

The extracted anomalies and anomalous areas are largely arranged in the WNW-ESE direction and this is inferred to reflect the macro geologic structure of the region.

Comparing the data of this area to those of the B soil horizon of the Tavua Caldera, the gold producing area, the Cu, Pb, Zn contents of this area are somewhat smaller as seen in the following table. The data for the Tavua area in this table have been quoted from the Report of the First Phase Survey of this project. The original data have been recalculated by anti-logarithmic figures after eliminating the singular values (abnormally high values).

Contrast of Soil Assay between Tavua Caldera and Sigatoka

	Number of Samples		Average		Maximum		Minimum		unit
	Tavua	Siga'	Tavua	Siga'	Tavua	Siga'	Tavua	Siga'	
Au	56	660	85.5	2.6	*3420	20	<1	<5	ppb
Ag	57	660	-	-	<0.2	<0.2	<0.2	<0.2	ppm
Cu	52	660	131	36	406	500	59	2	ppm
Pb	52	660	9	2	120	250	<2	<1	ppm
Zn	52	660	89	81	154	800	32	1	ppm
As	56	660	6.4	0.6	*500	10	1	<1	ppm
Sb	55	660	0.2	0.4	*58.0	4.0	<0.2	<0.2	ppm
Hg	52	660	42	28	*5800	140	10	10	ppb
Mo	51	660	3.1	0.60	*4190	5	<1	<1	ppm

Siga': Sigatoka

\* : abnormal high values (excluded from calculation)



## **PART III CONCLUSIONS AND RECOMMENDATIONS**



## PART III CONCLUSIONS AND RECOMMENDATIONS

### Chapter 1 Viti Levu Island

#### 1-1 Conclusions

(1) The major geologic units of Viti Levu are mainly of Late Eocene-Early Oligocene volcanic and plutonic rocks, Late Oligocene-Middle Miocene volcanic and sedimentary rocks, Middle to Late Miocene plutonic rocks, Late Miocene-Early Pleistocene volcanic, plutonic and sedimentary rocks, and Pleistocene-Holocene sediments.

(2) Vein, network, dissemination, porphyry copper, replacement, skarn, and sedimentary type mineralization occur in this island.

The vein and dissemination types are grouped into epithermal gold and meso-hypothermal base metal mineralization. The epithermal group is further classified into adularia-sericite type and acidic sulfate type. The epithermal gold mineralized zones in the Mba and Koroimavua Volcanic Groups occur near the volcanic centers which were the source of the volcanic rocks or near the zones where these centers are inferred to have existed.

The epithermal gold mineralized zones are distributed in the ENE-WSW direction from the northern to western Viti Levu.

(3) A total of 1,060 lineaments were extracted from SLAR imageries. Many of these lineaments are considered to have been formed associated with the lateral faults caused by maximum horizontal compressional stress in three main directions. Most of the mines and mineral prospects of Viti Levu, with the exception of bedded manganese, residual, placer deposits and those of the western part, occur within the zone of lineaments formed by ENE to ESE trending horizontal stress or in the vicinity.

(4) It is seen from SLAR studies, that annular structures and caldera structures occur in the vicinity of the epithermal gold deposits of the Emperor Mine, and that annular, caldera and dome structures exist near the Namosi porphyry copper deposit. These photogeologic structures were interpreted to reflect the intrusion of magma in the area. Working on this hypothesis, 15 areas which contain at least one of the SLAR structures, namely annular, caldera or dome structures were selected. From these 15 areas, Rakiraki, east of Vatukoula, upper reaches of the Mba River, north-

east of Nadi, and South of Mba area were selected as having strong geoscientific resemblance to the area near the Emperor Mine. Also northeast of Nadi and south of Mba area were selected as areas with geologic environment very similar to the Namosi Deposit area.

(5) The locations of the centers of the volcanism of the latest Miocene-earliest Late Pliocene Mba Volcanic Group were inferred from the distribution of the volcanic rocks and the photogeologic annular, caldera and dome structures. It is considered from the above that volcanic chains existed extending in the ENE direction in northern Viti Levu and in the NW direction in the eastern part of the island. These volcanoes are believed to have formed over the deep fissure zones.

(6) Many of the lineaments formed under latest Miocene-Early Pliocene NNW to NNE compressional field are distributed in the west and northwest to southeast Viti Levu. On the other hand, NW trending deep fissures are believed to have existed from northwest to southeast Viti Levu at that time. This is inferred from the distribution of the then active volcanic rocks, locations of the volcanic centers at the time and the distribution of the above lineaments.

(7) Medium-wavelength gravity features indicate that to the north of the NE-SW trending line joining Verevere in the northeast and Sigatoka in the southwest, large scale gravity highs with circular to oval shape occur isolated in a generally low gravity area, while to the southeast of the line, high and low anomalies elongated in the NE-SW direction in belt-form are distributed alternating each other. Thus the gravity features of the two areas are clearly different. The westernmost high anomaly southeast of Nadi in the northwestern part coincides well with the distribution of the Yavuna Group, but the other three highs cannot be correlated with surface geology. The zonal distribution of the high and low anomalies in the southeast more or less coincides with the distribution of the "Wainimala Group - Colo Plutonic Suite" and "Medrausucu Group - Verata Sedimentary Group".

(8) There are large medium-wavelength gravity highs at three localities, southwest of Mba, east of Vatukoula, and west of Rakiraki. Annular, caldera, dome structures identified photogeologically and collapsed structures, intrusive bodies, altered zones, and marked short-wavelength gravity anomalies are concentrated in the centers of these medium-wavelength gravity highs. The gravity gradient of the peripheral parts of these medium-wavelength highs is steep and the shape of these anomalies is circular to oval. These are considered to indicate the existence of subsurface high density igneous bodies and it is inferred that there was a large magma cham-

ber in the deeper parts.

(9) The Emperor Mine is situated at the periphery of the collapsed structure in the center of the medium-wavelength gravity high to the east of Vatukoula, the Kingston Mine in the center of the medium-wavelength gravity high to the southwest of Mba. Together with the high west of Rakiraki, the centers of these three medium-wavelength gravity highs are considered to be the localities where active volcanism occurred repeatedly. Therefore, these are listed as promising for epithermal gold exploration. The anomalies in the northern Mba-west are on the northern extension of the medium-wavelength gravity high to the southwest of Mba, and with the coincidence of SLAR annular and caldera structures and the short-wavelength gravity highs, it is believed that the northern Mba-west was the place of activity of small magma chamber branched out from large magma chamber. The area is listed as promising for epithermal gold deposit occurrence.

