram			120 150 °C	 120 150 ℃	140 170 °C	$\frac{1}{130}  \frac{1}{160} \text{ °C}$	<u>                                      </u>	$\begin{array}{c c} \hline \hline$	<u> </u>	$\begin{bmatrix} \hline 1 \\ 140 \\ 170 \\ 0 \end{bmatrix}$
מדובח מ		Std.	9.7		8.5 2	11.2	14.3	15.6	23. 5	16.3	15.6
זזז לתכו	Homogenization	Avg.	139.5		139.8	165. 0	151. 3	145.0	177.4	147.5	153.9
מדרב גם	omogeni emnerat	Min.	126	14 14 14	128	146	131	120	142	127	133
מה זיד זו	œ +	Max.	162		153	182	172	177	219	173	186
NENTOIT DINIT	Size of inclusions	$(\pi m)$	<2.5~ 5.0		<2.5~ 5.0	<2.5~20.0	<2.5~10.0	2. 5~12. 5	<2.5∼10.0	<2.5~ 5.0	<2.5~ 7.5
Result of Romogenization temperature mesured from figure inclusion in that is very vecaried area.	Number of		primary(20)		primary(12)	primary(20)	primary(10)	primary(20)	primary(20)	primary(15)	primary(14)
הדבימותוב	No. of 07 voin	1112A 7A	<del>,</del> .		<b>7~4</b>	23	62	ന	4	5	പ
CITZALIUI LEI	inate	Latit.100E	9.42	· · · · · · · · · · · · · · · · · · ·	9.44	6. 38	9. 41	9. 25	9. 58	10. 41	9.04
ROMON TO JIN	GPS Cordinate	Longit.46N	8.12		8.04	7.74	7. 67	7. 81	8. 58	9.10	8. 25
Iableu-2-3 Kes	Sample#		960827011		960827019	960830041	960830044	960826041	960919005	960806002	960813038
an	Ň				€3	ന	4	വ	9	2	8
					:				·····	•••••	•

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area, dril.core) (2)	Π <sup>4</sup> ο Φολακορικο	mp rZna stu		140 170 °C		$\begin{array}{c c} \hline \\ \hline \\ 100 \\ 130 \\ 160 \\ \hline \end{array}$		110 140 °C				100 130 °C	
tailed a		Std.	13.5		17.0		11.0		18.0		8.6		
ein (de	Homogenization	AVS.	157.9		128.1		125.3		122.2		116.4		/
uartz v	Homogen	Lengera	141		108		113		96		103		
o ui no		Мах.	189		165	· ·	144		152		131		
fluid inclusion in quartz vein (detailed	Size of	SIDTSTATAT	<2.5~ 5.0		<2.5~10.0		<2.5~ 7.5		2. 5~32. 5		<2.5~ 7.5		
Result of homogenization temperature mesured from	Number of	SHOTENTOIT	primary(20)		primary(20)		primary(20)		primary(20)		primary(20)		
mperature	No. of		ß		L		<i>`L</i>		6		ß		
genization te	GPS Cordinate	Latit.100E	9.86		11.03		10.42		11.29		11.42		
ult of homog	GPS COI	Longit.46N	7.97		7.27		7.37		9, 08		8. 15		*
Table -2-3 Res	Sample#		960901012		960907008		60060696		960912005	•	960912032		
Tabl	No.		റ		10		1		.12		со Н	· · · ·	

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Table I-2-3 Result of homogenization temperature mesured from fluid inclusion in quartz vein (detailed area, dril.core) (3)

