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REPORT ON THE MINERAL EXPLORATION IN THE PACHAPIRIANA AREA REPUBLIC OF PERU

(PHASE Ⅲ)



MARCH 1991

JAPAN INTERNATIONAL COOPERATION AGENCY
METAL MINING AGENCY OF JAPAN

国際協力事業団

26046

Preface

In response to the request of the Government of the Republic of Peru, the Japanese Government decided to conduct a Mineral Exploration in the Pachapiriana Area Project and entrusted the survey to the Japan International Cooperation Agency(JICA) and the Metal Mining Agency of Japan(MMAJ).

The JICA and MMAJ sent to the Republic of Peru a survey team headed by Mr. Hiroshi Hama from July 7 to December 7, 1990. The team exchanged views with the officials concerned of the Government of the Republic of Peru and conducted a field survey in the Pachapiriana area. After the team returned to Japan, further studies were made and the present report has been prepared.

We hope that this report will serve for the development of the Project and contribute to the promotion of friendly relations between our two countries.

We wish to express our deep appreciation to the officials concerned of the Government of the Republic of Peru for their close cooperation extended to the team.

February, 1991

Kensuke YANAGIYA

President

Japan International Cooperation Agency

Gen-ichi Fukuhara

President

Metal Mining Agency of Japan

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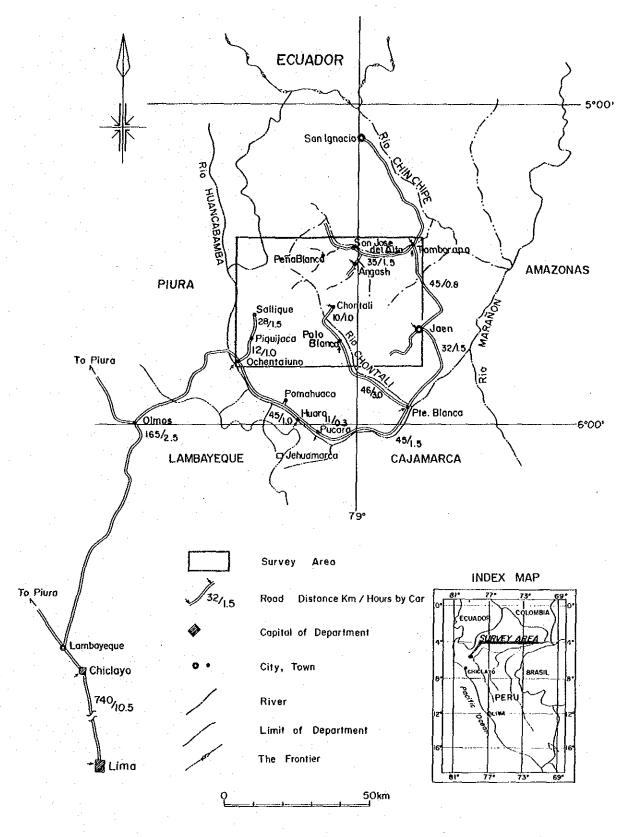


Fig.I-1 (1) Location and Accessibility of the Survey Area

SUMMARY

This report summarizes the third year results of the surveys conducted in the Pachapiriana area, Peru. The surveys aimed to reveal potentiality of an existence of useful mineral resources in the area through clarification of geological setting in the area. The field survey was carried out from July to November, 1990.

The third year survey included drilling survey (16 holes, total length 2334.06m), geophysical survey (gravity survey, 33 km2, 305 points) and detailed geological survey (2.5 km2) combined with geochemical survey (100 rock samples).

The drilling survey was conducted in Chontali (6 holes, total length 1332.51m) and Jehuamarca (10 holes, total length 1001.55m) in order to verify the expected mineralization conditions in the deep underground for large-scaled quartz veins and for silicified breccia and quartz zone, respectively. Geophysical survey was performed in Chontali to clarify the feature of gravity basement, which has been inferred to be closely connected with mineralization. Detailed geological survey was conducted in Jehuamarca to clarify the mode of occurrence of silicified breccia and to get the geological informations assisting the interpretation of drilling data.

Through the gravity prospecting it is clarified that the gravity basement has higher density of 2.8g/cm3, and three higher density zones were extracted above the gravity basement. It has also been clarified that drilling cores from northern most high density zone have undergone distinct carbonatization rich in Mn, Fe and Mg and it makes the density higher. Basement granitic rocks may also undergo regional carbonatization to make the density higher. In this case, the density can be as high as 2.84g/cm3.

As a result of the drilling survey, it has been clarified that quartz veins could tend to plunge. If it is the case, the mineralized zones could also plunge.

The analyzed results by homogenization temperature for quartz veins ranged from 102 to 194°C, and rather lower values are predominant. It is inferred that a zone most adequate for gold mineralization can exist deeper than the depth of altitude 1700m, until which this year survey has reached.

It can be inferred that the mineralization in Jehuamarca was progressed at first as argillization under neutral to alkaline environment, then as acidic mineralization, by fluids ascended through the NE-SW to NW-SE subsidiary fault-fissure system branched from regional fault systems and/or through the bedding

boundary. '

Silicified breccia and quartz vein, which have been expected to associate gold-silver mineralization and auriferous base metal mineralization respectively, are lower graded than expected and discontinuous, therefore it must be concluded that there is a small possibility that they develop to a large scale high-grade ore body.

For the fourth year survey, it is necessary to conduct drilling survey in Chontali to confirm the plunge of quartz veins as well as mineralization zones, and also to be carried out detailed geological survey combining with detailed mapping and systematic sampling of quartz veins in the quartz vein distributed zone.

PART I GENERAL REMARKS

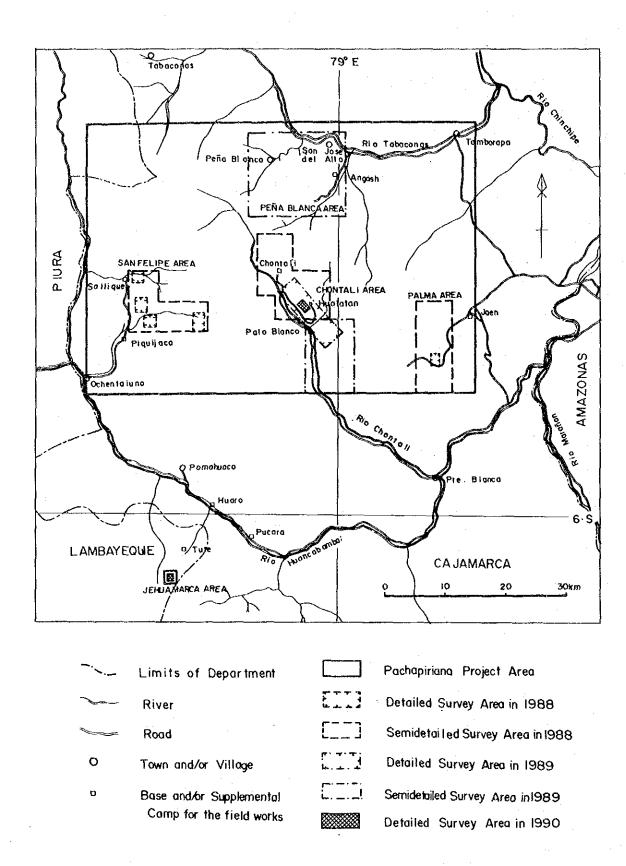


Fig. I-1 (2) Location of the Survey Area

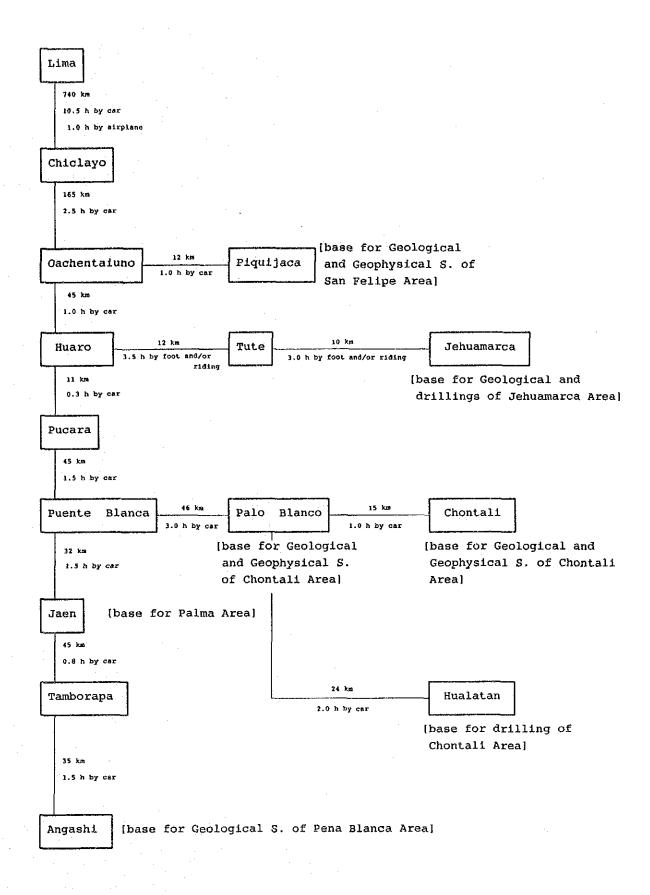


Fig. I-2 Summarized Accessibility of the Survey Area

CHAPTER 1 INTRODUCTION

1-1 Antecedents of the Survey

The survey area is situated in a part of the region for which a geochemical survey using stream sediments was carried out under the Northern Geochemical Project(Proyecto Geoquemico del Norte) sponsored by the U.K. The detailed survey for the extracted geochemical anomalous zones was realized partially by INGEMMET (Instituto Geologica Minero y Metalurgico) itself and partially by German and French organizations. However, the major part has remained pending due to shortage of funds.

Under these circumstances, INGEMMET requested, through the Ministry of Foreign Affairs of the Republic of Peru, a technical cooperation from the Japanese Government for the follow-up survey in March 1988. In August 1988, a delegation for the preliminary survey and agreement negotiations for this purpose was organized among the Ministry of International Trade and Industry (MITI), Japan International Cooperation Agency (JICA) and Metal Mining Agency of Japan (MMAJ), and sent to Peru. On August 15, 1988, the scope of work to the Pachapiriana area Project was signed between the parties.

According to the scope of work concluded among INGEMMET, JICA and MMAj, the survey encompasses an area of 2,820 km2.

The first year survey included a semi-detailed geological survey which was conducted over a total area of 300 km2 in regions among the geochemical anomalous zones assessed prospective by INGEMMET, combined with the LANDSAT image analysis of entire survey area. An additional detailed geological survey was also conducted over a total area of 25km2, in five regions (21km2) extracted from semi-detailed survey area and one region (4km2) extracted by INGEMMET. Geophysical survey using CSAMT method was implemented over 25km2 in two regions. During the second year survey, following the recommendation of the year survey, detailed geological survey (42km2), semi-detailed survey (80km2) and geophysical survey using CSAMT method (35km2) were conducted in Chontali area. Drilling survey was performed at three sites and total hole length reached 816.25m in Semi-detailed geological survey (220km2) Jehuamarca. ducted in Pena Blanca area.

- 1-2 Conclusion and Recommendation of the Second Year Survey
- 1-2-1 Conclusion of the second year survey

The conclusion of the second year survey is summarized as follows:

1) This survey area consists of rocks from Precambrian to Cenozoic.

2) Mineralized alteration occurs in Mesozoic Leche Formation

and Oyotun Volcanics and Cenozoic Porculla Volcanics.

3) The mineralized alteration occurs through the NE-SW trending fault systems and its subsidiary NW-SE trending fissure systems, relating to the skarnization and the epithermal mineralization.

- of 4) The skarnization zones suggest the existence auriferous base metal dissemination type ore deposits and epithermal mineralization zones suggest the existence of epithermal gold and silver vein type and epithermal auriargentiferous base metal vein type ore deposits.
- 5) Geophysical survey implemented in the Hualatan west in Chontali area, where a high possibility of an existence of epithermal gold and silver vein type has been verified, revealed that the alteration zone was characterized by low resistivity zones and that the silicified and silicified-argillized zones with closely developed quartz veins was extracted as high resistivity within the low resistivity zone.
- 6) The high resistivity zone continued toward deeper underground as far as the highest part and overlapped to the part part of large-scaled quartz veins within the area with closely developed quartz veins extracted in Hualatan west.
- 7) Based upon geological setting and the analyzed results by homogenization temperature for quartz veins, the highest resistivity zone is inferred to be connected with the granitic intrusion feature.
- 8) Average width and length of six quartz veins in the area with large-scaled quartz veins are 3.04m and 80m, respectively. The average grade of veins was 2.54g/ton Au, and 13.99g/ton Ag. The analyzed results by homogenization temperature for the large-scaled quartz veins was lower than 150°C.
- 9) In Jehuamarca area, it is revealed that the silicified zone is ubiquitously associated with mineralization of gold, silver, copper, lead and zinc, and that the zone can be characterized by a mushroomed structure as interpreted.
- 10) A layered ore body (quartz zone) of high grade gold, silver, copper, lead and zinc was found in the silicified zone. An existence of high grade gold and silver mineralization zone was confirmed in silicified breccia widely distributed at the surface.

1-2-2 Recommendations of the second year survey

The following surveys are proposed for the third year survey based on the results obtained in the second year:

1) Jehuamarca area

(1) Drilling survey for silicified zone with high grade layered quartz zone and for silicified breccia zone.

(2) More detailed geological survey in the area where the drilling survey will be conducted (with a scale of 1/2000, for example).

2) Chontali area

(1) Drilling survey for Hualatan west zone where a high possibility of an existence of high grade gold deposits has been verified.

(2) More detailed geological survey in the area where the drilling survey will be conducted (with a scale of

1/2,000, for example).

(3) Geophysical survey in the area including Hualatan west to clarify the geological structure of basement rocks (gravity prospecting, for example).

3) Tuna area

- (1) Semi-detailed geological survey for the geochemical anomalous zones extracted by INGEMMET.
- 4) Pena Blanca area
- (1) Detailed geological survey for mineralization zone extracted in Oyotun Volcanics.
- (2) Geophysical prospecting for the mineralization zone (by the CSAMT method, for example).
- 1-3 Outline of the Third Year Survey
- 1-3-1 Area and Purpose of the second year survey

The third year survey was conducted according to priority of recommendations of the second year.

1) Jehuamarca area

Drilling survey was performed at ten sites and total hole length reached 1000m. The purpose of the drilling was to verify the mineralized conditions in underground silicified zone and silicified breccia widely distributed at the surface, which have been confirmed to contain a large quantity of sulfide minerals during the second year phase (recommendation 1-1 of the second year).

The propose of drilling at each site is as follows:

MJPJ-4,5,6,7 and 12; to verify the extension of gold and silver mineralization in silicified breccia extracted at MJPJ-3 and to verify the mineralized conditions in deeply underground silicified zone

MJPJ-8,9,10 and 11; to verify the underground extension of the silicified veins extracted at MJPJ-1 and -2 and to verify the mineralized conditions of deeply underground silicified zone

MJPJ-13: to verify the extension and the mineralized conditions of the silicified breccia associated with a fault extracted at MJPJ-3

In order to assist the drilling survey, detailed geological surface mapping combined with geochemical survey was carried out over a total area of 25km2, using the route map on the scale of 1/2,000 (recommendation 1-2 of the second year).

