Republic of Indonesia

Master Plan for Land Erosion and Volcanic Debris Control in the Area of Mt. Merapi

SUPPORTING I

- A. GEOMORPHOLOGY
- B. GEOLOGY AND LAHAR DEPOSITS
- C. HYDRAULICS AND HYDROLOGY
- D. RIVER SURVEY
- E. RECORD OF DISASTER AND HAZARD AREAS
- F. SOCIO-ECONOMY

March 1980

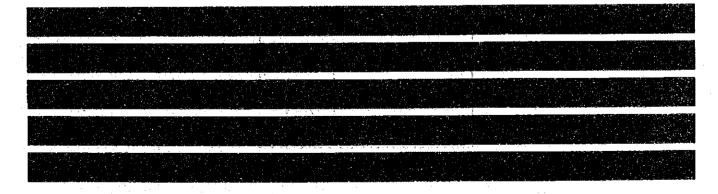
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JAPAN INTERNATIONAL COOPERATION AGENCY

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SUPPORTING REPORT A GEOMORPHOLOGY

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1. INTRODUCTION

1.1 Topography

- G. Merapi is located on 30km north from Yogyakarta in central Java and is 2,968m high above the sea level. The studied area spread approximately $850 \rm km^2$ on the south and west slopes of G. Merapi.
- G. Merapi, one of the most active volcano in Indonesia, is a volcanologically very young Quarternary strato volcano with a beautiful conical shape. It erupted many times on a large scale even in recent times. The south-western part of mountain body, or the so-called older Merapi, was almost totally destroyed in 1,006 A.D. The present active younger Merapi was formed on the old Merapi. The younger Merapi is surrounded by the older Merapi except the south-western slope.

Nine main tributaries flowing down radially dissect the mountain slope. Five of them are tributaries of K. Progo: K. Pabelan, K. Blongkeng, K. Putih, K. Batang and K. Krasak. Three of them are tributaries of K. Opak: K. Boyong, K. Kuning and K. Gendol. K. Opak and K. Progo pour into the Indonesian Sea 30km south from Yogyakarta. K. Woro is a branch of K. Dengkeng, which is one of the tributaries of Bengawan Solo.

- G. Merapi can be divided to three areas vertically as shown in Figs. 1 and 2 in accordance with its geomorphological characteristics: namely, the upper slope area above 2,000m the middle slope area between 2,000 500m and the foothill area below 500m. All slopes are described in detail below.
 - (1) The upper slope area is situated on the older Merapi body and is always affected directly by the present volcanic activity. This area is composed of pyroclastic materials, lava and ash, and has an incline of more than 30 degrees. Owing to the regular supply of the new volcanic products, this area has become almost bare. Sometimes thin gullies develop on the slope composed of pyroclastic materials.

As a great drop occurs between the crater in the south-west side of this area, a large part of the eruptive products flow down along this drop and cause damage in the south-western foothill area.

(2) The middle slope area has somewhat complex topography because it is composed of the remaining of the older Merapi and of the slope composing of the eruptive products of the younger Merapi.

The older Merapi forms a horseshoe shape which is higher in northeast side and opened and lower on the south-east side. The older Merapi is very steep due to severe erosion. Especially in the northeastern side, valleys are deep with steep walls because the eruptive products of the younger Merapi were not distributed around this area.

In the south and north area, the products of the younger Merapi flow down along the vaileys dissecting the mountain body of the older Merapi so that the area can be divided into visibly separate topographic units: upper, middle and lower slopes. Products of the younger Merapi, flowing down along the older valleys, expanded and were deposited in a fan shape along the outlets of the valleys.

In the upper slope area, a valley formation is not pronounced; however for some small gullies in the middle slope area, below 2,000m where, water has easily concentrated to cause the rill erosion to develop into radial deep valleys. In the middle slope area, mainly consisting of deposits of nuce ardente, deep valleys are formed at the present time. During a large scale eruption, nuce ardente flows down into this area, fills up valleys and even changes the direction of the valleys. When the eruption becomes weak, valley formation immediately starts because the gradient deposition of nuce ardente is steeper than the gradient of streams below it.