	· · ·								
ature mesured iron initia inclusion in quariz vern quetatieu area, unit. core) vol-	Wistodram.	1170 COST CHIII			100 130 °C	$100  130  160  \mathbb{C}$		$\begin{bmatrix} \hline \\ 100 \\ 130 \\ 0 \end{bmatrix}$	
מזזבח מ		Std.	5.0		8.5	15.7	15.1	9.6	13.9
חבה וחבר	Homogenization	Avg.	114.4		115.2	133. 6	154.1	118.4	147.1
מורל עם	omogeni	Min.	105		102	101	119	105	126
חה ווז וו	FFI +	Max.	126		131	161	621	134	172
OTSNIDUT DINIT	Size of inclusions	(m m)	<2.5~12.5		<2.5~10.0	<2.5~12.5	<2.5~12.5	<2.5~ 2.5	<2.5~ 5.0
mesured trom	Number of	SHOTSHIDJIT	primary(20)		primary(20)	primary(20)	primary(20)	primary(12)	primary(20)
	No, of Og vein	TT2A 7A	ත ප		10	10	10	10	10
Result of nomogenization temper	dinate	Latit.100E	11.83		11.46	11.50	11.60	11. 89	12. 23
gomon To Tlu	GPS Cordinate	Longit.46N	7. 43		<b>9.</b> 09	8 96	8.09	7. 63	7.11
lablel-2-3 Kesi	Sample#		960912062		960903007	960903009	960903044	960903064	960903092
labi	No.		14	1	15	19	17	18	61
			2		· · · ·				

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			· · · ·						*			
fluid inclusion en quartz vein (detailed area, dril.core) (4)		niscogram				100 130 °C	120 150 180 °C		130 160 °C	140 170 °C		$120  150  180  \mathbb{C}$
ailed a		Std.	13. 5		10.1		16.1	8. 9		9.6	11.1	
ein (det	eization	. Min. Avg.	127.0		115.6	<sup>-</sup>	155.8	146.6		156. 2	156.2	
uartz ve	Homogen	tempera	103	· · · · ·	104		127	132		138	138	
on en q		Max	143		134		184	166		172	181	
fluid inclusion	Size of inclusions $(m \mu)$		<2.5~ 5.0		<2.5~12.5		<2. 5~65. 0	<2.5~ 5.0		<2. 5~25. 0	<2.5~10.0	
Result of homogenization temperature mesured from	Number of inclusions		primary(20)		primary(10)		primary(14)	primary(13)		primary(10)	primary(20)	
mperature	No. of 07 vein	11724 72	none No.		none No.		<b></b> 1	<del>,</del> 1		<b>p≖-i</b>	10	
genization te		Latit.100E	13.01		10.63	-	I	1			1	
ult of homog	GPS Cordinate	Longit.46E	7.05		6. 83		1000 - 10000 - 10000 - 1000 - 1000 - 1000 - 1000 - 1000 - 1000 - 1000 -	1		• • • • • • • • • • • • • • • • • • •	ŀ	
Table1-2-3 Res	Sample #		96084040		910606096		MJMT-2 152.60m	MJMT-2 152.95m ~153.09m		МЈМТ-2 153. 50m 153. 65m	MJMT-1 202.78m ~202.88m	
Tabl	No.		20		21		22	53		24	25	

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Table1-2-3 Result of homogenization temperature mesured from fluid inclusion en quartz vein (detailed area, dril.core) (5)

{ }

[				
Histogram			120 150 °C	
	Std.	25. 1	14.8	
ization Tre(°C)	Avg.	197.1	147.6	
Homogeneization	Min.	156	123	
ĔĂŦ	Мах.	229	182	
Size of inclusions	(m m)	2. 5~22. 5	<2. 5~15. 0	
Number of	STICTOTOTOT	primary(20)	primary(20)	
No. of		10	10#	
	Latit. 100E	1	t i t i t i t i t i t i t i t i t i t i	
# GPS Cordinate	Longit.46E	I		
ampie		MJMT-1 203.08m ~203.20m	NJNT-1 298.84m ~299.04m	
No.		26	27	

The homogenization temperatures of field samples, drilling core and barren vein are as follows:

Field quartz vein	165°C to 118°C
Core quartz vein (Subsurfac	e depth of 150m or more)
	197°C to 147°C
Field barren veins	137℃ to 114℃

The core samples are the highest and the field quartz vein and the barren vein had trends to be lower in this order. As a whole, these temperatures are extremely lower than those of Japan's epi-thermal gold vein deposits which range from 240 to 180°C (Enjoji, Takeuchi, 1976).

The reason for this is not yet known. The homogenization temperature is related also to the pressure condition and the saline composition in inclusion when veins were formed. This will be the subject of future studies.