2) Chontali area

Drilling survey was performed at six holes and total hole length reached 1332.51m in Hualatan west, where a high possibility of an existence of high grade gold deposits has been extracted in quartz veins (recommendation 2-1 of the second year). The purpose of the drilling was to verify the mineralized conditions in the deep underground.

In order to verify the change of mineralized conditions and mineralization temperature in the deep underground, the survey was conducted by vertically fan-shaped method, in which two holes were drilled at each of three sites.

As the rise of high resistivity basement overlap with the distribution of quartz veins in the alteration zones, it is inferred that the basement structure and mineralization are connected with each other. Therefore, gravity prospecting was performed at 305 points to clarify the geological structure of basement rocks (recommendation 2-3 of the second year).

1-3-2 Survey procedure

The procedure of each survey will be outlined as follows.

1) Detailed geological survey

Detailed geological survey was conducted only in Jehuamarca area this year. A topographical map on the scale of 1/2,500 measured by INGEMMET was used as the base map for the survey. The survey was carried out along measuring lines, which were drawn using string measures and pocketcompasses with a scale of 1/2,000. Measuring lines were closed with each other, except for in some part of outer margin of the survey area. The revision of error of closure was made in terms of horizontal distance. The altitude was measured based on the that of altitude of 3230.0m at triangulation station J-2. The revision of error of closure was made in terms of vertical distance.

Trench survey was also performed to confirm the mode of occurrence of silicified breccia and to draw up a trench map on the scale of 1/100. Route maps on the scale of 1/2,000 were compiled into a geological map on the scale of 1/2,500.

2) Geochemical survey

Rock samples for geochemical survey were taken together with geological survey only in Jehuamarca area. Sampling at the outcrop was carried out as a rule every 100m along the measuring line for detailed geological survey. Sampling density was heightened in the silicified and silicified breccia zones.

3) Geophysical survey

Gravity prospecting was performed in Chontali area using LaCoste gravimeter. Measuring points were set on the measuring line network, which covered whole survey area. The survey was carried out as a rule every 100m in the central part and every 200 to 300m in other parts. Total prospecting points is 305.

The gravity value at standard station was confirmed by the comparison with that at existing gravity station, and the accuracy of measurement was less than 0.2 mgal for closing error.

The altitude at each measuring point was determined by leveling using automatic level of Sokki-sha. The accuracy of measurement was less than $200\sqrt{D}(\text{mm})$ for closing distance D km. If the specification of the measuring point is difficult, stadia measurement and/or distance measurement using esron tape were also applied to make the specification accurate.

The measured gravity values for each point were reduced for such factors as tide, topography, latitude and altitude (Bouguer). Isogravitational map for most adequate assumed density was made using the correlation between gravity anomaly and topographic feature, the results of rock density measurement and G-H correlation method.

Based on the isogravitational map, using Filter analysis and two dimensional analyses, the results of gravity prospecting was indicated on plane figure and cross section, each of which were on the scale of 1/10,000.

4) Drilling survey

Drilling was performed in Jehuamarca and Chontali areas taking the local drilling company GEOTEC, S.A. into employment.

Jehuamarca area

As no roads were passable for motorcars, drilling rig was transported by a helicopter chartered from the Third Peruvian Air Force. The base camp was set up on the summit of Mt. Jehuamarca. The supply bases were set up in Pucara, Huaro and Tute to replenish mending parts, fuel oil, mud, cement and food. Due to the steep configuration, it took two days on outward way and one to two days on homeward to transport the materials from Huaro to Jehuamarca.

Chontali area

The base camp was set up at Hualatan to replenish mending parts, fuel oil, mud, cement and food. Two transit bases were set up at the junction to the national road in Puente Blanco and at the edge of the bridge with a load limitation in Pena Blanca.

5) Quality of the survey

The survey of this year is summarized as follow:

Quality of the Survey

AL IN SEC. 344 AN AN AN AN AN AN		ological hemical	l and survey	geophysical survey		drilling survey	
, 	area km2	route length	geochemical samples			hole No.	hole length m
Jehuamarca	2.5	22.176	100			10	1,001.55
Chontali				33	30	6	1.332.51
total	2.5	22.176	100	33	30	16	2.334.06

The rock samples taken for laboratory test are summarized as follows:

Quality of geochemical samples

			X-ray diffraction		polished section	density
Α	12		23	5	2	•
В	153		5	4	5	
C						22
D	162	12	20	16	15	

- A: detailed geological survey in Jehuamarca
- B: drilling survey in Jehuamarca
- C: gravity prospecting in Chontali
- D: drilling survey in Chontali

1-3-3 Organization of the survey group and period of the survey

The representatives from Japanese government for the third year survey and agreement negotiation were dispatched to Peru during the period from July 3 to 6, 1990. The delegation members and their counterparts in Peru are shown below:

From Japan:

Mr. Hideya METSUGI,

Metal Mining Agency of Japan

From Peru:

Mr. Guillermo BALCAZAR RIOZA INGEMMET Mr. Gregorio FLORES NANES INGEMMET

The representatives from Japanese government for the deliberation of the survey progress and future development were dispatched to Peru during the period from November 22 to 26,1990. The delegation members and their counterparts in Peru are shown below:

From Japan:

Mr. Katsumi YOKOKAWA ; Metal Mining Agency of Japan Mr. Hideya METSUGI ; Metal Mining Agency of Japan

From Peru:

Ms. Juana Rosa Del CASTILLO ; INGEMMET Mr. Rafael Del AGUILA Del AGUILA ; INGEMMET Mr. Nestor CHANCON ABAD ; INGEMMET

The survey group was organized by geological and geochemical survey team, geophysical survey team and drilling survey team, the first being sent during the period from July 9 to December 7, 1990, the second from July 9 to August 24, 1990 and the third from July 9 to November 16, 1990.

The group members from Japan and their counterparts from Peru are as shown below:

From Japan:

Mr. Hiroshi HAMA ; leader of the survey team MINDECO Mr. Kazuhiko KINOSHITA; geophysical survey MINDECO Mr. Mitsuyoshi SAITO ; geophysical survey MINDECO Mr. Haruo HARADA ; drilling survey MINDECO

MINDECO: Mitsui Mineral Development Engineering Co., Japan

From Peru:

Mr. Carlos JIMENEZ VELASCO; general review and INGEMMET drilling survey(Jehuamarca)
Mr. Carlos A. GAMARRA ROMERO; geophysical survey INGEMMET
Mr. Walter PARI PINTO; geophysical survey INGEMMET
Mr. Luis QUISPE ARANDA; drilling survey(Chontali)

CHAPTER 2 GEOGRAPHY OF THE SURVEY AREA

2-1 Topography and Drainage System

The third year survey was conducted in the central part of

the survey area as shown in Fig. I-1(2).

In Chontali area, the Chontali (called as the Huallabamba or the Chunchuca, on different topographic maps) runs along the west boundary southward in the direction of NW-SE. Each tributary tends to meet nearly at right angles to the Chontali, namely they are arranged in the direction of NE-SW.

In Jehuamarca area, the semi-basin opened northwestward develops and the drainage system is dendritic, the main river of

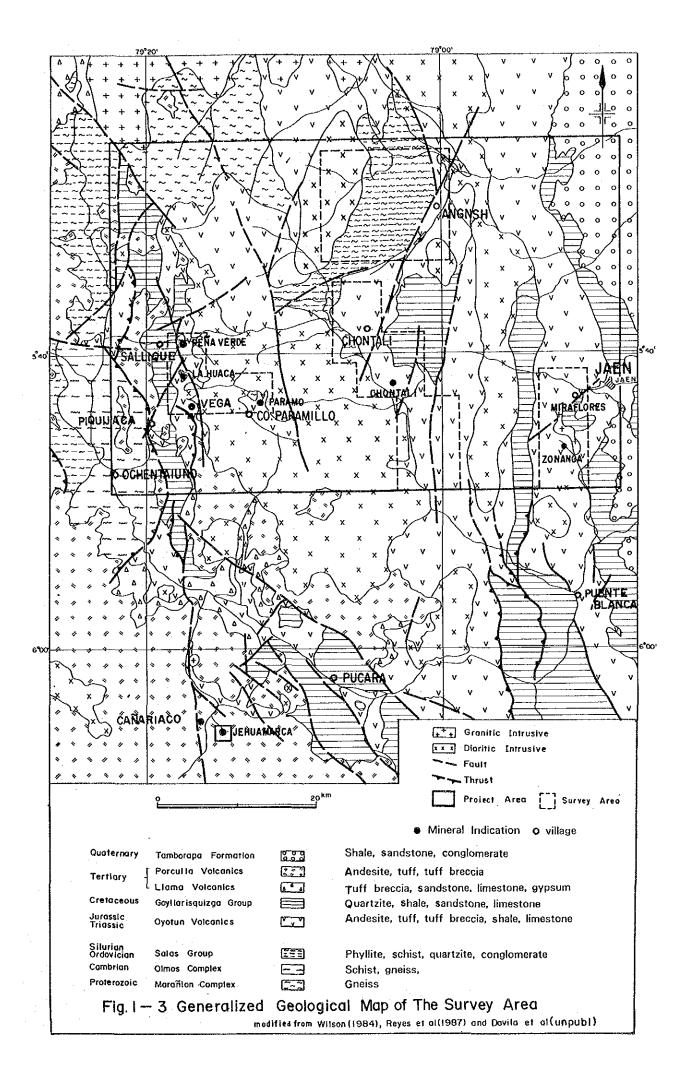
which runs northwestward.

Both of the areas are topographically so steep. In Chontali area, the elevation is 1650m between the Chontali(1050m above the sea level) and the highest peak(2700m). In Jehuamarca area, though the area is narrow, the elevation is as high as 320m between the river-side in the central west (3085m) and the highest peak in the central north (3405m). Cliffs are so steep with the average slope angle more than 60° in the north and east that climate is locally changed.

2-2 Climate and Vegetation

As reported last year, vegetation in the survey area show a particularly remarkable variation due to the altitude difference. Another critical factor for the variation is whether the land is reclaimed or not. The area belongs essentially to tropical to subtropical rainforest zones characterized by thick virgin forests with trees of 10 to 15m high. In the progress of disaforestation, the desert has been expanded but the irrigated area are preserved as agricultural lands but non-agricultural lands become dry shrubbery zones.

In Chontali, whole survey area has been reclaimed to be used as agricultural or pasture lands. In Jehuamarca, whole survey area has not yet been reclaimed due to the cool climate. It is covered with paja (needle-like grass) and is used as natural pasture land. Northern, eastern and almost of southern parts are covered with virgin forest, thus the survey is very difficult, Although climate in the surcoupled with the steep topography. vey area seems to show remarkable local variation, it is difficult to systematize the variation because meteorological observation stations are few and meteorological data are poor. Temperature shows a remarkable variation due to the altitude. The average values decrease as the altitude becomes higher. The average temperature variation for each month is so small that it can be said the seasonal variation of temperature is slight. However, it is inferred that the diurnal variation is very large, as the difference between maximum and minimum temperatures for



Geo	lgical Age	Stratigraphic Unit					Geological Event		. y (11)
		Formation	Columnar Section	Lithology	Thick- ness	Area	Tecto- genesis		Minerali zation
	Holocene	Alluvium		sand, gravel	(m)				
Quat	Pleistocene	Tamborapa F		-552 cgl -=5h->	150				_
	Pliocene	Shimbe V.	V V V V V	and by	200				Blanca
Tertiary	Miocene	Porculla V.		and flbr 11.	600	Jehuamarca		(x)	Peña Bla
	Oligocene { Palaeocene	Llama V.	\\ \a\\\ \a\\\\\\\\\\\\\\\\\\\\\\\\\\\	1f- br -1s ss	500		dean	\(\hat{x}\\ \\x\\\ \\\ \\\\\ \\\\\\\\\\\\\\\\\	Palma Chontali
	Upper	Pulluicana F.		and iv	150		Ą	(+\frac{1}{2})	
snoac		Pariatambo F. Chulec F. Inca F.		is 11-br mri is 55	330 ?	tali		E L	San Felipe,
Creta	Lower	GoyHarisquizga G.		55 55 51 51 51 51 51	525 ?	Sout	1	$\sqrt{(\cdot)}$	
	Jurassic	Oyotun F.	Δ Δ Δ Δ Δ Δ Δ Δ Δ Δ Δ Δ Δ Δ Δ Δ Δ Δ Δ	tf. Is at sh tf.br ss and - Iv tf.br ond - Iv sh abo	1500	T Cha		granodio	
Trids-	Upper	Leche F		and-ty ts	100+		ercynic an(?)	diorite	
_	Silurian	Salas G.		phy 55 ff	1400+	j	T Loga		
	Ordovician	Olmos Cpx.	~ ~ ~ ~ ~	sch gn_	?		\$ 8		
Pre	cambrian	Marañon Cpx.	~ 0~0~0~	gn.	?				
an ac br cg	d andesite d acidic breccia l conglome	gp Is erate Iv	lava	mri phy qt	marl phyllite quartzit		sh ss tf	shale sandstone tuff	
	Dred on or	Miocene Oligocene Palaeocene Upper Jurassic Upper Silurian Ordovician Precambrian Abbreviations and andesite acd acidic br breccio cgl conglome	Miocene Porculla V. Oligocene S Lama V. Palaeocene Upper Pulluicana F. Pariatambo F. Chulec F. Inca F. Lower Goyllarisquizga G. Upper Leche F. Silurian Salas G. Ordovician Olmos Cpx. Precambrian Marañon Cpx. Abbreviations. and andesite acd acidic gp br breccio Is cgl conglomerate ly	Miocene Porculla V. Oligocene Palaeocene Upper Pulluicana F. Pariatambo F. Chulec F. Inca F. Oyotun F. Silurian Salas G. Ordovician Olmos Cpx. Precambrian Marañon Cpx. Abbreviations. and andesite gn gneiss acd acidic gp gypsum br breccia Is limestone cgl conglomerate ly lava	Miocene Porculla V. Oligocene Separation V. Palaeocene Upper Pulluicana F. Pariatambo F. Chulec F. Inca F. Inca F. Silurian Salas G. Ordovician Oligocene Silurian Salas G. Ordovician Oligocene Salas G. Salas G. Ordovician Oligocene Salas G. Salas	Miocene Porculla V. ANA ANA III. BY OND III. SS OND II	Miocene Porculta V. And And It. br. and IV. an	Miocene Porculla V. Oligocene Separation of the property of	Milocene Porculla V. Oligocene Separation of the proculla V. Diligocene Separation of the procullation of the proculation of the proculat

Fig. I -4 Generalized Stratigraphic Column of the Survey Area

each month is more than 10°C.

Relative humidity tends to be higher as the altitude becomes high in contrast to the case of temperature. Judge from the average humidity for each month, the variation coincides with that of average temperature, independently to dry or rainy period.

The precipitation is not dependent on the altitude but fairly on the presence of virgin forest in the surrounding.

Annual variation of precipitation is not so remarkable and there is not a clear distinction between the rainy and dry periods. If a month belonging to the rainy period is defined as with more than 100mm precipitation, the rainy period covers December to March and the dry one June to November. Among the periods, January to April and July to September are relatively much and less precipitated, respectively.

In Jehuamarca, no meteorological data are available. The whole survey area is always covered with nimbus grown by an ascending current to characterize a typical local climate. The current is formed when the atmosphere with high humidity supplied from northern and eastern parts covered with virgin

collides with steep cliffs in northern and eastern parts.

