

3.4 HYDROGEOLOGICAL FEATURES OF AQUIFERS

3.4.1 Objective Aquifers to the Existing Wells

Most of the screens of existing wells in the Study Area are installed in the layers of scoria, sand and gravel intercalated in Kaybubutong Formation and the layers of sandstone or conglomerate intercalated in Talisay Formation.

In case of the well located in the low elevation zone (lower than 50 m), the screens are installed in the conglomerate layer intercalated in the lower horizon of Kaybubutong Formation and in conglomerate or sandstone layer intercalated in Talisay Formation.

In case of the well located in the middle elevation zone (50 – 200 m), the screens are installed in the layers of scoria or gravel intercalated in the middle and lower horizons of Kaybubutong Formation and Talisay Formation.

In case of the well located in the high elevation zone (higher than 200 m), the screens are installed in the andesite lava or scoria bed intercalated in the middle of Kaybubutong Formation. The major springs in the Study Area also gush out from these layers.

3.4.2 Hydrogeological Characteristics of Main Aquifers

Table 3.4.1 shows the hydrogeological constants obtained from electric logging and pumping test conducted on the four test wells.

According to the drilling results of the four test wells, three aquifers, shown in Table 3.4-2, were distinguished.

Middle Aquifer is the best aquifer among the three judging from its large specific capacity (Sc) and transmissivity (T). The sediments consisting this aquifer were deposited at the end of Cavite Slope as alluvial fan conglomerate.

Electric conductivity is low in Upper Aquifer and Middle Aquifer but high in Lower Aquifer. Usually, confined groundwater shows high electric conductivity because pressure holds dissolved minerals in solution. Consequently, groundwater in Upper and Middle Aquifers is considered as unconfined and that in the Lower Aquifer is as confined.

3.4.3 Distribution of Specific Capacity

Fig. 3.4-1 shows the distribution pattern of specific capacity based on existing well data and the results of the pumping test carried out in this Study.

TABLE 3.4-1 DETAILS OF THE TEST WELLS

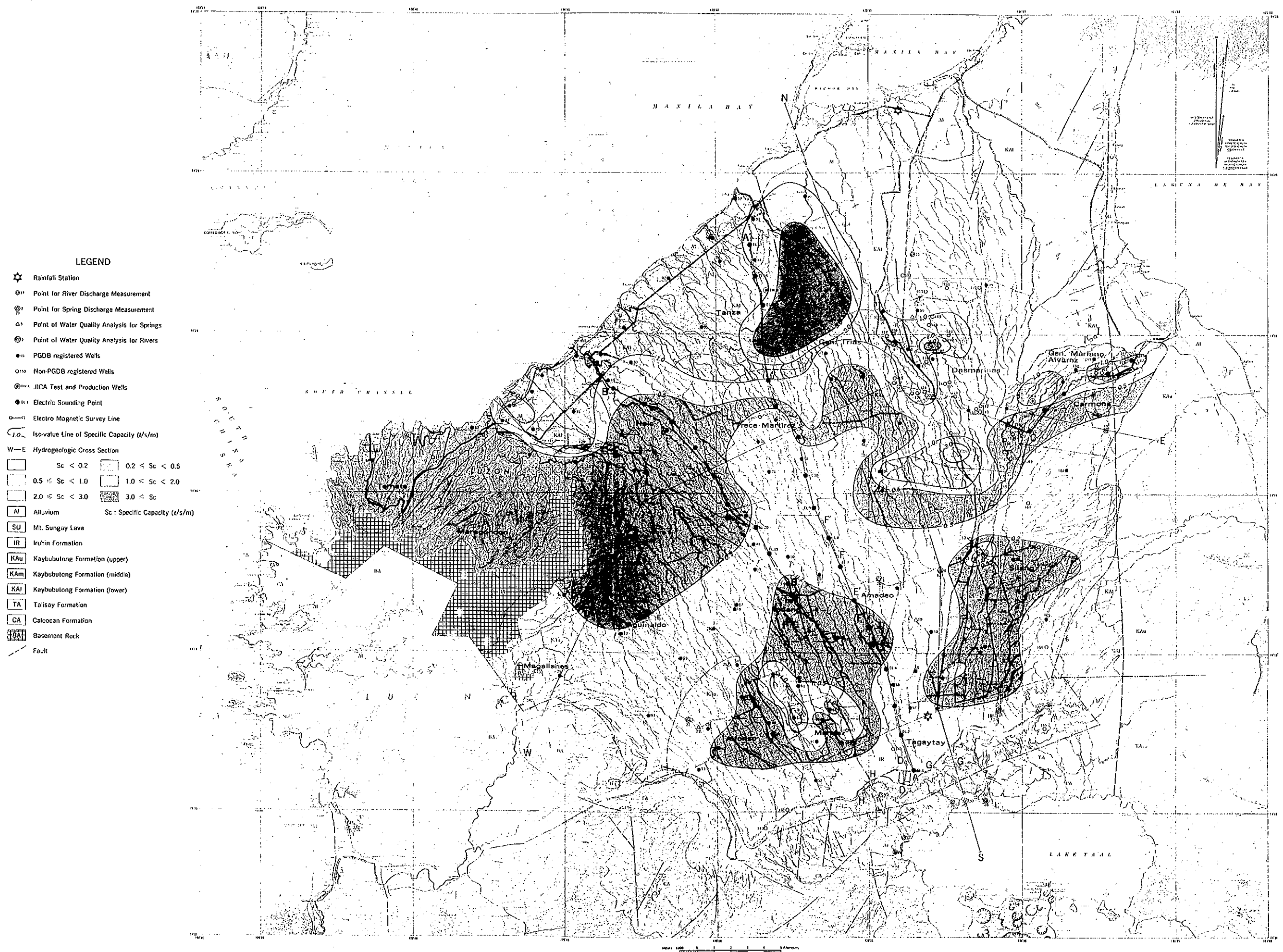
No.	LOCATION		WELL STRUCTURE							SWL GL-m	PUMPING TEST DATA						WELL LOGGING DATA				Stratigraphic Horizon												
	Site	Ground Level (m)	Borehole		Casing		Screen				STEP DRAWDOWN TEST			CONSTANT DISCHARGE TEST			Temp (°C)	Cond. (S/cm)	Main Aquifer	Resistivity (ohm-m)													
			Depth (m)	Ø (mm)	Depth (m)	Ø (mm)	Ø (mm)	TOP GL-m	BOTTOM GL-m		Qmax (l/s)	Drawd. (m)	Sc (l/s/m)	Q (l/s)	Drawd. (m)	Sc (l/s/m)						Sc (m³/s)											
A	GMA	163.65	100	350	100	200	150	106	154	99.45	1.88	1.59	1.18	1.74	1.62	1.07	34.00	26.5	370	Scoria tuff	30~50	Kaybubutong Formation (lower)											
								(Total 48 m)													~C.S.S (Scoria rich)		20~30										
			220	320	160	150																											
B	TANZA	30.65	60	400	60	250	200	66	84	10.45	34.10	19.43	1.76	14.54	9.06	1.60	35.95	29.5	450	(Upper) Conglomerate ~Gravely C.S.S (Lower) Gravely C.S.S ~Pumice tuff	20~60	Kaybubutong Formation (lower)											
			150	390	150	200	200	138	144																								
C	NAIC	5.40	60	400	60	250	200	66	72	6.28	36.00	33.20	1.08	33.08	36.38	1.05	22.20	30.2	510	MSS ~FSS (scoria rich)	30~50	Talisay Formation											
			150	370	150	200	200	102	114																								
D	MENDEZ	542.00	200	350	200	200	150	206	212	112.1	8.24	12.60	0.65	8.27	9.85	0.84	15.30	24.5	272	Scoria tuff ~ Volcanic conglomerate	100~180	Kaybubutong Formation (middle)											
			290	320	270	150	150	251	257																								

C.S.S.: Coarse Sandstone

TABLE 3.4-2 CLASSIFICATION OF THE AQUIFERS DISTINGUISHED IN THE STUDY AREA

	Stratigraphic Horizon	Lithology	Sc (1/s/m)	T (cm ² /s)	Temperature (°C)	Conductivity (μ s/cm)	Resistivity (ohm-m)	Test Well
Upper Aquifer	Middle horizon of Kaybubutong Formation	Scoria, Tuff, Volcanic Conglomerate lava	0.84	15.30	24.5	272	100~180	Mendez
Middle Aquifer	Lower horizon of Kaybubutong Formation	Conglomerate ~Coarse sand	1.60 1.007	36.0 34.0	29.5 26.5	450 370	20~60 20~40	Tanza G.M.A
Lower Aquifer	Talsay Formation	Coarse sandstone with gravel ~Medium Sandstone	1.05	22.2	30.2	510	30~50	Naic

Fig. 3.4-1 EVALUATION MAP OF GROUNDWATER RESOURCES



Specific capacity is as high as more than 0.5 l/s/m in the low elevation zone, the northern part of Cavite Slope, while it is as low as less than 0.5 l/s/m in the middle to high elevation zone in general.

In the areas along the National Road between Dasmarinas and Silang and around Mendez, specific capacity rarely exceeds more than 1.01 l/s/m.

The high specific capacity zone around Tanza roughly coincide with the valley on the surface plane of Talisay Formation. High specific capacity in this zone may reflect thick fan deposits (lower horizon of Kaybubutong Formation) supplied from the high elevation zone of Cavite Slope.

The wells tend to be deeper in the middle to high elevation zone with high specific capacity. For example, the wells deeper than 120 m in Mendez and 200 m in Silang show more than 1.0 l/s/m in specific capacity. This indicates that the aquifer consisted of coarse-grain sediments of Kaybubutong Formation which lies in the depth of the middle to high elevation zone.

CHAPTER 4

WATER BALANCE AND GROUNDWATER POTENTIAL

CHAPTER 4

WATER BALANCE AND GROUNDWATER POTENTIAL

Using the records of seven PAGASA rainfall stations in and near the Study Area, mean annual areal rainfall was determined at 2,505 mm.

Many springs are distributed in the Study Area, especially between 300 m and 400 m in elevation and their total discharge amounts to 580 lps which is equivalent to 4.6% of the annual rainfall.

Based on the results of river survey, annual total surface runoff was estimated as 1,082 mm in depth, of which about 70% is discharged in wet season and 30% is in dry season.

Actual evapotranspiration was estimated as 931 mm/yr from pan evaporation value obtained from the existing data and the modified Blaney-Criddle Method.

Using the results of the well inventory survey and the pumping discharge survey, the present annual groundwater withdrawals in the Study Area was estimated as 32.573×10^6 cum.

As for the whole Study Area, 37.2% of the annual rainfall is consumed as actual evapotranspiration, 43.2% as surface runoff, 15.7% as groundwater discharge and 3.9% is groundwater recharge. In addition, about 33% of the total groundwater is pumped up for human activities. Water balance was also computed for some major river basins.

4.1 STUDY METHODS

In its general form, the water balance equation may be represented by

$$P = Et + R + O + \Delta G + \Delta S + I + M$$

where P is precipitation, Et is evapotranspiration, R is surface runoff, O is the groundwater from aquifers, ΔG is the groundwater storage increment, ΔS is the soil moisture increment, I is the amount of water intercepted and M is the water recharge or depletion due to human activity, all values are in millimeters.

P , R , I and M were measured or surveyed in this Study as outlined below. Other variables were lumped or estimated empirically.

4.1.1 Rainfall Measurement

Though an automatic rain gage (STAGELOG 2D) was installed at Malinta Elementary School, Brgy. Malinta, Dasmariñas, it malfunctioned and was not used in the Study. Therefore, rainfall data observed during the Study period were obtained from seven (7) PAGASA stations in and near the Study Area.

4.1.2 River Survey

The river survey measured the dry season and wet season discharge, Ph, EC and temperature, and sampled the river for water quality analysis. A total of 30 measurement points were established in the six (6) major river basins in the Study Area: Maragondon, Labac, Canas, San Juan, Imus and Binan. Of which, at least five (5) were in the basin selected for groundwater simulation. The measurement points are plotted in Fig. 4.1-1.

4.1.3 Spring Survey

The spring discharge measurement aims to estimate quantitatively the groundwater outflow in the Study Area. As in the river survey, this estimate shall be used in making a detailed water balance analysis of the area.

Twenty (20) representative springs were selected and measured for discharge, temperature, pH and EC. Of which, 10 were sampled for water quality analysis. The measurement points are plotted in Fig. 4.1-2.

A questionnaire survey was also conducted to determine the utility condition of the springs.

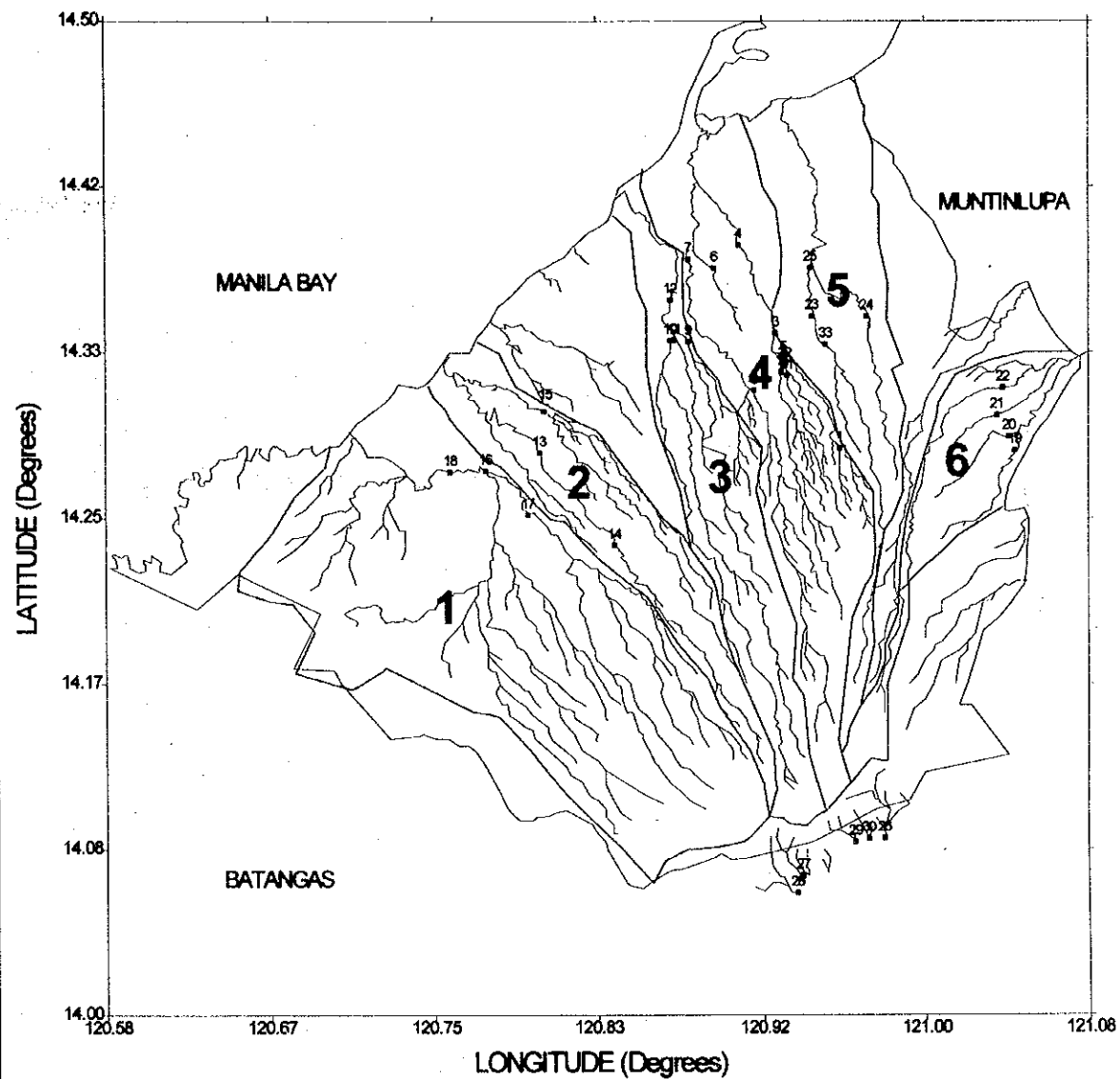
4.1.4 Well Inventory Survey

Well inventory provides the basic data to carry out the various investigations and surveys of this Study, e.g., selection of sites for test wells and core drilling, measurement and monitoring of groundwater levels and quality, and conduct of geological and geophysical surveys. It involves the collection of all available data on existing wells including: location, well design and construction, static water level, pumping test results, electric logs, lithologic logs and water quality. In this Study, the well inventory survey specifies collection and arrangement or transcription of data and information on existing 150 deep wells with sufficient hydrogeological information. This number represents 10% of the total number of deep wells believed to be drilled in the Study Area.

Through a UNDP grant, LWUA, together with other government agencies, had established the Philippine Groundwater Data Bank (PGDB). At present, about 639 wells located in the Study Area are registered in this PGDB. Table 4.1-1 details the availability of well data from the data bank.

The well inventory survey has retrieved 81 new deep well data which shall be supplemented to the data bank to form an updated well inventory database for Cavite. These new well data were arranged using PGDB's well inventory form and the use of which was confirmed and fixed during the survey.

Considering not only the sufficiency of hydrogeological data but also the distribution over the Study Area, 167 deep wells were selected from both PGDB-registered wells and the 81 non-PGDB-



MAJOR RIVER BASIN

- 1 MARAGONDON RIVER BASIN
- 2 LABAC RIVER BASIN
- 3 CAÑAS RIVER BASIN
- 4 SAN JUAN RIVER BASIN
- 5 IMUS RIVER BASIN
- 6 BINAN RIVER BASIN

LEGEND :

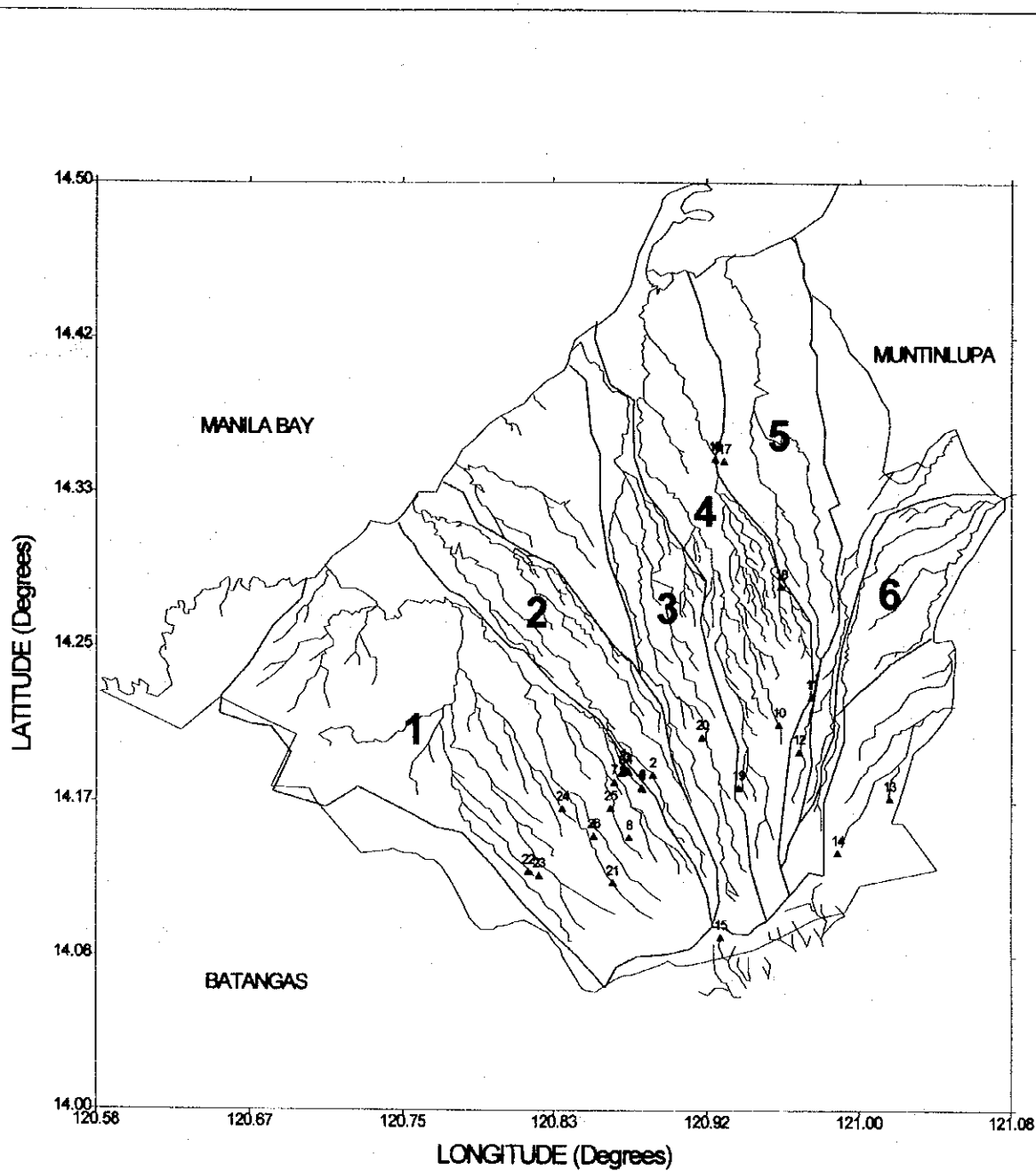
- RIVER DISCHARGE MEASUREMENT STATION
- DRAINAGE BOUNDARY

CAVITE WATER SUPPLY DEVELOPMENT STUDY

JAPAN INTERNATIONAL COOPERATION AGENCY

Fig. 4.1-1

LOCATIONS OF RIVER SURVEY STATIONS



MAJOR RIVER BASIN

- 1 MARAGONDON RIVER BASIN
- 2 LABAC RIVER BASIN
- 3 CANAS RIVER BASIN
- 4 SAN JUAN RIVER BASIN
- 5 IMUS RIVER BASIN
- 6 BINAN RIVER BASIN

LEGEND :

- ▲ SPRING NO.
- DRAINAGE BOUNDARY

CAVITE WATER SUPPLY DEVELOPMENT STUDY

Fig. 4.1-2

LOCATIONS OF SPRING SURVEY STATIONS

JAPAN INTERNATIONAL COOPERATION AGENCY

Table 4.1-1 EXISTING PGDB AND NON-PGDB WELL DATA

MUNICIPALITY	NUMBER OF WELLS		LOCATION DATA (1)		(1) + WELL CONSTRUCTION DATA		(1) + LITHOLOGIC LOG DATA		(1) + PUMPING TEST DATA		(1) + WATER QUALITY DATA		(1) + ELECTRIC LOG DATA	
	PGDB	NON-PGDB	PGDB	NON-PGDB	PGDB	NON-PGDB	PGDB	NON-PGDB	PGDB	NON-PGDB	PGDB	NON-PGDB	PGDB	NON-PGDB
1 DASMARIÑAS	67	36	49	27	6	24	22	16	22	12	9	7	23	
2 INDANG	65		43		9		32		23		5			
3 GMA	2	3	2	3	1	3	1	3	0	2	0		2	
4 MENDEZ	24	1	22	1	2	1	20	1	17	1	3			
5 SILANG	70	13	44	13	4	10	32	13	17	12	1	1	5	
6 TANZA	52	1	40	1	5	1	25	1	10		8		1	
7 TAGAYTAY CITY	13	5	12	4	0	3	3	2	1		1		1	
8 AMADEO	19		12		2		8		5		1			
9 MAGALLANES	14		4		0		3		2		0			
10 MARGACONDON	39	1	30	1	0	1	14	1	6	1	6	1	1	
11 TERNATE	23	2	11	2	0	2	8	2	3	1	0	2	1	
12 ALFONSO	51	1	44	1	3	1	41	1	19		5		1	
13 NAIC	54		41		2		30		6		5			
14 GEN. E. AGUINALDO	5		2		0		0		0		0			
15 CARMONA	78	12	77	6	1	4	44	6	12	2	5		4	
16 TRECE MARTIRES CITY	17	5	6	2	1	2	4	1	4	1	3	1	2	
17 GEN. TRIAS	46	1	33	1	4	1	19	1	8	1	3	1	1	
TOTAL	639	81	472	62	40	53	307	48	155	33	55	13		42

registered wells (or supplementary wells). Of these 167 wells, 20 wells were chosen for simultaneous groundwater leveling and water quality sampling.

4.1.5 Groundwater Level Monitoring

(1) Simultaneous groundwater leveling

The purposes of simultaneous groundwater leveling or measurements of static water levels are as follows: to prepare a groundwater contour map of the Study Area and analyze regional groundwater flow and to investigate the short-term and long-term changes in groundwater levels for groundwater modeling and water balance analysis. The fluctuation of groundwater level is the best indicator of groundwater storage development.

Fig. 4.1-3 shows the approximate locations of the 20 wells measured for static water levels. These active wells were selected from the 166 deep wells of the well inventory survey. Static water levels were measured at least 6 hours after the operation stops to allow the water level to recover. Just after the restart of the operation, water sample is taken from the blow-off pipe or tap of the same well.

(2) Continuous groundwater level monitoring

Five (5) automatic water level gages (STAGELOG 40) were installed in five (5) existing deep wells in the simulation area to measure water pressure in mBar. The groundwater level monitoring points are plotted in Fig. 5.2-1. The behavior of groundwater levels in the simulation area was monitored for six month, from June to November, 1994. After the measuring period, the logger was taken out from the well, connected to the PC to read the measured groundwater level data. Fig. 4.1-4 presents the groundwater level hydrographs measured at JICA monitoring stations.

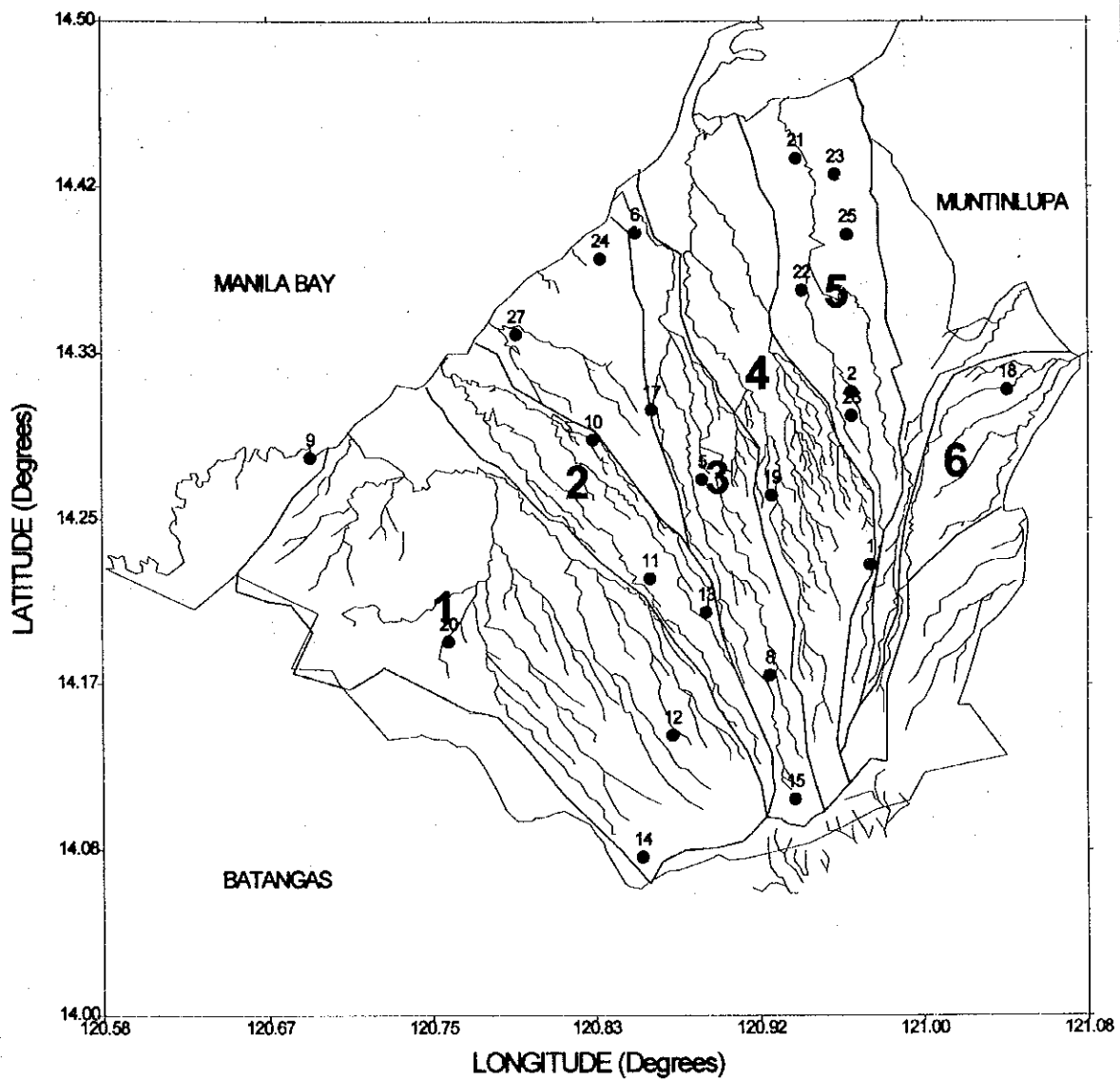
4.1.6 Pumping Discharge Survey

Pumping discharge survey was conducted to estimate the amount of groundwater use in the Study Area. Locations of the surveyed wells are shown in Fig.4.1-5.