Examples of gold bearing quartz vein deposits having temperature ranges identical value to this area are as follows:

Todoroki mine	150 to 122°C (Fujiwara, 1961)
Hikaritaki mine	180 to 140°C (Fujiwara, 1961)
Teine mine	186 to 160°C (Fijiwara, 1961)
Charters Towers gold	140 to 192°C (S.G. Peters and S.D.
field, Australia	Golding, 1988)
(Host roak) Ordenisian	Devenin Constitution

(Host rock: Ordovician - Devonin Granitoids)

Particularly, Charters Towers gold field is in the gold ore deposit area and one sample out of 11 measurement samples of 4 mines was 220 to 230°C and others are approximately identic to the temperature range of the core quartz veins described above.

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#### Chapter 3 Supplementary survey

In detailed area, altered zone survey by POSAM (Portable spectro radiometer for mineral identification) and sketch of trenches excavated by former East Germany were carried out.

The details of the surveys are as follows,

Alternation zone survey area:

(200m grid)

Trench survey:

### 220 locations, total length of 1,540m

282 samples of measurement

#### 3-1 Altered zone survey

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( )

Samples were collected along the geophysical survey line (will be explained later) for identifying altered minerals. The results are shown in Table II-3-1.

Identified clay minerals are chlorite, montmorillonite and (sericite). Although the sericite is not distinguishable from the rock-forming muscovite such as the granodiorite. Distribution map of the identified clay minerals shown in Fig. II-3-1.

In the detailed survey area I as shown in the figure, the sericite zone distributed in the granodiorite and the adamellite areas, and the chlorite or montmorillonite zone distributed mainly in the metamorphic areas. Particularly the diorite area is covered by chlorite zone.

In the sericite zone, distribute locally the chlorite or the montmorillonite zones but these have no structural relation to the quartz veins.

Detailed area II is mostly of the sericite zone corresponding to the geology and there are also the chlorite or the montmorillonite area locally but have not relation to quartz veins.

From the above mentioned, in the area of metamorphic rock, it is estimated that the chlorite or the montmorillonite zone was formed before the intrusion of granodiorites. The altered zone in granodiorite is also estimated that formed by regional alteration after intrusion of quartz veins.

The clay mineral survey by POSAM in the granitic field is made sure that may be not suitable.

#### 3-2 Trench survey

The trench re-surveys were carried out in quartz veins No. 1, 2, 3, 6,7, 8, 9, and 10. They are classified as follows:

1) Those have excavated mainly for collecting samples from the waste at the old mining zone.