CHAPTER 3 GENERAL GEOLOGY

The survey area is situated in a southernmost of the tectonically disturbed zone, so called Huancabamba deflexion zone. This causes a great variation of rock facies of each geological unit among the former surveys. In addition most of the survey area is uncivilized and the accessed of the survey routes is not easy, thus bringing some confusion to correlate geological formation and it requires further study.

An geological outline of the area is given below after Wilson(1984), Reyes y Caldas(1987) and Davila et al(unpublished), revised after the results of first and second years geological survey. A geological map and a stratigraphic column are generalized as shown in Figs. I-3 and I-4, respectively.

The survey area consists of metamorphic rocks correlative with Precambrian to Paleozoic, Mesozoic sedimentary and volcanic rocks and Cenozoic volcanic and intrusive rocks.

The metamorphic rocks, consisting of the basement Maranon Complex, Ordovician Olmos Complex and Silurian Salas Group, have such a wide lithofacies as gneiss, schist and phyllite. They are developed in the western half of the survey area.

Mesozoic rocks are the main constituent of the survey area and consist of the following units in ascending order: Leche Formation (mainly calcareous rocks), Tinajones Formation (arenaceous rocks intercalated with tuffaceous rocks), and Inca, Chulec, Pariatambo and Pulluicana Formations (mainly calcareous rocks).

Cenozoic rocks composed mainly of volcanic rocks which in ascending order are Llama, Porculla and Shimbe Volcanics, distributed in western and southwestern parts of the survey area. Tamborapa Formation consists of conglomerate, with loose consolidation, being correlative with the Quaternary sediments. This Formation occurs at the eastern flange of the survey area.

Intrusive rocks consist of gabbros, diorites and granites. Generally, gabbros and diorites are older than granites which are intruded even Porculla Volcanics.

The absolute ages determined using K/Ar method are 119±6 million years for quartz diorite, 106±5 million years for quartz monzonite, 82.5±4 million years for granodiorite, 78±3.9 million years for monzonisyenite and 47.6±2.4 for adamelite. The intrusive trends are tend to be NW-SE and N-S, reflecting the geological structure.

Geological structure of the survey area is characterized by its situation of which located at the south flange of a distorted zone of general Andean Trend. This distorted zone, so called the Huancabamba Deflection Zone, corresponds to the area at which the general NNW-SSE direction, the basic characteristic of the Andes, changes direction to the NE-SW trending of the Colombia-Venezuela area. This deflection zone is assumed to have been formed during the Mesozoic tectogenetic movement.

Two combined fault systems are observed in the survey area.

One is E-W with NE-SW trending caused by an east-west lateral compression, and another is N-S with NNW-SSE trending by northwest-southeast one. Both these systems reflect the tectonic movement at the time when the Huancabamba Deflection Zone was formed.

CHAPTER 4 SURVEY RESULTS, COMPREHENSIVE ANALYSIS

4-1 Geological Structure, and Characteristics and Controlling Factors of the Mineralization

This survey area is regionally located in the Huancabamba Deflection Zone, that is, the area belongs to a distorted zone of the general Andean trend in NW-SE changes to NE-SW trending.

1) Chontali area

It was concluded by last year survey that alteration and quartz zones occur on the secondary branched NW-SE fissure zone developed between the regional NE-SW fissure systems. As a result of microscopic observations for drilling cores taken through this year survey, it is concluded that the tectonic movement has continued even after the formation of quartz vein, because quartz grains in quartz vein and in fracture zone extended from the quartz vein commonly show undulatory extinction. Moreover, as fracture zone without quartz vein is in large scale and continuous, and a large-scaled quartz vein is pinched out (MJPJ-2) or branched (MJPJ-4) in fracture zone, it is assumed that quartz vein develops in fracture zone controlled by the structure of fracture zone (unidentified up to now). In other words, quartz vein could tend to plunge, and if it is the case, the mineralization zone could also plunge.

As ore minerals containing such valuable metals as gold, silver, zinc, copper, lead and so on are closely associated with each other, the mineralization environment is inferred to be xenothermal. It is worthy of note that the alteration is characterized by the occurrence of carbonate minerals suggesting the neutral to alkaline environment and that the minerals are composed mainly of such heavy metals as Fe, Mn, Mg and so on.

2) Jehuamarca area

Though distinct fissure systems trend NE-SW and NW-SE, the displacement along the systems are a few to ten and a few meters. Additionally, the fissure systems are other discontinuous. Therefore, they are concluded to be the secondary or third systems branched from major tectonic systems. Along the fissure systems, it is schematically assumed that southern and eastern sides have been displaced downward.

It can be inferred that the mineralization in Jehuamarca was commonly characterized by argillization under neutral to alkaline environment, then overlapped acidic mineralization by fluids ascended through the above-mentioned fissure systems and/or through the bedding boundary. This mineralization process could make the mineralization zone to distribute as stratiform.

4-2 Potentiality of an Existence of Ore Deposits

The survey area contains numerous geochemical anomalous zones with using the stream sediments which were previously sampled by INGEMMET in the "Proyecto Geoquimico del Norte" and the "Proyecto Integral Chinchipe". Semi-detailed geological survey has been continued to implement up to last year to find out to what origin these anomalies were attributed. As a result, obvious mineralized alteration was verified in the backland of the rivers where geochemical anomaly is marked. Therefore, we can conclude that the geochemical anomalies extracted by INGEMMET suggest an existence of mineralized indications.

While no economical ore deposits have been found as yet, the follow-up study of the aforementioned "Proyectos" verified that the existence of La Granja ore deposits and Canariaco ore deposits. In addition, it is reported that a stratiform gold deposit and an epithermal gold deposit have been found in the northern part of this survey area by the "Proyecto Integral Chinchipe: Cordillela del Condor" which is currently under way. It is certain that the area contains promising mineralized area.

Detailed geological survey, which was conducted last year for the Chontali alteration zone extracted in the first year survey, has confirmed the existence of predominant silicified alteration zone with abundant quartz veins and veinlets. Moreover, the geophysical survey conducted last year confirmed a three-dimensional distribution of mineralized alteration zones with quartz veins. Drilling survey performed this year has confirmed the underground fracture zone in large scale (yet unidentified on the surface) and native gold is extracted from the quartz vein in the zone, therefore it becomes more certain that the area contains promising mineralized area.

Drilling and trench surveys performed this year for Jehuamarca alteration zone, where three-dimensional distribution of mineralized alteration zone had been assumed up to last year, have confirmed the frequent existence of distinct mineralization zone, though in small scale. As affairs now stand, the whole average grade of the mineralization zones is relatively low, there is a small possibility that it develop to a large scale ore body, but it still has a potential for resources in the future or a small-scaled mining.

4-3 Relation between Geochemical Anomaly and Mineralization

Geochemical anomalies obtained up to last year was assumed to be classified into driving from mineralized alteration and indicating horizons. The latter is inferred to correspond to the anomalous values found by geochemical survey of the first year, whose distributions were detected in the Salas Group, Chontali. It is clarified, however, by geochemical survey conducted in Jehuamarca this year that zinc anomaly exists along the andesitic sheet in the southwestern part, suggesting the close connection

of geochemical anomaly with a specific rock. In this connection, andesite itself contains zinc as high as 5045ppm (two specimens, arithmetic mean).

Geochemical anomalies obtained this year tend to be distributed overlapping to or around the silicified breccia and fis-

sures, except for a part of zinc anomaly mentioned above.

Due to the weathering, silicified breccia tends to contain no sulfide minerals. Through the preparation of the drilling site of MJPJ-5, however, a sulfide ore body with a large quantity of sphalerite associated with tetrahedrite at the base of silicified breccia has been extracted. Thus, gold and silver mineralization zone now extracted is originally derived from auri-argentiferous base metal mineralization zone. It is concluded, therefore, that geochemical anomaly in this area is closely connected with original mineralization. Regarding the geochemical anomalies distribution of each element, gold anomaly tends to be distributed in the central part and silver, lead and silver anomalies surround it in order outward, to reflect the thermal gradient during the mineralization.

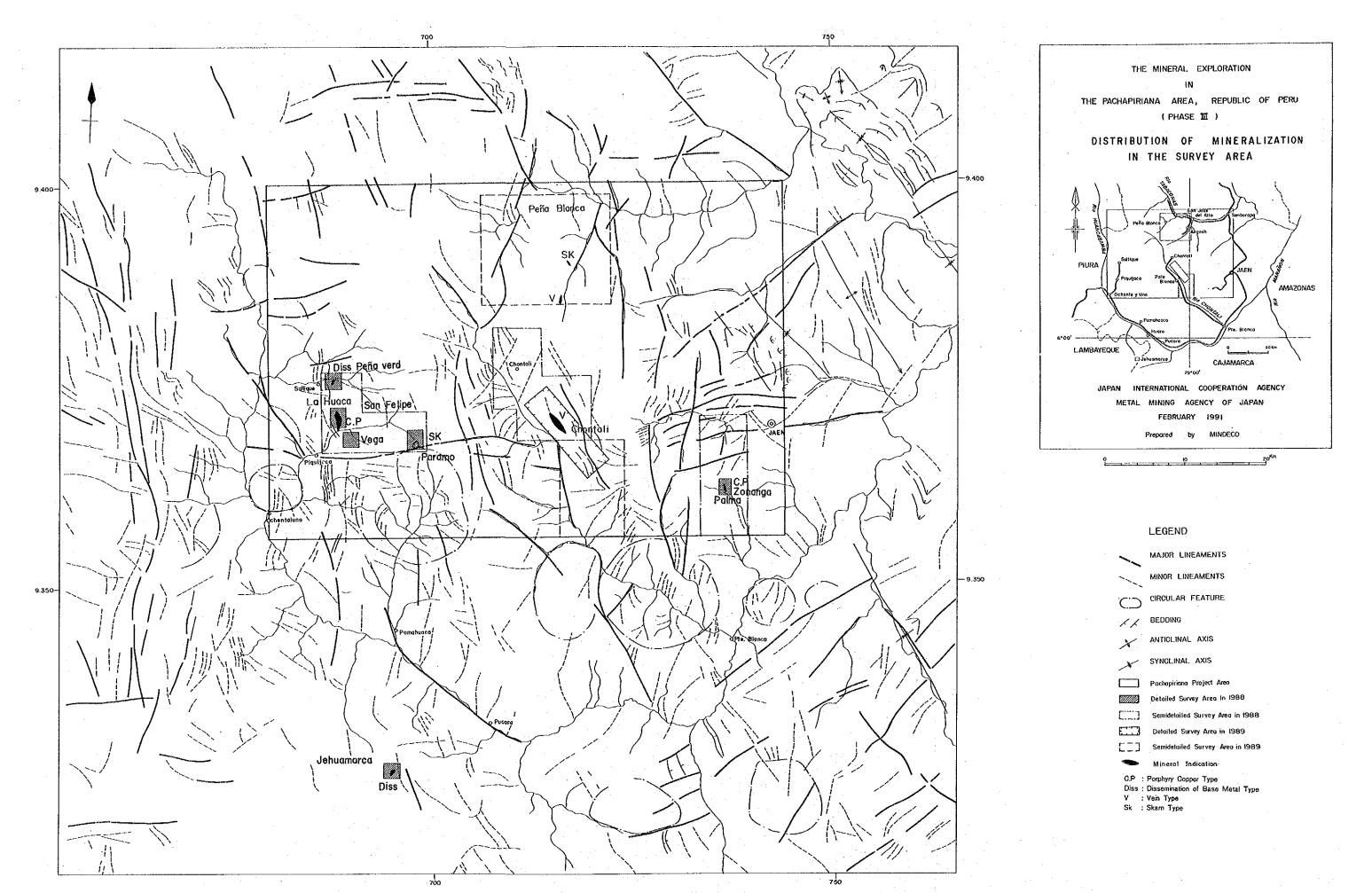


Fig. I -5 Distribution of Mineralization in the Survey Area

CHAPTER 5 CONCLUSION AND RECOMMENDATION

5-1 Conclusion

1) Chontali area

It is clarified, through the gravity prospecting that the gravity basement with the density of 2.8g/cm3 is widely developed underground. The value is rather higher than observed density (2.68g/cm3) of granitic rocks on the surface which is inferred to constitute the basement. Meanwhile, it has been clarified through this year drilling survey that drilling cores have undergone distinct carbonatization rich in heavy metals, which could not been extracted through the surface survey. The arithmetic mean density of drilling core and the specimen taken from surface is 2.71g/cm3 and 2.56g/cm3 respectively, namely the former is 5.9% higher than the latter. If it is assumed that basement granitic rock has also undergone regional carbonatization, its density can be as high as 2.84g/cm3. It, therefore, is possible to suppose that the high density basement is composed of carbonatized granitic rocks. As compared with the distribution of high resistivity zone extracted through last year geophysical survey by CSAMT method, the high density basement (gravity basement) is, though showing similar undulation, rather shallower. It is solely due to the lack of control points to regulate the relative depth of gravity basement. Among the high density zones just above the gravity basement, the northernmost one is just situated in the area for drilling survey this year. just situated in the area for drilling survey this year. The area extracted as high density can almost coincide with localized zone such as highly carbonatized mineralization-alteration zone. Drilling survey was performed this year based on the assumption that the enriched mineralization zone on the surface can continue directly under the outcrop. Quartz vein extracted by drilling survey exists in breccia (fracture) zone in which distinct undulatory extinction, the zone quartz grains show having been confirmed to continue even when quartz vein has been pinched out. It is suggested that quartz vein could tend to plunge as it is pinched out or branched within the fracture zone. If quartz vein tends to plunge, it must be interpreted that the enriched mineralization zone could also plunge. Although the analyzed results this year shows that quartz vein is lower grade than expected, native gold grain is confirmed in one specimen by microscopic observation to suggest the possibility of an existence of high grade gold mineralization zone. Moreover, ore minerals containing such metals as silver, copper, zinc and lead are closely associated with each other, thus the mineralization environment is assumed to be xenothermal. Namely, it is possible that the mineralization zone could change into base metal ore deposit deep underground.

The analyzed results by homogenization temperature for quartz veins ranged from 102 to 194°C, having rather lower

values with a mean of 137°C. It is inferred that a zone most adequate for gold mineralization (180° to 230°C) can exist deeper than the depth of altitude 1700m, until which this year survey has reached. The temperature distribution on geological sections show that higher and lower temperature zones develop western and eastern parts of survey area, respectively. It supports the conclusion of last year that in the west part of the survey area there exists the highest center, say heat source of granitic rocks.

2) Jehuamarca area

Based upon the results of detailed geological survey and drilling survey this year, the mineralization model was assumed to be as follows. Along the secondary or third NW-SE and NE-SW fissure systems, which are branched from regional fault system, associated with small displacement and along the bedding boundary, argillization under neutral to alkaline environment has progressed followed by silicification with mineralization under acidic condition. Therefore it is inferred that three types of mineralization (i.e. base metal mineralization with silicification, high grade layered base metal mineralization within quartz zone and gold and silver mineralization within silicified breccia) shown last year are essentially a series of mineralization. Namely gold-silver mineralization in silicified breccia is a residual deposit, which is formed from auri-argentiferous base metal mineralization zone altered by meteoric water. High grade layered base metal mineralization in quartz is a typical base metal mineralization zone formed by silicification.