(10) The zones where volcanic centers are inferred to have existed in northern Viti Levu correspond to the zones of short-wavelength high gravity anomalies related to basaltic activities. It is believed that since the contents of the magmatic chambers of that time has changed from basalt to olivine-gabbro of higher density, there are positive gravity anomalies near the altered volcanic centers. However, even in cases of Kilauea type caldera, the interior of the caldera is filled with thick compact lava which is more dense than the whole volcano and thus the center of eruption would show somewhat higher density.

(11) The Tavua Caldera whose upper parts are filled by low density formations such as andesitic pyroclastics and lacustrine sediments show short-wavelength low gravity anomaly surrounded by gravity lineaments. The SLAR annular structure zone near Rakiraki in northeast Viti Levu and the vicinity of the volcanic centers west of Mba are the zones which have Tavua Caldera type gravity structure among the possible collapse caldera zones extracted photogeologically.

(12) The geologic environment necessary for the formation of epithermal deposits is the existence of magmatic heat, subsurface fractures and circulating water. The magmatic heat and the subsurface fractures are mostly likely to exist in volcanic collapsed and volcanic dome structures. The circulating water formed the mineralized and altered zones. Structures which are likely to be volcanic collapse and volcanic domes were extracted through photogeologic studies of annular, caldera and dome structures; short-wavelength gravity anomalies; and field survey. Of these zones; vicinity of Rakiraki, Tavua Caldera zone, area west of Mba to the

southern part, Sabeto Range, south of Lautoka and Namosi area are considered to contain high potential for locating mineralized and altered zones.

Regarding the selection of the survey areas for the second phase and further work, the area from west of Mba to southern Mba was chosen because it was relatively unexplored; the Sigatoka area because it was considered to have relatively high mineral potential from the occurrences of alteration and mineralized zones (porphyry copper, skarn and other types) around plutonic bodies.

(13) It was shown by the geochemical orientation survey that gold tends to be concentrated in the A soil horizon. Also it was clarified that zones of gold concentration are continuous vertically in the A and B soil horizons. It was shown by this survey that Au has high positive correlation with As, Te, Sb, Hg, F, Tl in the A horizon, while As, Hg, Pb have positive correlation with Au in the B horizon.

## **Chapter 2 Mba-west Area**

### **2-1 Conclusions**

(1) The main geologic unit of Mba-west are Miocene-Pliocene andesitic/basaltic volcanic products and limestone; Pliocene basaltic/andesitic volcanic products, sandstone, and conglomerates; Holocene alluvium; and intrusive rocks (monzonite, dacite, andesite, basalt) penetrating the Pliocene formations. The Miocene and Pliocene formations largely dip northward at low angles and are superposed. Thus the strata become younger northward.

(2) A total of 95 lineaments was extracted photogeologically in the Mba-west. Many of these are concentrated in the southern and northern parts of the area. The directions of the maximum horizontal compressional stress axes were inferred to be NNW to NNE and ENE to ESE from the en echelon arrangement of the lineaments.

(3) Many lineaments of Mba-west are developed near the inferred volcanic centers in north and south, also lineaments with various trends are developed within the photogeologic annular structures. Also in short-wavelength low gravity zones and in parts of the short-wavelength high gravity zones, lineaments parallel to the elongation of the zones are developed in and near the zones. This is interpreted as reflecting fractures which were developed as the result of the vertical block movement accompanying the



rise of magma.

(4) The circular depression extracted in southern Mba-west as photogeological annular structure is believed to have been the center of volcanic activity from the distribution of the volcanic products and intrusive bodies. A large scale medium-wavelength gravity high is distributed throughout this area. This gravity high is believed to reflect high density rock bodies (deep-seated bedded basic intrusive bodies) formed by the solidification of the magma chamber which supplied the volcanic products of this area. The above annular structure is located near the center of this high gravity anomaly.

(5) Propylitized zones and sericitized zones are developed near the southern photogeological annular structures, and geochemical anomalies related to Au mineralization and auriferous quartz veins occur overlapping some of these altered zones. These features regarding geologic structure and mineralization/alteration are very similar to those of the Emperor Mine area. It is anticipated that low sulfidation epithermal gold mineralization akin to that of the Emperor Mine would exist in this area.

(6) Photogeological caldera structures were extracted at three localities in northern Mba-west and volcanic products are distributed in the vicinity. These calderas all occur in short-wavelength high gravity zones. This reflects the fact that these calderas are crater and/or volcanic collapsed structures and that the short-wavelength highs are caused by shallow high density rocks. These shallow bodies are considered to be small magma chambers formed as offshoots of the large, deeper chamber whose existence is inferred from medium-wavelength gravity high.

(7) Acidic alteration zones accompanied by silicification are developed in some of the photogeological calderas in northern Mba-west. Geochemical anomalies related to Au mineralization occur overlapping these altered zones. This is of the high sulfidation epithermal gold mineralization. This type is considered to form under shallower environment than the low sulfidation type. The results of drilling at Namosau Alteration Zone of this year, showed that the deposits could have been eroded out. The conditions of the lower parts of the Raviravi Alteration Zone is not clear, and the possibility of the occurrence of gold deposits has not died.

(8) Marked Au, As, Te geochemical anomaly zones which coincide with the altered/mineralized zones on the surface were extracted at four localities (Raviravi, Nalotawa, Nanuku-Yaloku and Tavanasa Creek) in Mba-west area. Aside from the above, small geochemical anomalies not associated with

alteration/mineralization were confirmed at several localities (Namosau Creek, Lololo Creek, Nayanggali Creek, Tauarau Creek, Koroniviria and Karawa) and blind buried altered/mineralized zones were anticipated to occur in shallow subsurface parts.

(9) The mineralization of Mba-west was brought about by hydrothermal activities related to Pliocene volcanism. And it is considered that high sulfidation type epithermal gold mineralization occurred above the shallow small magma chamber while low sulfidation type occurred near the volcanic center in the central part of the deep and large scale chamber

(10) Drilling was conducted at four localities of Mba-west Area. The following conclusions were obtained.

Namosau Creek Alteration Zone: Two holes drilled in this zone penetrated through basalt lava and basaltic pyroclastics of the Pliocene Namosau Volcanics belonging to the Ba Volcanic Group and confirmed wide argillized zone accompanied by pyrite dissemination, but promising Au mineralization could not be confirmed.