2) To confirm unclear exposed outcrops

3) To confirm the continuity of veins where there are no outcrops

# Table 1-3-1 Identification of clay minerals by POSAM

) No.		Rock name	Clay mineral	NO.	Line-No.	Rock name	Clay mineral
1	A-13	Granodiorite	Ser	72	K-23	Granodiorite	Chl?
2	A-15	Granodiorite	Epi, Chl, Ser	73	<u>L-13</u>	Granodiorite	Ser>Hont
3	A-17	Granodiorite	Ser	74	L-15	Granodiorite	Ser, Hont
4	A-19	Granodiorite	Ser	75	L-17	Granodiorite	Ser
5	A-21	Granodiorite	Ser	76	L-19	Granodiorite	Ser
6	A-23	Basalt-dyke	Chl?	77	L-21	Granodiorite	Chi, Hont, Ser?
7	A-23 B-13	Granodiorite	Ser, Hont?	78	L-23	Granodiorite(si-alt)	Chl?
9	B-13 B-15	Granodiorite	Ser	79	H-13	Granodiorite	
10	B-13 B-17	Quartz porphyry-dyke Granodiorite	Sor	80	H-15	Granodiorite	
ii i	B-19	Basalt-dyke	Ser Chl?	81	H-17	Granodiorite	Ser
12	B-19	Granodiorite	10	82	<u>H-19</u>	Granodiorite	Ser
13	8-21	Granodiorite	l Ser /	83	H-21	Gneiss	Epi
14	B-23	Granodiorite	Ser	85	<u>H-23</u> N- 1	Gneiss	?
15	C-13	Granodiorite	Ser	86	N- 3	Granodiorite	Ser
16	C-15	Basalt-dyke	Ch1?	87	N-5	Granodiorite Granodiorite	Ser, Chl
17	C-15	Granodiorite	Ser	88	N 7	Granodiorite	Ser Ser
18	C-17	Basalt-dyke	Chi?	89	N- 9	Quartz porphyry-dyke	
19	C-17	Granodiorite	Ser, Hont	<u> </u>	<u>N-9</u>	Granodiorite	Ser, Hont
20	C-19	Granodiorite	Ser	91	N-11	Granodiorite	Ser, Hont
21	C-21	Granodiorite	Ser	92	N-13	Granodiorite	Ser, Hont
22	C-23	Granodiorite	Ser>Hont	93	N-15	Granodiorite	Ser?
23	0-13	Basalt-dyke	Chl	94	N-17	Granodiorite	Ser
24	0-13	Granodiorite	Ser, Hont	95	N-19	Granodiorite	Ser
25	D-15	Granodiorite	Ser	96	N-21	Granodiorite	Ser, Ch
26	<u>D-17</u>	Granodiorite	Ser	97	N-21	Gneiss	i Chi
27	D-19	Granodiorite	Ser	98	N-23	Gneiss	
28	D-21	Granodiorite	Ser, Chl	99	K-25	Granodiorite	Hont, Chl
2 <u>9</u> 30	D-23	Granodiorite	Ser, Hont, Chl	100	N-25	Gneiss	Epi>Chl
11	E-13 E-19	Quartz porphyry-dyke	Ser	101	0-1	Granodiorite	Ser
2	E-19 E-21	Granodiorite	Ser	102	0-3	Granodiorite	Ser
33	E-23	Granodiorite	Ser, Chl	103	.0- 5	Granodiorite	Ser, Chl
34	F-13	Granodiorite	Ser	104	0-7	Granodiorite	Ser
35	F-15	Quartz porphyry-dyke Granodiorite	Ser Ser	105	0-9	Granodiorite	Ser
<u>86</u> 1		Granodiorite	Ser	106	0-11	Granodiorite	Ser, Hont
37	F-19	Granodiorite	Ser	107:	0-13	Granodiorite	Ser, Chl
88	F-21	Granodiorite	Ser	108	0-15	Granodiorite	Sor, Chl
19	F-23	Granodiorite	Ser, Hont	110	0-17	Granodiorite	Ser, Chi?
10	G-13	Quartz porphyry-dyke	Ser	111	0-15	Granodiorite Granodiorite	Ser
11	G-13	Granodiorite	Ser	112	0-23	Granodiorite	Ser Chi Hant
12	G-15	Granodiorite	Ser, Kont	113	0-23	Gneiss	Ser, Chl, Hont
3	G-17	Granodiorite	Ser>Hont	114	0-23	Gneiss	Chi Chi
4	6-19	Granodiorite	Ser, Chl	115	0-25	Granodiorite	Ser?
5	G-21	Granodiorite	Ser	116	0-25	Gneiss	Ser, Chl
6	G-23	Gneiss	Chl	117	P- 1	Granodiorite	Ser, Hont, Chl
7	11-13	Granodiorite	Ser	118	P- 3	Granodiorite	Ser
8	ll- <b>1</b> 5	Granodiorite	Ser <hont< td=""><td>119</td><td>P- Š</td><td>Granodiorite</td><td>Ser</td></hont<>	119	P- Š	Granodiorite	Ser
9	H-17	Granodiorite	Ser, Hont	120	P- 7	Granodiorite	Ser
0		Granodiorite	Ser	121	P- 9	Quartz porphyry-dyke	Ser
1	11-21	Granodiorite	Ser	122	P- 9	Granodiorite	Ser
$\frac{2}{2}$	11-23	Gneiss	Ser>Epi	123	P-11	Granodiorite	Ser
$\frac{3}{4}$	<u>1-13</u> 1-15	Granodiorite	Ser	124	P-13	Granodiorite	Ser
$\frac{4}{5}$		Granodiorite	Ser	125	P-15	Granodiorite	Ser
6		Granodiorite	Ser Chil?	126	P-17	Granodiorite	Ser
7		Granodiorite Granodiorite	Ser, Chl? Chl, Ser	127	P~19	Granodiorite	Ser
8		Gneiss	un ser	128	P-21	Granodiorite	Ser, Hont, Chi
ğt		Granodiorite	Ser <hont< td=""><td>129</td><td>P-23</td><td>Granodiorite</td><td>Ser, Hont, Chl</td></hont<>	129	P-23	Granodiorite	Ser, Hont, Chl
ŏt		Granodiorite	Ch1?	130 131	P-23 P-25	Gneiss	Chl.Ser
Ť		Granodiorite	Ser	132	P-25	Granodiorite	Ser
21	J-17	Granodiorite	Ser	133	Q-1 Q-3	Granodiorite	Ser
3		Granodiorite	Ser	134	Q-3 Q-5	Granodiorite Granodiorite	Ser
1	J-21	Granodiorite	Ser	135	Q-7	Granodiorite	Ser Ser
5	J-23	Granodiorite	Ser	136	Q- 9	Granodiorite	Ser
6	J-23	Granodiorite		137	Q-11	Granodiorite	Ser
7	K-13	Granodiorite	Ser	138		Granodiorite	Ser
8	K-15	Granodiorite	Ser	139	0-15	Granodiorite	Ser
9	K-17	Granodiorite	Ser	140		Granodiorite	Ser
0	K~19	Granodiorite	Scr, Hont, Chi	141		Granodiorite	Ser <hont< td=""></hont<>
1	K-21	Granodiorite	Ser, Hont, Chi	142			Ser, Ch1?
	<b>Usilici</b>	cation Ser:sericite	Chl:chlorite Epi;				