The slope surface of this area is relatively more undulated than that of the footslope area. This may mean that the deposited materials in this area are more viscous than the materials in the lower part. Gentle and step slopes are numerous in this area and their gradient becomes more gentle in the lower area. These steep parts can be visualized as many small circles steps surrounding G. Merapi.

(3) The foothill area stretches from the place where the deep valleys of the middle slope disappear. This area stretches from a location of 500 - 600m in height with gradient of 3 degrees for K. Progo. The slope of the south side of Yogyakarta is approximately 1 degree. This area is poorly undulating and forms a very gently inclined smooth surface. The radially developed valleys from G. Merapi have become shallow by the time they reach this area. However, they form deep valleys before they flow into K. Progo and K. Opak because the channels of K. Progo and K. Opak are deep in this flat area.

K. Woro, flowing into K. Dengkeng, forms a raised riverbed in the lower reach area where the riverbed is $5 \sim 6m$ higher than the general surface. K. Dengkeng generally forms a marsh area on the right bank side.

However, the foothill area is also considered to be composed of volcanic materials, and shows only small undulation due to its more fluid origin.

Along K. Boyong, K. Kuning and K. Woro, alluvial fans are formed in this area. These fans are composed of old volcanic deposits and fan gravels.

1.2 Objectives and Scope of This Study

The purpose of this study is to determine the geomorphological land conditions in order to forecast the endangered area to Lahar and Banjir flooding and to design the masterplan for sediment and erosion control.

The main work of the geomorphological survey is to prepare a land condition map in order to locate the endangered points and areas.

The scope of work of geomorphological study was limited to the Type-I area since the greatest amount of recent damage from volcanic eruption has taken place there and since Type-II and Type-III areas have similar features. (for definition of terminology see Section 2.1)

1.3 Validity of Methodology

In the study area, the relationship between landforms and Lahar/Banjir occurances was studied. The probable endangered areas were analysed by using the landform classification map. For a more accurate and quantative estimation, other maps and data (isoclinal slope distribution, relative height of the banks/cliffs of valleys, stream system map, etc.) were also prepared and analysed.

In order to determine the accuracy of the estimation, two methods were employed: One was to study the data of Lahar/Banjir occurances in 1975 and 1976. The other was to simulate a model of the endangered area in the K. Krasak and K. Putih.

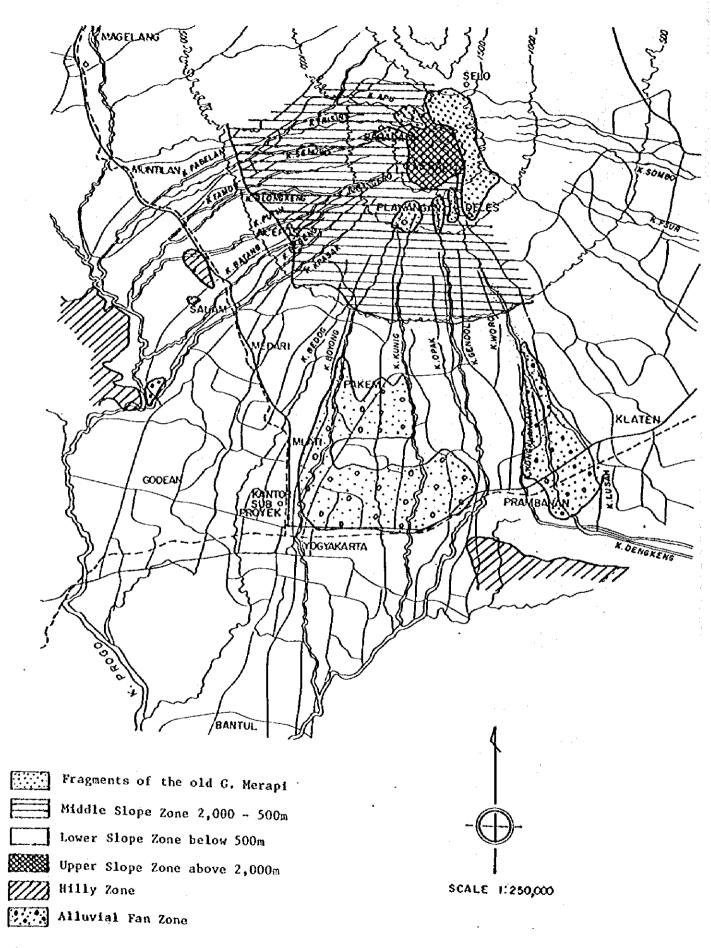
According to the results of this investigation, the accuracy and the forecast of endangered area are valid for use a basis for the master-plan for the following reasons