Nayanggali Creek Geochemical Anomalous Zone: The subsurface geology confirmed by drilling (MJF-3) comprises basalt lava, basaltic pyroclastics, and sedimentary rocks of the Ba Volcanic Group, basalt lava and basaltic pyroclastics of the Namosau Volcanics, and basalt dykes.

Mineral showings and alteration of significance are not found in this zone. The central part of volcanic activities probably occurred in this zone and NE-SW trending fractures are inferred to exist in the deeper parts. The Au, As, Hg geochemical anomalies are inferred to be the products of ascending post volcanic hydrothermal fluids along the NE-SW trending fissures. Subsurface gold mineralization of this zone, if any, is concluded to be small.

Nalotawa Alteration Zone: The subsurface geology comprises basalt lava, basaltic pyroclastics of the Koroyanitu Volcanic Products and intrusive bodies (basalt, hornblende andesite, altered andesite).

There are many clay-pyrite veins, but evidences of gold mineralization do not exist on the surface.

On the other hand, in the subsurface parts, occurrence of gold in quartz-calcite veins, calcite veins, and clay-pyrite-(calcite) network is confirmed by drilling. The best part contains Au 0.176g/t in 18.10m of the

drill core (include 1m of Au 0.52g/t).

In this zone, the assemblage of major gangue minerals (quartz, calcite, adularia, smectite, sericite) and that of the major alteration minerals in the host rock near the veins (quartz, calcite, pyrite, smectite) is very close to that of the low sulfidation epithermal veins.

The potential for gold occurrence in the deeper parts of this zone is concluded to be high.

Yaloku Alteration Zone: The subsurface geology comprises andesite lava, andesitic pyroclastics, basalt lava of the Sabeto Volcanics, and basalt dykes.

Quartz veins, clay-pyrite veins, and calcite veins occur on the surface of this zone. These veins are divided into the western and eastern groups. Auriferous quartz veins occur in both groups and the highest grades are 12.10g/t (15cm wide) in the western group and 4.52g/t (3cm wide) in the east.

In the east, calcite-quartz network with Au 0.114g/t (sampling width 40cm) was confirmed by drilling, but generally the development of the veins is poor. The auriferous quartz veins exposed on the surface deteriorates downward. The potential for gold vein occurrence is concluded to be poor in the eastern side.

In the west, although of low grade, a large number of auriferous veins was confirmed by drilling. Regarding N-S trending veins, a group of relatively wide auriferous veins (Au 0.055g/t, sampling width 400cm, clay-calcite-dolomite vein; Au 0.20g/t, sampling width 15cm, pyrite-calcite-dolomite vein; others) was confirmed almost at the lower extension of the exposed gold-bearing quartz vein (Au 12.10g/t). With ENE-WSW to E-W veins, the grade of the downward extension of the exposed vein (Au 2.19g/t) deteriorates, but a different group of auriferous veins (Au 0.375g/t, Ag 880g/t, Cu 6.76%, sampling width 3cm, chalcopyrite vein; others) were found by drilling.

The common ore minerals of this zone are chalcopyrite and pyrite, with rare association of molybdenite, bornite, galena, and stromeyerite in the west. This mineral assemblage corresponds to those of the high temperature epithermal deposits formed at relatively deeper parts.

The assemblage of the main gangue minerals is quartz-smectite-chlor-

ite-calcite in the east while it is quartz-potash feldspar in the west. Adularia is associated in both groups at times.

The alteration mineral assemblage of the host rock near the veins also differ between the western and the eastern groups. The assemblage common for both groups is quartz-chlorite-calcite-smectite with sericite in the east and potash feldspar in the west.

The above assemblages of gangue and alteration minerals are very similar to those of the low sulfidation epithermal mineralization. It is concluded that the veins in the west were formed under higher temperature.

The mode of occurrence of the veins confirmed by drilling in this zone corresponds to the quartz + adularia + illite + Ag sulfides + base metal sulfide zone of the low sulfidation (quartz-adularia type) epithermal model (Berger and Eimon, 1983). Therefore, it is believed that the three boreholes were drilled below the bonanza.

Regarding the drill holes in the west, it is inferred from the above model that the bonanza lies higher than the gold showings confirmed by drilling. This gold occurrence is only about 70m below the surface. Therefore, the potential of these veins are controlled by the topography and the direction of the ore shoots. There is not sufficient data for determining the direction of the shoots.

## **2-2 Recommendations for Future Exploration**

### **(1) Nalotawa Alteration Zone**

A total of three holes is recommended to be drilled in order to confirm the state of gold mineralization of the veins located by MJF-4. These veins are inferred to extend in the NNE-SSW direction and two holes should be drilled westward from the eastern side of MJF-4. Also one drilling should be made south-westward from MJF-4 in order to explore the lower parts of the NE-SW veins which exist in this zone.

### **(2) Yaloku Alteration Zone**

A total of three holes is recommended to be drilled as follows in order to confirm the state of gold mineralization of the auriferous veins located by MJF-6 and -7 in western part of this zone. One hole should be drilled westward for exploring the N-S trending veins to the south of MJF-6. Also one northward hole should be drilled each from the east and west of MJF-7 in order to explore the ENE-WSW to E-W veins.

(3) It is recommended that one hole be drilled westward from the eastern side of the geochemical anomalies of eastern Yaloku in order to confirm the subsurface mineralization of the anomalies to the north of MJF-5.

(4) It is recommended that one vertical hole be drilled at the gold anomalies of Raviravi Alteration Zone for confirming the high sulfidation epithermal gold mineralization.

### **Chapter 3 Sigatoka Area**

#### **3-1 Conclusions**

(1) The geology of Sigatoka consists of; Miocene basaltic and/or andesitic volcanic products, and detrital sediments; Pleistocene (?) fluviatile sediments; and intrusive bodies (granodiorite porphyry-diorite porphyry bodies, granodiorite, diorite, diorite porphyry, quartz porphyry, aplite, basalt, andesite, dacite, and rhyolite) penetrating Miocene Series. The Miocene units largely dip southwestward are superposed.

(2) Most of the mineralized/altered zones occur near the above faults, near the Colo Plutonic Suite bodies, near the SLAR lineaments, and near the en echelon dykes. Also some of them occur within the Colo Plutonic Suite bodies.