Silicified breccia occur continuously as a stratiform, but the mineralization zone of gold and silver is localized and their grade seems to change frequently, therefore it must be concluded to be lower grade than expected. Quartz zone expected to associate high-grade base metal mineralization is discontinuous and changes its thickness and grade frequently, therefore there is a small possibility that it develops to a large scale highgrade ore body.

5-2 Recommendation for the Fourth Year Survey

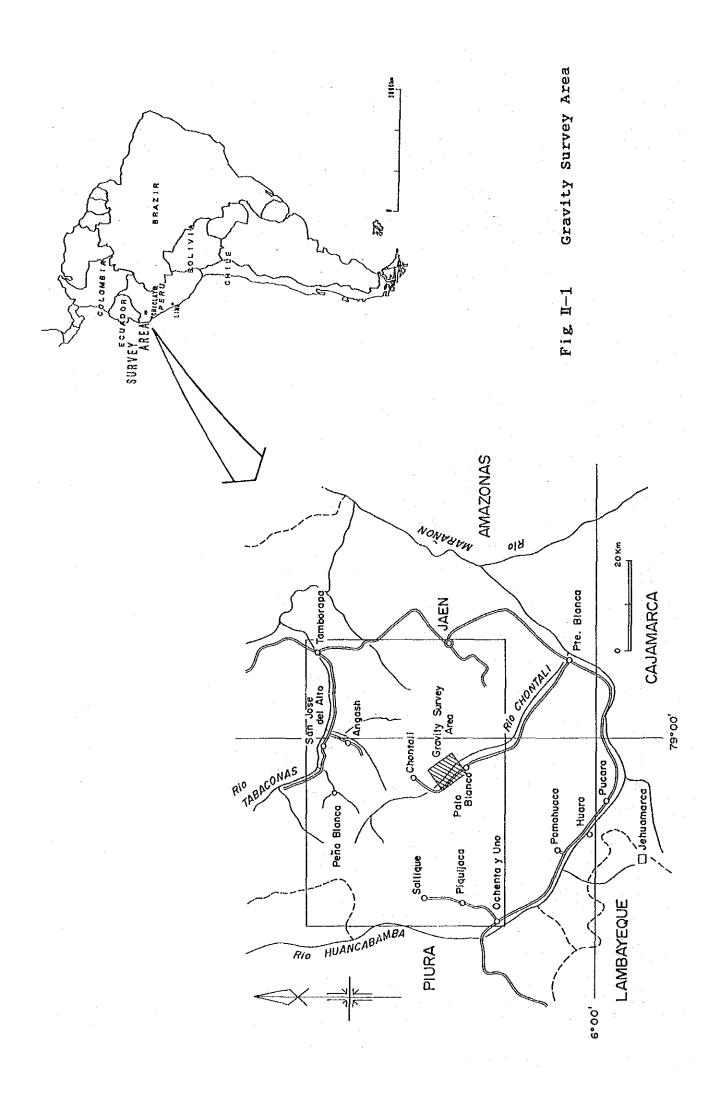
The following surveys are proposed in the order of priority for the fourth year survey based on the results obtained in third year:

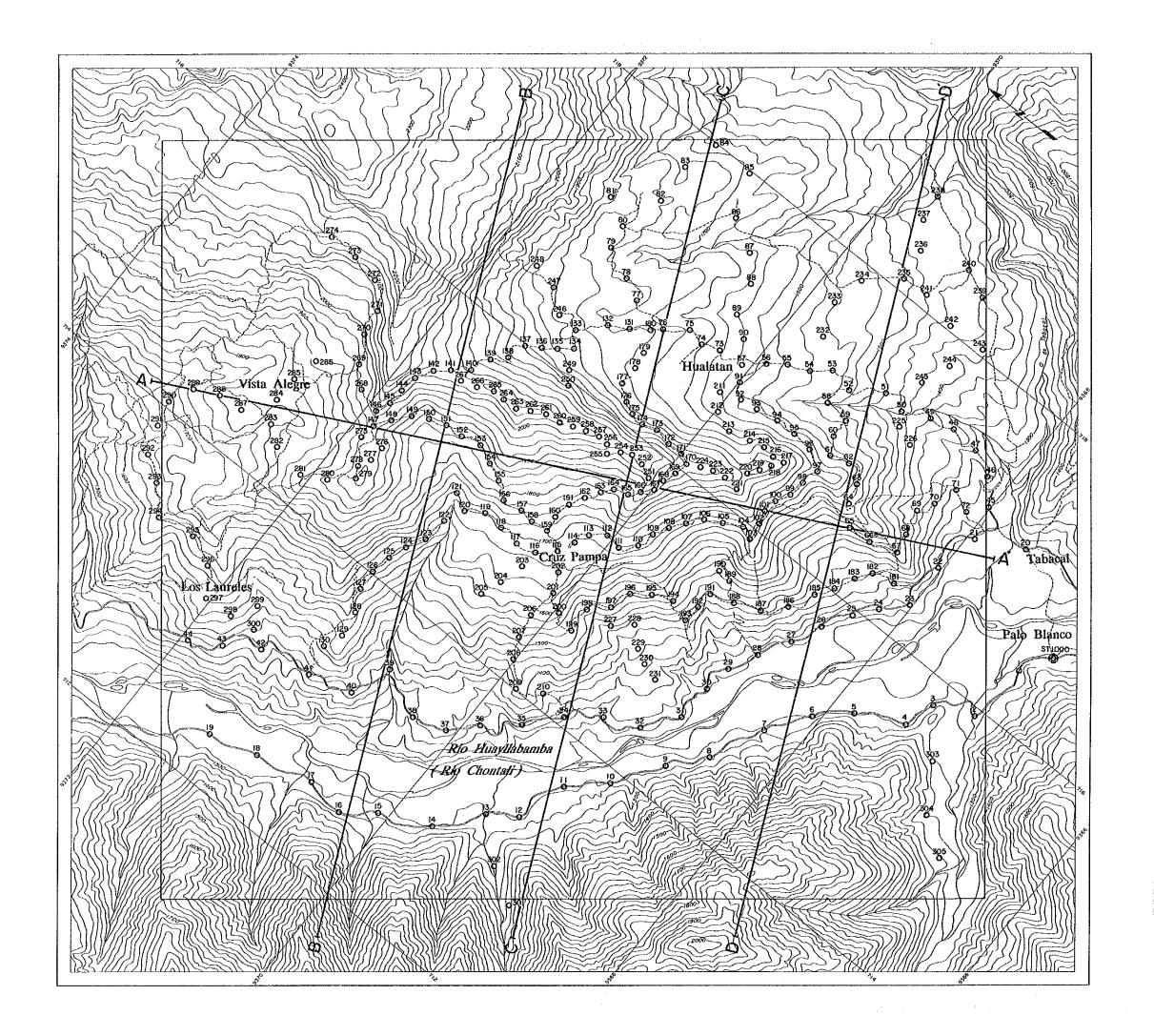
1) Chontali area

(1) Drilling survey at more than two holes to confirm the plunge of quartz vein.

(2) More detailed geological survey combining with detailed mapping and systematic sampling of quartz veins in the quartz vein distributed area will be conducted (with a scale of 1/2000, for example).

PART II PARTICULARS





LEGEND

Base station

305 Station number Gravity station

A-A' Cross section

THE MINERAL EXPLORATION
IN
THE PACHAPIRIANA AREA
REPUBLIC OF PERU
(PHASE II)
GRAVITY SURVEY

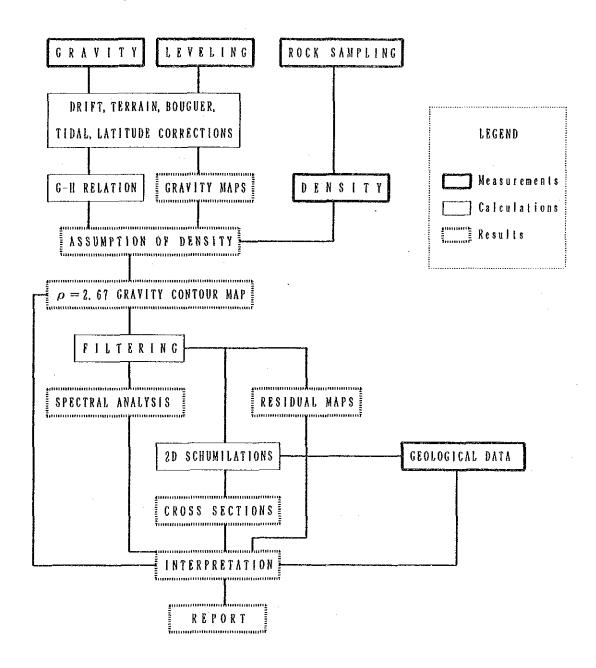
Fig. II-2
Location of
Gravity Stations
FEBRUARY 1991

1000 E

CHAPTER 1 CHONTALI AREA

1-1 Geophysical Survey

The procedure followed in this gravity survey is shown schematically as follows:



1-1-1 Gravity stations

This gravity survey area and the location of gravity stations within the survey area are shown in Fig.II-1 and Fig.II-2. Hualatan is located in the center of the area in a geological district of quartz vein occurrence.

Gravity measurements were made at a total of 305 locations with a dense distribution of gravity stations surrounding

Hualatan.

1-1-2 Gravity measurement

1) Gravimeter

The instrument employed in this survey was a G-283 Lacoste & Romberg, Inc. gravimeter. The specifications of this device are as follows:

ТҮРЕ	La Coste & Romberg. INC Model G Geodetic Gravimeter
No. Operation Range Temperature Reading Line Date of Production Dimensions	283 0~7,386.54 mgal 51.7℃ 2.80 Jan., 1972 14×15×20

This gravimeter has a wide operating range and is capable of measuring gravity from values of 0 mgal to more than 7000 mgal. Very precise measurements can be made with this device in that it has an accuracy of +0.01 mgal and an instrument drift rate of only 1 mgal/month. This allows the collection of precise data while requiring only daily closure.

2) Standard Absolute Gravity Value

We utilized a standard absolute gravity value at a base station identified as CQCH-1. The base station is located on the edge of a bridge near Chamaya in southern Jaen. The standard absolute gravity value is taken as 977,853.54 mgal.

A temporal base station, ST.1000, was established in Palo Blanco where the survey crews were housed. Measurements were made twice per day at the temporal base station to establish a standard absolute gravity value of 977,687.179 mgal.

Closure of gravity values measured during the course of this survey was made by the Potsdam system.

Table II-1 Calculations of Gravity Standard Value

									Observer	Observer : K.KINOSEITA
date of		reading	x constant	height of	н.G.	tide	drift	total	difference	absolute
observation	time	value	(mgal)	gravitymeter	correction	correction	correction	correction	from Base	gravity
scatton				(E)	(mgal)	(mgal)	(mgal)	(mgal)	(mgal)	(mgal)
1990.07.27										
ST.1000	9:15	1301.588	1372.134	0.27	0.083	0.069	0.0	1372.286	0.0	977,687.165
1990.07.30										ally yell Alfa (Maria III)
CQCH-1	12:11	1459.415	1538.550	0.27	0.083	990.0	-0.038	1538.661	+166.375	977,853.54
СОСН-1	12:28	1459.393	1538.527	0.27	0.083	0.062	-0.038	1538.634	+166.348	977,853.54
ST.1000	18:04	1301.705	1372.258	0.27	0.083	-0.014	-0.041	1372.286	0.0	977,687.192

absolute gravity of ST.1000(average): 977,687.179 mgal

1-1-3 Leveling

The direct leveling method was employed for all 305 stations using an Auto level B-2 made by Sokkisha of Japan. Terrain corrections were made by cross section leveling.

An elevation correction value was determined by measuring the elevation of 88 stations relative to the temporal base station by leveling. The elevation of the base station and the leveling stations were then taken from topographic maps and the mean error determined. This shift value (1214.95m) was then added to the elevations measured at gravity stations by leveling.

The accuracy of leveling is expressed as follows:

$$\varepsilon \leq 20 \sqrt{D}$$

where, ε : error of leveling(cm) D: closure distance(km).

The error of closure is divided among readings in relation to the horizontal distance from the measurement point to the leveling base station.

1-1-4 Corrections and data processing

In order to be most useful in prospecting, gravity data obtained in the field must be corrected for tide, height of instruments, drift, latitude, influence of topography, altitude, free-air, Bouguer, and so on.

The results of these corrections are presented in Table II-3.

1) Tidal correction

Tidal forces result from the attraction of the sun and moon at earth's surface. These vary in direction and intensity with time and place of observation.

Tidal forces due to the sun and moon are calculated by the following formula:

$$\Delta g = -3/2 \, G \, M \, a \, / \, r^3 \, \{3 \, (\sin^2 \delta - 1/3) \, (\sin^2 \delta - 1/3) + \sin^2 \delta \, \sin^2 \phi \, \cos \theta + \cos^2 \delta \cos^2 \phi \, \cos^2 \theta \} \times 1.7 \, (g \, a \, l)$$

where, G: gravitational constant = 6.67×10^{-11} (m¹/kg·sec²)

M: mass of the planets (moon: 7.348×10^{22} kg, sun: 1.9891×10^{10} kg) a: distance from the center of the earth to the station = $6378388(0.99832+1.6835\times 10^{-3}\cos 2\phi - 3.5\times 10^{-6}\cos 4\phi)$ + altitude of the station(m)

r: distance from the center of the earth to the planet (moon: 3.844×10^8 m, sun : 1.496×10^{11} m)

 ϕ : latitude of the station

 δ : declination of the planet θ : hour angle of the planet.

2) Drift correction

In the case of spring-balance type gravimeter, drift is caused by creep of the spring material. The influence of the drift is accounted for by assuming that the drift values vary proportionally with time. In this survey, the drift rate is less than 0.2 mgal/day.

3) Terrain correction

A topographic irregularity(hill, knoll, slope, valley, etc.) exerts an attraction directly proportional to its density. The vertical component of this attraction is directed upwards and reduces the observed gravity. Therefore, a term of equal magnitude must be added to the observed gravity.

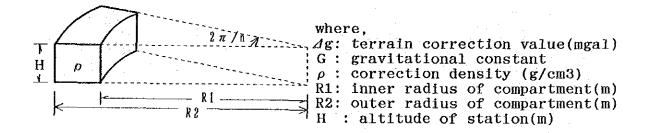
It is necessary to compensate for such effects when topographic features are sufficiently close to an observation point to cause distortions in the observed gravity which are large enough to affect the interpretation of anomalies. For this reason, topography near measurement stations has to be considered in more detail.

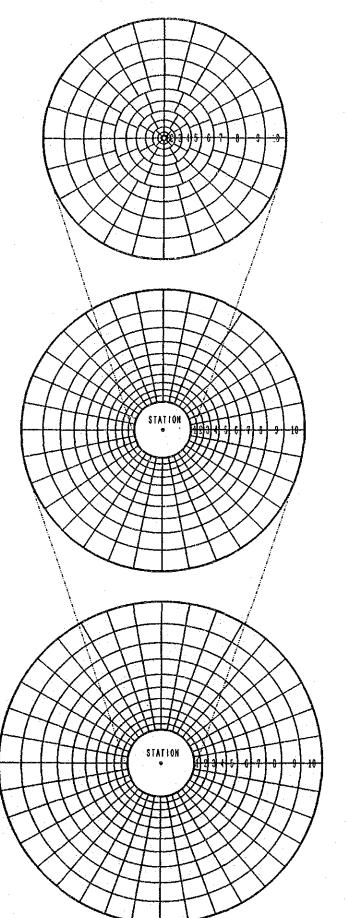
The usual procedure in making corrections for such distortions is to estimate the average altitude by computer within each compartment shown in Fig.II-3 using gridded altitude data from topographic maps. The terrain correction is calculated by dividing the area around the station into three groups, far, middle, and near area.