- 1) The area covered by Lahar/Banjir in 1975 and 1976 correspond to the probable endangered area estimated by geomorphological method.
- 2) The cross-sections estimated as endangered points by the simulation study are located in the geomorphologically probable endangered area.
- 3) The area estimated by the geomorphological survey and the simulation are very similar with each other.

Although the investigation of the probable endangered area was undertaken only in Type I areas along the tributaries of K. Krasak and K. Putih, the characteristics of flood, geomorphology and volcanic materials are more or less the same in Type II and Type III areas.

Consequently, the masterplan can be developed for the whole project area based on the assumptions verified by this study.

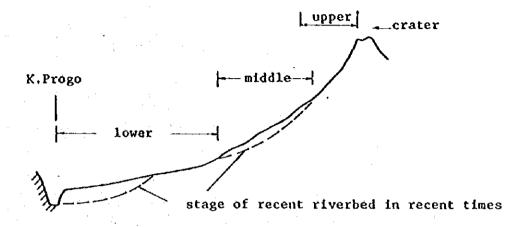
Fig. 1 Map of Main Morphological Areas



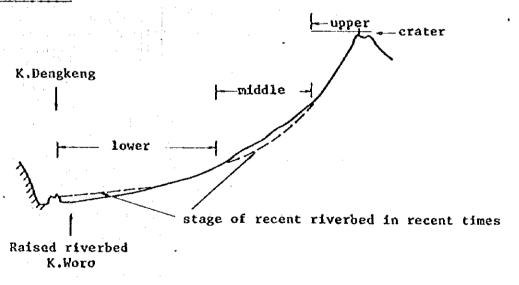
(Note: Main cities along the national road are shown in capital letters)

Fig. 2 Idealized Profile of the River and Slope

K. Krasak Profile



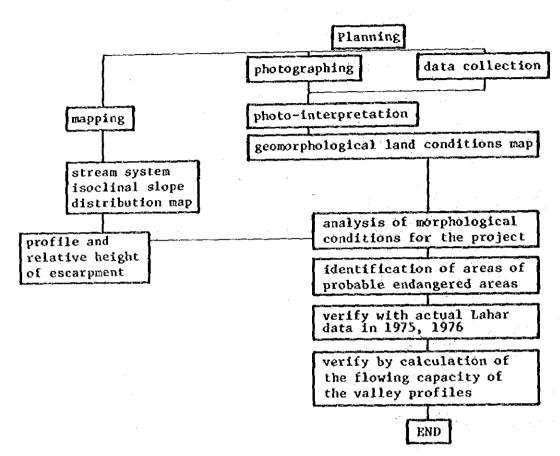
K. Woro Profile



1.4 Geomorphological Study Work Flow

An outline of the flow of work is shown in Fig. 3.

Fig. 3 Work Flow



2. GEOMORPHOLOGICAL STUDY

2.1 Land Condition Map

The land condition map was prepared using a scale of 1/25,000. Photo-interpretation techniques and field survey were employed for this work. Geomorphological features and their explanations used in this map are as follows (See Appendix-1; photo Interpretation Cards):

2.1.1 Items of geomorphological classification

(1) Volcanic morphology

- Recent lava flow deposits (Young valcanic deposits of G. Merapi)
- Former lava flow deposits (Young valcanic deposits of G. Merapi)
- 3) Old valcanic deposits of G. Merapi.
- 4) Volcanic debris: around the crater
- 5) Deposits of nuée ardante d'avalanche (1954, 1969)
- 6) Deposits of nuée ardante d'avalanche (Uncertain dates)
- 7) Mountain side composed of valcanic debris (Upper Zone)
- 8) Mountain side composed of valcanic debris (Lower Zone)
- 9) Breaking zone of slope caused by valcanic debris flow
- 10) Flowing units of valcanic debris
- 11) Former valley buried by nuée ardante d'avalanche
- 12) Poorly defined shallow valley