(3) Four large geochemical anomalous zones were extracted in the Sigatoka area. These four zones coincide with the surface mineralized/altered zones, and many other small anomalies were also extracted. These are believed to be anomalies related to the activities of the Colo Plutonic Suite which occur extensively in the deeper parts.

(4) The mineralization of the Sigatoka area is closely related to the activities of the Colo Plutonic Suite and they are emplaced in fractured zones in the vicinity of the plutonic and porphyry bodies. And the mineralization took the form of veins, replacement, porphyry, and other types of meso- to hypothermal activity.

(5) The intensity of the mineralization/alteration is weak with some exceptions. The intensity of the geochemical anomalies is also generally low. Many of the altered zones and anomalies have been drilled without significant success. There are two undrilled localities where multi-component anomalies are noted. If large deposits are to be anticipated, the weak surface manifestation indicates deep occurrences.

## REFERENCES

Ahmad M., Solomon M. and Walshe J.L. (1987): Mineralogical and studies of the Emperor gold telluride deposit, Fiji: *Econ. Geol.*, **82**, 345-370.

Anderson W.B. and Eaton P. (1989): Gold mineralization at the Emperor Mine, Vatukoula, Fiji: *Journal of Geochemical Exploration*, **36**(1990), 267-296.

Baltis E.J. and Levy I.W. (1985): The 1985 gravity survey of the Tavua Basin caldera contact: T.B.J.V. Technical Report, No.10.

Band R.B., B.Sc., A.R.S.M (1968): The Geology of Southern Viti Levu and Mbengga: Ministry of Natural Resources, Department of Geological Surveys.

Berger B.R. and Eimon P.I. (1983): Conceptual models of epithermal precious metals deposits; in Sanks, W.C. (ed.), *Cameron Volume on Unconventional Mineral Deposits*, Society of Mining Engineers p.191-205.

Colley H. and Greenbaum D. (1980): The Mineral Deposits and Metallogenesis of the Fiji Platform: MRD, *Econ. Geol.*, **75**, 807-829.

Colley H. (1976): Mineral Deposits of Fiji (metallic deposits): *Mem. Miner. Resour. Div. Fiji, Legts. Counc. Pap.* 1910(19).

Colley H. (1986): Epithermal Gold Mineralization associated with Mio-Pliocene Volcanism in Fiji: *International Geological Congress, 1986, Proceedings of Symposium 5*, p.29-35.

Drake R.E., Kollman E., Whelan P.M. and Gill J.B. (1985): Radiometric Dating of Magmatic Stages in Fiji: *Economic Geology*, p.415-440.

Gill J. and Whelan P. (1989): Early Rifting of an Oceanic Island Arc (Fiji), Produced Shoshonitic to Tholeiitic Basalts: *Journal of Geophysical Research*, **94**(B4), 4561-4578.

Gill J. and Whelan P. (1989): Postsubduction Ocean Island Alkali Basalts in Fiji: *Journal of Geophysical Research*, **94**(No.B4), 4579-4588

Gill J.B. and Stork A.L. (1979): Miocene Low-K Dacites and trondhjemites of Fiji: *Trondhjemites, Dacites, and Related Rock* (ed F.Barker), Elsevier, p.629-650.

Gill J.B.(1987): Early Geochemical Evolution of an Oceanic Island Arc and Backarc, Fiji and the South Fiji Basin. *Journal of Geology*, **95**, 589-615.

Gill J.B.: Sr-Pb-Nd Isotopic Evidence that Both MORB and OIB Sources Contribute to oceanic island arc magmas in Fiji.

Hamburger, M.W., Everingham I.B. and Isacks, B. and Barazangi, M. (1988): Active tectonism within the Fiji platform, southwest Pacific. *Geol.*, **16**, 237-241.

Hawkes, H.E. and Webb, J.S., (1962): *Geochemistry in mineral exploration*. Harper and Row, New York, N.Y. 28-31.

Hayaba, D.O., Bethke, P.M., Heald, P. and Foley, N.K. (1985): Geologic, mineralogic, and geochemical characteristics of volcanic-hosted epithermal precious metal deposits. In *Geology and geochemistry of epithermal system: Rev. Econ. Geol.*, v. 2, 129-167.

Heald, P., Foley, N.K. and Hayaba, D.O. (1987): Comparative anatomy of volcanic-hosted epithermal deposits. *Econ. Geol.*, **82**, 1-26.

Hedenquist, J.W. (1987): Mineralization associated with volcanic-related hydrothermal systems in the circum-Pacific basin. In *Transactions of the Fourth Circum Pacific Energy and Mineral Resources Conference*, Singapore. Horn, M.K. Ed., *Am. Assoc. Pet. Geol.*, 513-524.

Hirst J.A. (1965): *Geology of east and north-east Viti Levu*: *Bull. Geol. Surv. Fiji*, 12.

Ibbotson P., B.Sc., Ph.D., A.R.C.S., F.G.S. (1967): *Petrology of the Tertiary Caldera, Tavua Goldfield*: Geological Survey Department.

Ishihara S. and Urabe T. (1989): Gold mineralization of immature island arcs, Fiji: *Chishitsu News*, No. 415, p. 18-31 (in Japanese).

Jezeq, P. (1976): *Gravity base stations in Indonesia and in the southwest pacific*. Technical Report, Woods Hole Oceanographic Institution.

JICA and MMAJ (1991): *Report on the Cooperative Mineral Exploration in the Viti Levu Area, the Republic of Fiji, Phase I*.

JICA and MMAJ (1992): *Report on the Cooperative Mineral Exploration in the Viti Levu Area, the Republic of Fiji, Phase II*.

Koide H.(1982): Analysis of mechanism of diapir formation and tectonics: Gekkan Chikyu, 4, p.15-22.(in Japanese).

Kouda R. and Suwijanto (1989): Volcanic collapse structure and gold-silver exploration-discovery: Chishitsu News, No.423, p.13-26. (in Japanese).

Kwak T.A.P.(1989): Geochemical and temperature controls on ore mineralization at the Emperor gold mine, Vatukoula, Fiji: Journal of Geochemical Exploration, 36(1990), 297-337.

La Porte,M.(1962): Elaboration rapide de cartes gravimetriques deduites, del anomalie de Bouguer a laide dune calculatrice electronique:Geophys. Prosp.,10, 238-257.

Lawrence J.L. and Savage E.N.(1976): Ore Genesis in the Wainivesi Area, Fiji, and some Exploration Implications, p.59-68.