4.2 WATER BALANCE IN THE STUDY AREA

4.2.1 Rainfall

Fig. 4.2-1 exhibits the isoyethal map of mean annual rainfall for the Study Area based on the records of seven (7) PAGASA rainfall stations. Annual rainfall varies from 2,000 mm to 3,800 mm with the southern part experiencing much rainfall. Mean annual areal rainfall was determined at 2,505 mm.



MAJOR RIVER BASIN

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LEGEND :

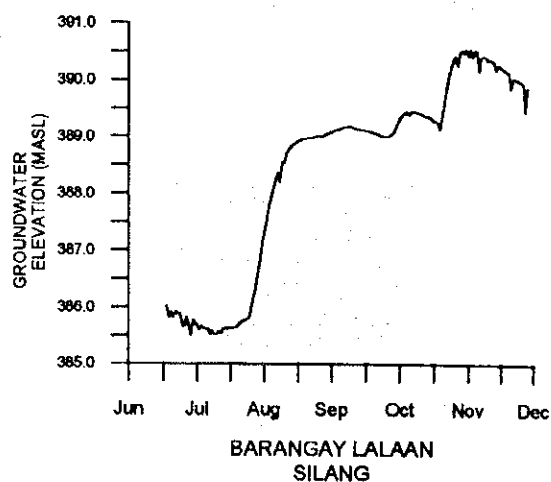
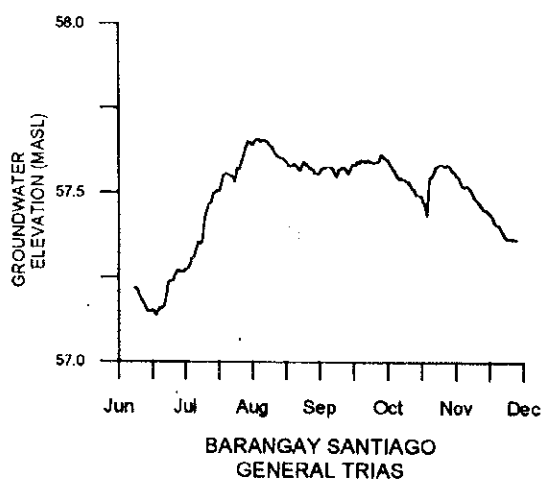
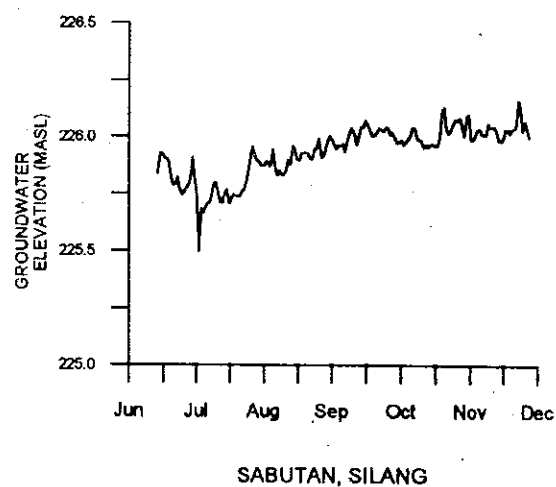
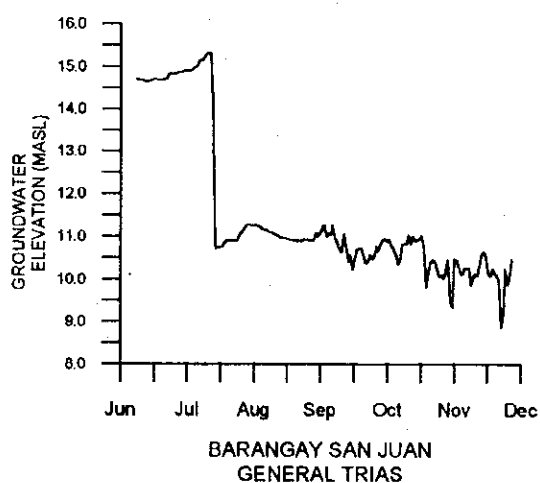
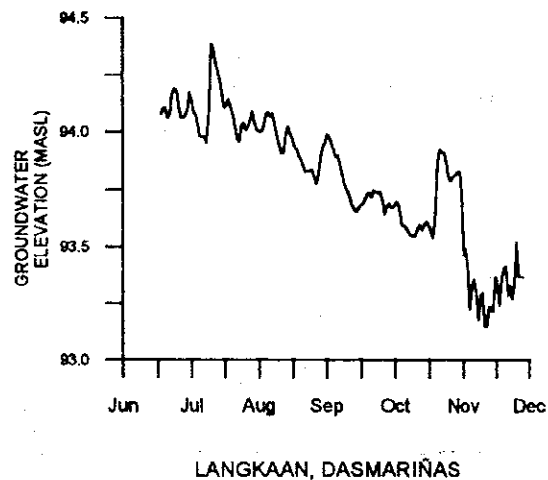
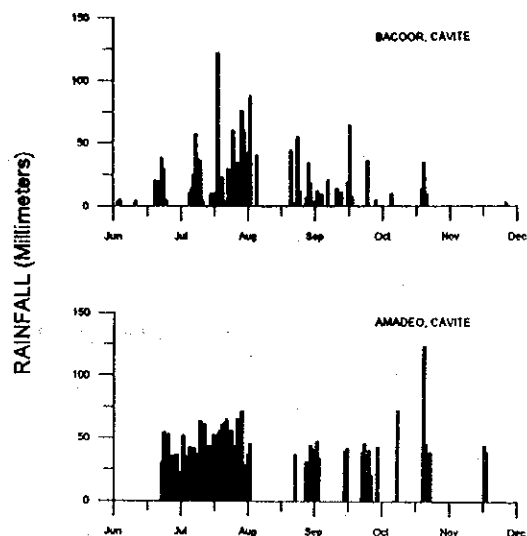
- WELL I.D. NO.
- DRAINAGE BOUNDARY

CAVITE WATER SUPPLY DEVELOPMENT STUDY

Fig. 4.1-3

JAPAN INTERNATIONAL COOPERATION AGENCY

LOCATIONS OF GROUNDWATER LEVEL
MONITORING STATIONS

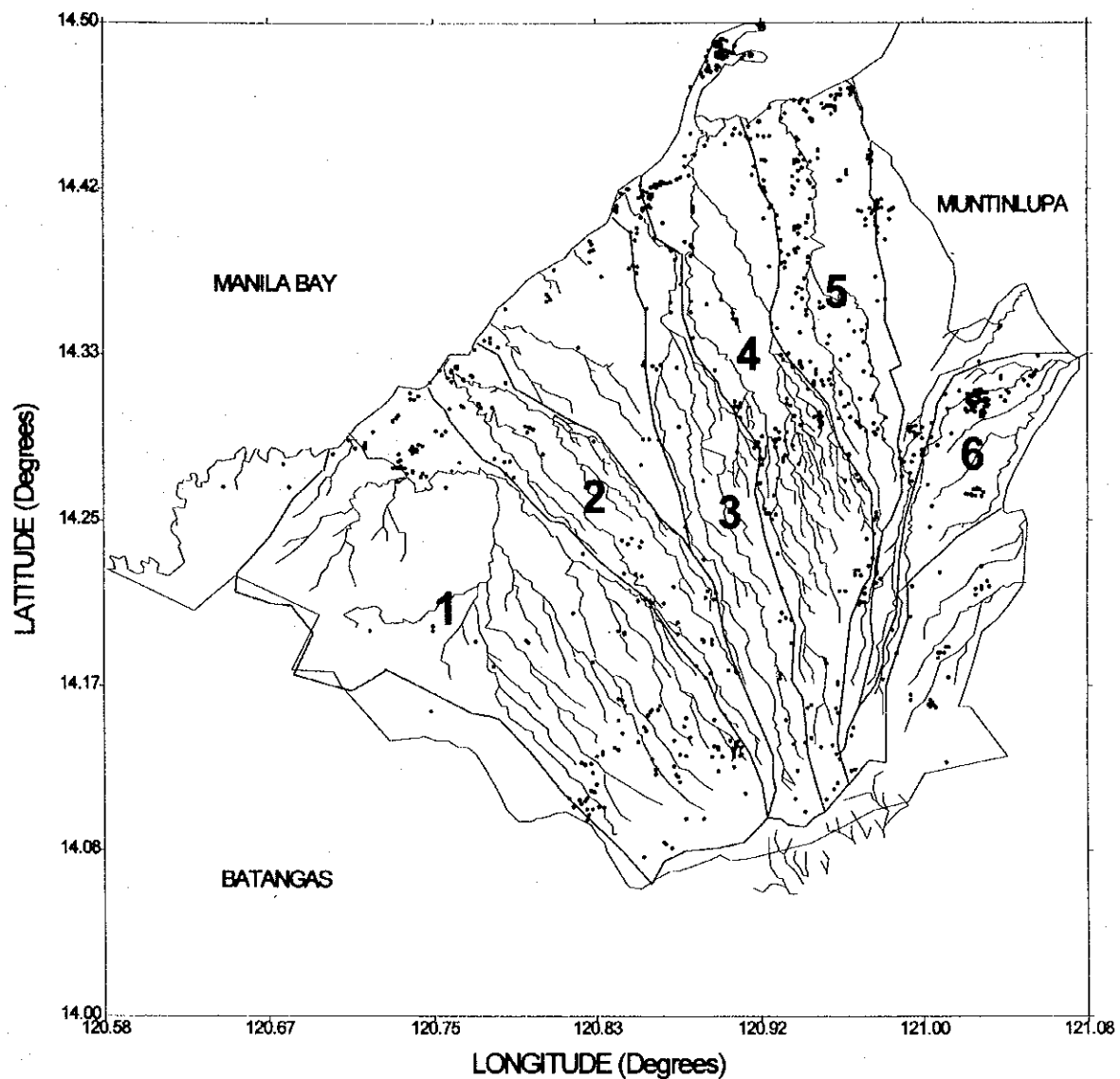


CAVITE WATER SUPPLY DEVELOPMENT STUDY

JAPAN INTERNATIONAL COOPERATION AGENCY

Fig. 4.1-4

**GROUNDWATER LEVEL HYDROGRAPHS
OF JICA MONITORING STATIONS**



MAJOR RIVER BASIN

- 1 MARAGONDON RIVER BASIN
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- 6 BINAN RIVER BASIN

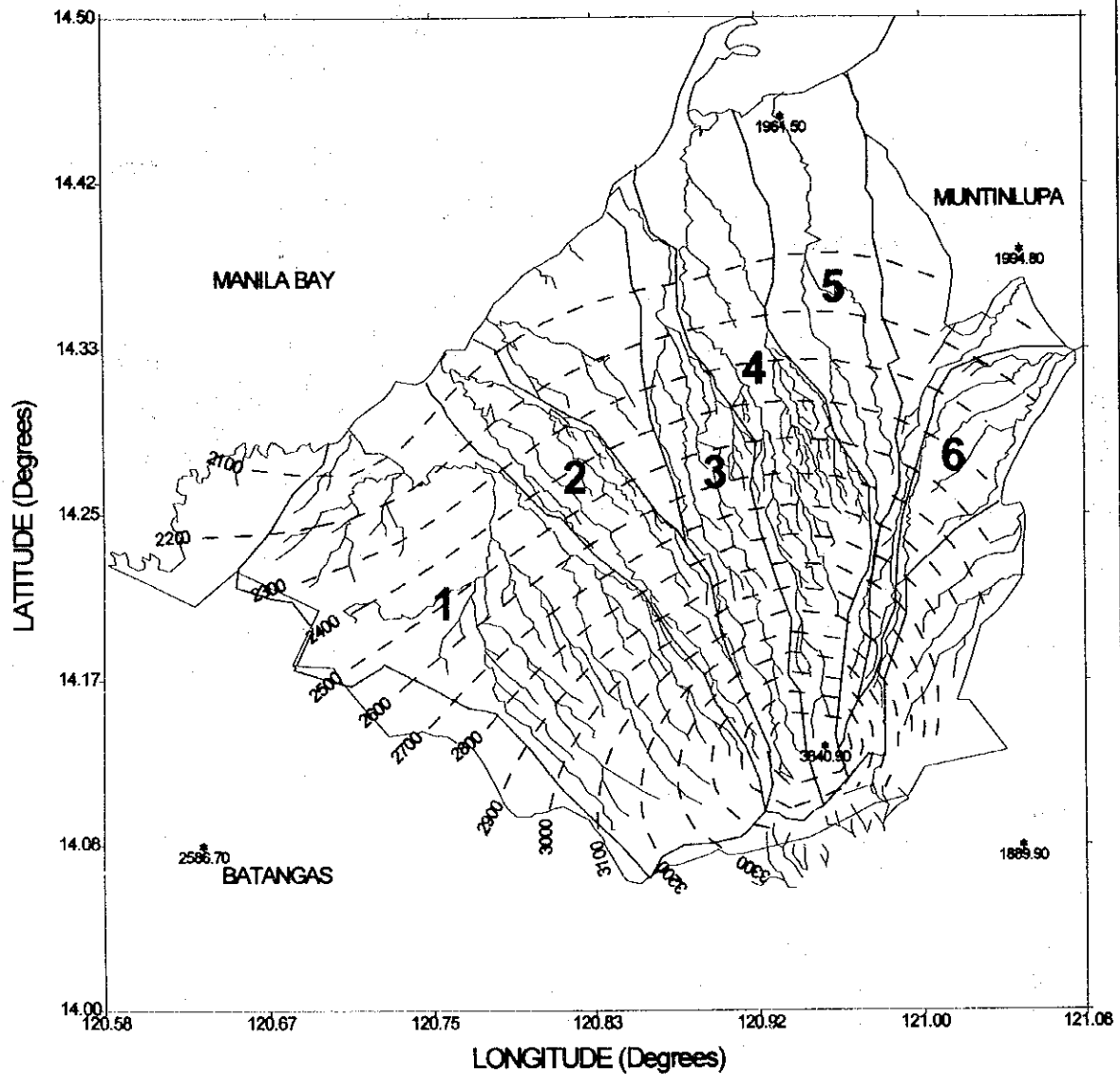
NOTE : EXCLUDES WELLS WITH-OUT COORDINATES

CAVITE WATER SUPPLY DEVELOPMENT STUDY

JAPAN INTERNATIONAL COOPERATION AGENCY

Fig. 4.1-5

LOCATIONS OF WELLS FOR
PUMPING DISCHARGE SURVEY



MAJOR RIVER BASIN

- 1 MARAGONDON RIVER BASIN
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- 3 CANAS RIVER BASIN
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- 5 IMUS RIVER BASIN
- 6 BINAN RIVER BASIN

LEGEND :

- 2500 - ANNUAL RAINFALL IN MILLIMETERS
- DRAINAGE BOUNDARY
- * PAGASA RAINFALL STATION

CAVITE WATER SUPPLY DEVELOPMENT STUDY

Fig. 4.2-1

JAPAN INTERNATIONAL COOPERATION AGENCY

ANNUAL RAINFALL DISTRIBUTION IN THE STUDY AREA

4.2.2 Runoff through Rivers

Many springs recharge rivers in the Study Area. Discharge amount and the utility condition of these springs obtained from the spring survey are summarized in Tables 4.2-1 and 4.2-2. Table 4.2-3 shows the results of the river discharge survey.

Fig. 4.2-2 illustrates the measurement results in the flow scheme for each river basin. In general, almost all rivers have higher discharge in wet season than in dry season, except for three (3) rivers namely, Ilang-Ilang River (R-2), Cañas River (R-10) and Balsahan River (R-13). This is due to the NIA diversions (for irrigation) located upstream of the measurement points. The dry season discharge is only 28% of the annual river discharge, while wet season discharge is 72%. The Electric Conductivity (EC) and pH values obtained for wet season are lower compared to the measured values during dry season.

Streamflow of the major rivers is being diverted as irrigation water for paddy fields by NIA. The water requirement for irrigation in the Study Area amounts to 185×10^6 cum per year, of which about 77% is required in the wet season. Diversion capacity, irrigated area and water requirement for each river are shown in the S/R.

As shown in Table 4.2-4, these data were needed to adjust the discharge records to obtain a better estimates of surface runoff for water balance analysis.

4.2.3 Evapotranspiration

In this Study, the pan evaporation is considered as the potential evapotranspiration and assumed to include 75 mm from residual wet season soil moisture.

The pan evaporation data from Ambulong, Tanauan, Batangas was considered for the Study Area. The data consist of records from 1959 to 1975. Monthly mean evaporation values range from 97 mm to 153 mm over a 16-year period. The observed pan evaporation at this station is 1,401 mm/yr.

In the FRAMEWORK PLAN SOUTHERN TAGALOG TAAL LAKE BASINS (NWRC,1983), evapotranspiration rates were also computed using the modified Blaney-Criddle Method. Temperature data of Ambulong as well as relative humidity, sunshine and wind conditions were considered. Evapotranspiration was estimated at 1,400 mm/yr.

4.2.4 Amount of Water Use

The annual amount of water use by type of user in the Study Area as a whole in 1993, which totals 173.427×10^6 cum, is summarized as:

Table 4.2-1 SPRING SURVEY STATIONS AND SURVEY RESULTS FOR DRY AND WET SEASONS

SPRING NO.	NAME OF SPRING	BASIN NO.	LOCATION	DISCHARGE (lps)		WATER QUALITY SAMPLING POINT
				DRY	WET	
S-1	Ikdoy	1	Kaytambog, Indang	181.2	209.0	Y
S-2	Niyog	2	Buna Cerca, Indang	6.6	11.7	Y
S-3	Ipie I	1	Kaytambog, Indang	57.8	70.0	N
S-4	Ipie II	1	Kaytambog, Indang	17.5	159.0	N
S-5	Ulo	1	Kayquit, Indang	15.3	21.5	N
S-6	Buho	1	Kaytambog, Indang	9.8	16.0	N
S-7	Pulo	1	Pulo, Indang	3.0	44.8	N
S-8	Siloy	1	Carasuchi, Indang	38.7	68.5	N
S-9	Silid-silirán	1	Kayquit, Indang	6.0	16.3	N
S-10	Lucsubin	4	Lucsubin, Silang	11.7	11.7	Y
S-11	Ilog ng Bayan	5	Ilog Bayan, Silang	3.0	6.8	Y
S-12	Sulsugin	5	Balete II, Silang	0.8	0.9	N
S-13	Pulong Usiw	7	Iruhin, Tagaytay City	0.6	0.6	N
S-14	Matang Tubig	7	Francisco, Tagaytay City	8.2	26.3	Y
S-15	Kaybubutong	7	Sambong, Tagaytay City	168.4	177.0	Y
S-16	Sabang I	4	Sabang, Dasmariñas	1.0	1.0	N
S-17	Sabang III	5	Sabang, Dasmariñas	0.2	0.8	N
S-18	Bucal	4	Bucal, Dasmariñas	3.5	5.0	Y
S-19	Bucal	4	Bucal, Amadeo	0.2	0.6	N
S-20	Balete	3	Halang, Amadeo	42.5	44.8	Y
S-21	Alibandog	1	Mangas I, Alfonso	2.2	4.0	Y
S-22	Sta. Theresa I	1	Sta. Theresa, Alfonso	11.4	18.7	Y
S-23	Sta. Theresa II	1	Sta. Theresa, Alfonso	0.2	1.6	N
S-24	Salvacion	1	Taywanak Ilaya, Alfonso	1.1	5.2	N
S-25	Taywanak II	1	Taywanak Ilaya, Alfonso	0.4	4.3	N
S-26	Taywanak I	1	Taywanak Ilaya, Alfonso	0.7	2.5	N
S-27	Kaybubutong II	7		72.6		N
S-28	Kaybagal	3		49.8		N

Table 4.2-2 UTILITY CONDITION OF THE MAJOR SPRINGS

SPRING NUMBER	NAME OF SPRING	LOCATION	ELEVATION ABOVE POBLACION, m	TYPE OF SPRING	USER	POPULATION SERVED	FACILITIES
S-1	IKLOY**	Kaytambog, Indang	27	Depression	Indang Water District	6,490 (1990)	Spring Box with 18" dia. pipe to Pumping Station, 2" dia. and 4" dia. transmission pipes
S-2	NIYOG**	Buna Cerca, Indang	60	Depression	Privately Owned Factory		Spring Box with 3/4" dia. intake pipe (2) 2" dia. and 6" dia. overflow pipes
S-3	IPIE I	Kaytambog, Indang	30	Depression	Indang Water District	590 (1990)	Spring Box with 4" dia. intake pipe
S-4	IPIE II	Kaytambog, Indang	24	Depression	Indang Water District	5,900 (1990)	Spring Box with 4" dia. supply pipe to Ikloy and 4" dia. overflow pipe
S-5	ULO	Kayquit, Indang	70	Contact	Brgy. Kayquit	3,670 (1990)	None
S-6	BUHO	Kaytambog, Indang	30	Depression	Brgy. Kaytambog	590 (1990)	Spring Box with 4" dia. overflow pipe
S-7	PULO	Pulo, Indang	80	Depression	Brgy. Pulo	470 (1990)	3" dia. and 4" dia. intake pipes
S-8	SILYO	Carasuchi, Indang	140	Depression	Brgy. Banaba	890 (1990)	Spring Box with 6" dia. supply pipe and (2) 6" dia. overflow pipes
S-9	SILID-SILIRAN	Kayquit, Indang	70	Depression	Brgy. Kayquit	3,670 (1990)	None
S-10	LUCSUHIN**	Lucsuhin, Silang	23	Depression	Silang Water District	23,746 (1994)	Pumping Station, overflow pipes
S-11	ILOG NG BAYAN**	Ilog Bayan, Silang	-7	Depression	Brgy. Ilog Bayan	1,922 (1994)	Spring Box with (3) 3" dia. overflow pipes
S-12	SULSUGIN	Balete II, Silang	105	Depression	Brgy. Balete II	None	None
S-13	PULONG USUI	Iruh-in, Tagaytay City	-122	Depression	Tagaytay City Water District		Pumping Station
S-14	MATANG TUBIG**	Francisco, Tagaytay Cit	-142	Depression	Tagaytay City Water District		Pumping Station
S-15	KAYBUBUTONG**	Sambong, Tagaytay City	-348	Impervious Rock	Tagaytay City Water District		Reservoir and Pumping Station
S-16	SABANG I	Sabang, Dasmariñas	-30	Artesian	Brgy. Sabang	2,174 (1994)	2" and 1" dia. overflow pipes
S-17	SABANG III	Sabang, Dasmariñas	-32	Artesian	Brgy. Sabang	2,174 (1994)	None
S-18	BUCAL**	Bucal, Dasmariñas	50	Depression	Dasmariñas Water District	705 (1994)	Spring Box with 6" dia. transmission pipe, (3) 4" dia. overflow pipes and (2) 4" dia. drain pipes
S-19	BUCAL	Bucal, Amadeo	-40	Depression	Brgy. Bucal	None	None
S-20	BALETE**	Halang, Amadeo	-90	Depression	Brgy. Halang	None	None
S-21	ALIBANDOG**	Mangas I, Alfonso	75	Depression	Poblacion		Spring Box with 6" dia. supply pipe and 3" dia. overflow pipe
S-22	STA. THERESA I**	Sta. Theresa, Alfonso	-30	Depression	Brgy. Sta. Theresa		Spring Box with 4" dia. supply pipe and 3" dia. overflow pipe
S-23	STA. THERESA II	Sta. Theresa, Alfonso	-20	Depression	Brgy. Sta. Theresa		None
S-24	SALVACION	Taywanak Ilaya, Alfonso	-65	Depression	Brgy. Taywanak		Spring box with 4" dia. intake pipe
S-25	TAYWANAK II	Taywanak Ilaya, Alfonso	-50	Depression	Brgy. Taywanak		None
S-26	TAYWANAK I	Taywanak Ilaya, Alfonso	-80	Depression	Brgy. Taywanak		Spring box with 3" dia. intake pipe
S-27	KAYBUBUTONG II						
S-28	KAYBASAL						

* VALUES TAKEN FROM PREVIOUS STUDIES

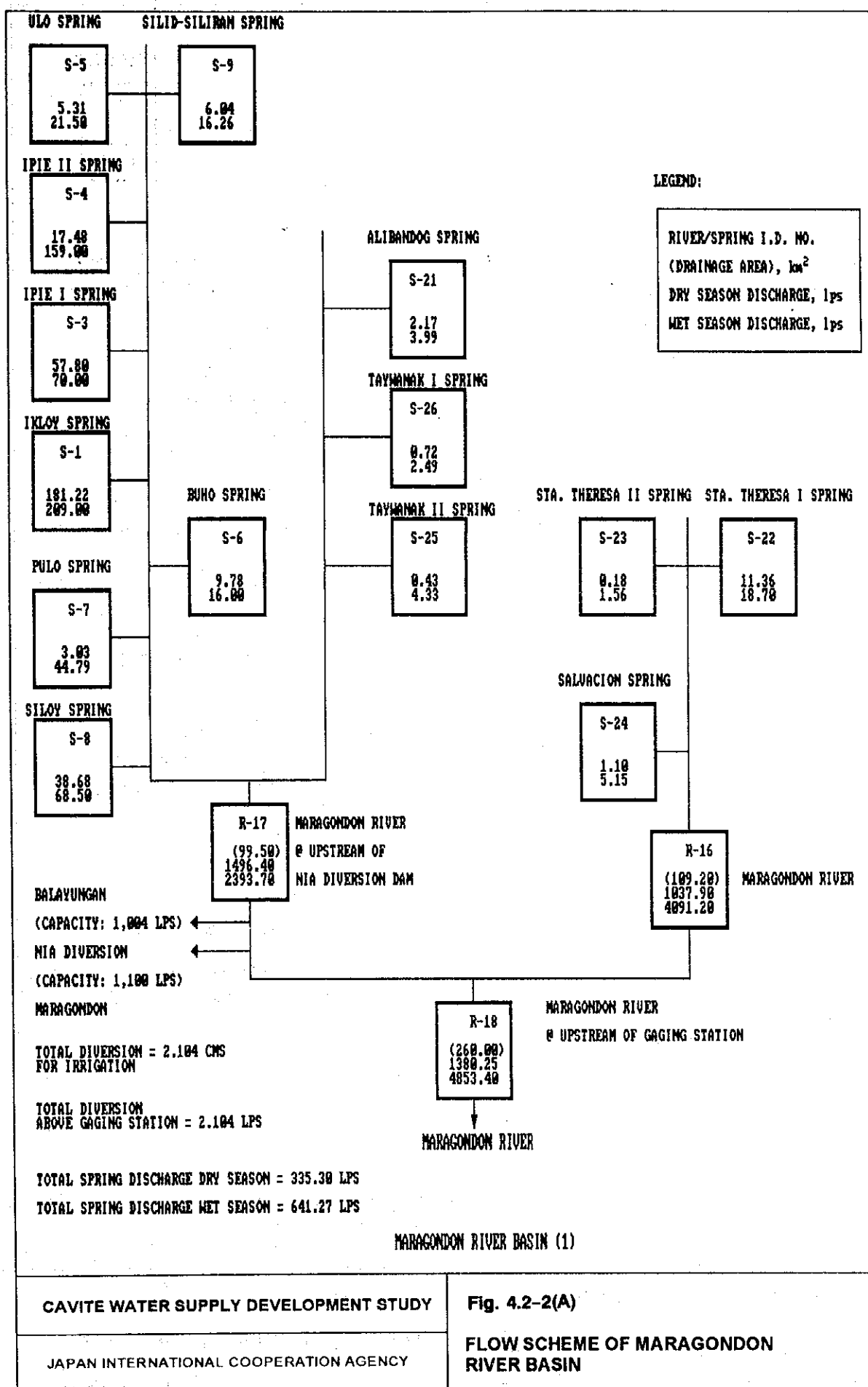
** SAMPLING POINTS FOR WATER QUALITY ANALYSIS