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143	Line-No. Q-23	Rock name	Clay mineral	NO.		Rock name.	Clay minera
144	Q-25	Granodiorite	Ser?	183	T-11	Granodiorite	Ser
45	R~ 1	Granodiorite	Ser, Hont, Chl	184	1-13	Granodiorite	Ser, Hont
146	R-3	Granodiorite	Ser	185	1-15	Granodiorite	Ser
47		Granodiorite	Ser	186	1-17	Granodiorite	Ser
48	<u>R-5</u> R-7	Granodiorite	Ser	187	T-19	Granodiorite	Ser
	R-7	Quartz porphyry-dyke	Ser	188	I-21	Granodiorite	Ser
19		Granodiorite	Ser, Chi	189	1-21	Granodiorite(pelitic)	Ser
50	R- 9	Granodiorite	Ser	190	1-23	Greiss	Ser
51	<u>R-11</u>	Granodiorite	Ser	191	T-25	Granodiorite	Ser
52	R-13	Granodiorite	Ser	192	I-25	Gneiss(pegmatite)	Ser
53	R-15	Granodiorite	Sor	193	U- 5	Quartz porphyry-dyke	Ser
54	1-17	Granodiorite	Ser	194	U- 5	Granodiorite	Ser
55	R-19	Granodiorite	Hont	195	U- 7	Granodiorite	Ser
56	11-19	Gneiss		196	U- 9	Granodiorite	Ser
57	<u>R-21</u>	Granodiorite	Chl.Ser?	197	U-11	Granodiorite	Ser
58	R-21	Gneiss	Chl,Epi,Hont	198	U-13	Granodiorite	Ser, Hont
59	<u>R-23</u>	Granodiorite	Ser	199	U-15	Granodiorite	Ser
60	R-23	Gneiss	Ser, Hont, Chl	200	Ŭ-17	Granodiorite	Ser
61	R-25	Granodiorite	Ser, Hont, Chl	201	<u> </u>	Granodiorite	Ser
62	R-25	Gneiss	Chl, Ser	202	Ŭ-21	Granodiorite	Ser
63	S- 1	Granodiorite	Ser	203	<u>Ŭ-23</u>	Diorite	
64	<u>S- 1</u>	Gneiss(pelitic)	Chl	204	Ŭ-25	Diorite	
65	<u>S-3</u>	Granodiorite	Ser	205	V- 5	Granodiorite	Ser
66	R- 5	Granodiorite	Ser	206	V- 7	Granodiorite	Ser
67	R-7	Granodiorite	Ser	207	V- 9	Granodiorite	Ser
68	R- 9	Granodiorite	Ser	208	V-11	Granodiorite	Con
69	R-11	Granodiorite	<u>Ser</u>	209	V-13	Granodiorite	<u>Ser</u>
70	R-13	Granodiorite	Ser	210	V-15	Granodiorite	Ser
71	R-15	Granodiorite	Ser	211	V-17	Granodiorite	Zeo, Ser
72	R-17	Granodiorite	Ser	212	V-19	Granodiorite	Ser
73	R~19	Granodiorite	Ser, Hont	213		Granodiorite	Ser, Hont
74	R-21	Granodiorite	Ser	214	V-23	Diorite	Ch ?
75	R-21	Gneiss	Ser	215	V-25	Diorite	<u>Ch17</u>
76	R-23	Granodiorite	Ser, Hont	216	H-1	Gneiss(marble)	Cal
77	R-23	Gneiss	Ser	217	₩-11	Granodiorite	
18	R-25	Gneiss	Ser	218	W-13	Granodiorite	Ser
79	T-3	Granodiorite	Ser	$\frac{210}{219}$	W-15	Granodiorite	Ser
30	T- 5		Ser	220	H-17	Chapadianita	Ser
31	Ī-7		Ser	221	H-21	Granodiorite	
32	1-9		Ser	222	<u></u>	Diorite	
				1666	W-25	Diorite	Ser