Lepeltier,C.(1969): A simplified statistical treatment of geochemical data by graphical representation. Econ. Geol., 64, 538-550.

Mallick D.I.J.and Habgood F.(1987): Interpretation of SLAR imagery of the main islands in Fiji: British Geological Survey, p.1-9.

MMAJ (1986): Gold Deposits of the World, Information Center of Metal Mining Agency of Japan, p.4-122 (in Japanese).

Okuda Y.(1989): Geology of Fiji: Chishitsu News, No.415, p.6-17 (in Japanese).

Otsu,H.,Kubota,R., and Matsuda,Y.(1983):Determination of statistical frequency of geochemical data. Mining Geol., 33, 427-431.

Rodda P. B.Sc.(1969): Analysis of Rocks from Fiji: Ministry of Natural Resources Department of Geological Surveys.

Rodda P.and Duberal,R.(1966): Specific gravity of Viti Levu rocks. G.S. Note:23/66, Geological Survey Department.

Rodda P.(1967): Radiometric Age Data on Rocks from Viti Levu, Fiji: Geological Survey of Fiji, p.1249-1259.

Rodda P.(1976): Geology of northern and central Viti Levu: Bull. Miner. Resour. Div. Fiji, 3.



Rodda, P. (1989): Geology of Fiji. MRD.

Rugless C.S. (1983): Lithogeochemistry of Wainaleka Cu-Zn Volcanogenic Deposit Viti Levu, Fiji, and Possible Applications for Exploration in Tropical Terrains: Journal of Geochemical Exploration, p.563-586.

Scheibner, E., SATO, T., Douth H.F., Addicott W.O., Terman, M.J. and Moore, G.W. (1991): Tectonic Map of the Circum-Pacific Region, South Quadrant, scale 1:10,000,000. U.S. Department of the Interior, U.S. Geological Survey.

Setterfield T.N. (1990): The Tavua Caldera, Fiji: A Complex Shosonitic Caldera formed by concurrent faulting and downsagging: p.1-43.

Silberman M.L. and Berger B.R. (1985): Relationship of trace-element patterns to alteration and morphology in epithermal precious-metal deposits: Reviews in Economic Geology, 2, 203-232.

Sinclair, A.J. (1974): Selection of threshold values in geochemical data using probability graphs. J. Geochem. Explor., 3, 129-149.

Stephen T. (1986): Fluid Inclusion, Alteration and Ore Mineral Studies of an Epithermal Vein System, Mount Kasi, Vanua Levu, Fiji: International Geological Congress, 1986, Proceedings of Symposium 5, p.87-94.

Stoffregen, R. (1987): Genesis of acid-sulfate alteration and Au-Cu-Ag mineralization at Summitville, Colorado. Econ. Geol., 82, 1575-1591.

Talwani, M., Worzel, J.L. and Landisman, M. (1959): Rapid gravity computations for two dimensional bodies with application to the Mendocino Submarine fracture zones: Jour. Geophys. Res., 64, p.49-59.

The Geodetic Society of Japan (1989): Accurate positioning by GPS-satellite system: Japanese Association of Surveyors.

Thomas G. and Jones D.G. (1989): South Pacific Deposits: The Geology Department & University Extension, The University of Western Australia, Publication, No.17, p.1-32.

Yokoyama I. (1963): Structure of caldera and gravity anomaly: Bull. Volcanol., 26, p.67-72.



## APPENDICES



**Table 1 Results of Radiometric Age Determination**

(Phase I, Viti Levu Area)

Sample No.	Locality	Rock Name	Sample Type	Potassium (K wt%)	Rad. <sup>40</sup> Ar (10 <sup>-6</sup> cc/g)	K-Ar Age (Ma)	Air Cont. (%)
		Formation					
A-2	NE of Singatoka N of Tuvu	Ilb-Ad	Whole rock	1.00±0.03	24.2±0.7	6.23±0.26	46.7
		(Nva)					
A-5	NE of Singatoka Korolevu	Ilb-Ad	Whole rock	0.94±0.05	26.3±0.8	7.20±0.49	48.7
		(Ta)					
C-2	W of Nanukuloa	Micro-Dio	Whole rock	3.97±0.08	56.8±1.4	3.68±0.12	45.7
		(ND)					
C-7	Vaturu Dam Site	Ad	Whole rock	0.86±0.05	16.7±0.7	5.00±0.37	61.1
		(Ks)					
C-8	E of Vaturu Dam Mbukuya	OI-Bs	Whole rock	1.87±0.06	33.9±1.4	4.69±0.23	56.1
		(Ks)					

Abbreviations: Ilb-Ad; Hornblende Andesite, OI-Bs; Olivine Basalt, Dio; Diorite

**Table 2 Results of Whole Rock Analysis**

(Phase I, Viti Levu Area)