(1) Hammer's method

The terrain corrections for far, middle, and near areas are calculated by Hammer's method as shown the following form:

$$\Delta g = 2\pi G \rho / n \{R_2 - R_1 + (R_1^2 + H^2)^{1/2} - (R_2^2 + H^2)^{1/2}\}$$



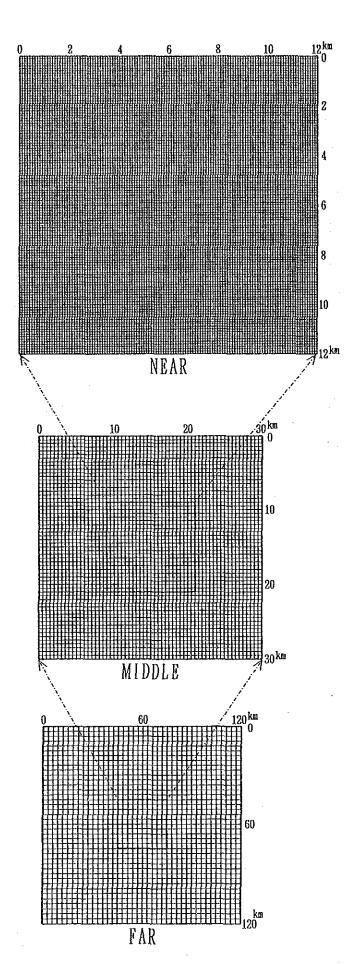


[NEAF	Data Grid It	nterval = 100 m
No.	Distance from Station	Number of Block
1	30~ 90 ^m	8
2	90~ 180	8
3	180~ 300	12
4	300~ 450	12
5	450~ 600	16 Σ=168
6	600~ 800	16
7	800 ~ 1, 050	24
8	1, 050 ~ 1, 350	24
9	1, 350~1, 650	24
10	1, 650 ~ 2, 000	24

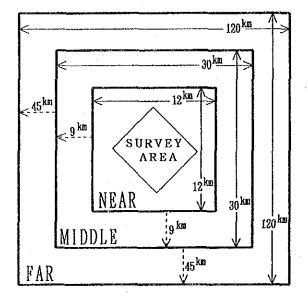
MIU	OLE Data Grid II	nterval = 500 m
No.	Distance from Station	Number of Block
1	2. 0 ~ 2. 4 km	32
2	2. 4~ 2. 8	32
3	2. 8 ~ 3. 3	32
4	3. 3 ~ 3. 9	32
5	3. 9~ 4. 6	32 Σ=320
6	4.6~ 5.4	32
7	5. 4 ~ 6. 3	32
8	6. 3 ~ 7. 3	32
9	7.3~8.5	32
10	8. 5~10.0	32

[FAR	Data Grid II	nterval = 3,000 n
No.	Distance from Station	Number of Block
1	10. 0~12. 0 ^{km}	36
2	12. $0 \sim 14.5$	36
3	14. $5 \sim 17$. 0	36
4	17. $0 \sim 20$. 0	36
5	20. $0 \sim 23.5$	36 Σ=360
6	23. 5~27. 5	36
7	27. $5 \sim 32$. 0	36
8	32. $0 \sim 37. 0$	36
9	37. $0 \sim 43. 0$	36
10	43. $0 \sim 50$. 0	36

Fig. II-3 Annular Segment for Terrain Correction



DIVISION OF CORRECTION AREA





	NEAR	MIDDLE	FAR
Λrea	12×12 =144 km²	30×30 =900 km²	120×120 =14. 400km²
Grid interval	100 m	500 m	3, 000 m
Number of grids	120×120 = 14, 400	60×60 = 3,600	40×40 = 1,600
Scale of map	1:25, 000	1:100.000	1:100.000

Fig. II-4 Division of Terrain Correction Area

n: number of divided compartment in a same circular ring.

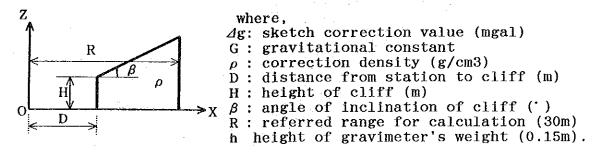
The average altitude of each compartment H is estimated from the surrounding four altitude data values, H1, H2, H3, and H4, as follows:

$$H_{1}$$
 S_{1} H_{2} H_{2} H_{3} H_{4} H_{4} H_{2} H_{1} H_{2} H_{3} H_{4} H_{2} H_{3} H_{4} H_{4} H_{2} H_{3} H_{3} H_{4} H_{4} H_{4} H_{4} H_{2} H_{3} H_{4} H_{5} H_{4} H_{5} H_{4} H_{5} H_{4} H_{5} H_{5

(2) Sketch correction

We sketched the topographic features along the direction of maximum topographic undulation in a radius of 30m. Using the sketches, we calculated the sketch correction by Hiroshima's method. The expression and the schematic map are as follows:

$$\Delta g = 2 G \rho \int_{0}^{R} \left(\tanh^{-1} \sqrt{(R^{2} - X^{2}) / (R^{2} + h^{2})} - \tanh^{-1} \sqrt{(R^{2} - X^{2}) / (R^{2} + X \tan \beta + H - h - D \tan \beta)^{2}} \right) dx$$



4) Elevation correction

This correction consists of three parts: (g1) the free-air correction, which accounts for the fact that each station is a different distance from the earth's center and the datum plane, (g2) the Bouguer correction, which removes the effect of a presumed infinite slab of material between the horizontal plane of each station and the datum, and (g3) the atmosphere correction, which removes the effect of the atmosphere.

Free-air correction: free-air correction is estimated by following expression:

$$\Delta g_1 = g_0 \{1 - R^2 / (R + H)^2\} = (2 g_0 H R + g_0 H^2) / (R + H)^2$$

 $= 0.3086 H (mgal)$

g₀: gravity at mean sea-level H: altitude of station (m)

R: average radius of the earth.

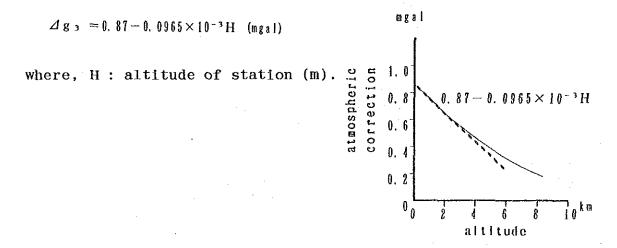
Bouguer correction: Bouguer correction can be expressed as,

$$\Delta g_2 = -2 \pi G \rho H = -0.0419 \rho \cdot H (mgal)$$

where, G: gravitational constant

H: altitude of station (m) ρ : density of slab (g/cm3).

Atmosphere correction: an atmosphere correction can be obtained from the average atmospheric density model by use of the integral for the range 0 to 50 km. As shown in following figure, the relation between atmosphere correction and the altitude of the station can be assumed to be one-order function less than 3km of the altitude. Therefore, we can write as follows:



Because free-air, Bouguer, and atmosphere corrections are functions of the altitude of the station, they are called elevation corrections and can be expressed as follows:

Elevation correction:
$$\Delta g = \Delta g_1 + \Delta g_2 + \Delta g_3$$

= (0.3086-0.0419 • ρ -0.0965×10⁻³) H + 0.87 (mgal)

5) Latitude correction

Because of its rotation, the earth is not actually spherical. The shape can be approximated as an oblate spheroid with an eccentricity of 1/297. Both its departure from sphericity and its rotation cause the earth's gravitational acceleration to have a maximum value at the poles and a minimum at the equator. Therefore, the variation of gravity values with the latitude of the station is reduced by using following International Formula on the International Ellipsoid.

$$\gamma = (a \gamma_E \cos^2 \phi + b \gamma_P \sin^2 \phi) / (a^2 \cos^2 \phi + b^2 \sin^2 \phi)^{1/2}$$
 (mgal)

where, a: equatorial radius of ellipsoid of revolution (6,378.14 km)

b: equatorial radius of ellipsoid of revolution

(6,356.18 km)
7: normal gravity value at the equator of ellipsoid of revolution(978.032 gal)

γ_P: normal gravity value at the polar of ellipsoid of revolution (983.218 gal).

This is the International Gravity Formula of 1967, which is adopted by IUGG(International Union of Geodesy and Geophysics) as the expression giving the normal gravity.

Actually, we use the convenient following expression:

$$\gamma = 978,031.85(1+0.005278895\sin^2\phi + 0.000023462\sin^4\phi$$
) (mgal)

6) Densities of rock samples

The terrain and Bouguer Corrections made in the reduction of gravity data require a knowledge of the densities of the rocks near the surface. To upgrade the precision of an analysis, an assumption of a detailed underground density structure is required.

In this survey, 35 rock samples were collected on the surface and from core samples and their wet densities were measured. The densities of rock samples can be obtained by using following expression:

$$\rho = \frac{\text{Wa}}{\text{Wa - Ww}}$$

where, ρ : density

Wa: weight in air (after a day in water)

Ww: weight in water

Table II-2 shows the results of measurement with main rocks.

7) Reduction density

Bulk densities for the thirty-five rock samples ranged from 2.43 to 2.70g/cm3. It is difficult to determine the formation density of these rocks from only their densities near the surface.

We determined a reduction density by the following methods in this survey.

i) Gravity contour maps were drawn for four densities, 2.2, 2.4, 2.67 and 2.80g/cm3. The correlation between gravity anomalies and topographic features are shown in following table. Fig.II-7 illustrates the gravity contour maps with densities of 2.67g/cm3, respectively.

	correlation between gravity anomaly and topographic feature														
	ρ =2.5	$\rho = 2.6$	$\rho = 2.67$	ρ =2.8											
High gravity and High altitude or Low gravity and Low altitude	9	7	4	0											
High gravity and Low altitude or Low gravity and High altitude	0	3	6	9											

The density value at which the topographic features have the least observable effect on the gravity contours is considered to be most nearly correct. Therefore, correct density appears to be about 2.67g/cm3.

ii) The average density of the thirty-five rock samples is estimated to be 2.60g/cm3.

iii) Gravity is a function of altitude of station. When the altitudes and the corrected gravity values (after latitude and terrain corrections and reduction of gravity trends) are plotted on X-Y coordinates, the inclination of fitting lines represent the average of rock densities in the area. Fig.II-5 shows the relation of gravity versus altitude (G-H relation) for this sur-

Table II-2 Rock Properties

alt. 2.18(1) £ 2.73(5) 2.76(2) ——	min.	nonmineralized	fields	sample	total
2.73(5) 2.76(2)		nonmin.	core	field) 5 4 1 1 2 3
2.73(5) 2.76(2) 2.36(1)	2.62(1)	2.38(4)		2.43(5)	2.43(5)
2.76(2)		2.59(9)	2.73(5)	2.51(6)	2.61(11)
2.36(1)		2.70(3)	2.76(2)	2.58(1)	2.70(3)
2.36(1)		2.71(4)	2.75(1)	2.70(3)	2.71(4)
		2.47(3)		2.47(3)	2.47(3)
granodiorite 2.68(2)		2.68(2)		2.68(2)	2.68(2)
quartz vein 2.61(7)	2.68(3)	2.57(4)	2.67(5)	2.48(2)	2.61(7)
					2.60(35)

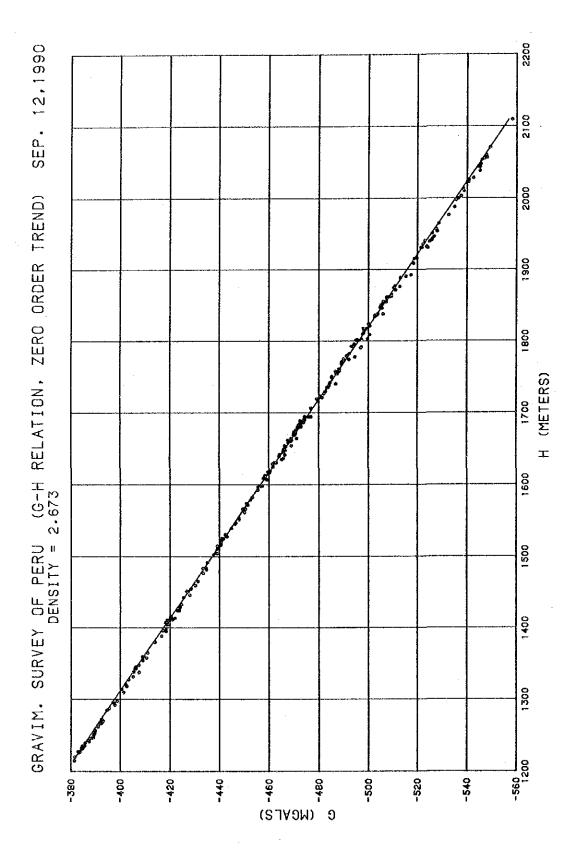


Fig. N-5 Gravity versus Height Relation

vey and a density was estimated as 2.673g/cm3 by least squares method.

Eventually, a correction density of 2.4g/cm3 was arrived at from the density derived from the correlation between topographic features and gravity anomalies (2.67g/cm3), the average density of rocksamples (2.60g/cm3), and the density determined from the G-H relation (2.673g/cm3).

1-1-5 Interpretation

The extraction of anomalies associated with individual subsurface sources from an observed gravitational field requires filtering operations. In this survey, we used the following techniques; surface fit analysis, and spectral analysis. Techniques developed by Talwani were also used to estimate two-dimensional density structures.

1) Surface fit analysis

Gravity contour maps frequently contain not only gravity anomalies produced by local geologic features but also gravity trends resulting from regional features.

In this area, gravity trends decrease from the northeast to the southwest. Therefore, a first-order approximate plane Z(X,Y) was adapted to remove the regional gravity trend.

Considering gravity data at grid points G(X,Y), we estimated a first-order surface fitting plane Z(X,Y) by the least squares method. Z(X,Y) and the residual values are presented in the following form:

$$Z(X, Y) = -141,85129 - 0.0403031 X + 0.0455303 Y$$

$$(Residual \ Value) = G (X, Y) - Z (X, Y)$$

The results are shown in Fig. II-8.

2) Spectral analysis

The gravity contour maps produced by this survey include anomalies of various wave lengths. By transforming such spatial gravity anomalies into the frequency domain, it is possible to estimate the average depth of the sources of anomalies and to eliminate gravity anomalies attracted by the source with spontaneous average depth (Spector and Grant, 1970).

Fig.II-6 shows the logarithmic energy spectrum of the gravity as a function of frequency, f(cycle/km), where the spectrum can be approximated by using two regression lines. From the gradients of these lines, the average depth (H) of buried sources can be expressed by following formula:

$$H = - \frac{1}{4 \pi} \cdot \frac{\Delta \log E}{\Delta f}$$

In this survey, the gravity field can be separated into a regional component (Hr) with an average source depth of 1,240m, a near-surface component (HN=170m), and a middle component (KX=340m).