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2)		L					
N	o. Line-No.	Rock name	Clay miner	al No.	Line-No.	Rock name	Clay mineral
·	<u>1 A-1</u>	Granodiorite	l Ser	31	C- 9	Granodiorite	Ser
	2 A-3	Granodiorite	Ser	32	Č-11	Granodiorite	Ser
	3 A-5	Granodiorite	Ser?	33	C-13	Granodiorite	Ser
	4 A-7	Granodiorite	Ser, Ch1?	34	D- 3	Granodiorite	Ser
	5 A-9	Granodiorite	Ser, Chl	35	D- 5	Granodiorite	Ser
	6 A-11	Granodiorite	Ser	36	D- 7	Granodiorite	Ser
-	7 A-13	Granodiorite	Ser, Chl	37	0-9	Granodiorite	Ser
	8 A-15	Granodiorite	Ser	38	D-11	Granodiorite	
	9 A'- 1	Granodiorite	Ser	39	D-13	Granodiorite	Ser
- 1	0 A'-3	Granodiorite	Ser, Ch1?	40	0-15 0-15		Ser
1	1 A'-5	Granodiorite	Ser	41	E- 3	Granodiorite	Ser
1	2 A'-7	Granodiorite	Ser, Chl	41	E-5	Granodiorite	Ser
1	3 A'- 9	Granodiorite	Ser, Ch	43	E- 7	Granodiorite	Ser
		Quartz porphyry-dyke	Ser	40		Granodiorite	Ser
1.		Granodiorite	Ser		E- 9	Granodiorite	Ser, Chi
1	6 A'-13	Granodiorite	Ser	45	E-11	Granodiorite	Ser
1		Granodiorite	Ser		E-13	Granodiorite	Ser
1	8 B-1	Granodiorite	Ser	47	E-15	Granodiorite	Ser
- h		Granodiorite	Ser	48	F- 3	Granodiorite	Ser
2	0 B- 5	Granodiorite		49	<u>F- 5</u>	Granodiorite	Ser
2	1 8-7	Granodiorite	Ser>Hont	50	<u> </u>	Granodiorite	Ser
2		Granodiorite	Ser, Hont, Chi	51	F- 9	Granodiorite	Ser
1		Granodiorite	Ser	52	F-11	Granodiorite	Ser
2	4 B-13	Granodiorite	Ser	53	F-13	Granodiorite	Ser, Hont
2			Ser	54	F-15	Granodiorite	Ser
20		Granodiorite	Ser	- 55	G- 5	Granodiorite	Ser
2		Granodiorite	Ser	56	G-7	Granodiorite	Ser
2		Granodiorite(si-alt)	Ser, Hont	57	G- 9	Granodiorite	Ser, Chl
29		Granodiorite	Ser	58	G-11	Granodiorite	Ser
$-\frac{23}{30}$		Granodiorite	Ser	59	G-13	Granodiorite	Ser, Chl
		Granodiorite(si-alt)	Chl, Eip, Ser	60	G-15	Granodiorite	Ser
51	alt:silicif	ication Ser:sericite	Chl:chlorite	Epi:epidot		montmorillonite	Zeo: zeolite Cal calcite