Sample No.	C038	BA102	C070	AA091	C002	C007	C008	NA052	A002	A005
SiO <sub>2</sub>	47.850	47.870	51.270	50.110	47.840	45.480	49.460	52.760	57.280	57.840
TiO <sub>2</sub>	0.600	0.620	0.860	0.740	0.750	0.620	0.770	0.800	0.560	0.580
Al <sub>2</sub> O <sub>3</sub>	12.640	12.730	17.060	17.090	18.040	17.860	18.540	19.170	19.100	16.470
Fe <sub>2</sub> O <sub>3</sub>	4.807	5.701	4.569	5.251	3.482	6.383	4.745	4.619	4.005	3.365
FeO	5.060	4.660	4.140	4.480	4.110	2.220	3.730	3.510	1.570	2.830
MnO	0.170	0.190	0.170	0.220	0.160	0.190	0.180	0.200	0.090	0.150
MgO	9.530	8.520	4.760	4.180	3.070	4.400	3.600	2.980	2.670	2.680
CaO	11.470	12.110	7.780	9.060	7.020	9.030	9.310	9.890	7.360	5.920
Na <sub>2</sub> O	1.660	1.580	3.280	3.120	3.200	5.130	3.040	2.870	4.110	3.200
K <sub>2</sub> O	2.420	3.150	2.860	4.470	5.450	1.420	2.400	1.910	1.370	1.450
P <sub>2</sub> O <sub>5</sub>	0.410	0.470	0.450	0.730	0.870	0.560	0.360	0.390	0.220	0.160
BaO	0.060	0.060	0.060	0.070	0.100	0.050	0.040	0.030	0.100	0.060
LOI	1.840	0.015	1.870	1.230	5.020	4.850	1.250	1.820	1.510	2.660
Total	98.557	97.676	99.129	100.751	99.112	98.293	97.425	100.949	99.945	97.355
FeO <sub>t</sub>	9.386	9.791	8.252	9.206	7.244	7.965	8.000	7.667	5.175	5.849
Fe/Mg	0.980	1.149	1.734	2.202	2.360	1.810	2.222	2.573	1.938	2.183
S. I.	40.745	42.494	43.087	43.888	38.199	42.108	46.949	49.699	38.835	44.383
Q	0.000	0.000	0.175	0.000	0.000	0.000	0.419	6.221	10.198	17.108
or	14.302	18.617	16.903	26.418	32.210	8.392	14.184	11.288	8.097	8.570
ab	13.954	11.426	27.738	17.506	13.455	24.318	25.709	24.271	34.757	27.062
an	19.938	18.341	23.386	19.430	18.768	21.795	29.859	33.789	29.630	26.300
ne	0.000	1.049	0.000	4.811	7.372	10.329	0.000	0.000	0.000	0.000
di-wo	14.320	16.150	5.132	8.679	4.353	8.087	5.839	5.319	2.275	0.846
di-en	10.779	12.530	3.775	6.123	2.743	6.990	4.316	3.921	1.966	0.609
di-fs	2.100	1.877	0.868	1.810	1.339	0.000	0.960	0.859	0.000	0.160
hy-en	3.162	0.000	8.075	0.000	0.000	0.000	4.646	3.498	4.680	6.062
hy-fs	0.616	0.000	1.857	0.000	0.000	0.000	1.034	0.793	0.000	1.587
ol-fo	6.944	6.063	0.000	3.002	3.434	2.778	0.000	0.000	0.000	0.000
ol-fa	1.491	1.004	0.000	0.978	1.847	0.000	0.000	0.000	0.000	0.000
nt	6.966	8.263	6.622	7.611	5.047	5.978	6.877	6.695	3.730	4.862
hp	0.000	0.000	0.000	0.000	0.000	2.257	0.000	0.000	1.431	0.000
il	1.140	1.178	1.634	1.406	1.425	1.178	1.463	1.520	1.064	1.102
ap	0.971	1.113	1.065	1.728	2.060	1.326	0.852	0.923	0.521	0.379
TOTAL	96.680	97.630	97.200	99.490	94.040	93.430	96.150	99.120	98.350	94.640
Fenic Total	48.489	48.199	29.027	31.336	22.247	28.594	25.987	23.557	15.668	15.607

FeO<sub>t</sub>: total iron

S. I.: Solidification index (Kuno et al. 1957)



Table 4 Results of Microscopic Observation of Polished Section

(Phase I, Viti Levu Area)

Sample No.	Locations		Material	Minerals								Note	
	No.	Mines/Prospects		Py	Cpy	Sph	Cov	Goe	Bes	Ba	Qz		
C-36	54	Baleboto	Py vein	○	*	*	*	*	*	*	*	*	
CA-115	-	Bakiraki	Qz vein					△	△			◎	* 1
EM-5	56	Eperor	Qz vein	△	*	*	*					◎	* 2
LC-17	53	Voda	Qz-Alu vein	○									
VA-126	41	Vailotu	Poedery sulfide	○	○	○							

Location No. denoted as the number in the list and map of the prospects and mines  
 Abundance of Minerals: ◎: abundant, ○: common, △: a few, \*: trace  
 Abbreviations: Py: Pyrite, Cpy: Chalcopyrite, Sph: Sphalerite, Cov: Covellite,  
 Goe: Goethite, Bes: Besmitite, Ba: Barite, Qz: Quartz, Alu: Alunite  
 \* 1: Goethite or lepidochroite.  
 Besmitite denote the pseudomorph of pyrite.  
 \* 2: Covellite are paragenetic with pyrite and are denoted as the pseudomorph of mundesite.

(Phase II, Drilling Cores of Mba-west Area)

No.	Location	Description	Cp	Po	Py	Mg	If	Goe	Bes	Remarks
PS-1	MJF-1, 87.9	Qz-Alu vein			○			△	△	
PS-2	112.9	Py dis brecc rock	△		○					
PS-3	232.8	Py dis brecc rock	△		○					
PS-4	MJF-2, 111.1	Qz-Alu vein			○					partly replaced by sphalerite
PS-7	287.8	Py dis brecc rock	△		○	△	△			

Abbreviations:  
 ◎: Abundant ○: Common △: Few ▲: Rare  
 Cp: Chalcopyrite, Po: Pyrrhotite, Py: Pyrite, Mg: Magnetite, If: Ilmenite, Goe: Goethite, Bes: Besmitite  
 Alu: Alunite, dis: disseminated, brecc: brecciated

(Phase III, Drilling Cores of Mba-west Area)

No.	Location	Description	Cp	Bo	Po	Py	Mg	If	Goe	Bes	Sph	En	St	Remarks
P4-1	MJF-4 51.0a	Qz-Cal-Py vein				○								
P4-2	100.5	Clay-Py network				○								
P4-3	117.0	Clay-Cal-Py network				○			△					partly replaced by Goe
P4-4	119.4	Clay-Cal-Py network				○			△					partly replaced by Goe
P4-5	144.3	Cal vein				○								
P4-6	145.0	Py-Cp dis rock	▲			○	○	○						
P4-7	163.4	Silica-Cal-Py vein with black band	△			○	○	○			△			
P4-8	196.5	Reddish gray mineral dis rock				△	△	△	△					only inclusions in Py
P4-9	250.0	Cal-Py-Clay vein				○			△					partly replaced by Goe
P4-10	249.0	Cal-Py-Lis vein				○	△	△						
P6-1	MJF-6 189.5a	Py vein	△			◎								
P6-2	181.8	Cal-Cp vein (network)	○			△	△						△	
P7-1	MJF-7 121.95a	Cal-Cp-Gn vein (network)	◎	△		△						○		▲
P7-2	123.5	Cp dis massive ore	◎											▲
P7-3	216.2	Py dis alt. rock	△			◎								

Abbreviations:  
 ◎: Abundant ○: Common △: Few ▲: Rare  
 Cp: Chalcopyrite, Bo: Bornite, Po: Pyrrhotite, Py: Pyrite, Mg: Magnetite, If: Ilmenite, Goe: Goethite, Bes: Besmitite, Sph: Sphalerite,  
 Gn: Galena, En: Enhydrosite, St: Stannoselite  
 Qz: Quartz, Lis: Lisovite, Cal: Calcite, dis: disseminated, alt: altered