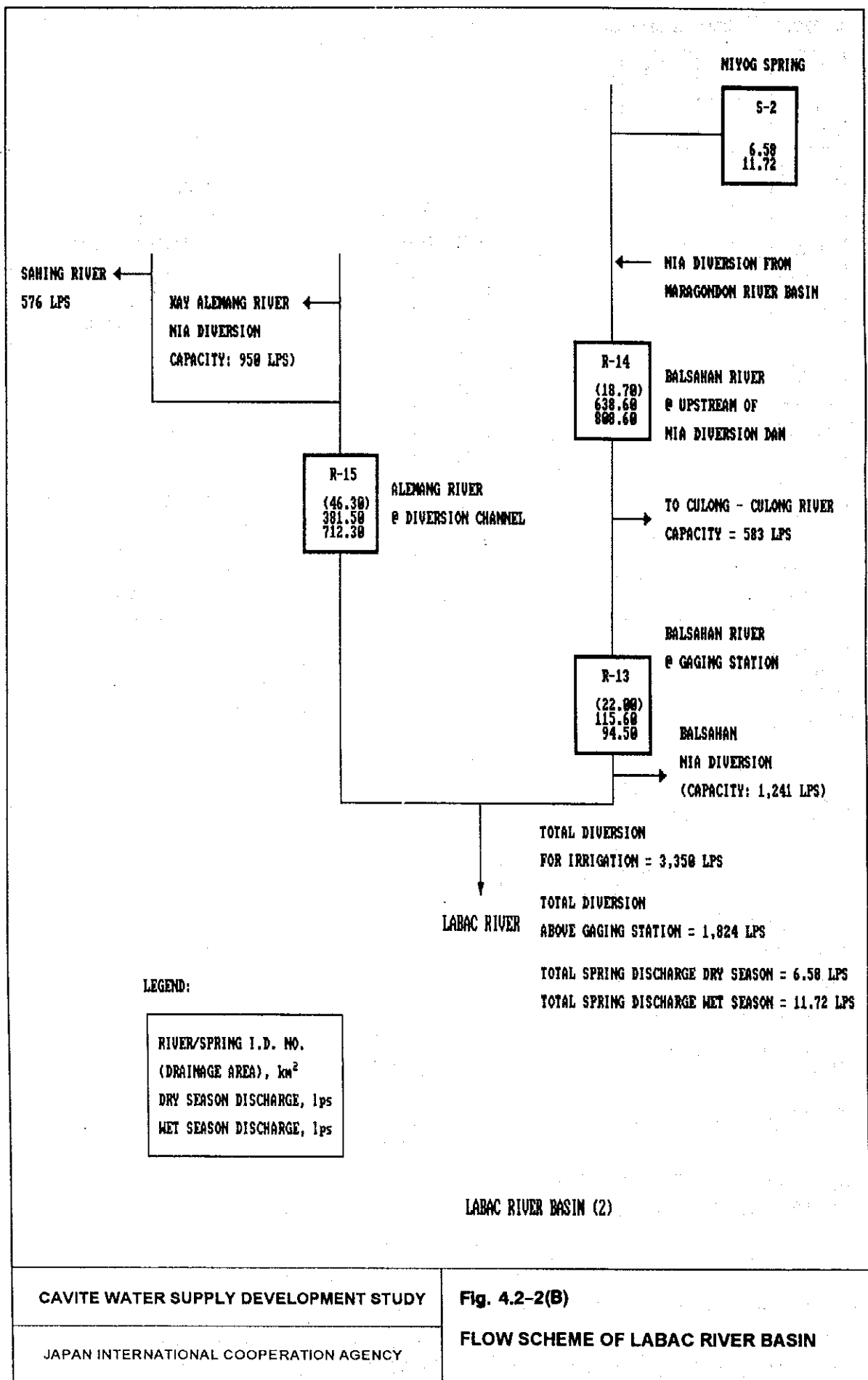
() MINOR SPRINGS

LATITUDE, LONGITUDE, ELEVATIONS AND DISTANCE WERE BASED FROM THE TOPOGRAPHIC MAP

Table 4.2-3 RIVER SURVEY STATIONS AND SURVEY RESULTS FOR DRY AND WET SEASONS

STA. NO.	RIVER NAME	BASIN NO.	LOCATION	DRAINAGE AREA (km ²)	DISCHARGE (lps)		WATER QUALITY SAMPLING POINT	REMARKS
					DRY	WET		
R-1	Dasmariñas	4	Bucal, Dasmariñas	9.4	123.7	269.9	N	Diversion Channel
R-2	Ilang-Ilang	4	San Francisco, General Trias	30.4	88.0	38.3	N	
R-3	Ilang-Ilang	4	San Jose, Dasmariñas	53.1	10.0	171.9	Y	
R-4	Ilang-Ilang	4	Alapan, Imus	60.0	35.6	69.7	N	
R-5	Dasmariñas	4	Fiesta Subd., San Jose, Dasmariñas	12.2	182.0	421.5	N	Kasundu Dam
R-6	Pasong Camachile	4	Pasong Camachile, General Trias	12.2	49.5	388.9	N	Pasong Camachile Dam
R-7	Rio Grande	4	Pinza, General Trias	50.4	157.4	557.9	N	Bayan Dam
R-8	Rio Grande	4	Buenavista, General Trias	38.9	544.3	969.6	Y	
R-9	Panaysayan	3	Pasong Cawayan II, General Trias	29.0	78.2	212.1	Y	BRS Gaging Station
R-10	Cañas	3	Pasong Cawayan II, General Trias	56.8	233.7	56.9	N	
R-11	Cañas	3	Pasong Cawayan II, General Trias	5.7	30.5	75.2	N	
R-12	Palububan	3	Pinagtipunan, General Trias	101.1	346.1	362.0	N	Placenta Dam
R-13	Balsahan	2	Palangui, Naic	22.0	115.6	94.5	N	BRS Gaging Station
R-14	Balsahan	2	Calumpang Lejos, Indang	18.7	638.6	808.6	Y	Culong-Culong Dam
R-15	Alemang	2	Sabang, Naic	46.3	381.5	712.3	N	Sahing Dam
R-16	Maragondon	1	Mabacao, Maragondon	109.2	1037.9	4091.2	Y	
R-17	Maragondon	1	Pantijan, Maragondon	99.5	1496.4	2393.7	N	
R-18	Maragondon	1	Bucal, Maragondon	260.0	1380.9	4853.4	Y	BRS Gaging Station
R-19	Biñan	6	Timbao, Carmona	12.7	12.8	137.5	N	
R-20	Biñan	6	Lantic, Carmona	9.5	13.0	31.2	Y	
R-21	Biñan	6	Lantic, Carmona	3.8	11.4	49.3	N	
R-22	Biñan	6	Mabuhay, Carmona	20.5	8.6	131.4	N	
R-23	Imus	5	Salitran III, Dasmariñas	32.3	67.1	407.6	Y	Salitran Dam
R-24	Baluctot	5	Salawag, Dasmariñas	9.6	16.0	402.8	N	
R-25	Imus	5	Anabu Segunda, Imus	50.7	50.7	810.4	N	
R-26	Unknown Creek	7	Calocan, Talisay	9.4	38.3	73.2	N	
R-27	Bubutong	7	Leviste, Laurel	2.1	76.0	78.6	N	Diversion Dam
R-28	Pata Creek	7	Pata, Laurel	5.9	16.6	26.2	N	
R-29	Benirayan Creek	7	Benirayan, Talisay	3.7	11.7	14.4	N	
R-30	Calocan	7	Sampaloc, Talisay	1.8	11.9	25.3	N	
R-31	Hasaan Creek	5	San Francisco, General Trias	9.1	0.0	504.5	N	Hasaan Dam
R-32	Ilang-Ilang	4	San Francisco, General Trias	9.3	0.0	212.7	N	
R-33	Imus	5	San Manuel, Dasmariñas	28.7	0.0	465.3	N	Justo Dam



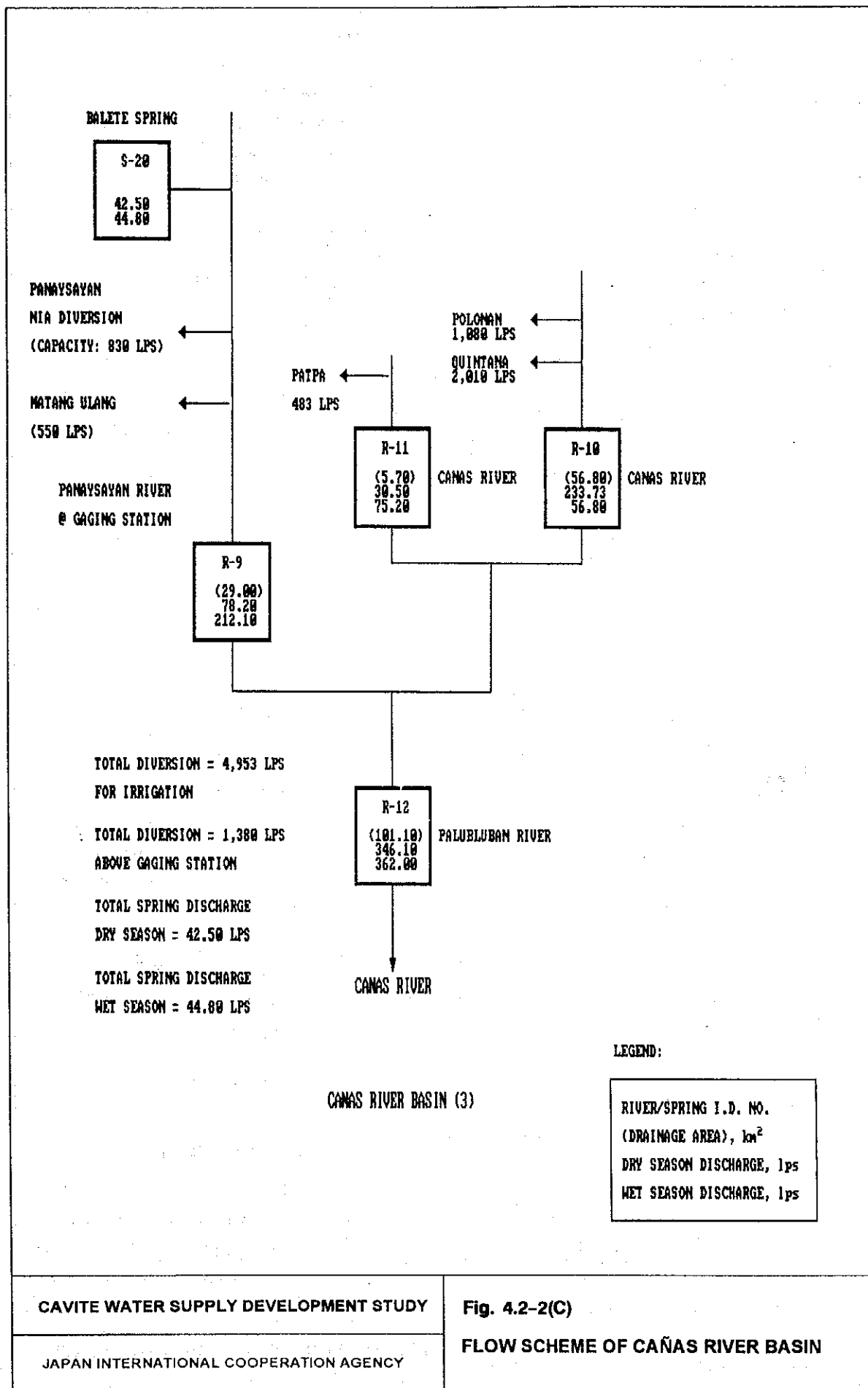


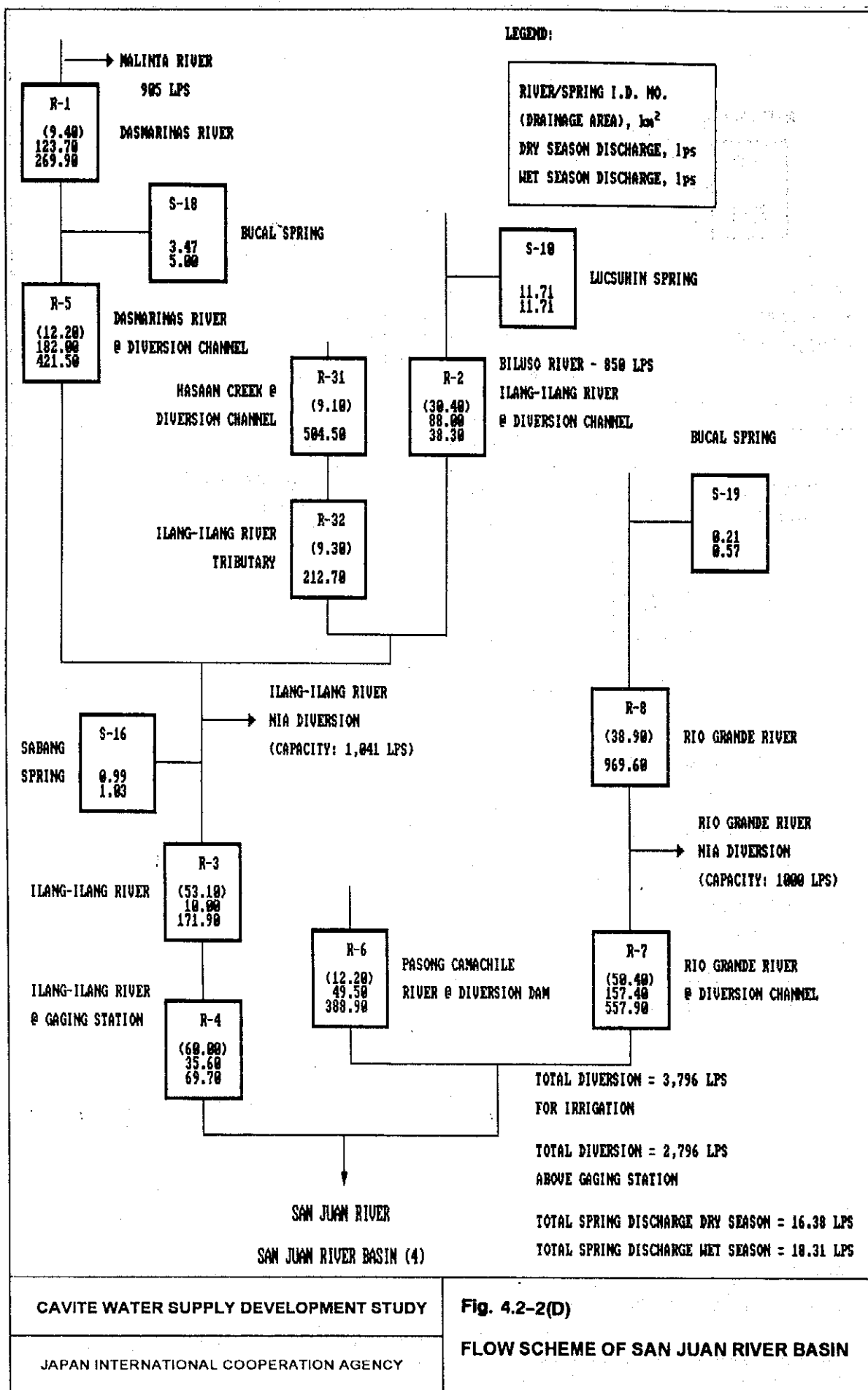
CAVITE WATER SUPPLY DEVELOPMENT STUDY

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Fig. 4.2-2(B)

FLOW SCHEME OF LABAC RIVER BASIN



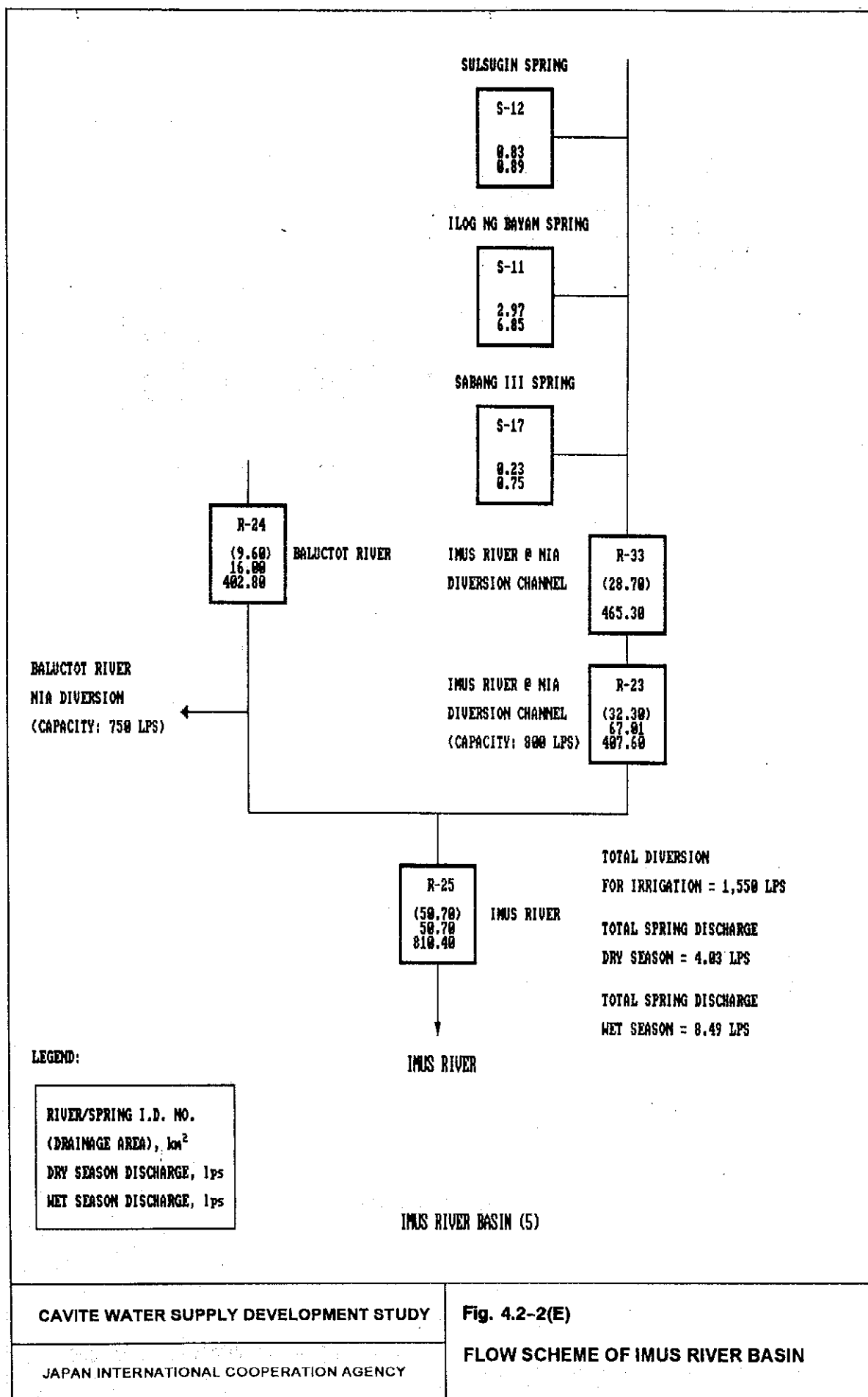


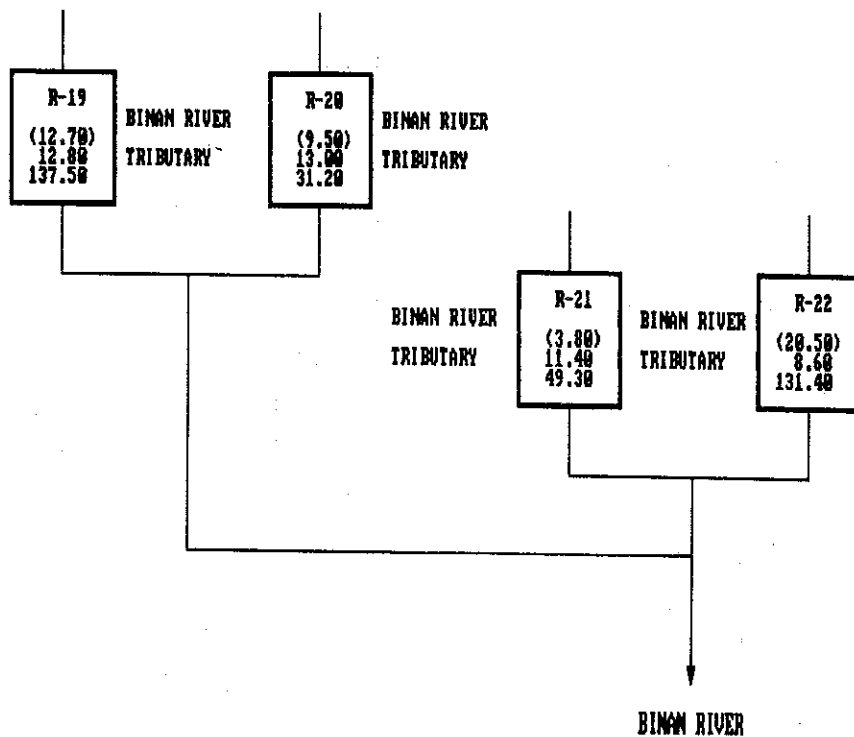
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Fig. 4.2-2(D)

FLOW SCHEME OF SAN JUAN RIVER BASIN





LEGEND:

RIVER/SPRING I.D. NO.
(DRAINAGE AREA), km ²
DRY SEASON DISCHARGE, lps
WET SEASON DISCHARGE, lps

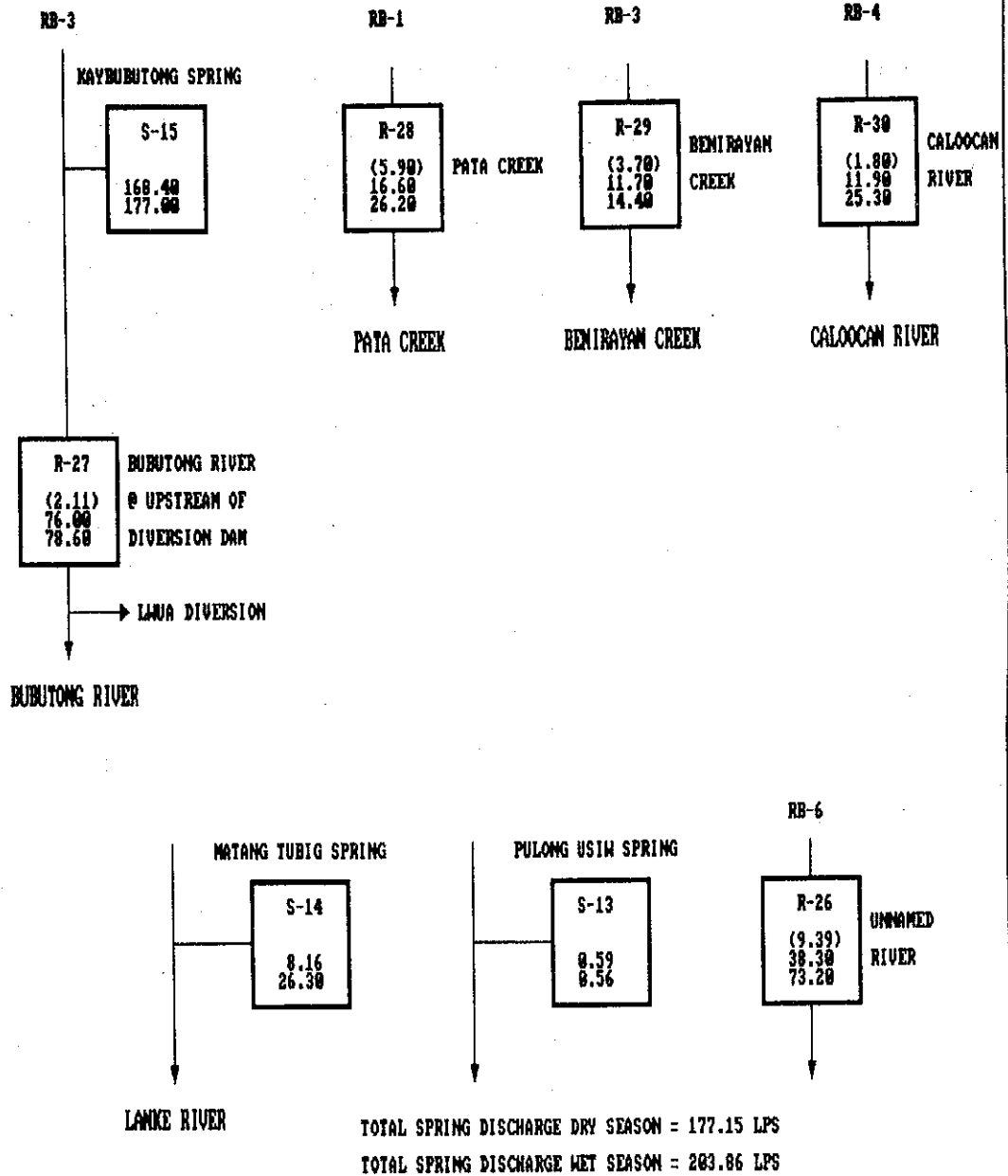
BINAN RIVER BASIN (6)

CAVITE WATER SUPPLY DEVELOPMENT STUDY

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Fig. 4.2-2(F)

FLOW SCHEME OF BINAN RIVER BASIN



CAVITE WATER SUPPLY DEVELOPMENT STUDY

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Fig. 4.2-2(G)

FLOW SCHEMES OF RIVER SYSTEMS BELOW
TAGAYTAY RIDGE ON BATANGAS SIDE

Table 4.2-4 SURFACE RUNOFF ESTIMATES BY RIVER BASIN

River Basin	MARAGONDON	LABAC	CAÑAS	SAN JUAN
River	Maragondon	Balsahan	Panaysayan	Ilang-Ilang
Drainage Area	260 km ²	22 km ²	29 km ²	60 km ²
	(mm)	(mm)	(mm)	(mm)
JAN	27.1	31.7	17.5	20.8
FEB	26.8	13.2	12.5	29.6
MAR	25.5	11.0	18.5	399.4
APR	22.3	8.2	17.9	166.6
MAY	98.1	67.0	24.9	24.9
JUN	149.2	205.0	139.4	492.6
JUL	237.3	438.3	215.2	53.4
AUG	474.2	303.1	234.6	44.1
SEP	283.5	453.6	241.3	23.3
OCT	180.8	174.1	106.2	16.6
NOV	141.1	180.3	37.5	17.2
DEC	87.9	63.3	24.0	11.9
ANNUAL, mm	1753.8	1948.8	1089.5	1300.4
MEAN, mm	146.15	162.40	90.79	108.37
Irrigation	79.8	0.00	470.00	0.00
Diversion, mm				
TOTAL RUNOFF, mm	1234.8	1571.10	1182.70	991.00
BASEFLOW, mm	598.8	377.7	376.8	309.4
SURFACE RUNOFF, mm	636.00	1193.40	805.90	681.60

Unit: $\times 10^6$ cum

Type of User	Annual amount
Domestic	18.089
Commercial	2.909
Institutional	0.436
Industrial	6.993
Agricultural (Irrigation)	145.000
Total	173.427

As for domestic use, the amount is calculated based on the estimated 1993-population and the unit consumption rate by level (I, II and III) of water supply facilities. Commercial (including cottage and small-scale industries) and institutional uses are estimated using the connection density ratio method. Industrial water use is estimated based on information gathered directly from interviews made in the industrial estates and in the local government units. Agricultural water use is calculated using the data supplied by NIA, i.e. the irrigable land area within the Study Area, and the pattern of water requirement shown in THE MASTER PLAN STUDY ON THE PROJECT CALABARZON (JICA, 1991).