From the results of spectral analysis, we can draw spectral maps for the regional and near-surface components. In practice, the regional component, GR, is expressed mathematically by using two average depths, HR and HN, and KX, where difference of crossing points between energy spectral axis and each regression line.

$$G_{R} (X, Y) = \frac{\int_{-\infty}^{\infty} \int_{-\infty}^{\infty} G(\xi, \eta) \cdot W(X - \xi, Y - \eta) d\xi d\eta}{\int_{-\infty}^{\infty} \int_{-\infty}^{\infty} W(\xi, \eta) d\xi d\eta}$$

where, the weight function, W, can be expressed as follows:

$$W(X, Y) = \frac{1}{2\pi} \int_{0}^{\infty} \frac{f \cdot J_0 (P \cdot f)}{1 + K_X \cdot e^{(H_R - H_N) \cdot f}} df$$

in which, $P^2 = X^2 + Y^2$

JO: Bessel function.

The gravitational effect of shallow structures, GN(X,Y), is obtained through spectral analysis, by subtracting the regional component of the observed field, GR(X,Y), from the total observed gravitational field, G(X,Y) as follows:

$$G_N (X, Y) = G(X, Y) - G_R(X, Y)$$

3) Two-dimensional simulation

A two-dimensional simulation was carried out along two sections where two-dimensionality could be assumed in the subsurface structural configuration.

Two-dimensional bodies were represented by polygons and the gravitational attraction caused by these bodies was compared

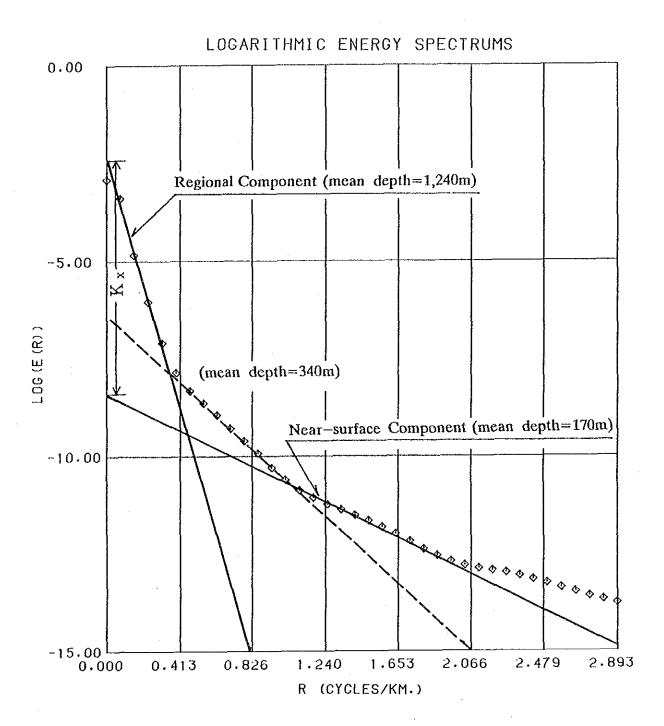


Fig. H-6 Energy Spectrum

Table II-3 List of Gravity Data

(density for correction ρ =2.67 g/cm³)

LEGEND

ANOM . B	Bouguer anomaly (mgal)	
ANOM . F	Free-air anomaly (mgal)	
υ •	normal	
C NORM	gravity (mgal)	
В.С.	Bouguer correction (mgal)	
D • E • E	Free-alr correction (mgal)	·
TERR . C	total terrain correction (mgal)	
ETC	leveling method and gravimeter	I
C.30M	sketch correction (mgal)	
ABS . G	absolute gravity (gal)	
LEVEL	altitude of station (m)	
rong •	longitude	
LAT .	latitude	
OBS . DAY	date of observation	
ST . NO	s tation number	

PAGE=	ANOM. F ANOM. B	5.674 -139.	4.156 -139.	-2.983 -159.617	3.040.130.	4.097 -139.	3.074 -140.	1.736 -140.	3.164 -140.	2.340 -140.	94 -140.	3.457 -141.	2.104 -140.	000	0.000 - 140.0	12 - 140	0.897 -140	0.703 -140.	3.846 -139.	3.165 -139.	3.491 -139;	3.601 -139.	2.656 -139.	3.225 -138.	33 -139.	2.432 -139.	5.775 -159.	0.640 H_0V.	2.000 -140. 2.000 -140.	204 1 40. 270 1140.	1.232 -140.	364 -139.	545 -139.	.38 -139	525 -140	728 -159	8.242 -139.46	C. 666 - 666 V	0.028 -1.59.46	707	10. KG11 040.0	70.401 177.07	2.069 Lito	772 -141-10	943 -140.84	494 -140	3 -139.
.67 (G/CM**3)	NORM. G	.08318	.08317	978.083129	0000	.08302	.08299	.08296	.08293	.08289	.08285	.08283	.08280	000000	77000	08264	08255	3.08249	3.08310	3.08304	3.08303	3.08303	3.08300	3.08299	3.08296	3.08295	0.0000		2000	0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	3.08283	3.08280	3.08277	3.08273	3.08270	3.08266	5.08260	0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	5.0825.5	0.0004	7000	000000	0000	00000	8.08294	8.08291	28
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ā	н. С	.71	64.		- 0		90	.08	.37	.37	. 32	.64	. 52		0 0	. υ Ο α	4	69	. 79	.67	.83	. 43	66	, 59	. 34	3.32	4 1		0 \ 0 \	4. C		3.36	86.	9.60	0	o i	2	<u> </u>	. i				א נו ה	7	9	71.6	4
	TERR.C	9.5	13.90	12.958	+ t	16.02	15.56	16.81	16.68	16.49	15.41	16.44	15.86		0.0	20.0	100	16.04	13.87	13.90	14.08	13.49	14.16	15.77	16.60		16.33	20.0	70.40	4.00	15.27	14.92	14.40	12.93	12.27	12.45	13.09	20.21	12.50	72.0	7.7.		- T	0 1 1	10.5	20	4
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T OF GRAV	LEVEL	215.0	227.3	0 n	7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7	226.3	242.1	255.1	246.3	252.8	252.6	247.2	253	0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	000	000	261.8	266.8	231.5	234.3	231.6	233.5	235.3	227.6	220.3	239.7	227.1	7.00.0	7. C	0 t t 0	256.7	272.2	287.2	295.7	319.8	342.6	338.1	4.40	7.604	415.4	0.1		- 0		9 1	3 (\ - 00	1428.02
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ο.	ANOM. F	1.61	7.97	.52	55.576	. a	, k	4 -			, ,	, ,	7.70	75.7	0.72	20.0	7.73	4.43	9	5	8	0.29	5,4	7.62	0.61	.89	2.58	5.55	5.66	2.95	2,44	7.22	2	9.25	8	0 r			10	0	5	3	6	2.57	4.1	5.37	4.65	5.35	5.66
67 (G/CM**3)	NORM.G	8.08285	8.08282	8.08279	978.082775	0.006/4	0.000.0	0.000.00 2000.00	78787 87	2000	100 C	8.08287	78.08289	8.08290	78.08292	78.08295	78.08298	78.08298	78.08297	78.08298	78,08299	78.08301	78.08267	78.08265	78.08263	78.08260	78.08256	78,08253	78.08250	78,08249	78.08246	78.08251	78.08251	78.08252	78.08257	70.000.00	70.0040.07	78.082	78.08268	78.08271	78.08273	78.0827	78.08277	78.08280	78.08282	78.08285	78.0828	78.0828	78.08283
ENSITY = 2.	ວ.ວ.ສ	61.51	62.00	70.88	175.401	70.07	7 7	47.00	77.89	74.87	77.04	75.96	74.24	72.50	68.63	64.38	57.57	54.02	48.71	45.81	45.99	40.24	82.78	86.32	89.25	90.17	90.75	93.46	92.38	94.59	99.31	94.32	89.0	85.73	900	- 100) a	77.62	80.46	77.97	77.88	8	4	м	õ	õ	ĕ	ŏ	ĕ
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OF GRAV	רפאפר	465.1	515.7	50.1	1500.10	7 7 7	, v	7 2 2 2	000	7 2 0	808	598.4	582.5	566.4	530.8	491.6	428.9	396.2	347.4	320.8	564.9	269.7	661.4	694.2	721.2	729.8	735.6	760.3	750.3	720.7	2.4	768.2	719.6	2.889	244.5) 0 0 0	000	0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	640.0	616.9	616.1	624.7	630.8	8.7.49	660.5	64.2	62.5	69.1	71.2
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* *	LAT.	542.	542.1	542.0	Ņα	4	74.1	177	142	542	542.2	542.3	542.3	542.4	542.4	542.5	542.6	542.6	545.6	542.6	542.7	542.8	541.6	541.5	541.5	541.4	541.2	541.1	541.0	541.0	240.9	541.0	241.0		7.7.7	1		547	541.6	541.7	541.8	541.9	541.9	542.0	542.1	542.2	545.2	542.2	542.1
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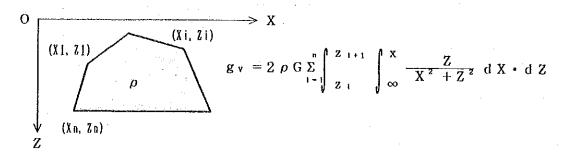
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with the measured gravity. The geometry of the disturbing body was then automatically adjusted from the starting model by using optimization techniques.

The gravitational attraction was calculated by Talwani's method, where the vertical component of the gravity, Gv, at an arbitrary point 0 on ground due to a polygon with a volume density p is given by following expression:



where, ρ : density
G: gravitational constant.

In the final analysis a multi-layered model based upon known geological structure was used. The results are shown in Fig.II-10(1), (2), (3), and (4).

1-1-6 Results of gravity data analysis

The analytical results obtained from this survey are presented in the following illustrations.

Fig. II-7 Bouguer Anomaly Map (2.67g/cm3).

Fig. II-8 First-order Residual Gravity Map.

Fig. II-9(1) Long-wave Gravity Map.

Fig. II-9(2) Short-wave Gravity Map.

Fig. II-10(1) Cross Section of A-A'.

Fig. II-10(2) Cross Section of B-B'.

Fig. II-10(3) Cross Section of C-C'.

Fig. II-10(4) Cross Section of D-D'.

Fig. II-11 Geophysical Interpretation Map.

1) Fig.II-7 Bouguer Anomaly Map (density=2.67g/cm3).

The Bouguer gravity in the survey area (Fig. II-1) is seen to vary from a low of -141 milligals (mgal) to a high of -136 mgal with a range of only 5 mgal. The maximum gravity gradient is 3 mgal/Km near Vista Alegre.

Major gravity highs are found in the central and north-eastern survey area with minor highs trending from Vista Alegre in the north to Tabacal in the south-east through the center of the area. The Bouguer gravity is found to be anomalously low along the south-eastern, southern, and south-western edges of the area. There are minor local lows in the interior of the area on the flanks of the major central high.

Most of the Bouguer highs occur on or near ridge lines with lows found in the valleys. An exception to this is seen near Los Laureles where a minor local high is located in a valley between two ridges and a minor low is located on a ridge. The Bouguer gravity contours conform roughly to topographic contours in the majority of the survey area, but there is only limited correlation between the two in the north-central quadrant of the area.

The Bouguer gravity values generally increase from the south-west to the north-east boundary of the survey area. This one-dimensional regional gravity trend is probably due to the influence of a broad geological structure at depth.

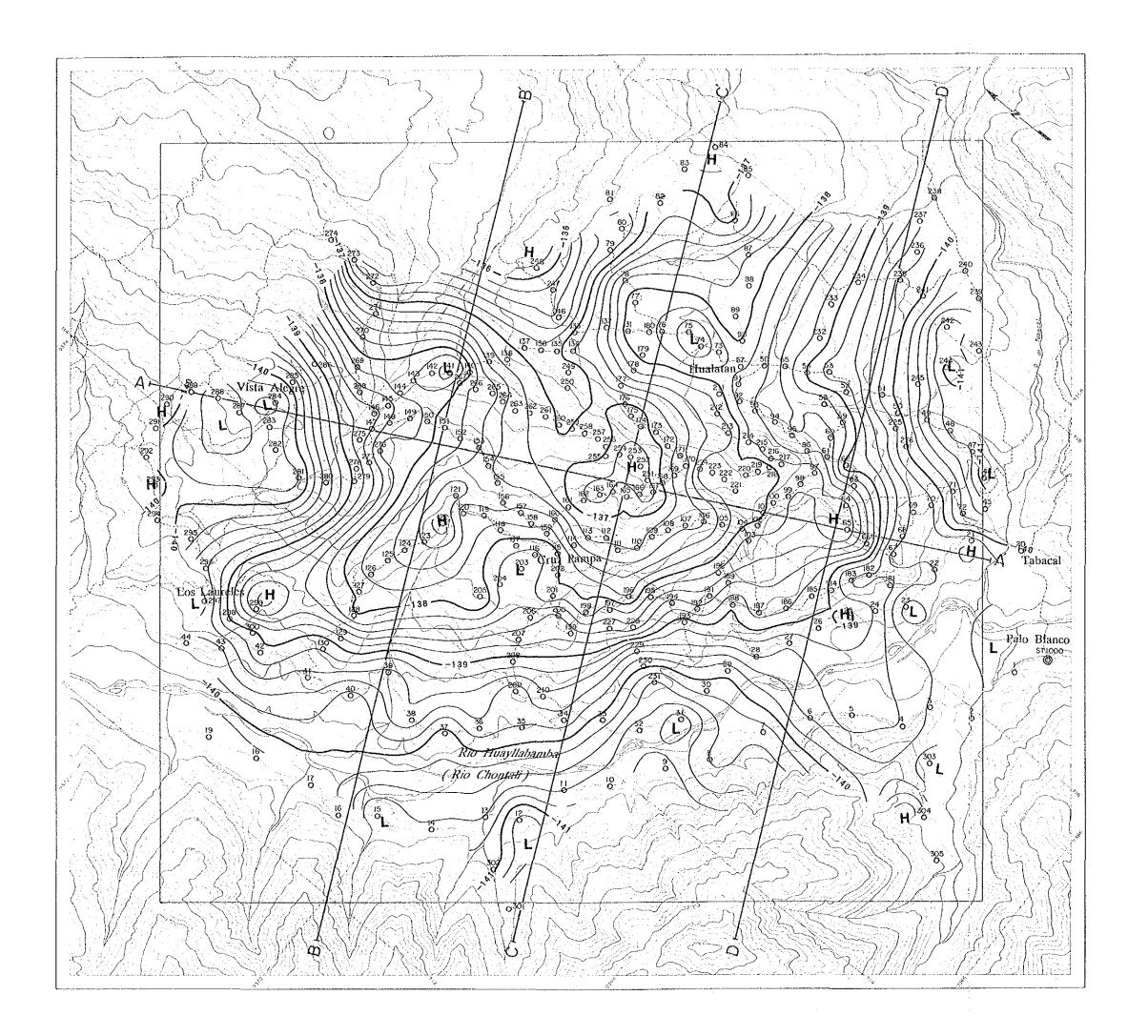
2) Fig.II-8 First-order Residual Gravity Map.

The first-order residual gravity map (Fig. II-8) is a representation of the residual gravity after removal of the one-dimensional regional gravity trend mentioned above. This data set represents the effect of shallow local features and ignores the effect of deep regional geological structures. There are only minor variations in the Bouguer Anomaly Map and the First-Order Residual Gravity Map. Residual gravity values range from a high of about 2.0 mgal to a low of -2.4 mgal. High residual gravity anomalies are observed in the middle of survey area, surrounded by low residual gravity zones. The gravity highs trend NNE-SSW and NW-SE through the center of the survey area.