Table 8 Results of Ore Assaying

(Phase I, Viti Levu Area)

Sample No.	Locations	Material	Assay Results					
			Au(g/t)	Ag(g/t)	Cu(%)	Pb(%)	Zn(%)	Mo(%)
C-36	54 Balebuto	Py vein	<0.07	<0.3	0.01	<0.01	<0.01	<0.001
CA-115	- Rakiraki	Qz vein	0.41	<0.3	0.01	<0.01	<0.01	<0.001
EF-5	56 Eaperor	Qz vein	0.14	<0.3	0.10	<0.01	<0.01	<0.001
LC-17	53 Vuda	Qz-Alu Vein	0.07	<0.3	0.06	<0.01	<0.01	<0.001
MC-48	- West of Nba	Py diss.	<0.07	<0.3	<0.01	<0.01	<0.01	<0.001

Location No. denoted as the number in the list and map of the prospects and mines

Abbreviations: Py: Pyrite, Qz: Quartz, diss: dissemination, Alu: Aluminite

(Phase I)  
(Nba-west Area)

Sample No.	Location	Description	Dip-strike	Width (cm)	Ore Grade							
					Au g/t	Ag g/t	Cu %	Pb %	Zn %	Mo %	Te %	
OK1	Naviravi	Qz vein	N6° W, 60° E	1	<0.07	<0.3	<0.01	<0.01	0.05	0.001	<0.001	
AK16	-	Silica-Gossan	-	-	<0.07	<0.3	0.04	<0.01	0.01	<0.001	<0.001	
AK17	-	Silica-Gossan	-	-	<0.07	<0.3	0.02	<0.01	<0.01	<0.001	<0.001	
SFS	-	Silica-Gossan	-	-	<0.07	<0.3	0.01	<0.01	<0.01	<0.001	<0.001	
ST1	Nanosau Creek	Silica-Gossan	-	-	<0.07	<0.3	0.01	<0.01	<0.01	<0.001	<0.001	
SM2	-	Qz-Alu vein	N27° W, 90° E	5	<0.07	<0.3	0.04	<0.01	<0.01	<0.001	<0.001	
SM4	-	Qz-Alu vein	N35° W, 70° E	100	<0.07	<0.3	0.04	<0.01	0.01	<0.001	<0.001	
SM101	-	Gossan float	-	-	<0.07	0.5	0.01	<0.01	<0.01	<0.001	<0.001	
SM102	-	Silicified rock	N10° E, 90° E	50	<0.07	<0.3	0.01	<0.01	0.01	<0.001	<0.001	
SM104	-	Silicified rock	N5° E, 90° E	170	<0.07	<0.3	0.01	<0.01	<0.01	<0.001	<0.001	
SM8	Nalotava	Lim. network	-	-	<0.07	<0.3	0.01	<0.01	0.01	<0.001	<0.001	
SM19	Tavanasa Creek	Qz-Alu float	-	-	<0.07	<0.3	0.01	<0.01	<0.01	<0.001	<0.001	
SM11	Valoku	Qz vein	N75° E, 80° S	15	<0.07	<0.3	0.02	<0.01	<0.01	<0.001	<0.001	
OK2	-	Qz vein	N40° W, 25° W	15	<0.07	2.0	0.05	0.05	0.01	<0.001	<0.001	
OK4	-	Qz vein	N77° W, 60° S	25	<0.07	1.0	0.04	0.07	0.01	<0.001	<0.001	
OK7	-	Qz vein	N15° E, 60° E	2	<0.07	1.0	0.04	<0.01	0.01	<0.001	<0.001	
OK10	-	Qz vein	N5° E, 80° E	15	12.10	2.7	0.03	0.02	0.01	<0.001	<0.001	
OK12	-	Qz vein	N66° E, 70° S	10	2.19	0.8	0.05	0.24	0.01	<0.001	<0.001	
OK16	-	Qz vein	N17° W, 80° W	5	<0.07	0.9	<0.01	<0.01	<0.01	<0.001	<0.001	
OK18	-	Qz vein	N90° W, 70° W	3	4.52	11.8	3.58	<0.01	0.01	0.020	<0.001	
OK1	-	Qz vein	N5° E, 45° W	5	<0.07	<0.3	0.01	<0.01	0.01	<0.001	<0.001	
AY8	-	Qz vein	N47° W, 90° S	5	<0.07	<0.3	0.10	<0.01	0.01	<0.001	<0.001	
AY10	-	Qz vein	N75° W, 80° W	3	<0.07	<0.3	0.02	<0.01	0.01	<0.001	<0.001	
ST2	-	Qz vein	N44° E, 15° S	2	<0.07	<0.3	0.01	<0.01	0.01	<0.001	<0.001	
ST3	-	Qz vein	N10° W, 80° E	3	<0.07	<0.3	0.07	<0.01	0.01	<0.001	<0.001	
ST4	-	Qz vein	N10° W, 80° E	15	<0.07	<0.3	0.04	<0.01	0.01	<0.001	<0.001	

(Phase I)  
(Sigatoka Area)

Sample No.	Location	Description	Dip-strike	Width (cm)	Ore Grade							
					Au g/t	Ag g/t	Cu %	Pb %	Zn %	Mo %	Te %	
ST201	Talasona Creek	Gossan	N60° E, 60° W	0.4	<0.07	<0.3	<0.01	0.01	0.01	<0.001	<0.001	
OK201	-	Quartz vein	N57° W, 70° S	1.0	<0.07	<0.3	<0.01	<0.01	0.01	<0.001	<0.001	
SM202	-	Silicified rock	N87° W, 55° S	1.2	<0.07	<0.3	<0.01	<0.01	<0.01	<0.001	<0.001	
SM203	-	Quartz float	-	1.3	<0.07	<0.3	<0.01	<0.01	<0.01	<0.001	<0.001	
OK104	Nalavaleva Creek	Gossan	-	-	<0.07	1.0	0.05	0.01	0.01	<0.001	<0.001	
IK204	Kule	Py dissemin. ore	-	-	<0.07	<0.3	<0.01	<0.01	0.01	<0.001	<0.001	
SM201	Nathlenga Creek	Sil-lim Gossan	N52° W, 75° W	0.5	<0.07	0.7	<0.01	<0.01	<0.01	0.001	<0.001	
ST204	Wavavakala Cr.	Argill. rock	-	-	<0.07	<0.3	<0.01	<0.01	0.01	<0.001	<0.001	
ST202	Tunayasi Creek	Quartz vein	-	-	<0.07	<0.3	<0.01	0.01	<0.01	<0.001	<0.001	
AT205	Korokitu	Py-sil. rock	-	-	<0.07	<0.3	<0.01	<0.01	<0.01	<0.001	<0.001	