Table 4.2-5 and Table 4.2.6 show the details of domestic and industrial water use respectively, and Table 4.2-7 summarizes the amount of water use by type of user in the Study Area and by municipality or city.

Using the results of the well inventory survey and the pumping discharge survey, the present annual groundwater withdrawals in the Study Area and in each river basin were estimated. Table 4.2-8 shows an annual total pumpage of 32.573×10^6 cum for the Study Area, and 53.724×10^6 cum for the water balance area. Among the six (6) river basins considered for water balance analysis, San Juan River Basin and Imus River Basin have the highest groundwater withdrawal of 12.221×10^6 cum and 11.657×10^6 cum, respectively. The total pumpage in the six (6) river basins is 31.415×10^6 cum within the Study Area and 52.464×10^6 cum within the water balance area.

4.3 GROUNDWATER LEVEL AND GROUNDWATER POTENTIAL

4.3.1 Estimation for the Whole Study Area

Table 4.3-1 arranges the results of simultaneous groundwater level measurements for dry and wet seasons. Figs. 4.3-1 and 4.3-2 show the groundwater level distributions in dry and wet seasons, respectively.

The groundwater level contours show steep to flat gradient reflecting the general topography of the Study Area. Groundwater gradient around or near Tagaytay Ridge is steeper, and it becomes flat towards the lower reach near Manila Bay.

Table 4.2-5 ESTIMATED DOMESTIC WATER USE BY WATER SUPPLY LEVEL BY MUNICIPALITY IN 1993

No.	City/Municipality	Total Demand	Level I		Level II		Level III		Doubtful Sources		Unit: 1000 cum
			Demand	%	Demand	%	Demand	%	Demand	%	
1	Dasmariñas	6,142	136	2.2	519	8.5	5,486	89.3	0	0.0	
2	Indang	1,038	19	1.8	618	59.6	401	38.6	0	0.0	
3	G. M. A.	1,921	32	1.7	954	49.7	935	48.7	0	0.0	
4	Mendez	465	70	15.1	62	13.2	334	71.7	0	0.0	
5	Silang	2,530	202	8.0	912	36.0	1,370	54.1	47	1.8	
6	Tanza	711	632	88.9	0	0.0	0	0.0	79	11.1	
7	Tagaytay City	911	8	0.8	27	2.9	856	93.9	21	2.3	
8	Amadeo	509	110	21.6	67	13.1	330	64.9	2	0.4	
9	Magallanes	212	48	22.7	142	67.1	0	0.0	22	10.2	
10	Maragondon	461	168	36.4	0	0.0	285	61.8	8	1.8	
11	Ternate	132	132	100.0	0	0.0	0	0.0	0	0.0	
12	Alfonso	853	59	6.9	214	25.0	578	67.8	2	0.3	
13	Naic	669	539	80.5	50	7.4	81	12.1	0	0.0	
14	Gen. Aguinaldo	231	48	21.0	79	34.4	100	43.4	3	1.2	
15	Carmona	329	327	99.2	0	0.0	0	0.0	3	0.8	
16	Trece Martirez City	348	118	34.0	61	17.7	168	48.4	0	0.0	
17	Gen. Trias	625	587	93.9	0	0.0	38	6.1	0	0.0	
Study Area Total		18,089	3,235	17.9	3,705	20.5	10,962	60.6	187	1.0	

Source: JICA Study Team

Note: Unit water consumption; 30 lpcd for Level I and doubtful sources, 60 lpcd for Level II and 110 lpcd for Level III

Table 4.2-6 INDUSTRIAL WATER USE IN 1993

No.	Name of Industrial Estate	Location	Number of Industries	Estimated Water Demand x1000cum/year
1	Gateway Business Park	Gen. Trias	2	180
2	First Cavite Industrial Estate	Dasmariñas	21	364
3	Dasmariñas Bagong Bayan-NHA Industrial Estate	Dasmariñas	4	182
4	G. M. A.-NHA Industrial Estate	G. M. A.	3	59
5	Cavite-Carmona Industrial Estate (PTC)	Carmona	46	432
6	New Cavite Industrial Center	Gen. Trias	5	17
7	First Cityland Heavy Industrial Center	Dasmariñas	3	7
8	Bulihan-NHA Industrial Estate	Silang	4	55
9	Cavite Export Processing Zone (CEPZ)	Gen. Trias/Rosario	6 *1)	55
10	Others 2)		97	5,644
Study Area Total			191	6,993

Source:

JICA Study Team; Basic data come from Environment Impact Assessment Report (Draft Final) for the Cavite Water Supply Development Study in November 1994 and interviews to the industrial estates and municipalities.

Notes:

- 1) Industries located in Gen. Trias Municipality
- 2) Major industries which are located outside the industrial estates shown above in the Study Area

Table 4.2-7 ESTIMATED WATER USE IN THE STUDY AREA BY MUNICIPALITY IN 1993

No.	City/Municipality	Population (Person)	Domestic		Commercial *1)		Institutional		Industrial *2)		Total
			Consumption	(%)	Consumption	(%)	Consumption	(%)	Consumption	(%)	
1	Dasmariñas	182,550	6,142	64.9	666	7.0	100	1.1	2,556	27.0	9,464
2	Indang	42,204	1,038	85.4	154	12.7	23	1.9		0.0	1,216
3	G. M. A.	73,703	1,921	51.5	269	7.2	40	1.1	1,499	40.2	3,730
4	Mendez	18,518	465	85.7	68	12.4	10	1.9		0.0	543
5	Silang	103,990	2,530	80.4	380	12.1	57	1.8	181	5.8	3,148
6	Tanza	68,542	711	56.0	250	19.7	38	3.0	271	21.4	1,269
7	Tagaytay City	26,565	911	80.2	97	8.5	15	1.3	113	10.0	1,136
8	Amadeo	22,718	509	84.2	83	13.7	12	2.1		0.0	605
9	Magallanes	13,569	212	70.0	50	16.4	7	2.5	34	11.2	303
10	Maragondon	24,489	461	81.8	89	15.9	13	2.4		0.0	564
11	Terate	12,748	132	71.2	47	25.1	7	3.8		0.0	186
12	Alfonso	31,435	853	77.3	115	10.4	17	1.6	118	10.7	1,103
13	Naic	56,482	669	72.6	206	22.4	31	3.4	16	1.7	922
14	Gen. Aguinaldo	11,407	231	82.8	42	14.9	6	2.2		0.0	279
15	Carmona	31,765	329	26.2	116	9.2	17	1.4	794	63.2	1,257
16	Trece Martirez City	18,799	348	26.6	69	5.2	10	0.8	881	67.3	1,308
17	Gen. Trias	57,624	625	44.7	210	15.0	32	2.3	530	38.0	1,398
Study Area Total		797,108	18,089	63.6	2,909	10.2	436	1.5	6,993	24.6	28,428

Unit: 1000 cum

Source: JICA Study Team

Notes:

1) Including small industries based on LWUA methodology manual

2) Large water users

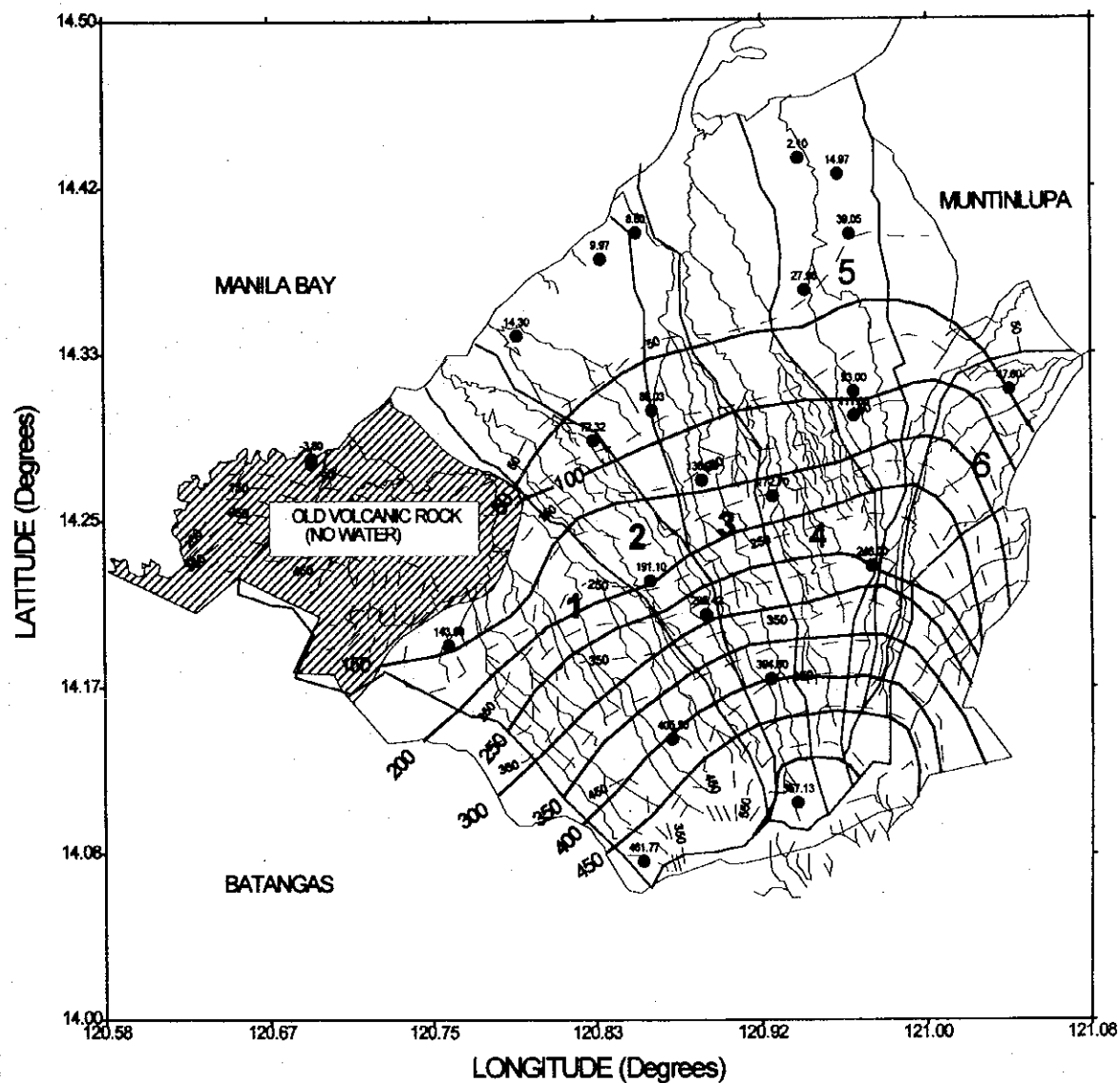
Table 4.2-8 PRESENT GROUNDWATER PUMPAGE (MCM/YEAR) BY TYPE OF USE AND BY RIVER BASIN

GROUNDWATER USE	WITHIN STUDY AREA	WITHIN WATER BALANCE AREA
Domestic	18.09	32.298
Institutional	0.436	0.436
Commercial	2.909	3.345
Agricultural	4.144	4.896
Industrial	6.993	12.748
TOTAL	32.573	53.724

RIVER BASIN	WITHIN STUDY AREA	WITHIN WATER BALANCE AREA
Maragondon	1.155	1.155
Labac	0.32	0.32
Canas	0.782	4.574
San Juan	12.221	20.153
Imus	11.057	20.383
Binan	5.879	5.879
TOTAL	31.415	52.464

Table 4.3-1 RESULTS OF SIMULTANEOUS GROUNDWATER LEVEL MEASUREMENT FOR DRY AND WET SEASONS

I.D. NO.	WELL INVENTORY NO.	LATITUDE	LONGITUDE	RIVER BASIN NO.	LOCATION	GROUND ELEV. (masl)	WELL DEPTH (m)	GW_DEPTH (mbgs)		GW_ELEV. (masl)		WATER QUALITY SAMPLING POINT
								DRY	WET	DRY	WET	
01	SIL-130	14 13 37	120 58 19	4	Poblacion, Silang PS #1	317	188.00	71.00	72.68	246.00	244.32	Y
02		14 18 49	120 57 47	5	San Luis, Dasmariñas PS#6	130		36.20	36.45	93.00	93.55	Y
03	GEN-162	14 17 30	120 00 16	6	Teacher's Village, GMA #1	180	213.41					Y
04	MEN-142	14 07 32	120 54 08	1	Payapa Mendez	522	243.90					Y
05	TRE-140	14 16 12	120 53 13	3	Trece Martirez	168	242.00	29.80	23.60	138.20	144.40	Y
06	TAN-94	14 23 39	120 51 13	3	Poblacion, Tanza	13	115.00	4.40	5.13	8.60	7.87	Y
07		14 16 32	120 44 03	1	Poblacion, Maragondon	20	33.84					Y
08		14 10 14	120 55 14	3	Poblacion, Amadeo	452	73.17	57.40	45.23	394.60	406.77	Y
09		14 16 53	120 41 17	1	Camandag, Puerto Azul	20	92.00	23.80	22.70	-3.80	-2.70	Y
10	TAN-139	14 17 23	120 49 55	7	Tanauan, Tanza	90	122.00	17.68	19.40	72.32	70.60	Y
11		14 13 12	120 51 38	2	Agos-os, Indang	223	76.22	31.90	31.25	191.10	191.75	Y
12		14 08 26	120 52 18	1	Guyam na Murti, Indang	447	91.46	41.50	40.15	405.50	406.85	Y
13		14 12 12	120 53 23	2	Alulod, Indang	318	33.54	22.58	20.43	295.42	297.57	Y
14	ALF-167	14 04 48	120 51 25	1	Esperanza, Alfonso	563	240.00	101.23	93.50	461.77	469.50	Y
15	TAG-143	14 06 34	120 56 01	4	Tagaytay City	652	213.41	67.87	65.00	557.13	560.00	Y
16		14 09 03	120 59 14	7	Ulat, Silang	510	157.00					Y
17		14 18 19	120 51 45	7	De Ocampo, Trece Martirez City	105	121.95	18.97	17.40	86.03	87.60	Y
18	CAR-160	14 18 53	121 02 31	6	Mla. Southwoods, Carmona	60	250.00	12.40	9.32	47.60	50.68	Y
19	TRI-10	14 15 45	120 55 21	4	Javalera, Gen. Trias	210	200.00	37.30	36.30	172.70	173.70	Y
20		14 11 17	120 45 27	1	Poblacion, Magallanes	172	51.83	28.10	25.20	143.90	146.80	Y
21		14 25 52	120 56 07	5	Poblacion, Imus	8	121.95	5.90	3.83	2.10	4.17	N
22		14 21 55	120 56 17	5	Anabu II, Imus	38	121.95	10.02	10.38	27.98	27.62	N
23		14 25 24	120 57 18	5	Vista Verde, Bacoor	20		5.03	4.93	14.97	15.07	N
24		14 22 53	120 50 09	7	Amaya, Tanza	15		4.28	8.33	10.72	6.67	N
25		14 23 55	120 57 41	5	Molino, Bacoor	55	182.93	15.95	12.95	39.05	42.05	N
26		14 18 06	120 57 48	5	St. Charbel, Dasmariñas	157	152.00	46.00	46.93	111.00	110.07	N
27		14 20 34	120 47 34	7	Tinalan, Naic	20		5.70	5.10	14.30	14.90	N
28		14 07 34	120 59 32	7	Sungay, Tagaytay City	578	140.24					N
29		14 18 49	121 01 08	6	Area-D, GMA	100	120.00					N
30		14 08 07	121 00 05	7	Iruhin, Tagaytay City	620						N

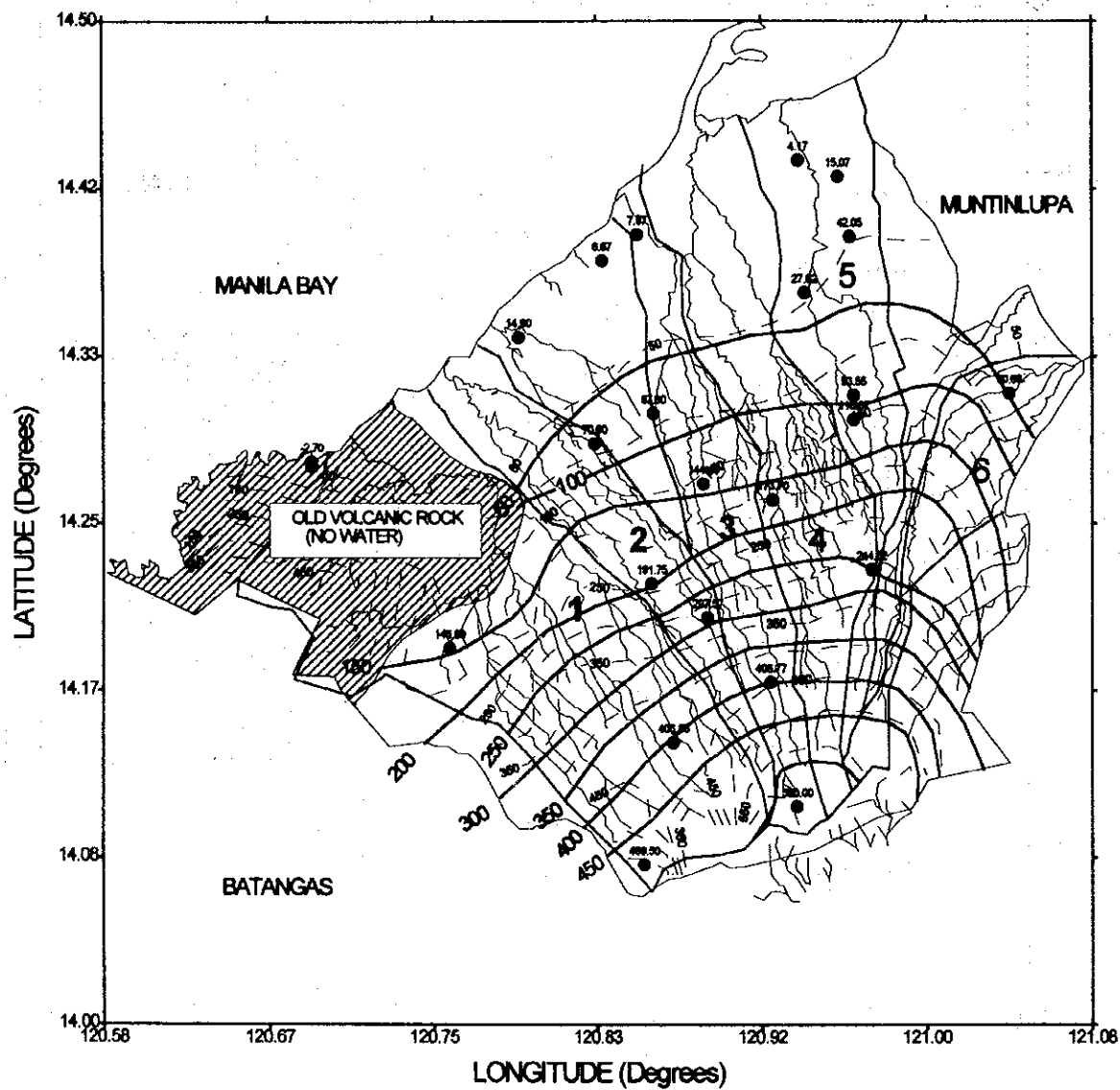


CAVITE WATER SUPPLY DEVELOPMENT STUDY

JAPAN INTERNATIONAL COOPERATION AGENCY

Fig. 4.3-1

DRY SEASON GROUNDWATER LEVEL



MAJOR RIVER BASIN

- 1 MARAGONDON RIVER BASIN
- 2 LABAC RIVER BASIN
- 3 CAÑAS RIVER BASIN
- 4 SAN JUAN RIVER BASIN
- 5 IMUS RIVER BASIN
- 6 BINAN RIVER BASIN

LEGEND:

- 100 -- TOPOGRAPHIC ELEVATION (Msl)
- 100 — GROUNDWATER LEVEL ELEVATION (Msl)
- /// OLD VOLCANIC ROCK

CAVITE WATER SUPPLY DEVELOPMENT STUDY

JAPAN INTERNATIONAL COOPERATION AGENCY

Fig. 4.3-2

WET SEASON GROUNDWATER LEVEL

The measured static water level varies from more than 450 above MSL as in Tagaytay, Mendez and Alfonso to less than 50 m above MSL as in Tanza, Maragondon, Ternate and Carmona. Lowest measured static water level was 3.8 m below MSL in Ternate.

In general, groundwater moves in the northeast direction from Tagaytay Ridge towards Laguna Lake and north, northwest direction towards Manila Bay.

In Fig. 4.3-3 the year-1994 fluctuation of groundwater levels between wet and dry seasons indicates a rise of about 7.5 meters in the south sector in Alfonso. In Silang, Dasmariñas, Tanza and Imus, groundwater level was continuously declining due to pumping.

The average annual water balance for the whole Study Area is presented schematically in Fig. 4.3-4(a). The water balance analysis involves the calculation of direct recharge or infiltration from rainfall. As discussed in 4.2, the annual rainfall depth over the Study Area is 2,504.9 mm; annual surface runoff depth is 1,475.3 mm; actual evapotranspiration is 931.0 mm, and the groundwater discharge is 393.2 mm. Hence, the annual recharge is estimated at 98.6 mm or 3.9% of the annual rainfall.

It could also be noted that the surveyed spring discharges shown in Table 4.2-1, which total 579.9 lps, seem to indicate a high recharge rate upstream of these springs. Converting this discharge into depth using 158.7 km², which is the computed area above the 400-meter contour line, gives 115.2 mm. It means that 115.2 mm or 4.6% of the annual rainfall is given off by these springs. However, this amount is part of the groundwater discharge mentioned above.

Comparing the present annual groundwater use of 32.6×10^6 cum as presented in Table 4.2-8 with the estimated groundwater recharge of 113.8×10^6 cum, the tentative estimate of the groundwater surplus of the Study Area is placed at 81.2×10^6 cum.

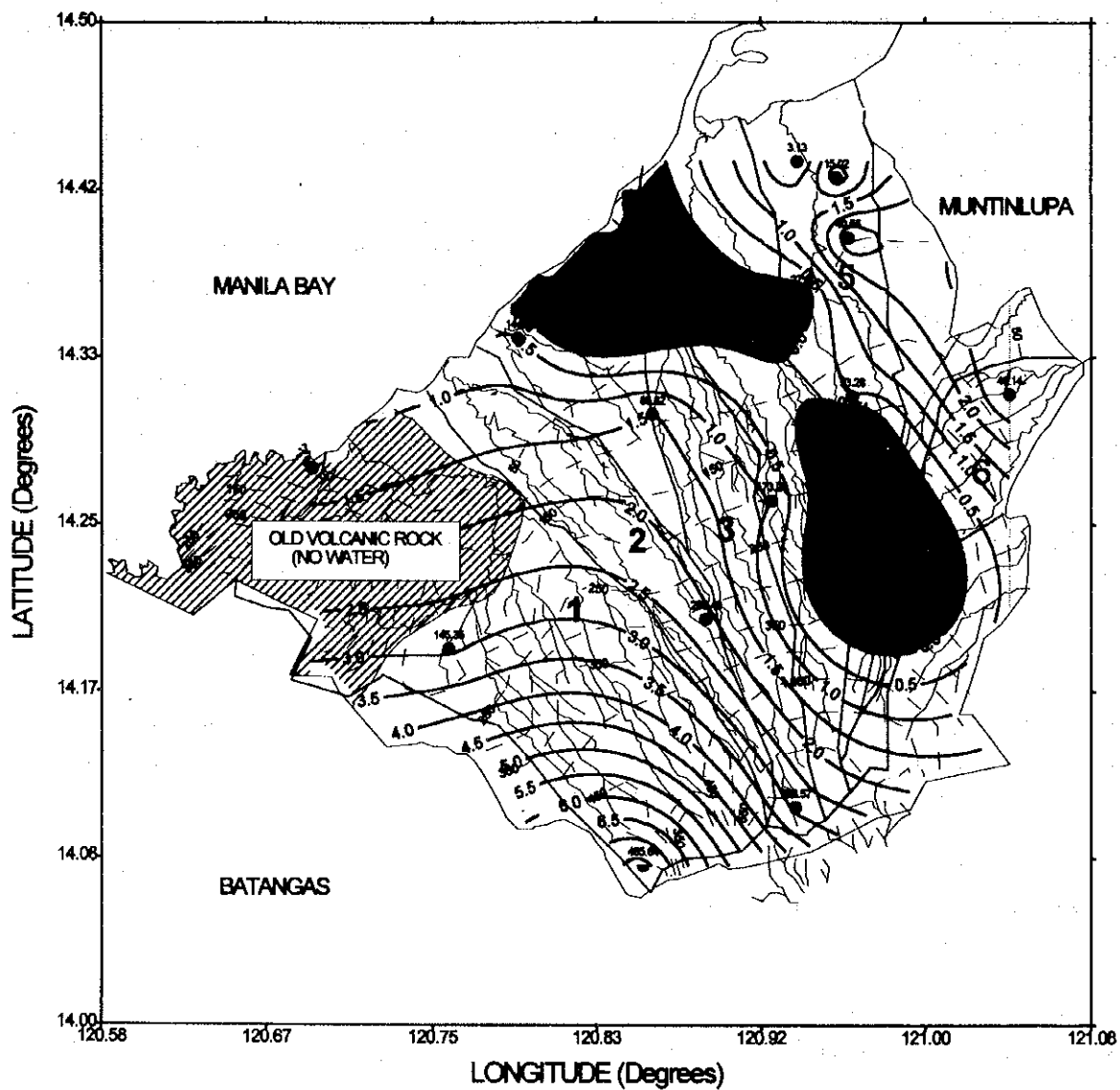
4.3.2 Estimation for Each River Basin

The results of annual average water balance computations are given for the following six (6) river basins.

(1) MARAGONDON RIVER BASIN

The basin has an area of 329 km². As presented schematically in Fig. 4.3-4(b), the annual rainfall depth over the basin is 2,605.2 mm; the annual surface runoff depth is 1,084.4 mm; actual evapotranspiration is 929.9 mm, and the groundwater discharge is 526.0 mm. Hence, the annual recharge is estimated at 64.9 mm or 2.5% of the annual rainfall.