The distribution of high and low residual anomalies may be summarized as follows:

i) A broad high residual anomaly is found in the middle of the survey area between Hualatan and Cruz Pampa with a maximum amplitude of about 2.0 mgal. Minor highs are distributed about the central high along conjugate axes running from NNE to SSW and from WNW to ESE through the center of the area. This distribution of highs may be the result of deformation and uplift of the basement rocks by regional conjugate stresses.

ii) There are low residual anomalies near Vista Alegre, north-east of Tabacal and along Rio Huallabanba. There is a granitic zone on the west side of Rio Huallabanba and, while granitic zones normally produce high gravity anomalies due to the relatively high density of granite, anomalously low values are



o Gravity station

A-A' Cross section

1.0 mgal interval

0.2 mgal interval

H High gravity zone

L Low gravity zone

THE MINERAL EXPLORATION
IN
THE PACHAPIRIANA AREA
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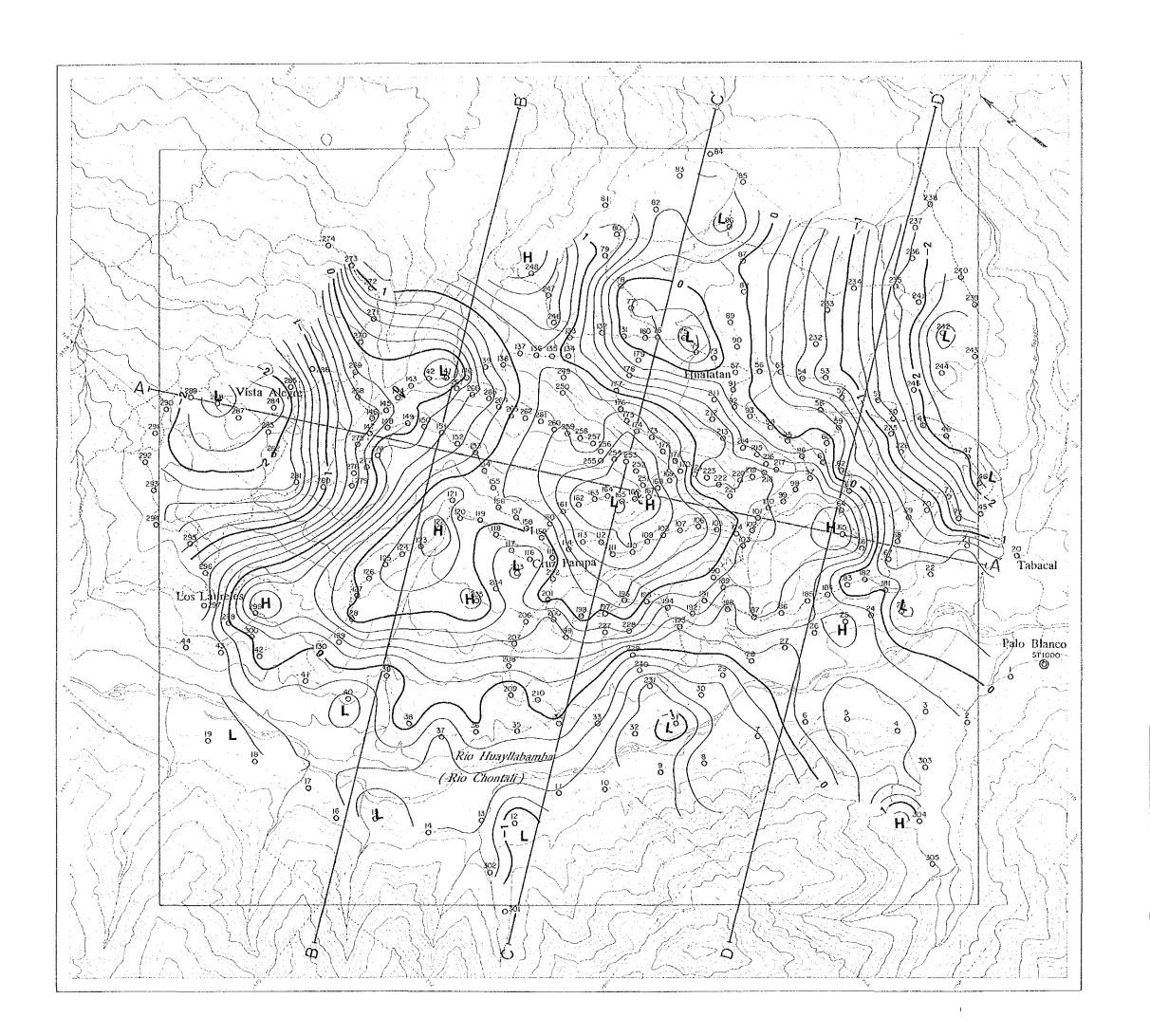
GRAVITY SURVEY

Fig. II-7
Bouguer Anomaly Map

FEBRUARY 1991

 $(\rho = 2.67 \text{g/cm}^3)$

0 500 1000 m



O Gravity station

A-A' Cross section

- 1.0 mgal interval

- 0.2 mgal interval

H High gravity zone

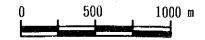
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THE MINERAL EXPLORATION
IN
THE PACHAPIRIANA AREA
REPUBLIC OF PERU
(PHASE I)

GRAVITY SURVEY

Fig. II-8
First-order
Residual Gravity
Map

FEBRUARY 1991



observed in this area.

3) Spectral analysis

The spectral behavior of the Bouguer gravity observed in the survey area will be analyzed in this section. The long-wave or low frequency variations in the gravitational field will be addressed first, then the short-wave or high frequency behavior of the field will be discussed.

(1) Fig.II-9(1) Long-wave Gravity Map.

The long-wave gravity map, Fig. II-9(1), generally shows variations in the gravity basement over the survey area. Values of long-wave gravity range from a low of minus 141.6 mgal in the south-west to a high of -136.8 mgal in the north-east. anomaly found at the north-eastern boundary of the survey area spreads to the south-west toward the middle of the area. The Bouguer gravity decreases from the center of the area to the south-east, south-west and north-west with a near constant gradient. Bouguer gravity values of less than -140 mgal occur on the west side of Rio Huallabanba, north-east of Tabacal, and north-west of Vista Alegre.

This implies that the gravity basement is shallow from the north-eastern quadrant to the center of the survey area and grows monotonicaly deeper to the south-east, south-west and north-west.

(2) Fig. II-9(2) Short-wave Gravity Map.

Local high and low short-wave gravity anomalies are distributed about the entire survey area, Fig. II-9(2). These anomalies reflect high and low density zones at a depth of about 170 meter below surface. The high short-wave gravity anomalies between Tabacal and Los Laureles are located on zones of anomalously high residual gravity, Fig. II-8. Remarkably high anomalies occur between Hualatan and Cruz Pampa, north-west of Tabacal and near Los Laureles. There is a high occurrence of quartz veins in the Hualatan and Cruz Pampa areas and this implies that high short-wave gravity anomalies may be correlated with mineralization.

There are numerous zones of anomalously low short-wave gravity in the Hualatan, Vista Alegre and Cruz Pampa areas which imply the probable presence of low density material at the surface.

4) 2 Dimensional simulation

In this section two-dimensional simulations derived from the processed gravity data will be addressed. The results of the simulations are presented in four illustrations, Fig. II-10(1-4). These illustrations include cross sections of the observed and calculated Bouguer gravity, Fig. II-10(1a), from which a one dimensional regional gravity trend has been removed by trial and error, long and short-wave gravity, Fig.II-10(1b), a density pseudosection, Fig. II-10(1c) and an assumed geological cross

section, Fig. II-10(1d). There are no control points to regulate the relative depth of gravity basement, such as deeply boring data. Therefore, the simulation was repeated until profile depths matched at points where sections lines crossed.

All sections were analyzed assuming a low density layer (2.67g/cm3) to overly a dense gravity basement (2.80g/cm3). From the data inversions, surface layers of anomalously low density (2.53-2.57g/cm3) and high density zones (2.80g/cm3) in the subsurface were found to exist. While low density units are present at the surface, geological investigations show the presence of no high density host rocks or zones of locally high density at the surface. This implies that there are zones of anomalously high density in the sub-surface rocks.

The analyzed gravity basement or high density zones are denser than rock samples on Table II-2. This implies that base rock or sub-surface rock are mineralized such as highly carbonatized.

(1) Fig. II-10(1) Cross Section of A-A'

Cross section A-A' runs from north-west of Vista Alegre south-south-east to Tabacal, Fig. II-7.

The gravity basement (2.80g/cm3) lies at an elevation of about 1000 meters at the northern end of A-A' and rises at about 45x to around 1400m south-east of Vista Alegre (Fig. II-10(1)). Here there is a break in slope and the basement rises more gradually (>30') to crest at an elevation of 1700m near the intersection of lines A and C. From the intersection of lines A and C, the basement dips monotonically to the south-east toward Tabacal at about 30'.

A layer of low density material (2.54g/cm3) overlies the upper layer (2.67g/cm3) at the north end of line A and is as thick as 450m in the depression near Vista Alegre. This layer quickly thins to a veneer south-east of Vista Alegre but persists to the intersection of lines A and B.

A high density zone (2.80g/cm3), which extends nearly to the surface, is implied in the upper layer from the crest of the basement to the south of line C and another occurs at the intersection of lines A and D.

These zones of anomalously high densities may be the result of alteration of the unit by intrusion with the possibility of mineralization. These anomalies will be discussed further in the analysis of lines B, C and D.

(2) Fig.II-10(2) Cross Section of B-B'

Cross section B-B' strikes WSW-ENE in the northern half of the survey area, Fig.II-7.

The basement rocks lie at an elevation of 1100m along Rio Huayllabamba (Rio Chontali) at the west end of line B, Fig. II-10(2). They rise to the east to an elevation of 1550m, following topography to the midpoint of the line. The basement then dips to an elevation of 1400m east of the intersection of lines B and A, then rises at about 40° to an elevation of nearly 2Km near the

east end of the line.

A thin (<100m) low density (2.56g/cm3) surface layer covers the denser upper section along Rio Huayllabamba. Another low density zone develops near the intersection of lines B and C which thickens to 300m in the eastern half of the line. This layer pinches out near the topographic high at the east end of line B.

No high density zone were found in the upper layer of this section.

(3) Fig. II-10(3) Cross Section of C-C'

Cross section C-C' is parallel to B-B' and passes between Cruz pampa and Hualatan through the center of the survey area,

Fig.II-7.

The general behavior of the basement along line C-C' is similar to line B. There is a basement low west of Rio Huayllabamba and the basement dome peaks with topography near the intersection of lines C and A. The west slope of the central basement dome is about 30' from horizontal and there is a depression east of the peak. The basement elevation at the east end of the line is only 1500m versus 1900m on line B and the east end of line C is flatter than line B.

The upper layer of line C is somewhat thicker than along line B (100-250m versus 50-200m). It is seen to thin on the flanks of the dome and thicken in the Rio Huayllabamba valley and

north-east of Hualatan.

The low density units encountered along line B (2.56 and 2.54g/cm3) are present in profile C-C' (2.53g/cm3 and 2.55g/cm3). They overlay the upper layer west of Rio Huayllabamba and north of Hualatan. Both of these units have a maximum thickness of about 200m. The western unit has migrated out of the river valley to the west of rio Huayllabamba and the east-central unit has thinned from line B and begun to pinch out.

There is a high density zone (2.80g/cm3) in the upper layer near the center of line C, just east of the peak of the basement dome and beneath the topographic high of the line. This zone is in an area of known quartz vein occurrence and from its

topographic relief appears to be erosion resistant.

(4) Fig.II-10(4) Cross Section of D-D'

Cross section D-D' parallels B-B' and C-C' near the southern boundary of the survey area, south of Hualatan and north of Taba-

cal, Fig.II-7.

The gravity basement has far less relief in this section than in the preceding two east-west cross sections and shows far less topographic control than on lines B and C, Fig. II-10(4). The basement elevation at the west end of the line, west of Rio Huayllabamba, is about 950m, which is comparable to its elevation on line C. It rises at about 10 to a maximum elevation of only 1250m near the intersection of lines A and D at the middle of line D. There is a basement syncline centered in the eastern half of profile D-D' with a trough depth of 900m.

The upper unit has thinned in the western half of the profile to a thickness of only 50 to 100m. In the eastern half of the line, however, it is as thick as 650m. This is a thickness comparable to that found in the depression north-west of Vista Alegre.

The low density unit observed on the west slope of the Rio Huayllabamba valley in cross section C-C' persists in profile D-D'. It has become marginally less dense (2.53g/cm3) and lies confined to the flank of steep topographic features to the west. The low density unit found near Hualatan on line C is not present in cross section D-D'.

The high density zone in the upper unit of line C, at the crest of the basal anticline, is again encountered in cross section C-C' and it again occurs upon a basement high. A second zone of high density has developed as well. It lies to the east of the basement crest and both of these zones have densities of 2.80g/cc.

5) Fig. II-11 Geophysical Interpretation Map

In the interest of developing an integrated geophysical interpretation, a geophysical interpretation map was developed, Fig.II-11. This maps incorporates the results of this gravity survey and those of the Controlled Source Audio Magnetotellurics (CSAMT) survey performed in this area in 1989.

The data sets presented in this illustration are:

- a) A contour map of the elevation of the gravity basement relative to mean sea level.
 - b) (Assumed) Zones of anomalously high densities.