(Phase II)  
Nba-west Outcrops)

Sample No.	Location	Description	Dip-strike	Width (cm)	Ore Grade							
					Au g/t	Ag g/t	Cu g/g	Pb g/g	Zn g/g	Te g/g	Mo g/g	
SM101	Nalotava	Py-sil vein	N42° W, 90° E	40	<0.005	<2	90	10	20	1.0		
SM102	Nalotava	Clay-Py vein	N 1° E, 60° W	20	0.005	<2	120	45	115	0.2		
SM103	Nalotava	Py-Clay alt.r.	-	100	<0.005	<2	105	10	115	0.1		
SM104	Nalotava	Py-Clay alt.r.	-	300	<0.005	<2	75	15	55	1.4		
SM105	Nalotava	Py-Clay alt.r.	-	300	<0.005	<2	95	15	30	0.4		
SM106	Nalotava	Py-Clay alt.r.	-	350	0.022	<2	80	10	100	0.3		
SM107	Nalotava	Clay-Py vein	N12° W, 90° E	35	<0.005	<2	35	40	35	1.3		
SM108	Nalotava	Clay-limo vein	N15° W, 85° W	45	<0.005	<2	80	45	115	<0.1		
SM109	Nalotava	Clay-Py vein	N22° W, 80° E	60	<0.005	<2	55	10	50	<0.1		
SM110	Nalotava	Py-Clay vein	N 7° E, 80° W	20	<0.005	<2	85	10	45	<0.1		
SM111	Nalotava	Clay-Py vein	N10° W, 15° E	50	0.003	<2	100	15	75	<0.1		
SM112	Nalotava	Clay-Py vein	N30° E, 65° W	40	0.016	<2	80	10	35	1.4		
SM113	Nalotava	Clay-Py vein	N18° W, 45° E	180	0.001	<2	95	20	45	1.2		
SM114	Nalotava	Clay-Py vein	N28° W, 10° W	100	0.006	<2	100	20	100	0.3		
SM115	Nalotava	Clay-Py vein	N50° W, 45° E	5	0.024	<2	115	15	280	0.1		
SM116	Nalotava	Clay-Py vein	N18° W, 45° W	30	<0.005	<2	55	15	35	<0.1		
SM117	Nalotava	Clay-Py vein	N11° E, 85° W	80	<0.005	<2	45	10	60	<0.1		
SM118	Nalotava	Clay-Py vein	N34° E, 15° W	30	<0.005	<2	80	10	80	0.5		
SM119	Nalotava	Clay-Py vein	N43° W, 80° E	100	<0.005	<2	45	10	125	0.4		
SM120	Nalotava	limo network	N25° W, 50° E	150	<0.005	<2	30	10	25	1.2		
Y-41	Yaloku	Qz vein float (ok. p.)	-	-	<0.005	<2	-	-	-	-		
Y-42	Yaloku	limo (Ql. banded p.)	-	-	<0.005	<2	-	-	-	-		
Y-5	Yaloku	Silicified rock float	-	-	<0.005	<2	-	-	-	-		
Y-6	Yaloku	Qz-limo-Py vein	N35° W, 55° E	8	0.022	<2	100	4	115	1.5		
Y-10	Yaloku	Qz-limo-Clay vein	N39° E, 30° W	15	0.014	<2	120	45	400	5.0		
Y-12	Yaloku	Qz-Py vein	N35° W, 50° E	4	0.010	<2	110	230	800	2.0		
Y-13	Yaloku	Qz vein	N15° W, 90° E	1	0.010	<2	110	150	120	3.1		
Y-14	Yaloku	Clay-Py vein	N33° W, 60° E	10	0.030	<2	120	23	43	3.4		
Y-15	Yaloku	Qz vein	N43° W, 75° E	3	0.010	<2	210	17	10	4.0		
Y-16	Yaloku	Qz vein	N18° E, 60° W	20	0.180	<2	150	40	200	10		
Y-17	Yaloku	Qz vein	N67° W, 60° S	3	0.194	<2	160	20	35	3.4		
Y-203	Yaloku	Qz vein	N63° E, 65° W	5	0.041	<2	150	12	25	4.5		
Y-4	Yaloku	Qz vein	N57° W, 70° S	1	0.008	<2	41	8	45	5.0		
Y-5	Yaloku	Qz vein	N27° W, 90° E	3	0.020	<2	250	5	88	4.3		
SM12	Yaloku	Qz-limo vein	N 7° W, 60° E	5	<0.07	0.4	200	600	<100	<10		
OK3	Yaloku	Qz-limo vein	N 7° W, 60° E	25	<0.07	1.0	300	200	<100	<10		
OK5	Yaloku	Qz-limo vein	N15° W, 90° E	5	<0.07	1.0	500	200	<100	<10		
OK6	Yaloku	Qz vein	N25° W, 85° E	1	<0.07	0.3	500	100	<100	<10		
OK8	Yaloku	Clay vein	N42° W, 50° E	5	<0.01	4.4	500	120	<100	<10		
OK9	Yaloku	Py-Clay vein	N53° W, 60° S	15	<0.07	0.3	100	150	<100	<10		
OK11	Yaloku	Qz vein	N65° W, 15° S	5	0.11	2.7	400	500	<100	<10		
OK12	Yaloku	Qz-limo vein	N62° W, 60° E	8	<0.07	<0.3	100	100	<100	<10		
OK14	Yaloku	Qz-Cal vein	N57° W, 50° W	2	<0.07	0.3	300	100	<100	<10		
OK15	Yaloku	Qz vein	N22° W, 15° W	2	<0.07	<0.3	160	<100	<100	<10		
OK17	Yaloku	Qz-Cal vein	N63° W, 10° S	2	<0.07	<0.3	<100	<100	<100	<10		
OK18	Yaloku	Qz vein	N 5° E, 81° E	1	<0.07	<0.3	200	100	<100	<10		
OK19	Yaloku	Silicified rock	-	-	<0.07	<0.3	100	200	<100	<10		





JICA