Comparing the present annual groundwater use of 1.2×10^6 cum as presented in Table 4.2-8 with the estimated groundwater recharge of 15.8×10^6 cum, the tentative estimate of the groundwater surplus of the basin is placed at 14.6×10^6 cum. It may also be noted that the measured spring discharge (15.4×10^6 cum) in the basin is almost equal in magnitude to the net



MAJOR RIVER BASIN

- 1 MARAGONDON RIVER BASIN
- 2 LABAC RIVER BASIN
- 3 CANAS RIVER BASIN
- 4 SAN JUAN RIVER BASIN
- 5 IMLUS RIVER BASIN
- 6 BINAN RIVER BASIN

LEGEND:

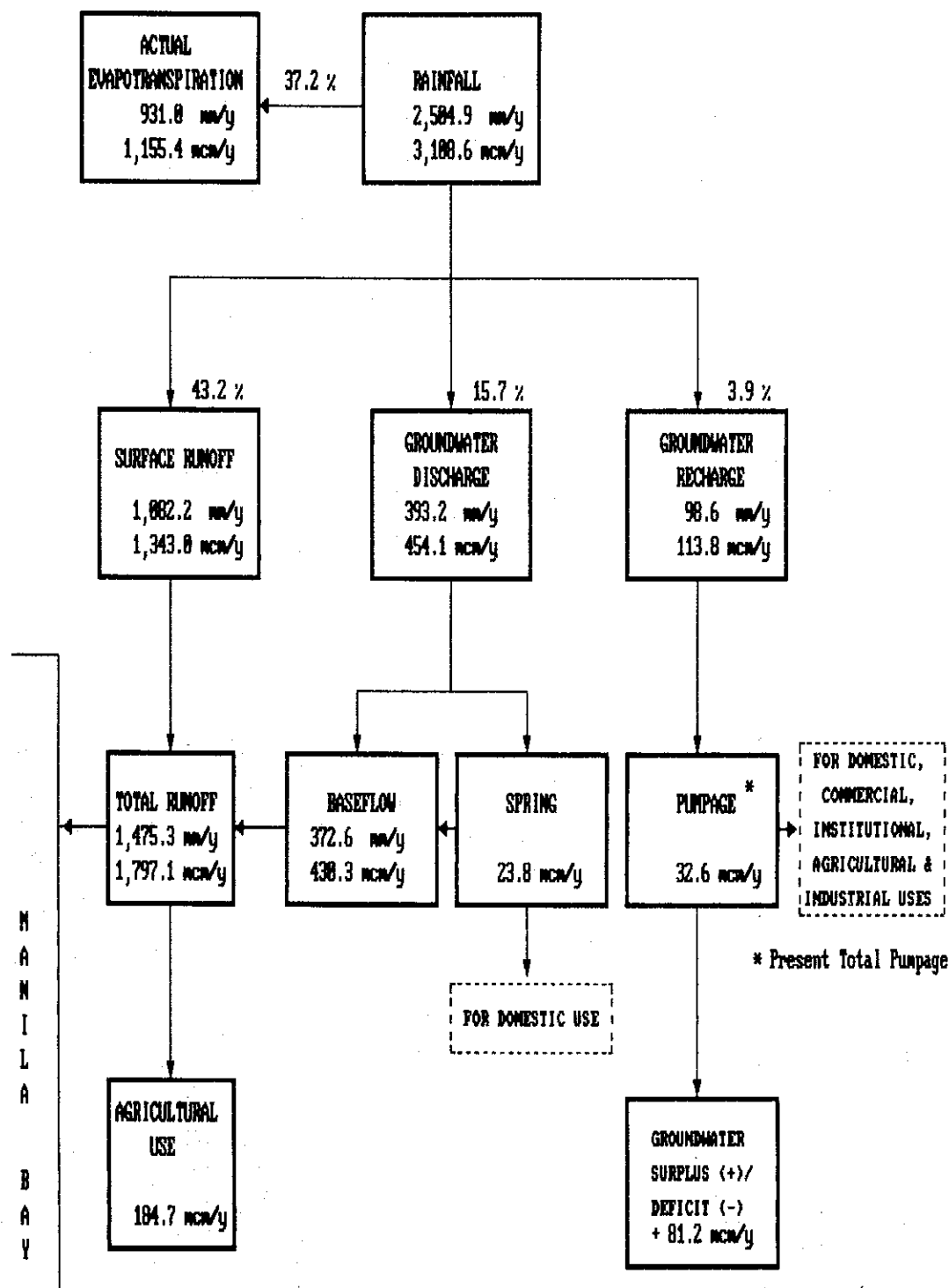
- 100 -- TOPOGRAPHIC ELEVATION (Masi)
- 100 — GROUNDWATER LEVEL ELEVATION (Masi)
- /// OLD VOLCANIC ROCK

CAVITE WATER SUPPLY DEVELOPMENT STUDY

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Fig. 4.3-3

GROUNDWATER LEVEL DIFFERENCE
BETWEEN WET SEASON AND DRY SEASON

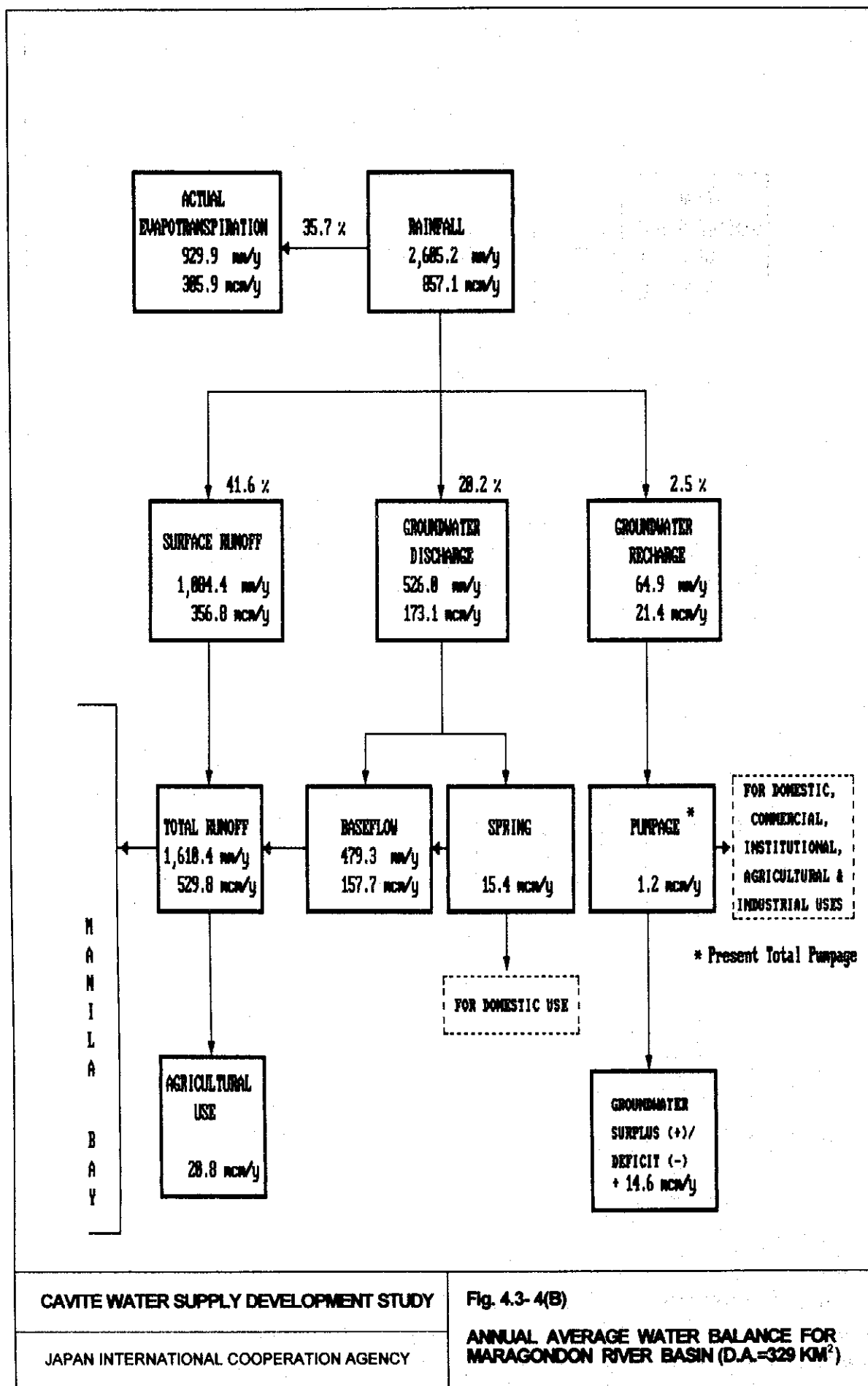


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Fig. 4.3- 4(A)

ANNUAL AVERAGE WATER BALANCE FOR
THE STUDY



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Fig. 4.3- 4(B)

ANNUAL AVERAGE WATER BALANCE FOR
MARAGONDON RIVER BASIN (D.A.=329 KM²)

recharge to its aquifers (15.8×10^6 cum). This may offer a promising type of water supply potential coming from springs, better than the recharge to its aquifers.

(2) LABAC RIVER BASIN

The basin has an area of 106 km^2 . The Balsahan River (22 km^2) is the representative sub-basin for the water balance analysis. As presented schematically in Fig. 4.3-4(c), the annual rainfall depth over the basin is 2,766.2 mm; the annual surface runoff depth is 1,402.6 mm; the actual evapotranspiration is 949.5 mm, and the groundwater discharge is 337.1 mm. Hence, the annual recharge is estimated at 77.6 mm or 2.8% of the annual rainfall.

Comparing the present annual groundwater use of 0.3×10^6 cum as presented in Table 4.2-8 with the estimated groundwater recharge of 8.2×10^6 cum the tentative estimate of the groundwater surplus of the basin is placed at 7.9×10^6 cum.

(3) CANAS RIVER BASIN

This basin has an area of 111 km^2 and is represented by Panaysayan River (29 km^2) for the water balance analysis. As presented schematically in Fig. 4.3-4(d), the annual rainfall depth over the basin is 2,699.9 mm; annual surface runoff depth is 1,231.1 mm; the actual evapotranspiration is 945.6 mm and the groundwater discharge is 392.0 mm. Hence, the annual recharge is estimated at 131.2 mm or 4.9% of the annual rainfall.

Comparing the present annual groundwater use of 4.6×10^6 cum as presented in Table 4.2-8 with the estimated groundwater recharge of 14.6×10^6 cum, the tentative estimate of the groundwater surplus of the basin is placed at 10.0×10^6 cum.

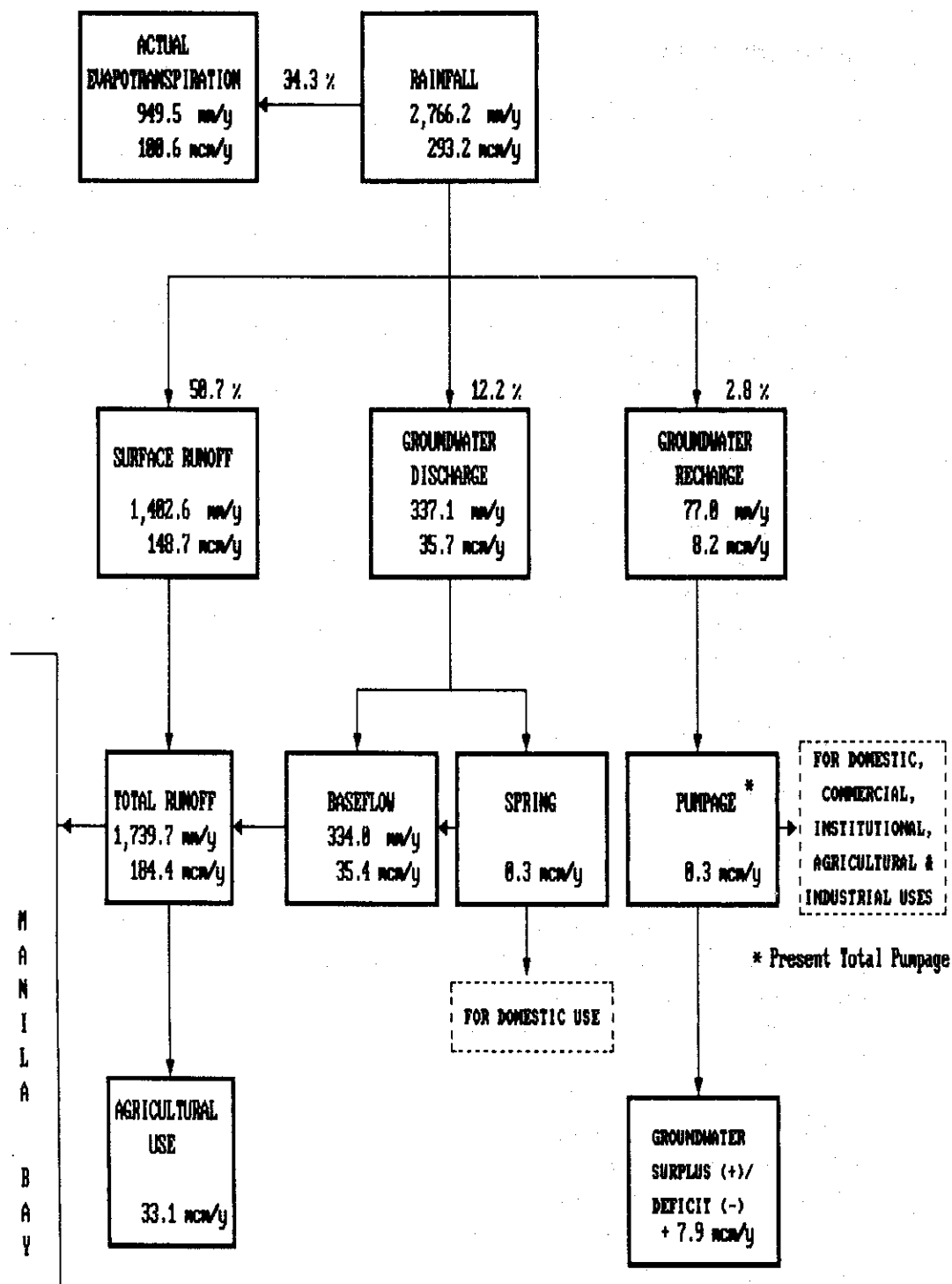
(4) SAN JUAN RIVER BASIN

This basin with an area of about 172 km^2 is represented by Ilang-Ilang River (60 km^2). As presented schematically in Fig. 4.3-4(e) the annual rainfall depth over the basin is 2,602.6 mm; annual surface runoff depth is 1,155.9 mm; the actual evapotranspiration is 938.7 mm; and the groundwater discharge is 360.5 mm. Hence, the annual recharge is estimated at 147.4 mm or 5.7% of the annual rainfall.

Comparing the present annual groundwater use of 20.2×10^6 cum as presented in Table 4.2-8 with the estimated groundwater recharge of 18.9×10^6 cum, the tentative estimate of the groundwater deficit of the basin is placed at 1.3×10^6 cum. This deficit confirms the annual decline of groundwater levels in the basin.

(5) IMUS RIVER BASIN

This basin has an area of about 128 km^2 . As presented schematically in Fig. 4.3-4(f), the annual rainfall depth over the basin is estimated at 2,355.6 mm; the annual surface runoff depth is 983.1

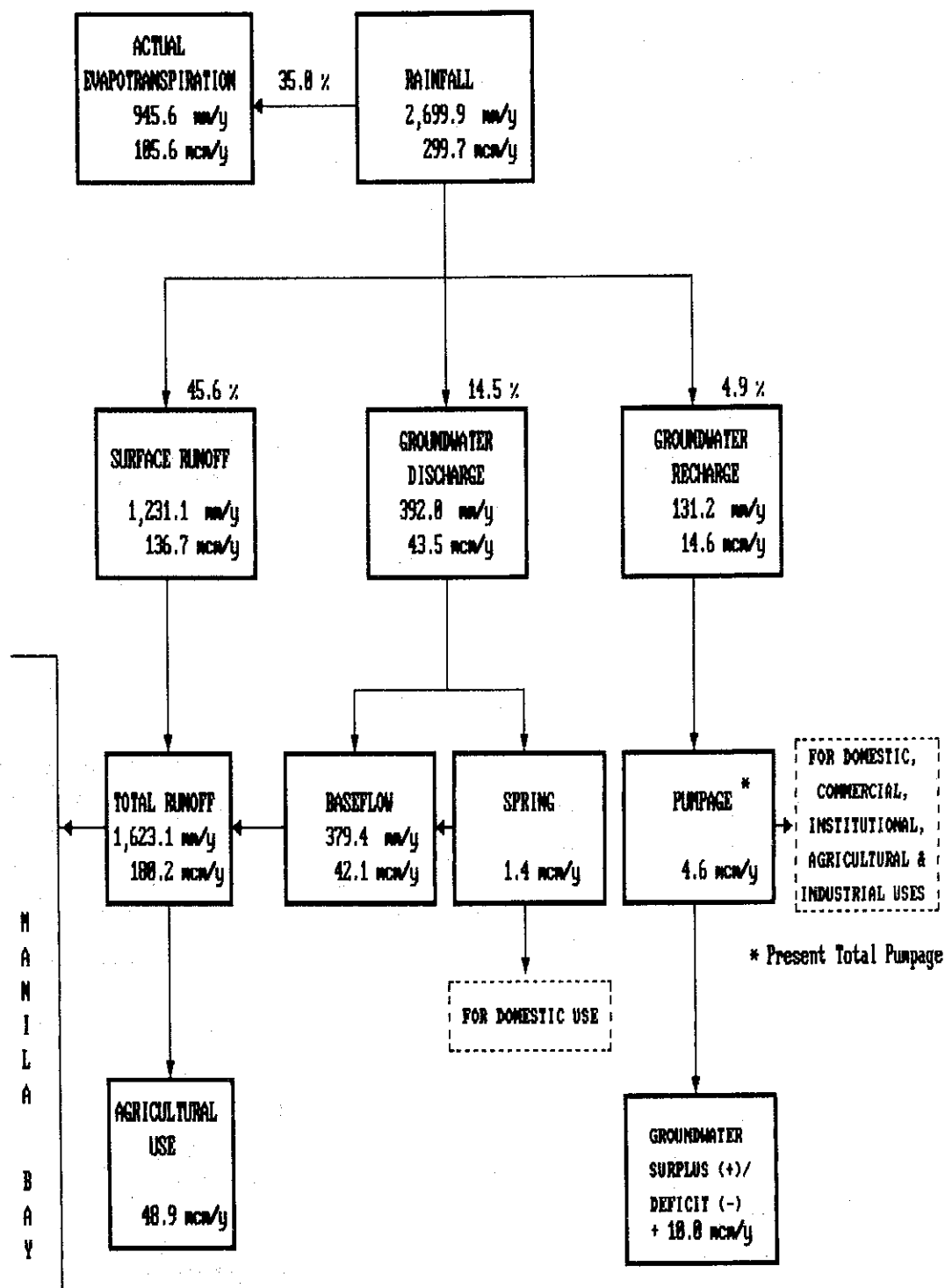


CAVITE WATER SUPPLY DEVELOPMENT STUDY

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Fig. 4.3- 4(C)

ANNUAL AVERAGE WATER BALANCE FOR
LABAC RIVER BASIN (D.A.=106 KM²)

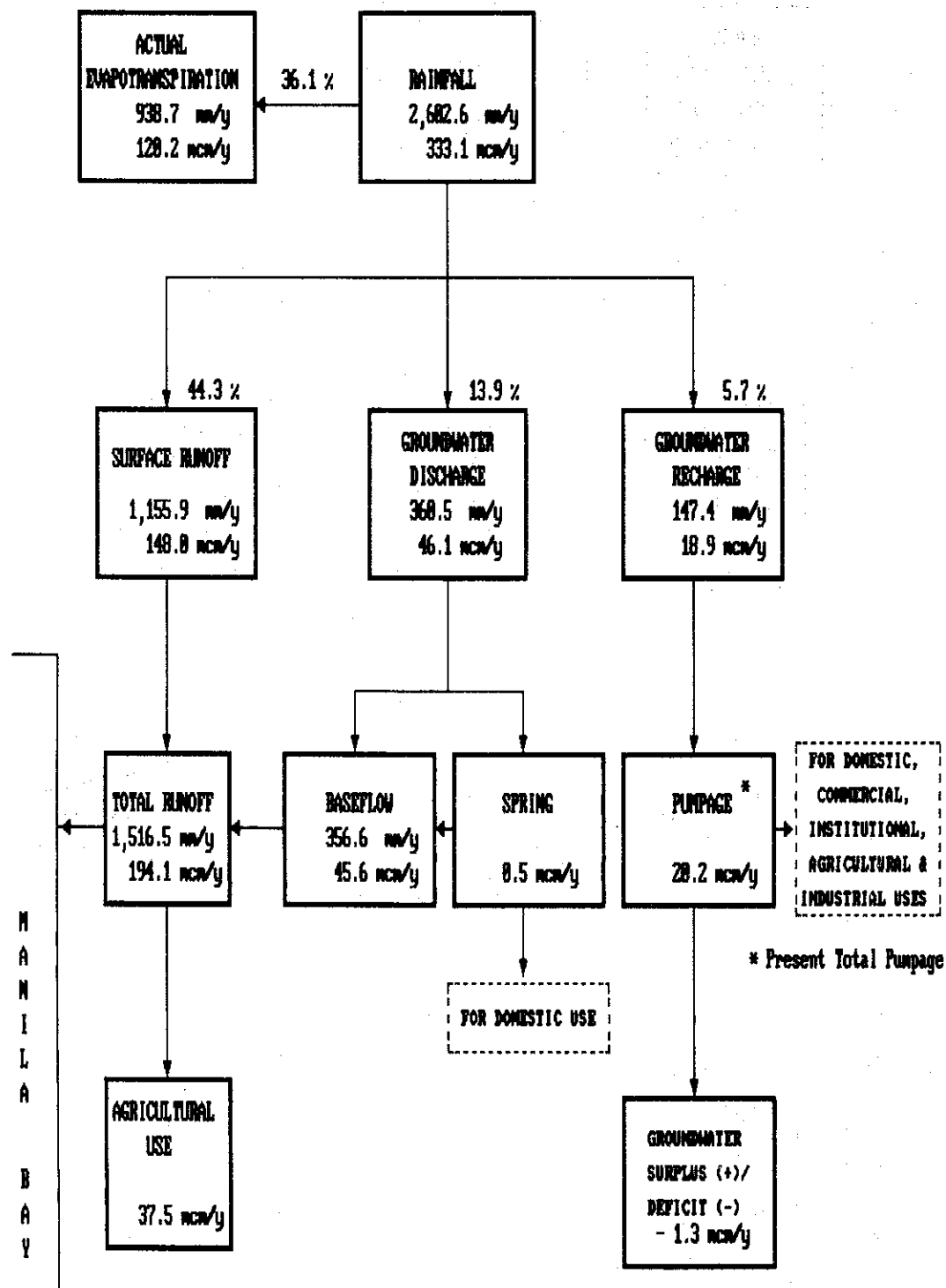


CAVITE WATER SUPPLY DEVELOPMENT STUDY

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Fig. 4.3- 4(D)

ANNUAL AVERAGE WATER BALANCE FOR
CAÑAS RIVER BASIN (D.A.=111 KM²)

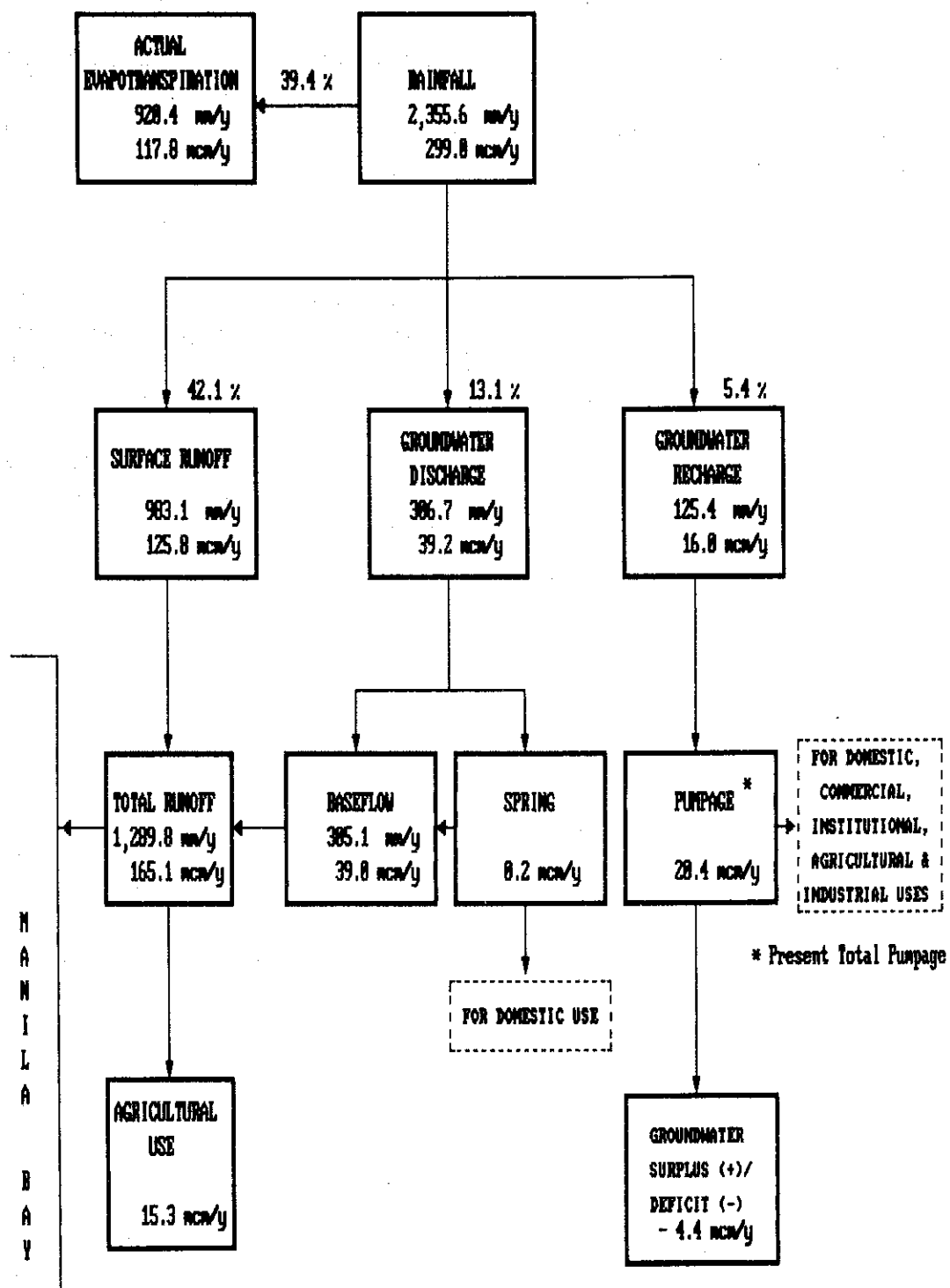


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Fig. 4.3- 4(E)

ANNUAL AVERAGE WATER BALANCE FOR
SAN JUAN RIVER BASIN (D.A.=127 KM²)



CAVITE WATER SUPPLY DEVELOPMENT STUDY

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Fig. 4.3- 4(F)

ANNUAL AVERAGE WATER BALANCE FOR
IMUS RIVER BASIN (D.A.=128 KM²)

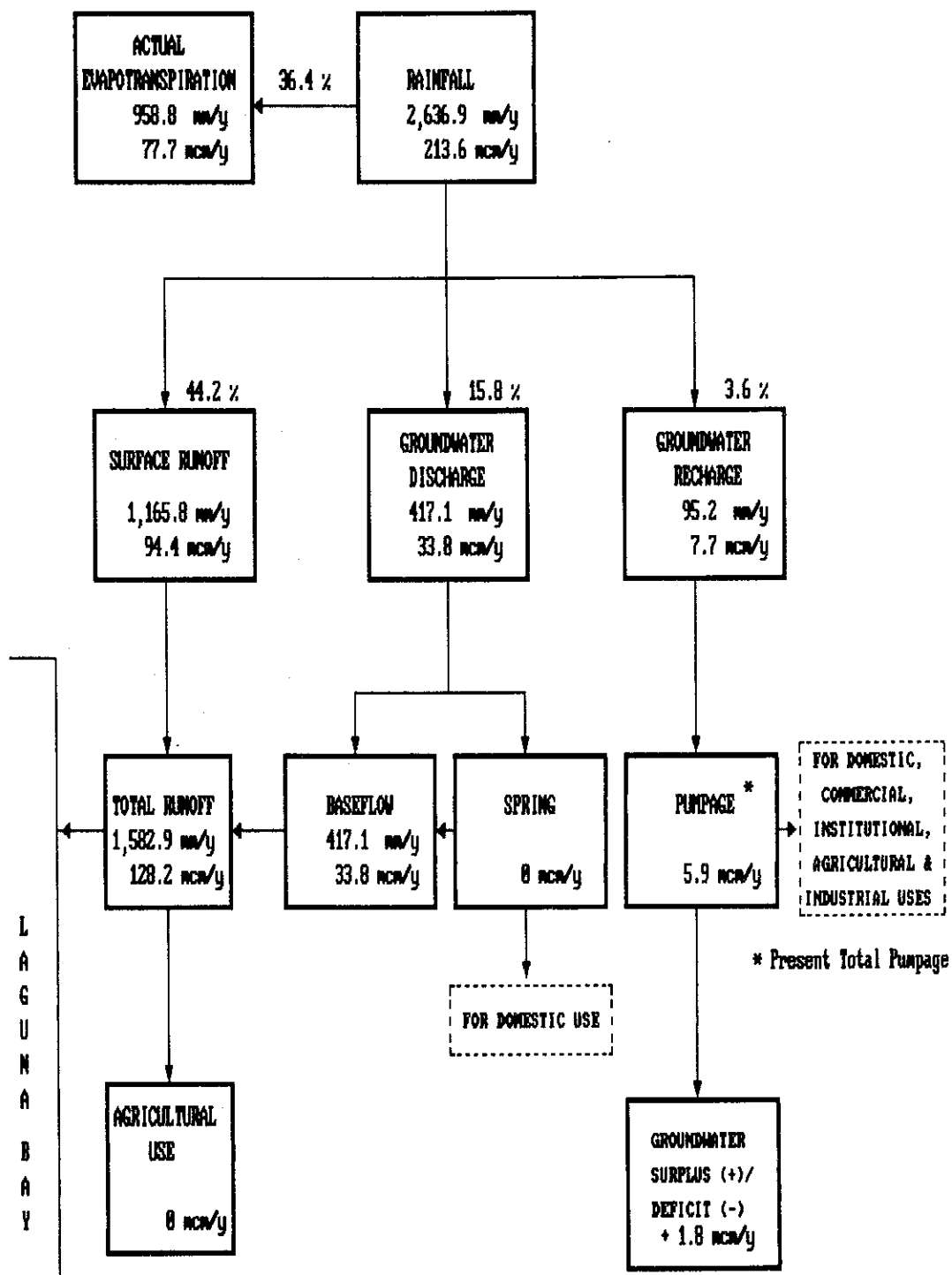
mm; actual evapotranspiration is 920.4 mm; and the groundwater discharge is 306.7 mm. Hence, the annual recharge is estimated at 125.4 mm or 5.4% of the annual rainfall.