Hont:montmorillonite Zeo:zeolite Cal;calcite

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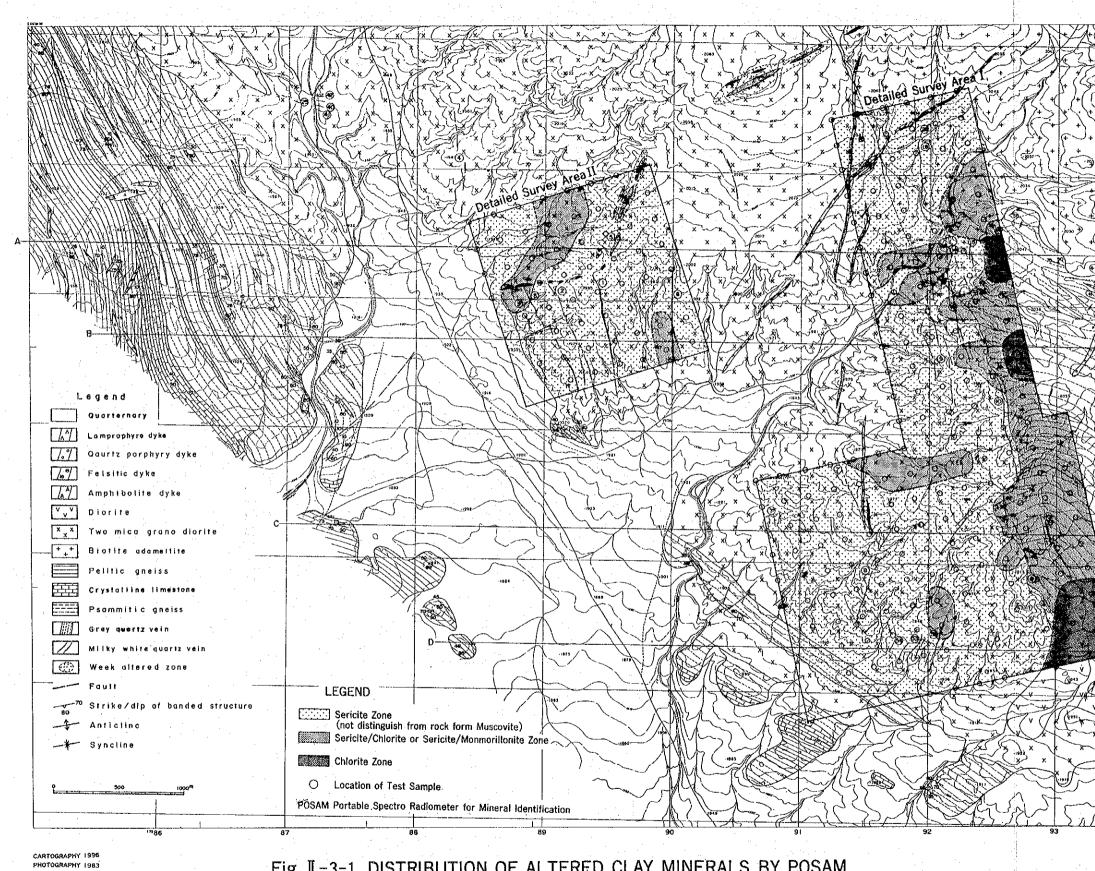
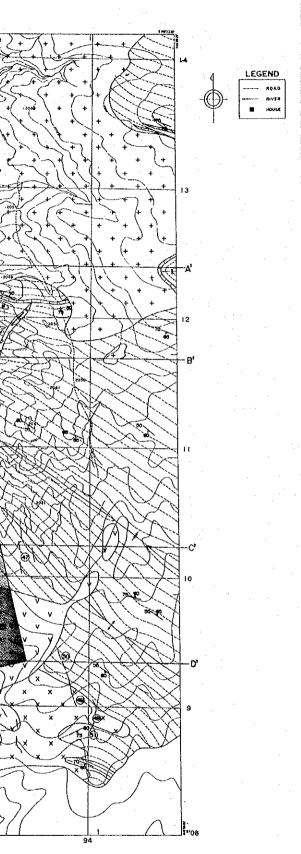


Fig. I-3-1 DISTRIBUTION OF ALTERED CLAY MINERALS BY POSAM



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The quartz vein No. 1 is mostly falls under a).

The survey were carried out at 220 locations that had been confirmed (212 locations according to the reports of the former East Germany). The presentative sketches are described in Fig. II-3-2 (1) to (8) for 6 to 10 locations of each veins.

No chronological relations were clarified between the quartz veins No.3 and No. 5 in the current survey (Fig. II-3-2(1)).