{data sets a) and b) were derived from this set of gravity data by two dimensional simulation}

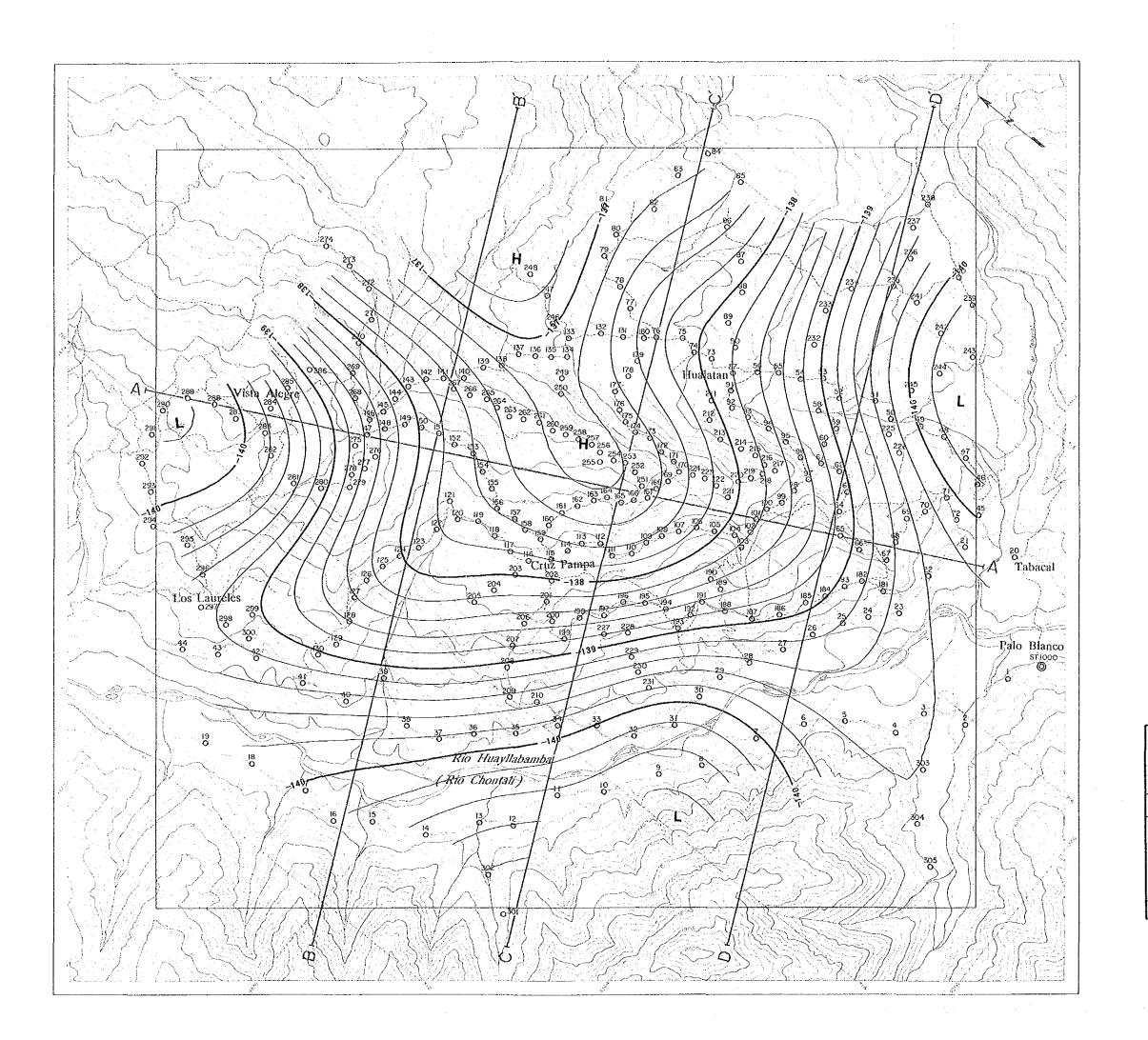
- c) Areas determined to have resistivities higher than 1,000 ohm-m.
 - d) Zones with resistivities lower than 20 ohm-m.

{ data sets c) and d) were derived from the CSAMT data contoured at an elevation of 1600m above sea level. }

e) Quartz veins are indicated as solid black lenses.

The basement depth map presented in Fig. II-11 is quite similar in shape to the long-wave gravity map, Fig. II-9(1). The gravity basement forms a closed dome at the center of the survey area where it is about 100m beneath the surface at an elevation of 1700m. There is a minor syncline in the basement to the north-east of the dome which plunges to the south-east in the direction of Hualatan and to the north-west toward Vista Alegre. The basement rises beyond the syncline in the north-eastern quadrant of the survey area to its maximum elevation in excess of 1700m above sea level. The gravity basement reaches lows west of Vista Alegre and south-west of Rio Huayllabamba where it drops to less than 900m above sea level.

A shallow zone of high density was found in the upper layer slightly south-east of the peak of the basement dome. Two more high density anomalies occur on the south-eastern flank of the



Gravity station

A-A' Cross section

-- 1.0 mgal interval

— 0.2 mgal interval

H High gravity zone

L Low gravity zone

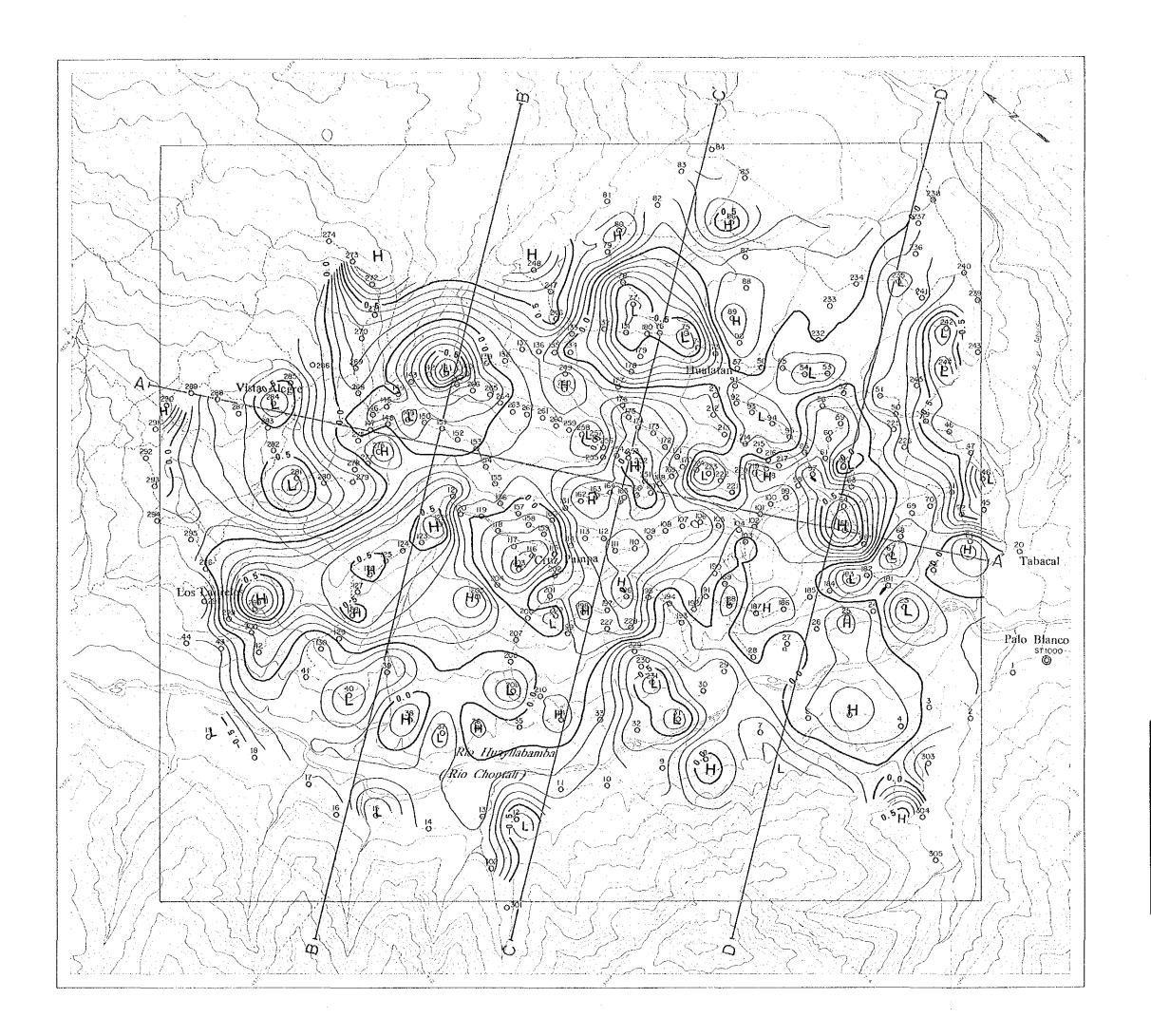
THE MINERAL EXPLORATION
IN
THE PACHAPIRIANA AREA
REPUBLIC OF PERU
(PHASE II)
GRAVITY SURVEY

Long-wave Gravity Map

Fig. II - 9 (1)

FEBRUARY 1991

0 500 1000 m



Gravity station

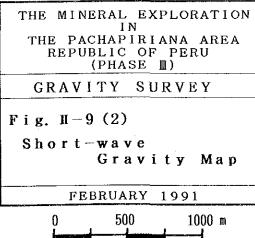
A-A' Cross section

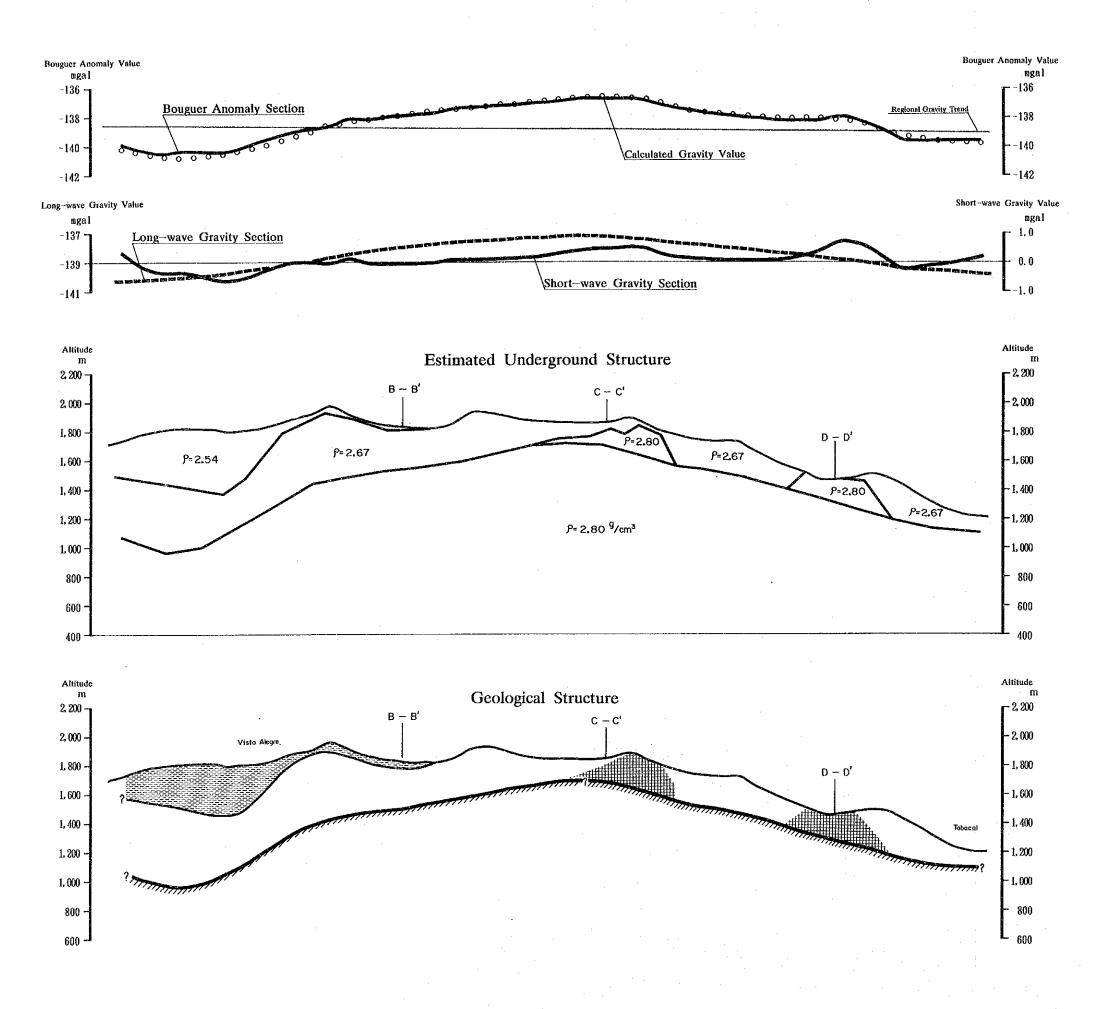
— 0.5 mgal interval

0.1 mgal interval

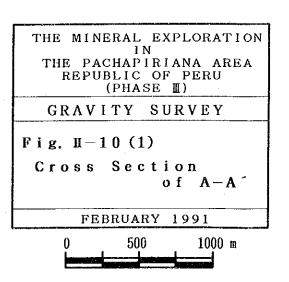
H High gravity zone

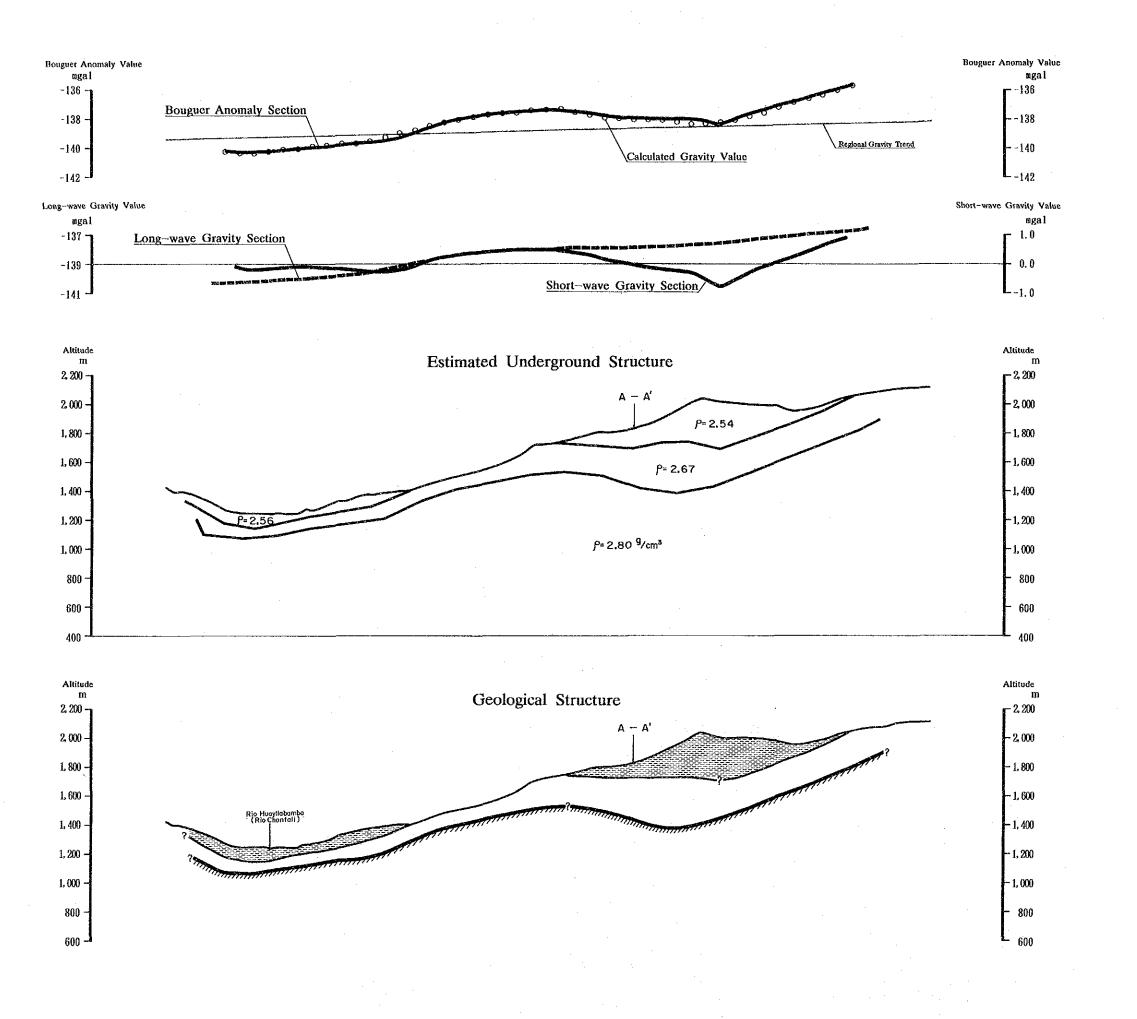
L Low gravity zone





- mm Gravity basement
- High density zone on basement
- Low density layer





- Gravity basement
- High density zone on basement
- Low density layer