Comparing the present annual groundwater use of 20.4×10^6 cum as presented in **Table 4.2-8** with the estimated groundwater recharge of 16.0×10^6 cum, the tentative estimate of the groundwater deficit of the basin is placed at 4.4×10^6 cum. This deficit confirms the annual decline of groundwater levels in the basin.

(6) **BINAN RIVER BASIN**

This basin has an area of about 81 km². As presented schematically in **Fig.4.3-4(g)**, the annual rainfall depth over the basin is estimated at 2,636.9 mm; the annual surface runoff depth is 1,165.8 mm; actual evapotranspiration is 958.8 mm; and the groundwater discharge is 417.1 mm. Hence, the annual recharge is estimated at 95.2 mm or 3.6% of the annual rainfall.

Comparing the present annual groundwater use of 5.9×10^6 cum as presented in **Table 4.2-8** with the estimated groundwater recharge of 7.7×10^6 cum, the tentative estimate of the groundwater surplus of the basin is placed at 1.8×10^6 cum.



CAVITE WATER SUPPLY DEVELOPMENT STUDY

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Fig. 4.3- 4(G)

ANNUAL AVERAGE WATER BALANCE FOR
BIÑAN RIVER BASIN (D.A.=81 KM²)

CHAPTER 5

RELATIONSHIP BETWEEN GROUNDWATER USE AND GROUNDWATER LEVEL FLUCTUATION

CHAPTER 5

RELATIONSHIP BETWEEN GROUNDWATER USE AND GROUNDWATER LEVEL FLUCTUATION

The groundwater simulation was conducted to examine the relationship between groundwater use and groundwater level fluctuation in the Study Area. The San Juan River Basin was selected as the objective area for groundwater simulation based on criteria set by the Study Team.

The groundwater simulation was conducted using MODFLOW. A 3-layer 3-D groundwater flow model was constructed based on the results of hydrogeological studies. Boundary conditions, initial conditions and parameter values were assigned to the model. Steady-state and historical calibrations were conducted, comparing the computed and observed piezometric levels.

The calibrated model was then used to predict future groundwater flow and piezometric heads based on three (3) future pumping schemes. The model response showed a maximum decline of 50 meters in the simulation area in ten years (1995 to 2005).

The accuracy and reliability of groundwater model installed at LWUA must be improved in the future by a more accurate and updated pumpage and aquifer parameter estimates and by a continuous monitoring of groundwater levels.

5.1 OBJECTIVE OF GROUNDWATER SIMULATION

Since groundwater utilization will continue as the Study Area develops, particularly no surface water supply system is feasible in the area at present, the problem of continual decline of piezometric levels will only get more serious in the future. This problem is basically due to unplanned groundwater resources development and management, far beyond the sustained capacity of the complex aquifer system underlying the Cavite area. Due to the complexity and seriousness of the problem, there is an urgent need for identifying rational solutions through systematic planning studies.

Computer models are essential for analyzing complex problems of groundwater flow because they provide a quantitative framework for synthesizing a large set of parameters that describe spatial variability within the subsurface environment, as well as spatial and temporal trends in hydrologic parameters and stresses, and historical trends in water levels. While some decisions can be made using best engineering or best geologic judgment, human reasoning alone is inadequate to synthesize the many factors involved in most complex problems.

A groundwater simulation was undertaken to predict spatial and temporal response of the aquifer system in terms of piezometric levels for the future pumping schemes based on the projected water demand in the Study Area from year-1995 to year-2005. The development and management strategy is ultimately aimed at regulating the groundwater pumpage to control the piezometric levels at allowable limits.

5.2 OBJECTIVE AREA FOR GROUNDWATER SIMULATION

5.2.1 Criteria for the selection of objective area

The San Juan River Basin, consisting of Ilang-Ilang River Basin and Malabon River Basin, was recommended by the Study Team as the objective area for groundwater simulation based on the following criteria.

- (1) The objective area must be a unit hydrological basin in which the actual groundwater system can be easily conceptualized and the boundary conditions and system parameters for numerical simulation can be properly set.
- (2) The objective area must have municipalities where groundwater levels have continually declined over the years due to overpumping of groundwater. This condition shall make the calibration of the mathematical model possible.
- (3) The water balance calculations for the objective area must be performed with high accuracy.

As per criterion (1), the San Juan River Basin is a unit hydrological basin defined by the hydrological watershed divides in the topographic map. Fig. 5.2-1 shows the objective area for groundwater simulation.

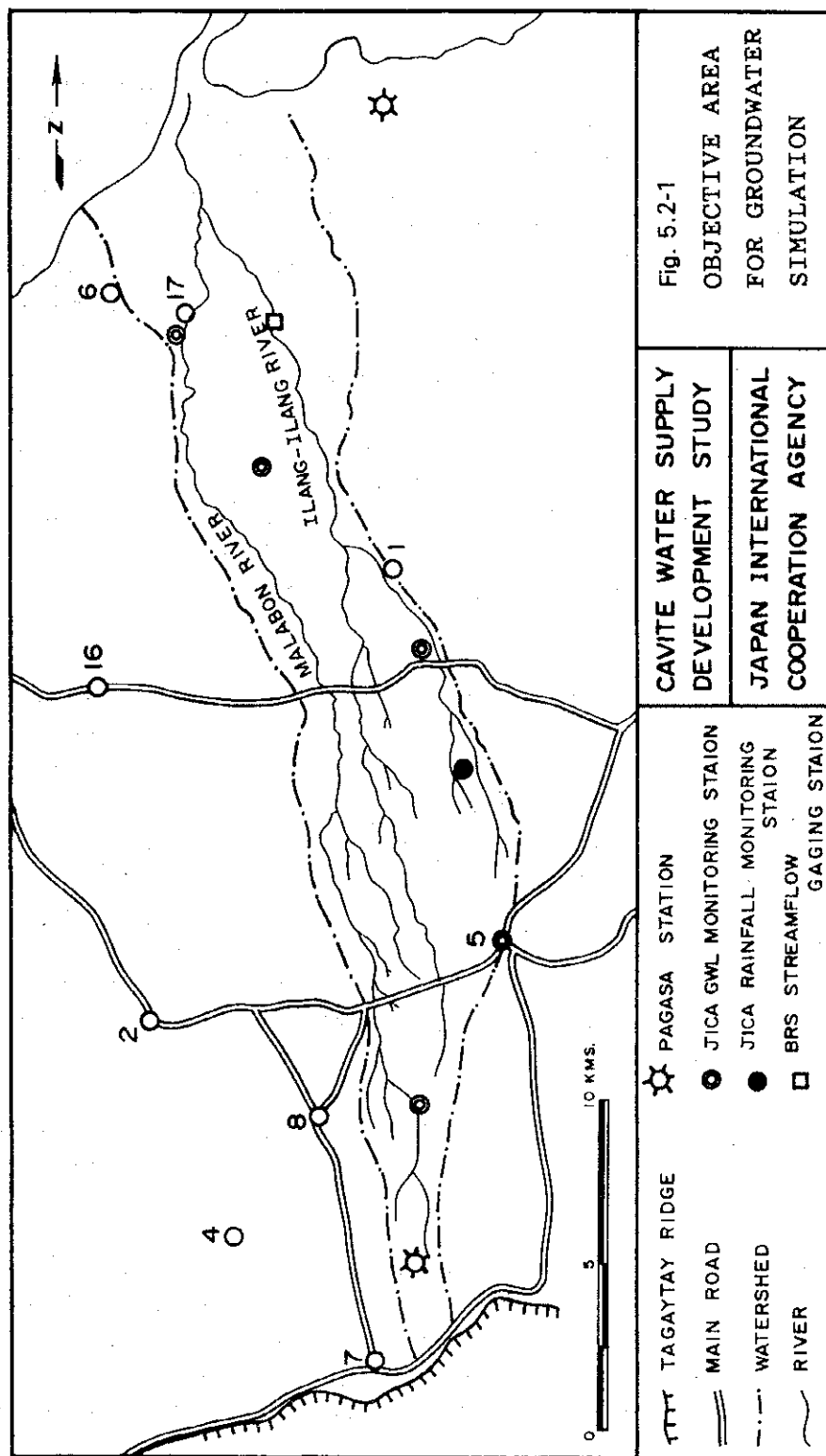
As per criterion (2), groundwater levels in Gen. Trias, Dasmarinas and Silang (portions of these municipalities are included in the San Juan River Basin) have declined significantly in recent years due to the rapid subdivision and industrial developments.

As per criterion (3), water balance analysis can be performed accurately for the San Juan River Basin based on the discussion in Chapter 4 regarding the monitoring of rainfall and groundwater levels, the dry and rainy seasons' survey of spring and river discharges, the availability of rainfall and runoff data, and the inventory survey of wells and groundwater use in the basin.

5.2.2 General Condition of the San Juan River Basin

The San Juan River Basin is bounded by Tagaytay Ridge to the south, Aguinaldo Highway to the east, municipalities of Noveleta, Kawit and Imus to the north, and Tagaytay-Gen. Trias Road to the west. It has a total land area of 155 km², which represents 12.5% of the total land area of the Study Area, and comprises portions of Tagaytay City, Amadeo, Silang, Dasmarinas, and General Trias. Portions of Imus, Noveleta, and Kawit, which are outside the Study Area, also form part of this basin.

The downstream portion of the basin is a small plain covered by Quaternary alluvium and Guadalupe formation, lying flat at slopes of 0 to 3%, facing Manila Bay, and below 100-meter elevation. The soil is typically sandy river deposits. The land use in these portions of Imus and Gen. Trias is largely prime agricultural lands while in Noveleta and Kawit at the southern end are built up areas, salt beds and marine ponds.



The central area is characterized by rolling and undulating terrain, sloping from 3% to 8% between elevations of 100m and 300m above MSL, and comprises portions of Dasmarinas and Silang. The middle reach of Malabon River and the upper section of Ilang-Ilang River are located in this area. The principal soil is made of loam, which includes tuff and pyroclastics as its base. Currently the land is used mainly to cultivate tree crops such as sugarcane and mango. Industrial and residential areas are expanding along main roads, especially Carmona-Trece Martirez-Naic Road, and water for these areas are supplied by deep wells. Although there have been no serious problems up to now, further urban and industrial developments in the future will correspondingly increase the dependency on groundwater which will result in uncontrolled decline of piezometric head.

The upland, with elevations higher than 300m above sea level, is hilly and mountainous, having slopes of 8% to 15%. The upland is geologically dominated by pyroclastics and tuff breccia. The upstream section of the Malabon River and several springs are located here. Portions of Silang, Amadeo and Tagaytay City are included in this southern part of the basin. The highest elevation of the basin, which forms part of the Tagaytay Ridge, is 610 meters above sea level. Tree crops such as coconut, coffee, mango, etc. are widely grown.

The San Juan River Basin is long and narrow, rolling and undulating with numerous moderately deep, narrow ravines and gulches. The direction of flow in the San Juan River System is generally dictated by the direction of slope in the upland areas. The Ilang-Ilang River which originates from the hills at Lucsuhin Spring more than one kilometer southwest of Silang flows north-northeast. Towards Dasmarinas, several creeks merge and contribute to its flow. Opposite this municipality, the river is joined by Dasmarinas River and meanders northwest to discharge finally into Manila Bay near Noveleta.

The Bureau of Public Works (BPW) constructed a gaging station in Alapan 2nd, Imus to monitor the flow of the Ilang-Ilang River. Details of this station is discussed in Chapter 4. This river is currently utilized by NIA for irrigation purposes. Several small dams, together with irrigation canals, were reportedly constructed by NIA across the river.

The Dasmarinas River originates from the highlands of Dasmarinas about 7.5 km southeast of the poblacion. It flows in a northwesterly direction rushing close to the poblacion. It then joins Ilang-Ilang River two kilometers northwest.

The Malabon River originates as Rio Grande River in Barangays Maitim and Matagbak, Amadeo about 10 km southwest of Poblacion, Silang, near Tagaytay Ridge. The Rio Grande River flows in a northwesterly direction and is joined by many tributaries towards Poblacion, Gen. Trias. The Pasong Camachile River joins this river to form the Malabon River which then flows in the north direction and drains into Manila Bay. The NIA dam across Rio Grande River diverts about 1,000 lps to irrigate the ricefields in the area.

The mean annual rainfall in the basin varies from 2,000 mm in the lowland to 3,800 mm in the upland.

5.3 METHOD OF GROUNDWATER SIMULATION

The groundwater simulation was conducted using the Modular Three-Dimensional Finite Difference Groundwater Flow Model (MODFLOW) developed by the U.S. Geological Survey (McDonald and Harbaugh, 1984). MODFLOW is a very popular, correctly formulated and versatile model with many options for hydrological computations and modular structure for efficient programming and modification purposes.

Fig. 5.3-1 shows the general flowchart of the groundwater simulation. After defining the objective of the groundwater simulation, the 3-D groundwater flow model was constructed based on the results of the hydrogeological investigations and analyses. Appropriate boundary conditions and hydrogeological parameters were assigned to the model. To understand the model behavior, a steady-state calibration was carried out with neither recharge nor pumpage for a period of 10 years. The assigned boundary conditions and the input parameters were checked and/or modified by evaluating the computed piezometric heads.

After preparing the historical pumpage data, historical calibration of the model was performed. In the process, some earlier assumed parameters and boundary conditions were finally fixed. It was carried out using the input historical annual and monthly pumpage and recharge data. The calibration was conducted several times until the computed piezometric heads agreed satisfactorily with the observed data.

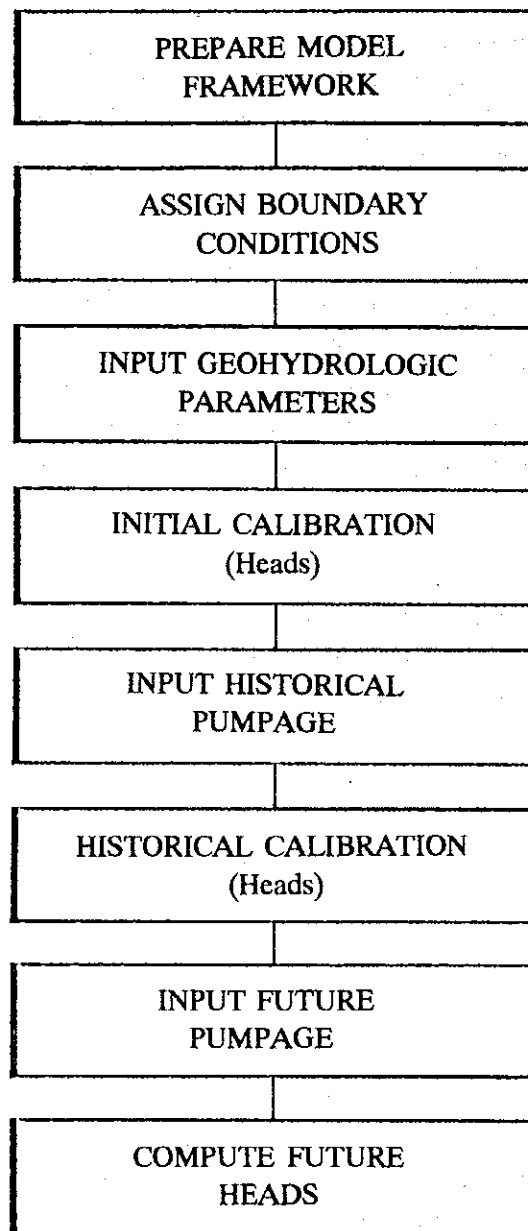
The calibrated model was then used to predict future groundwater flow and piezometric heads, based on the future groundwater pumpage plans. Three (3) future pumping schemes were generated from 1995 to 2005.

5.4 MODEL CONSTRUCTION AND BOUNDARY CONDITIONS

5.4.1 Model Grid

The simulation area covered the whole San Juan River Basin (SJRB). However, it was necessary to expand the model domain up to Maragondon River Basin on the west and beyond Binan River Basin on the east in order to reflect the structure of the groundwater basin itself in the model and to increase reliability of simulation at the marginal areas of the SJRB. For instance, it would be difficult to assign proper boundary conditions at the marginal areas if only the simulation area was taken into account as the model domain, because the aquifers continue outside the limits of the simulation area. Furthermore, the groundwater flow outside affects the groundwater flow inside the basin. Therefore, the model grid was located and established as shown in Fig. 5.4-1.

Fig. 5.4-2 illustrates the grid system of the 3-D groundwater flow model. The SJRB is located between columns 6 and 19 and rows 3 and 40. The grid size in the SJRB was fixed at 1 km \times 1 km. However, it varied outside SJRB from 1 km \times 2 km to 2 km \times 4 km as the grid becomes farther from the simulation area.



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Fig. 5.3-1

WORK FLOW OF GROUNDWATER SIMULATION

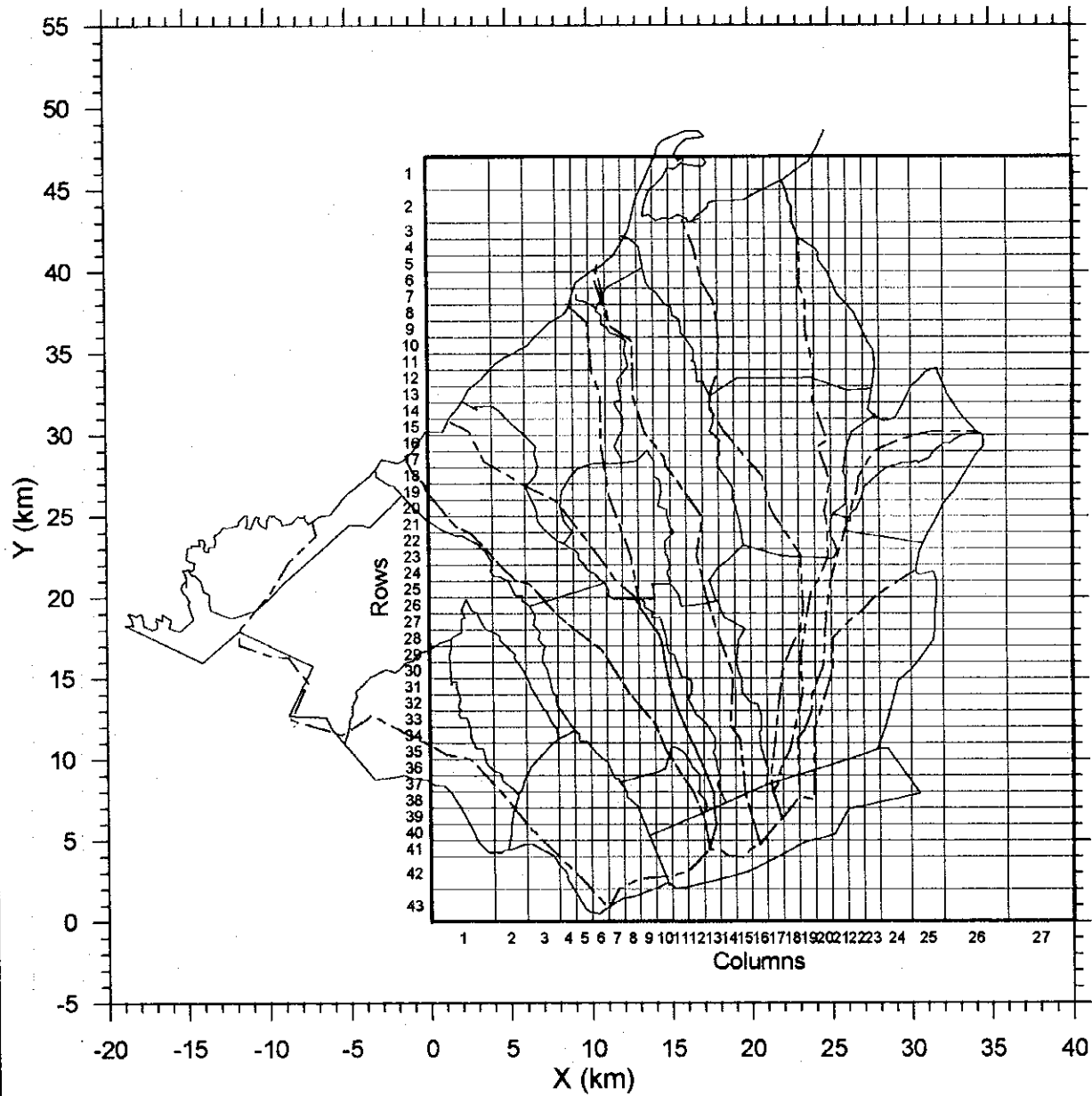


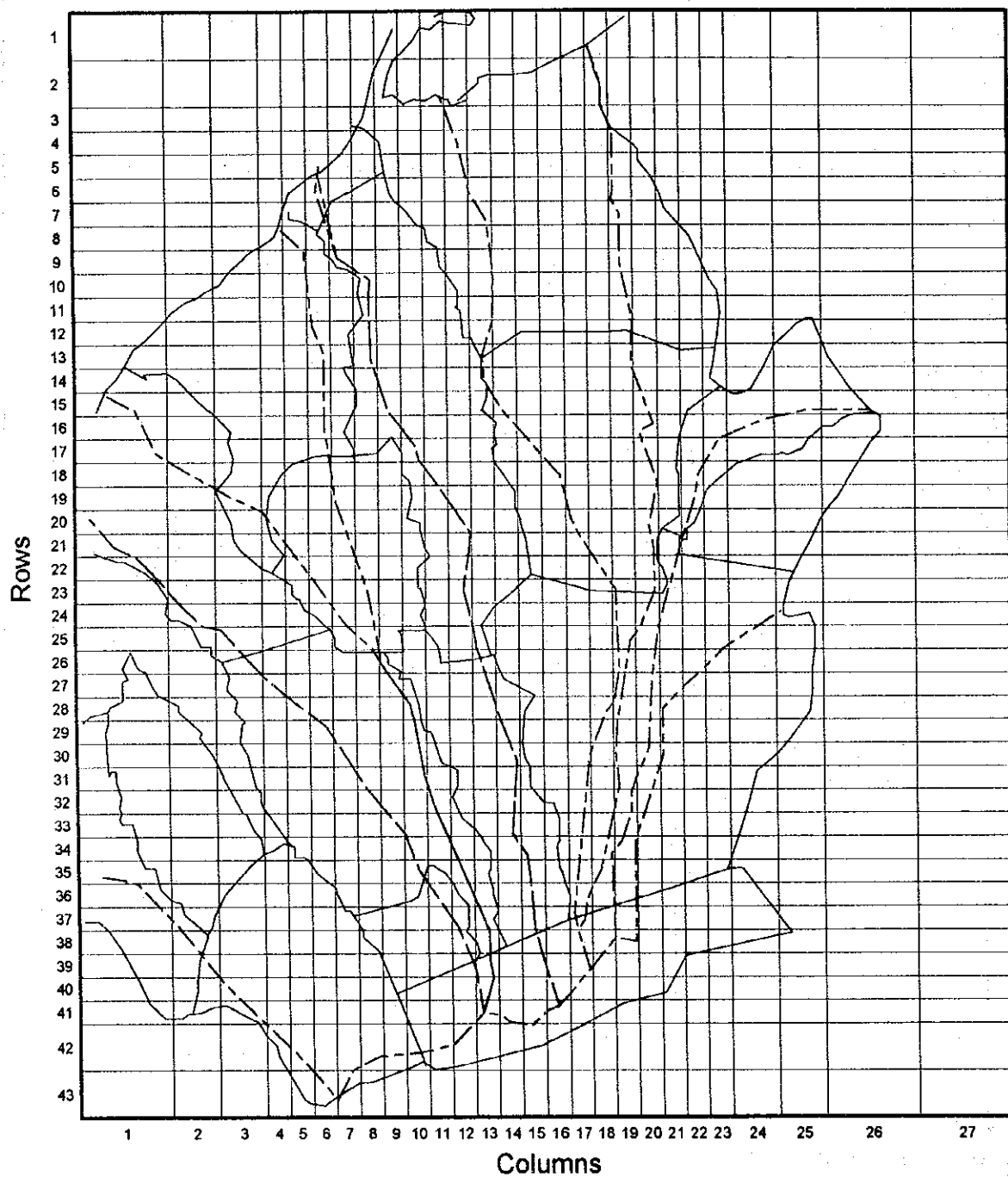
Fig. 5.4-1

LOCATION OF SIMULATION MODEL GRID

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0 5 10 15 20km

Fig. 5.4-2

GRID OF CAVITE 3-D GROUNDWATER
FLOW MODEL

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The model was divided into three (3) independent layers based on the number of aquifer units as shown in the hydrogeological map and sections prepared by the Study Team. The total number of modeled grids in one (1) layer is 1,161 (43 rows \times 27 columns), and the total number of 3-D cells is 3,483 (43 rows \times 47 columns \times 3 layers).

The modeled area was divided into three zones to define the structure of the 3-D model as presented in Fig. 5.4-3. The upper layer is composed of zones 1 and 2 and represents the unconfined aquifers, Kaybubutong Upper and Kaybubutong Middle; the middle layer is made up of zones 2 and 3, representing the confined aquifer, Kaybubutong Lower; and the lower layer consists of zones 1, 2, and 3 and represents the confined aquifer, Talisay Formation. The border between zones 1 and 2 indicates the upstream edge of the middle aquifer, while the boundary between zones 2 and 3 represents the downstream edge of the upper aquifer.

5.4.2 Boundary Conditions

Appropriate boundary conditions were specified for numerical calculation based on the hydrogeological information.

Fig. 5.4-4 shows the boundary conditions of the upper aquifer. The cells located outside the boundaries were treated as inactive cells in the model. Also constant heads, which are equal to the initial heads, were assigned from columns 10 to 23 of the southern boundary. The no flow boundary condition was assigned to the active cells at the northern border.

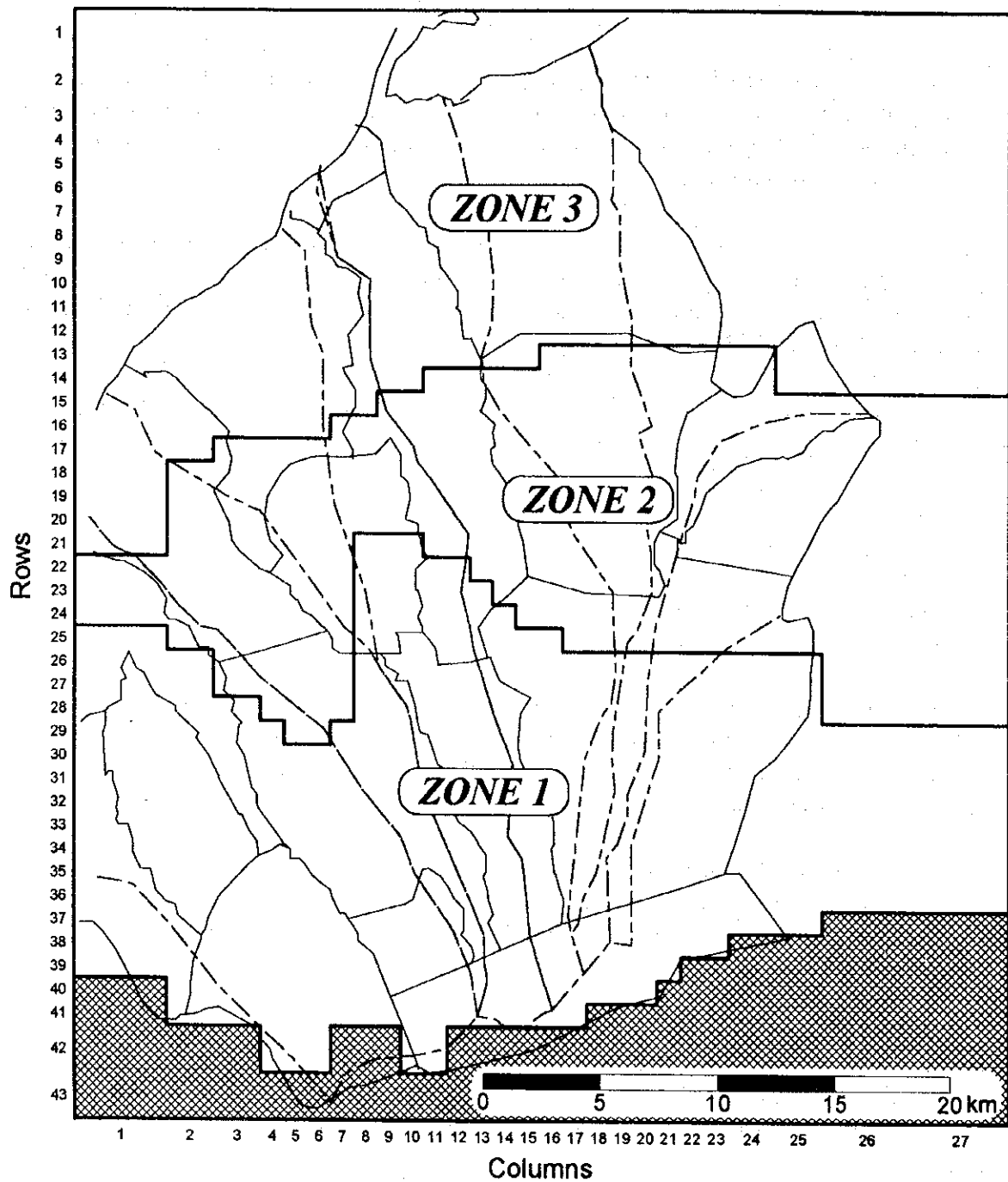
Fig. 5.4-5 shows the boundary conditions of the middle aquifer. The cells located outside the boundaries were treated as inactive cells in the model. Also constant heads, which are equal to 0 meter, were assigned as shown in this figure. The no flow boundary condition was assigned to the active cells at the southern border.

Fig. 5.4-6 shows the boundary conditions of the lower aquifer. The cells located outside the boundaries were treated as inactive cells in the model. Also constant heads, which range from -40 meters in Cavite City to -70 meters in Las Pinas (JICA, 1992), were assigned as shown in this figure. The no flow boundary condition was assigned to the active cells at the southern border of the lower aquifer.

5.5 HYDROGEOLOGICAL PARAMETERS

The MODFLOW program requires the following hydrogeological parameters:

- Top and bottom elevations of each layer
- Type of each layer
- Porosity
- Specific storage
- Hydraulic conductivity



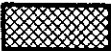

-  Inactive area
-  Boundary of zonation

Fig. 5.4-3

ZONING MAP OF MODELED AREA

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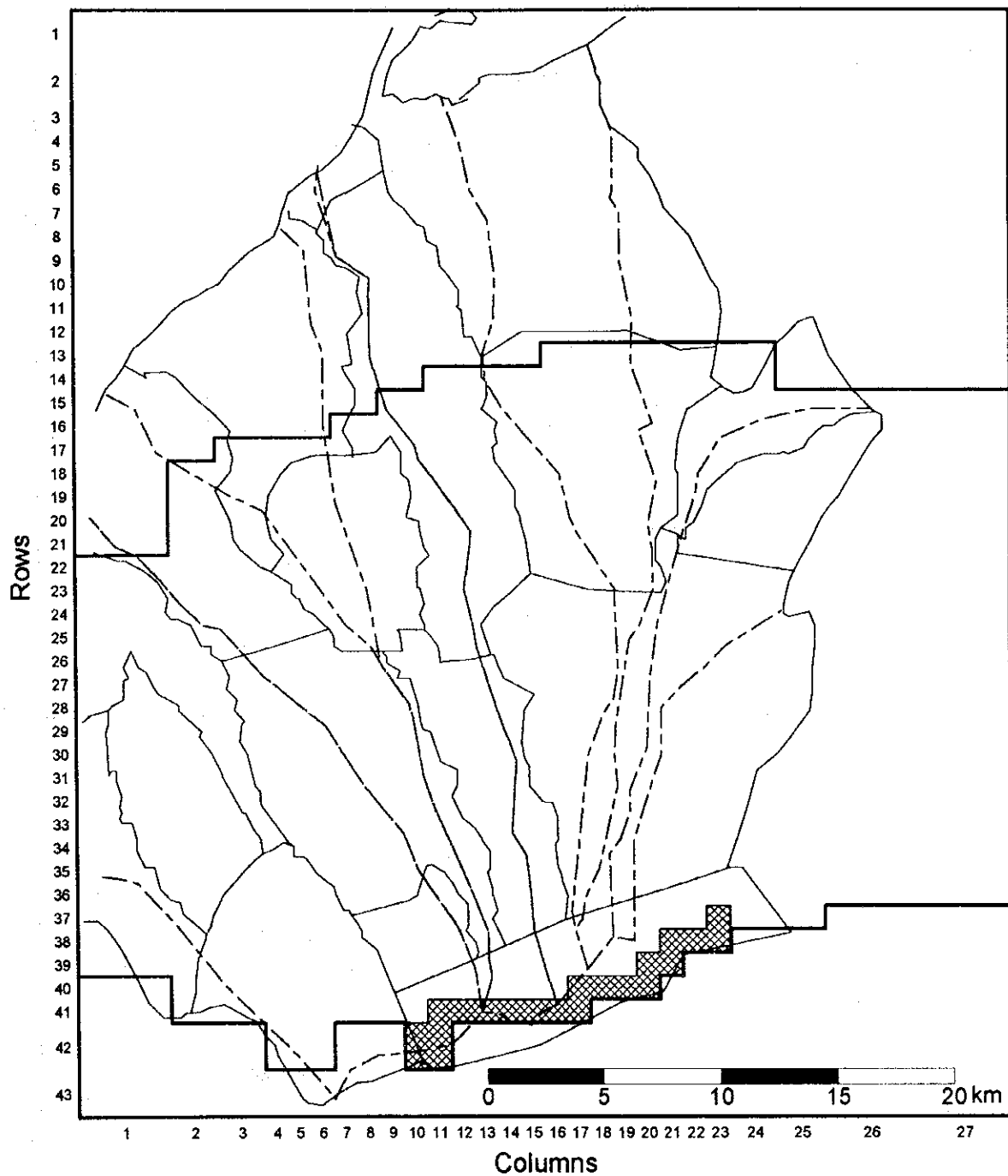


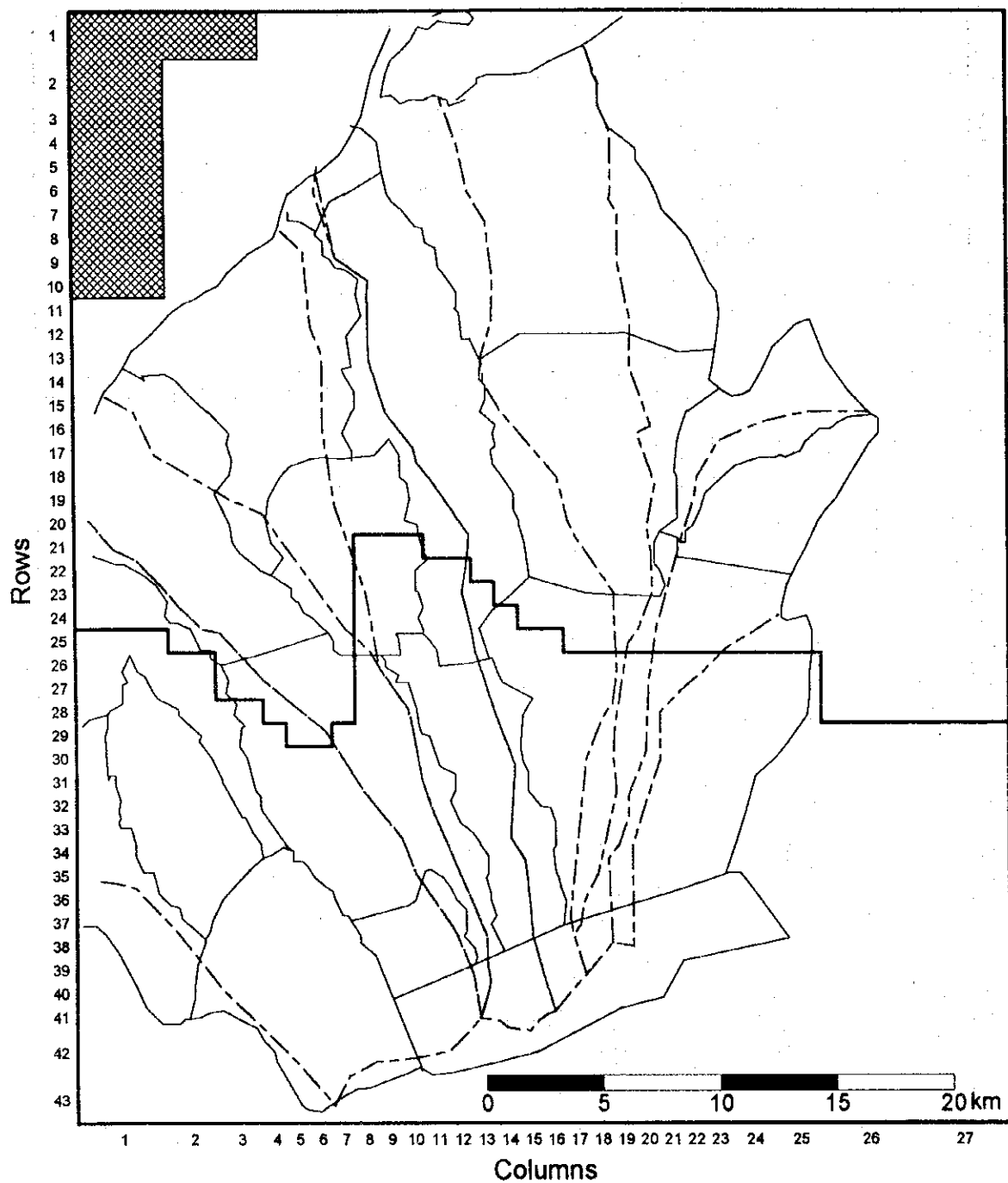
Fig. 5.4-4

BOUNDARY CONDITION OF UPPER AQUIFER

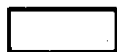
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Constant head boundary



Boundary of aquifer

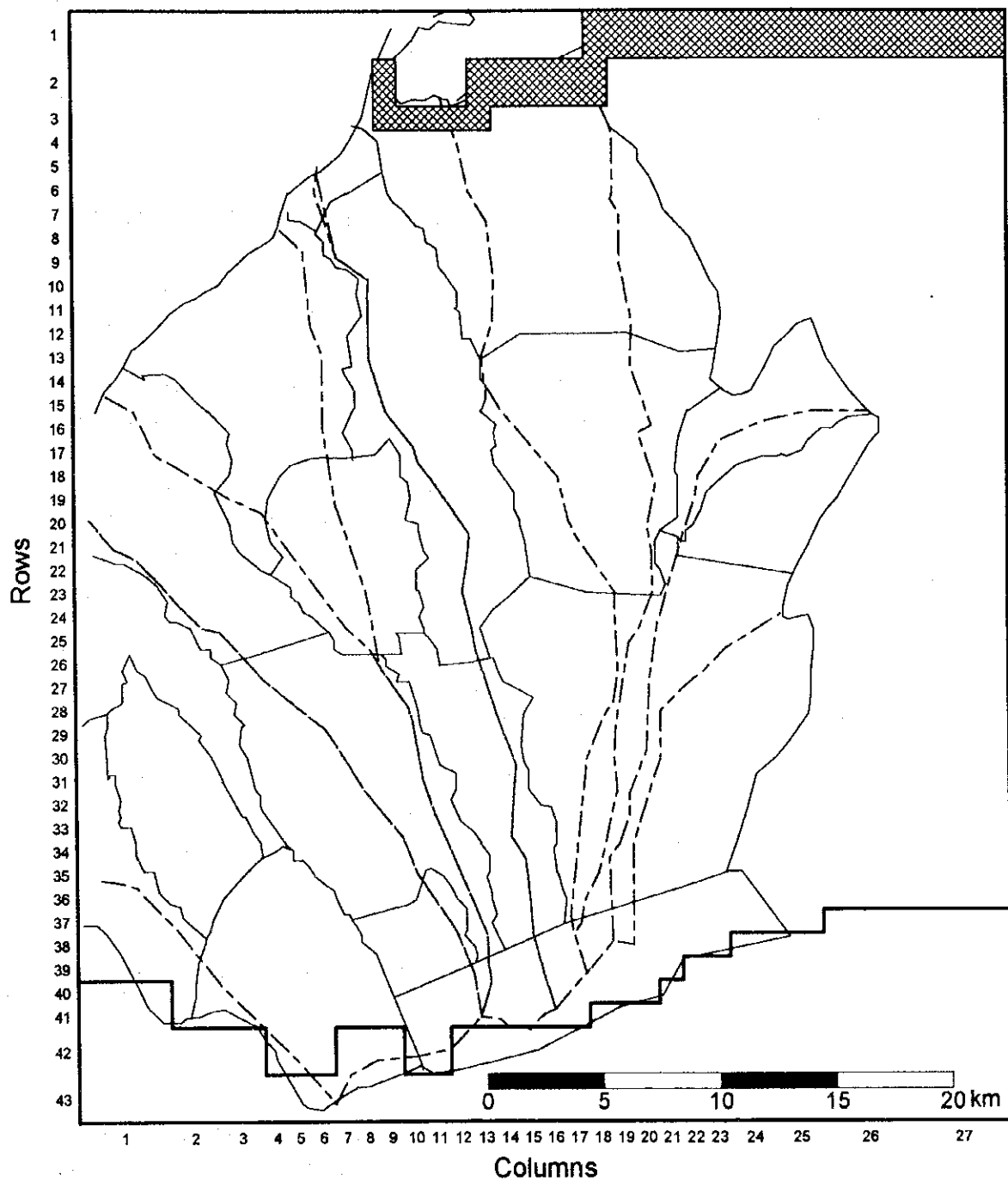
Fig. 5.4-5

BOUNDARY CONDITION OF MIDDLE AQUIFER

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-  Constant head boundary
-  Boundary of aquifer

Fig. 5.4-6

BOUNDARY CONDITION OF LOWER AQUIFER

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- Vertical hydraulic conductivity
- Initial piezometric heads
- Recharge rate

These parameters were prepared from the hydrogeological studies and the updated PGDB for Cavite.

Top and bottom elevations

Top and bottom elevations of each layer were determined from the ground elevation map (1:50,000 topo map of NAMRIA) and the hydrogeological map and sections prepared by the Study Team.

The top elevation of the upper unconfined aquifer (zones 1 and 2) is the same as the ground elevation, its bottom elevation is equal to the top elevation of the lower aquifer in zone 1 and the middle aquifer in zone 2.

The top elevation of the middle confined aquifer (zones 2 and 3) is the same as the bottom elevation of the upper aquifer in zone 2 and the ground elevation in zone 3, its bottom elevation is equal to the top elevation of the lower aquifer in zones 2 and 3.

The top elevation of the lower confined aquifer (zones 1, 2 and 3) is the same as the bottom elevations of the upper aquifer in zone 1 and the middle aquifer in zones 2 and 3, its bottom elevation is assumed equal to its top elevation minus 300 meters.

Type of each layer

Upper layer was specified as unconfined aquifer, while middle and lower layers were designated as confined aquifers.

Porosity

For the simulation of steady state flow, porosity is not required by MODFLOW. However, porosity is required by MODPATH program, which is a subprogram of MODFLOW for the calculation of the flow velocities. Therefore, this parameter is required by MODFLOW input program, PREMOD. The porosity value of 0.2 was uniformly given to the model (upper, middle and lower aquifers).

Specific storage

The Block-Centered Flow (BCF) Package of MODFLOW program requires dimensionless storage coefficient values in each layer of the model. PREMOD converts the specific storage value of the cell material into dimensionless storage coefficient by multiplying it by the layer thickness of the cell.

The values of storage coefficient obtained from pumping tests usually have a wide range of variation. However, Walton (1970) suggested that the range of storage coefficient values in confined aquifers should vary between 1.0×10^{-5} and 1.0×10^{-3} for all soils and between 5.0×10^{-5} and 1.0×10^{-2} in productive aquifers. Therefore, the initial storage coefficient of the upper layer was assumed at 1.0×10^{-1} , while that of middle and lower aquifers was set at 1.0×10^{-3} . These values were then converted into specific storage by dividing them by the thickness of the cell.

Hydraulic conductivity

Transmissivity is the most important parameter for the model. The PREMOD computes transmissivity from horizontal hydraulic conductivity (or permeability) by multiplying it by the layer thickness. An anisotropy factor (T_{yy}/T_{xx}) was specified in the model.

Generally, transmissivity values were obtained from pumping tests. The Study Team collected some pumping test data as a result of the well inventory survey and carried out pumping tests at the newly constructed test wells. However, the numbers of collected pumping test data and aquifer parameter information were insufficient to evaluate the regional aquifer characteristics. Therefore, the transmissivity values were obtained from the hydrogeological information and specific capacity data of the production wells stored in the PGDB database. The specific capacity map, showing the locations of the production wells with yield and drawdown data from pumping test, is presented in Fig. 5.5-1.

If the well has only the value of specific capacity, transmissivity can be estimated based on the following equation presented by Logan (1964):

$$T = 1.22 Sc$$

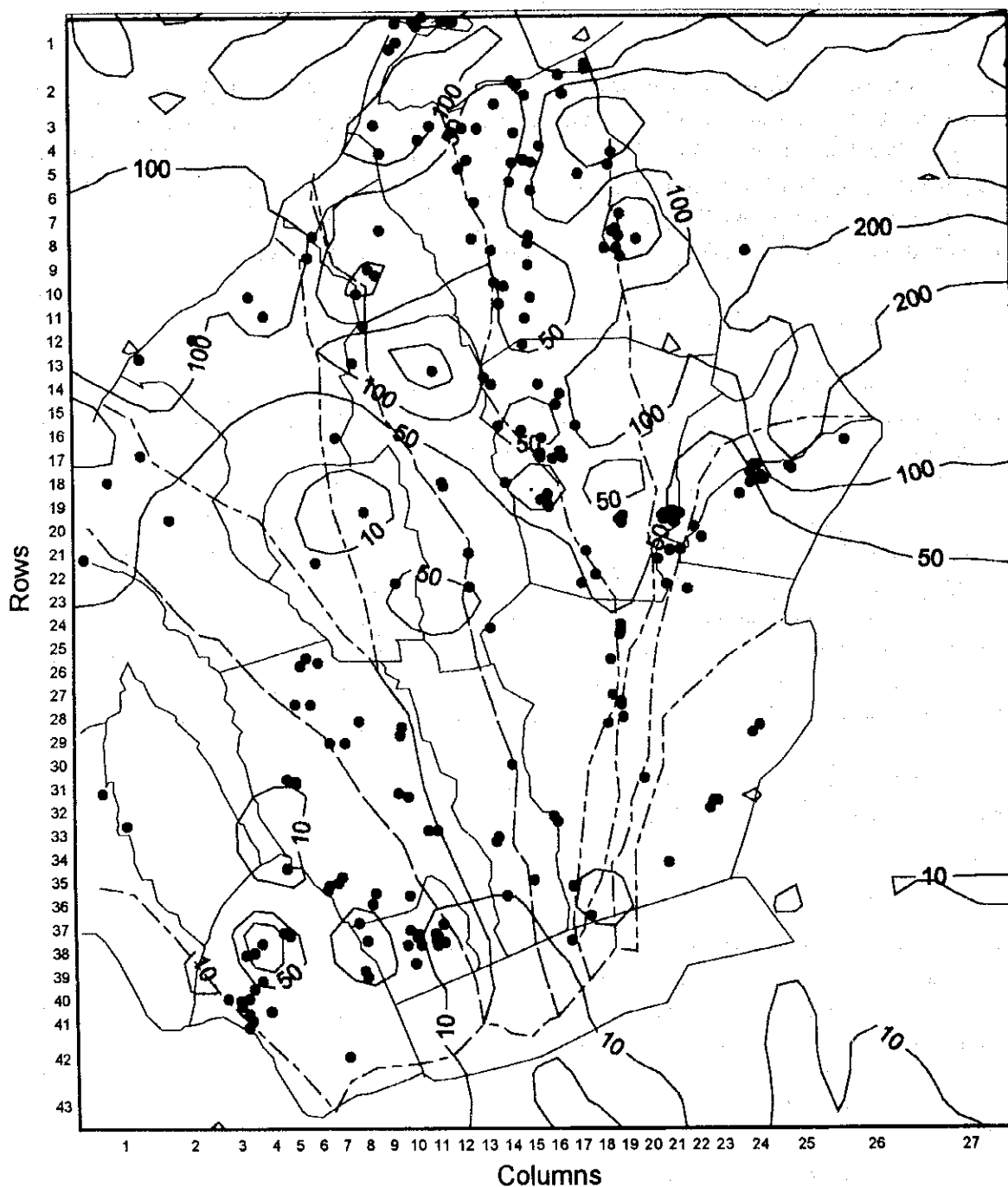
The estimated transmissivity can be called as "apparent transmissivity", because the estimated transmissivity describes the ability of the perforated portion of the aquifer to transmit water. After obtaining the apparent transmissivity of the well by the said methods, the permeability k value was computed based on the definition of transmissivity given by:

$$k = T/b$$

where, b is the thickness of the layer.

Vertical hydraulic conductivity

The MODFLOW needs an input data set of vertical hydraulic conductivity k' to calculate the vertical groundwater flow terms. Here, the vertical hydraulic conductivity of each layer was assumed equal to 10% of the permeability k .



- Line of equal specific capacity (m^2/d)
- Production well having specific capacity value

Fig. 5.5-1

DISTRIBUTION OF SPECIFIC CAPACITY

CAVITE WATER SUPPLY DEVELOPMENT STUDY

JAPAN INTERNATIONAL COOPERATION AGENCY (JICA)

KOKUSAI KOGYO CO., LTD.

Initial piezometric heads

The initial piezometric heads of the three (3) layers for the steady-state calibration were given as the assumed average static water levels in the Study Area in 1985. For the historical calibration period (1994), the computed heads by the steady-state calibration were input as the initial heads.

Recharge rate

The recharge rate expressed as a percentage of annual rainfall varies from basin to basin as shown in Chapter 4. Recharge rate was applied to the upper layer (zones 1 and 2) and the exposed downstream portion of the middle layer (zone 3).

5.6 MODEL CALIBRATION

5.6.1 Preparation of pumpage data

The estimated water use in the Study Area by municipality in 1993 (Table 4.2-7) was assumed equal to the groundwater pumpage for that year. Considering the type of use, the estimated amount for a certain municipality was distributed to the total number of wells in that municipality as registered in the updated PGDB for Cavite. The historical pumpage data were used for the model calibration. Each production well was located in the 1 km \times 1 km grid system using its latitude and longitude coordinates (if the coordinates are not available, the barangay and road map information were used instead). The pumpage of each model grid was computed by summing up the pumpage of each well located in each model grid. The pumpage data covers not only the simulation area but also the Study Area.

For the 3-D groundwater flow model, the vertical distribution of the pumpage was assumed as follows:

Zone 1: 100% for upper aquifer

Zone 2: 1/3 for upper aquifer; 1/3 for middle aquifer; and 1/3 for lower aquifer

Zone 3: 50% for middle aquifer and 50% for lower aquifer.

5.6.2 Steady-state calibration

The steady-state calibration was carried out prior to the historical calibration. The purposes of the steady-state calibration are to understand model behavior, to check boundary conditions, to estimate approximate values of assumed hydrogeological parameters, and to generate initial piezometric heads at the beginning of the historical calibration for each aquifer unit.

Neither recharge nor pumpage were used for the steady-state calibration. A duration of 10 years was taken for the steady-state simulation at a time-step of one (1) year. The initial piezometric heads for the upper, middle and lower aquifers were given as the average measured static water levels in the Study Area.

For the JICA monitoring stations shown in **Fig. 5.6-1**, the computed heads of the three aquifers were almost balance within the ten time-steps of the steady state calibration. The computed heads were then used as the initial heads for the historical calibration.

5.6.3 Historical calibration

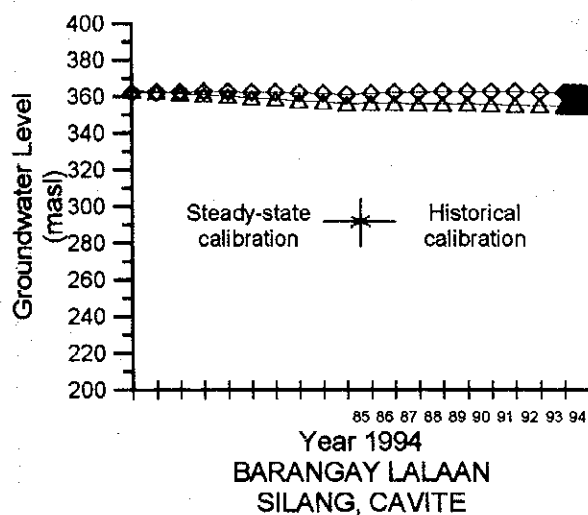
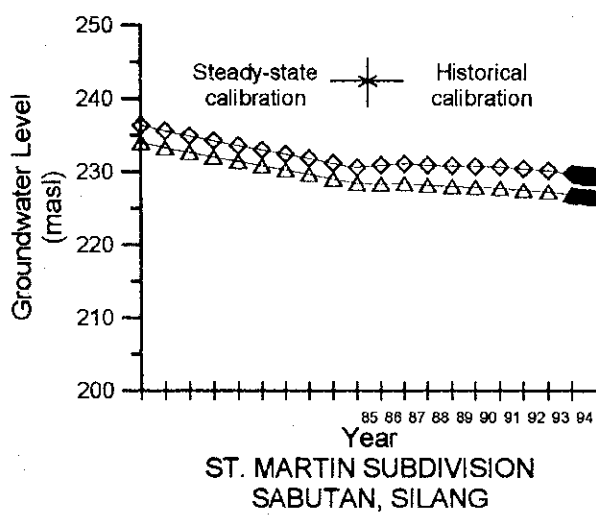
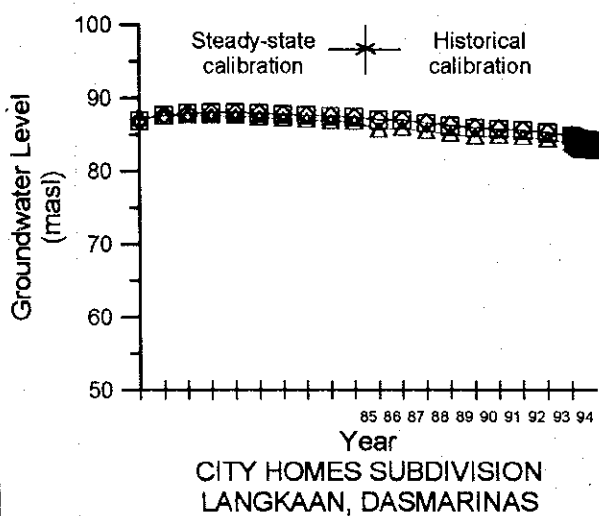
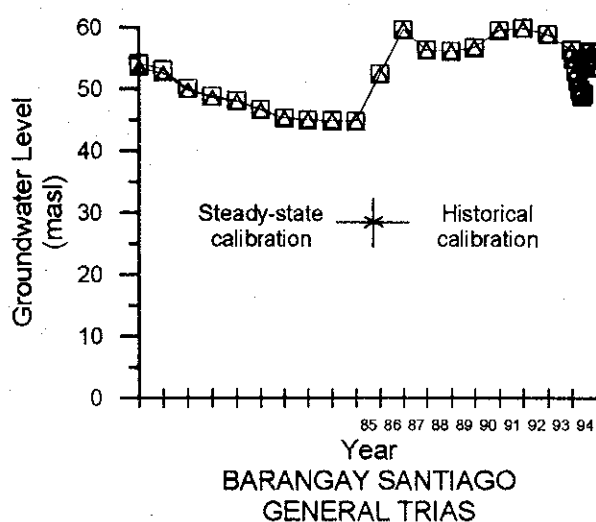
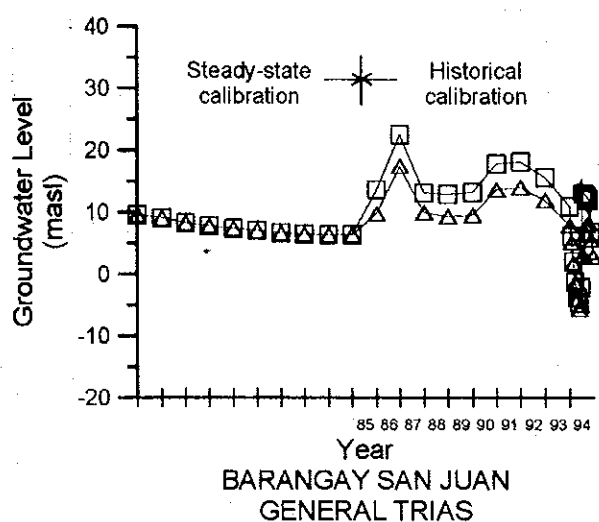
The historical calibration was performed in two stages for a total of 21 time-steps. The first stage was from 1985 to 1993 at one-year time-step (nine (9) time-steps in all) using annual recharge and a constant pumpage rate equal to the 1993 pumpage level, and the second stage consisted of 12 monthly time-steps from January to December, 1994 using monthly recharge and pumpage. The monthly historical pumpage data from January to December, 1994, which was assumed equal to 1994 pumpage amount divided by 12 months, were input to the model. **Fig. 5.6-2** shows the distribution of the 1994 groundwater pumpage in the Study Area. The recharge data were computed from observed rainfall data from PAGASA. During the historical calibration, computed piezometric heads were compared with the observed heads.

From the steady-state calibration to the historical calibration, the hydrogeological parameter values such as hydraulic conductivity, specific storage, etc. were not modified, only the boundary conditions were varied.

After modifying the boundary conditions, the historical simulation was re-run together with the steady-state simulation because the change affected the initial piezometric heads for the historical simulation. Therefore, the simulation of 31 time-steps, which consists of 10 time steps of steady-state simulation (1 time-step = 1 year) and 21 time-steps of historical simulation (9 time-steps at 1 time-step = 1 year and 12 time-steps at 1 time-step = 1 month), was carried out. **Fig. 5.6-1** also shows the results of the 21 time-steps historical calibration at the JICA monitoring stations.

Figs. 5.6-3 to 5.6-5 show the comparison between the simulated groundwater levels and the 1994 groundwater levels as surveyed in the simultaneous groundwater leveling in dry and wet seasons. Within the SJRB, the simulated groundwater levels of the upper, middle, and lower aquifers satisfactorily matched the 1994 observed groundwater levels.

Fig. 5.6-6 shows the comparisons between the simulated and observed piezometric heads at the JICA monitoring stations. The observed piezometric heads were reasonably simulated at the stations in Brgys. San Juan and Santiago, Gen. Trias and Bry. Sabutan, Silang. The significant difference between observed and simulated groundwater levels in the rest of the stations could be attributed to the assumed upstream boundary conditions.



LEGEND

- ◇— Simulated Groundwater Level of Upper Aquifer
- Simulated Groundwater Level of Middle Aquifer
- △— Simulated Groundwater Level of Lower Aquifer

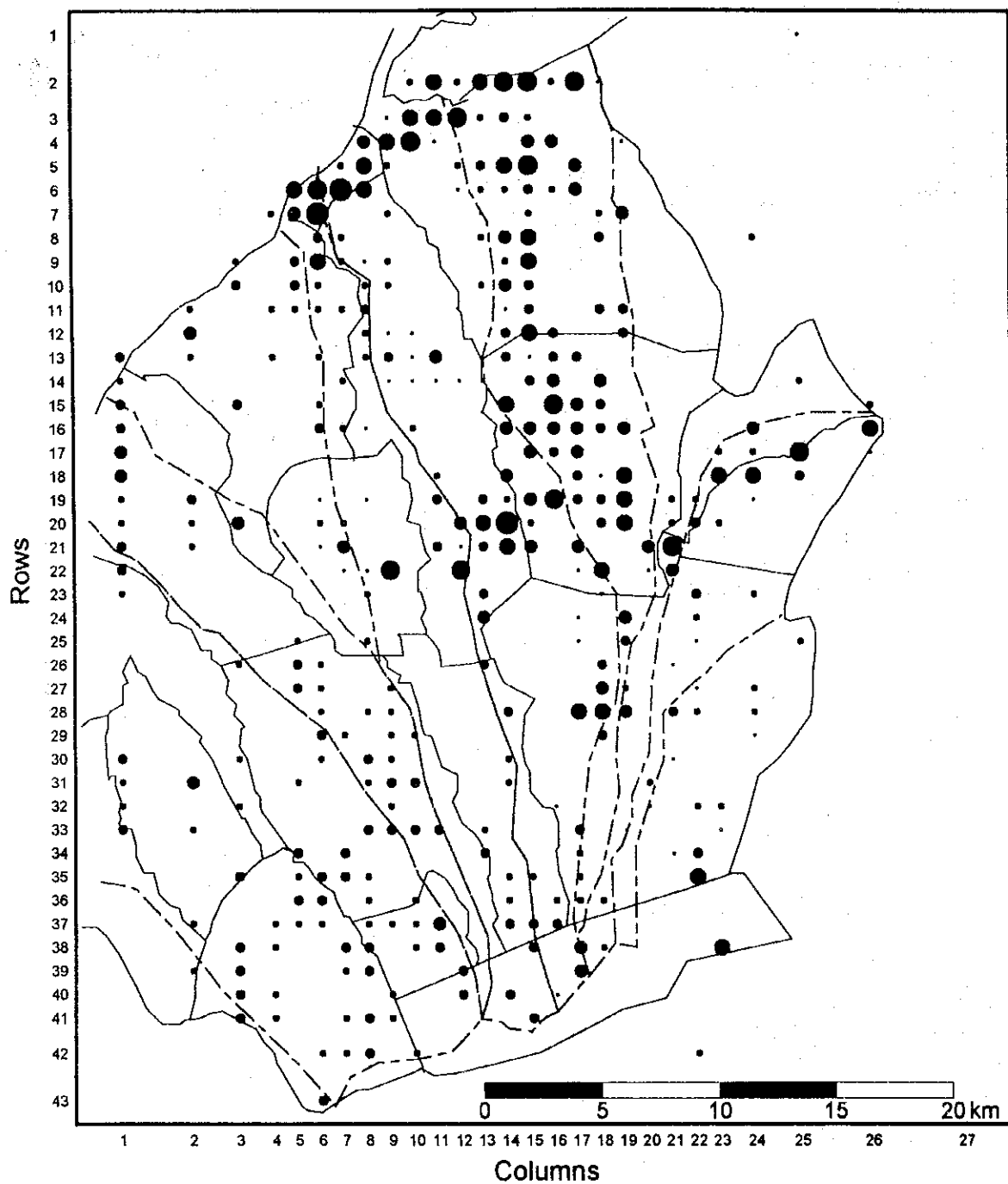
Fig. 5.6-1

RESULTS OF STEADY-STATE AND HISTORICAL CALIBRATION

CAVITE WATER SUPPLY DEVELOPMENT STUDY

JAPAN INTERNATIONAL COOPERATION AGENCY (JICA)

KOKUSAI KOGYO CO., LTD.



Groundwater Pumpage
(m³/day)

- 1.0 to 9.9
- 10 to 99
- 100 to 499
- 500 to 999
- 1,000 to 1,999
- 2,000 to 4,999
- 5,000 to 9,999

Fig. 5.6-2

**DISTRIBUTION OF GROUNDWATER PUMPAGE
IN 1994**

CAVITE WATER SUPPLY DEVELOPMENT STUDY

JAPAN INTERNATIONAL COOPERATION AGENCY (JICA)

KOKUSAI KOGYO CO., LTD.

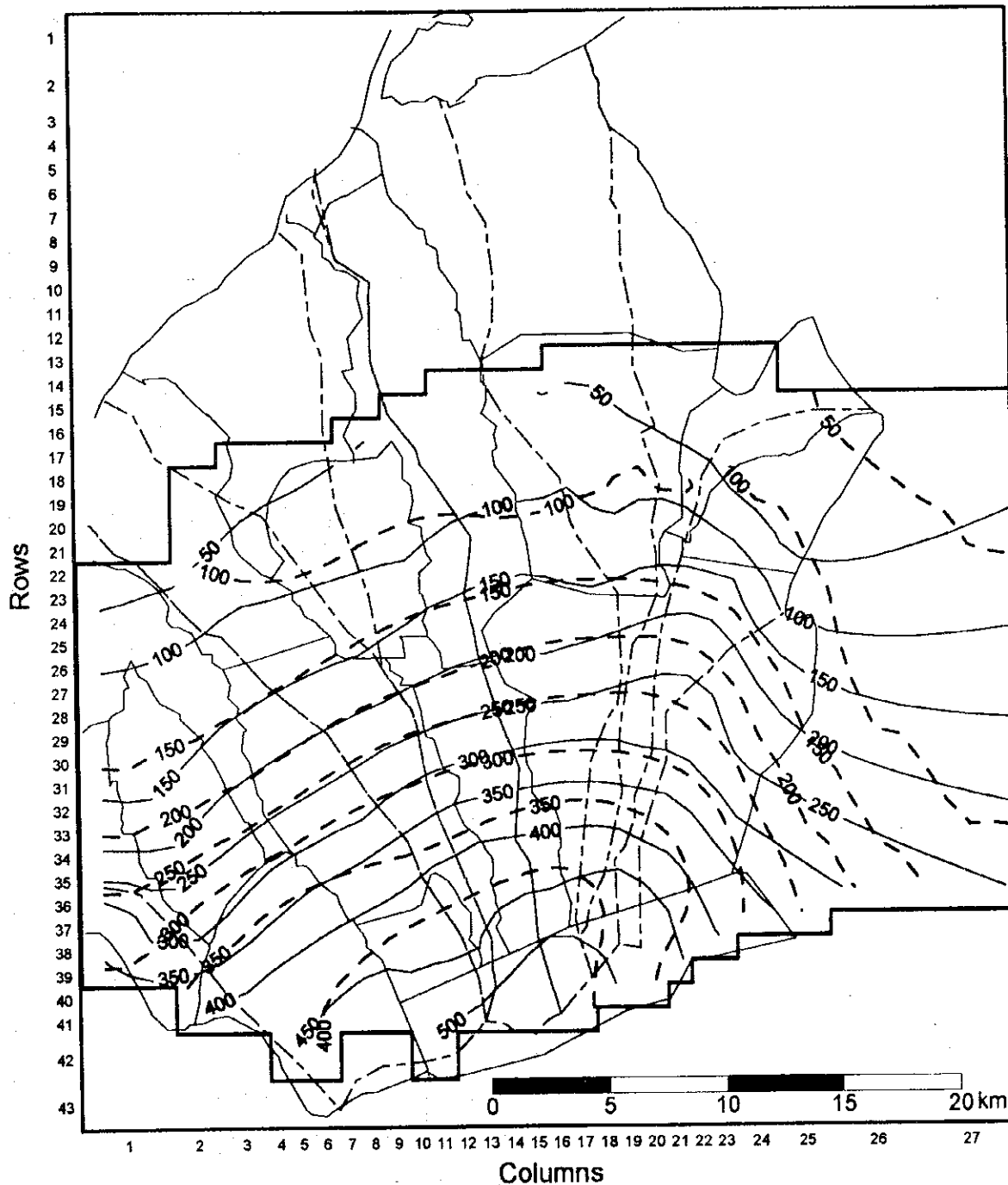


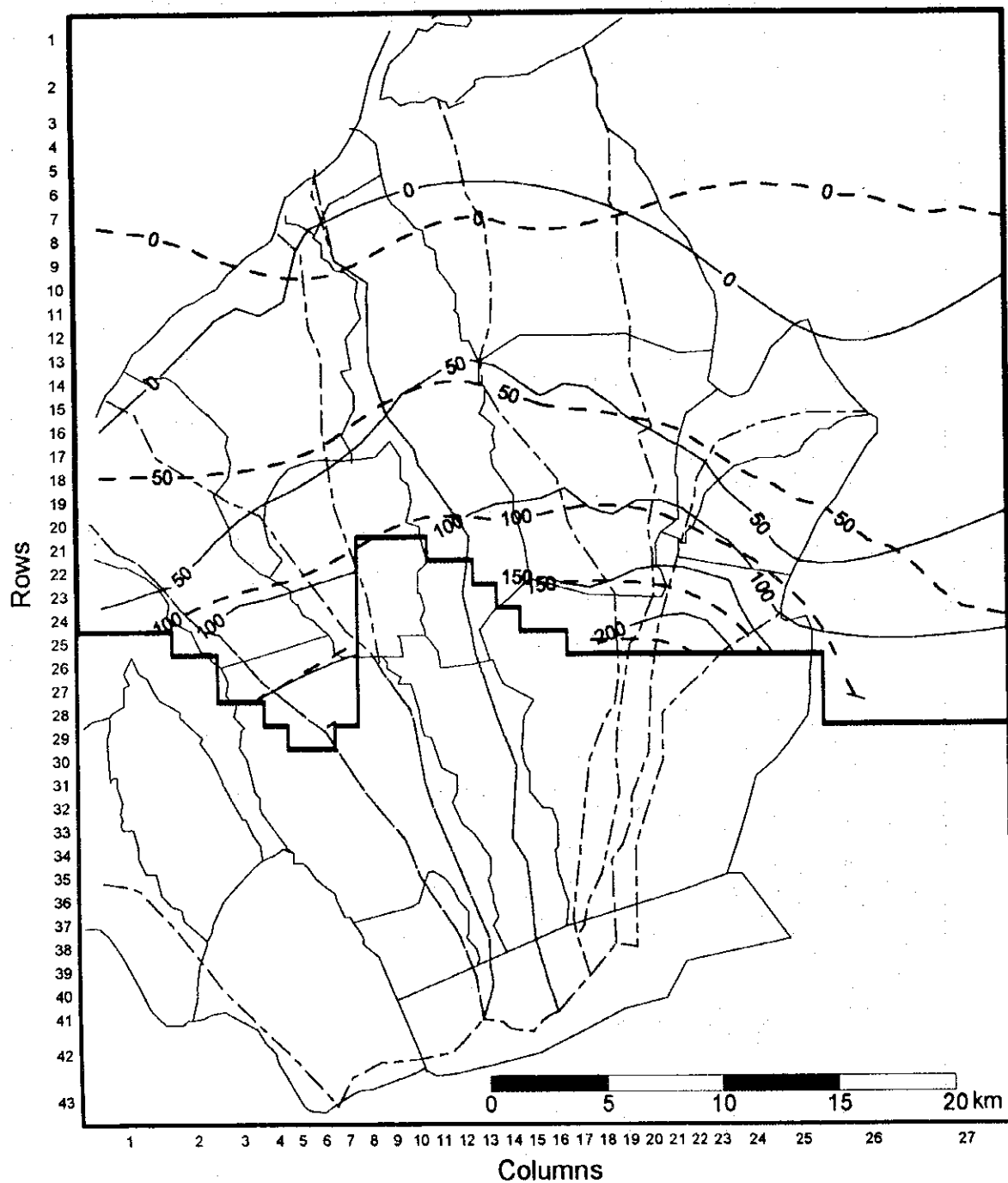
Fig. 5.6-3

**SIMULATED GROUNDWATER LEVEL OF
UPPER AQUIFER**

CAVITE WATER SUPPLY DEVELOPMENT STUDY

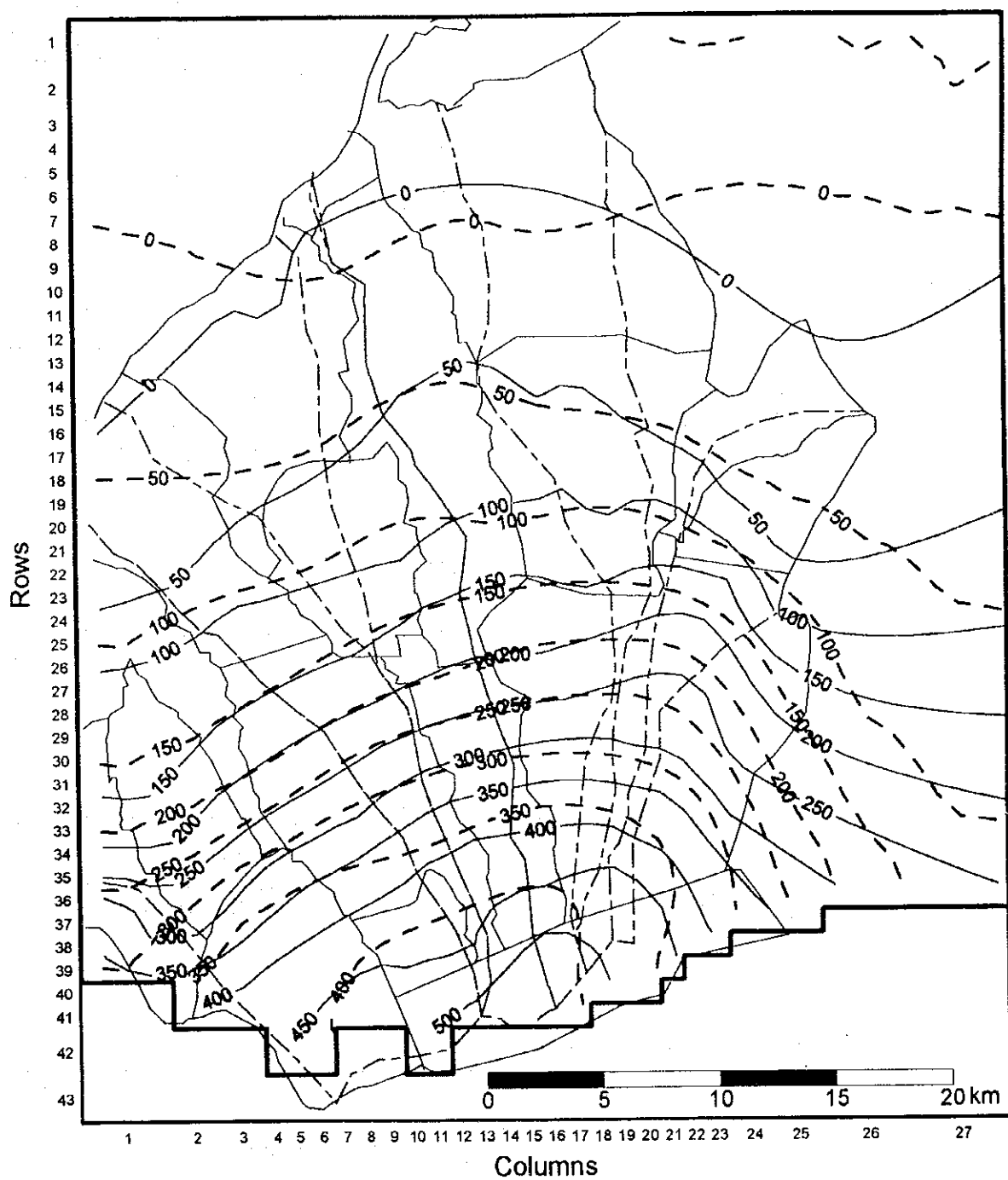
JAPAN INTERNATIONAL COOPERATION AGENCY (JICA)

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- Actual groundwater level in 1994 (masl)
- - - Simulated groundwater level in 1994 (masl)
- Boundary of aquifer

Fig. 5.6-4	SIMULATED GROUNDWATER LEVEL OF MIDDLE AQUIFER
CAVITE WATER SUPPLY DEVELOPMENT STUDY	
JAPAN INTERNATIONAL COOPERATION AGENCY (JICA)	KOKUSAI KOGYO CO., LTD.



- Actual groundwater level in 1994 (masl)
- - - Simulated groundwater level in 1994 (masl)
- Boundary of aquifer

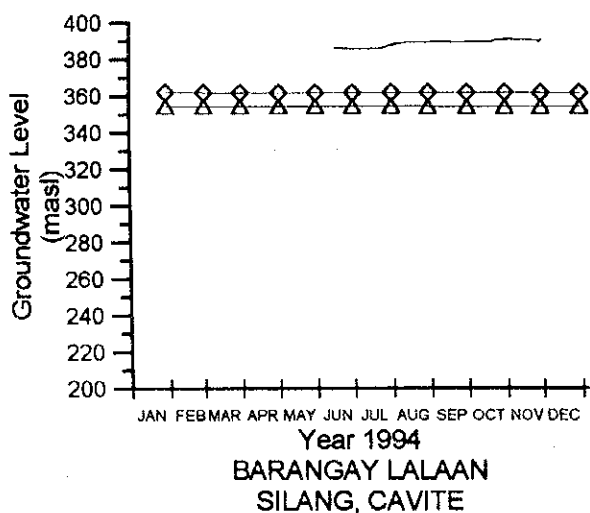
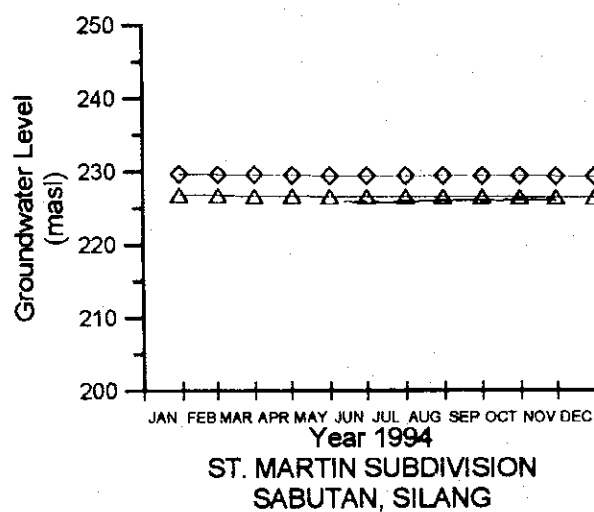
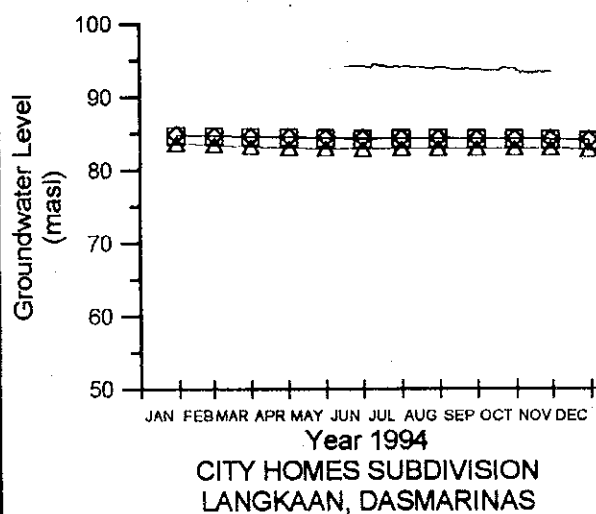
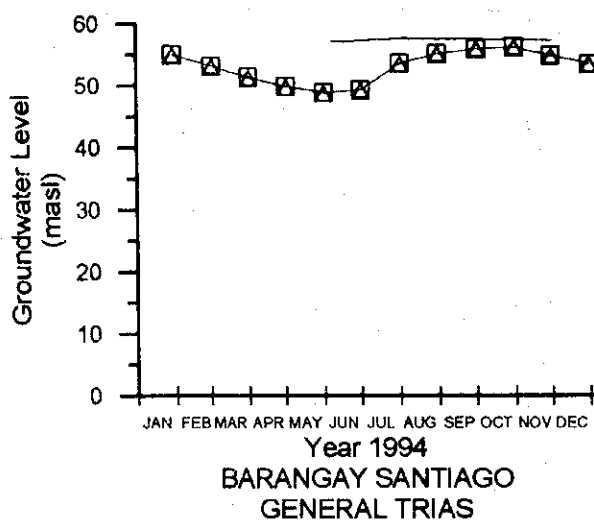
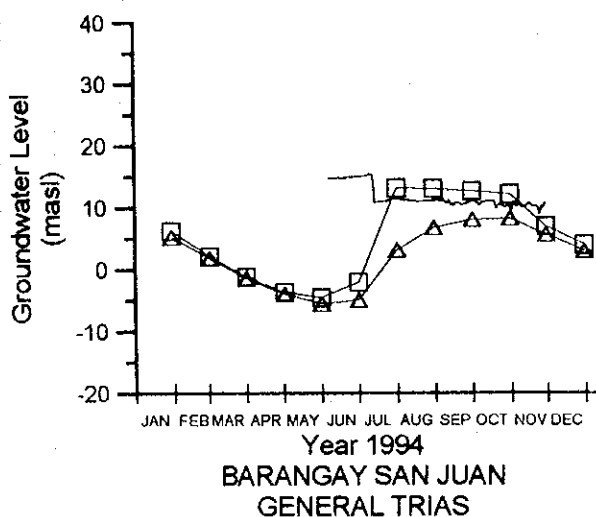
Fig. 5.6-5

**SIMULATED GROUNDWATER LEVEL OF
LOWER AQUIFER**

CAVITE WATER SUPPLY DEVELOPMENT STUDY

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KOKUSAI KOGYO CO., LTD.



LEGEND

- Actual Groundwater Level
- ◆ Simulated Groundwater Level of Upper Aquifer
- Simulated Groundwater Level of Middle Aquifer
- △ Simulated Groundwater Level of Lower Aquifer

Fig. 5.6-6

**SIMULATED AND ACTUAL GROUNDWATER
LEVELS AT JICA MONITORING STATIONS**

CAVITE WATER SUPPLY DEVELOPMENT STUDY

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