(2) River morphology

- 1) Riverbed
- 2) Point-bar deposits
- 3) Flood terrace
- 4) Accumulation terrace
- 5) Erosion terrace
- 6) Escarpment
- 7) Deposits of recent Lahar
- 8) Sand bars, Natural levees
- 9) Alluvial fan
- 10) Alluvial fan (dissected)
- 11) Raised riverbed
- 12) Local sediment by flood
- 13) Altered valley caused by nuée ardante

- 14) Former riverbed
- 15) Small valley (∐ shape)
- 16) Small valley (L______shape)
- 17) Small valley (V shape)
- 18) Alluvial plain
- 19) Swamp
- 20) Limit of sediment transportation
- 21) Knick point

(3) Others

- 1) Slopes of hill
- 2) Slump block
- Landslide
- 4) Talus
- 5) Piedmont plain
- 6) Gully, Rill
- 7) Spring

2.1.2 Explanation of Geomorphological items above

(1) Volcanic morphology

Recent lava flow deposits are divided into two parts. One is distributed on the upper part of G. Merapi. The other is developing on the middle and lower slope.

Old valcanic deposits of G. Merapi consist of the older G. Merapi which is surrounding the young Merapi crater.

Deposits of nuée ardente from the eruptions in 1954 and 1969 were deposited mainly around K. Blongkeng, K. Putih, K. Batang, K. Bebeng and K. Krasak. These materials filled up old river valleys and changed the river course in some cases for example, between K. Putih and K. Blongkeng.

Mountain sides are composed of valcanic debris rather than deposits from rivers. Distribution is on the upper, middle and lower slope. (Please refer Figs. 1 and 2). The breaking zone of a slope is caused by volcanic debris flow. When the flow stops on a slope, it makes a gap on the slope. Because the volcanic material have a high visconsity. This gap or breaking zone is quite significant for detecting the endangered points to nuée ardente and Lahar.

Former valleys can be filled up and buried by nuée ardante.

(2) River morphology

Flood terrace. Very low terraces which are submerged by big floods.

Sand bars, natural levees. A little higher place, composed of sand, which is distributing in the fan and riverbed.

Alluvial fan. Mainly sandy deposition. K. Krasak, K. Kuning and K. Boyong have big fans and fan-like areas. Only K. Krasak's fan is active nowadays. The other two rivers have dissected fans in the lower foothills.

Raised riverbed. A lot of sand deposited between the embankment has caused the riverbed to rise: This condition makes flooding over the levee very easy. Sandy sedimentation which was transported during the inundation can be seen along the embankment, and is identified as local sediment by flooding.

Altered valley caused by nuée ardente. After a valley has been filled by nuée ardente, the old valley capacity to hold water becomes very small. This area is located just downstream from dammed point on the old river.

Former riverbed. This area can become quite dangerous during nuée ardente flow, lahar and banjir.

Small valleys. They do not have big catchment areas, and ground water is usually the main water source. However, small valleys have possibility to become flowing courses for nuée ardente and so on. They are identified according to the shape of valley.

2.2 Relative Depth of Valley

The overflow point of Lahar/Banjir is related to the height difference of the mountain slope and valley bottom. The relative height of the side cliff was measured and plotted to obtain a longitudinal profile of each river.

2.3 Isoclinal Slope Distribution Map

By measurement of the contour interval, the gradient of the slope was estimated. Slopes in a uniform gradient zone where grouped to make the Isoclinal Slope Map. The zone where gradient clearly changes has the highest probability of occurance of Lahar/Banjir.

2.4 Stream System

The purpose of this map is to estimate the flow direction of Lahar and Banjir. When Lahar and Banjir overflow the riverbank, they spread over the foot of the mountain or flow down along the valleys which then fill up with volcanic materials.

3. RESEARCH OF ENDANGERED AREAS

3.1 Endangered Area Map Preparation

3.1.1 The purposes of this Survey

The purposes of this survey are to study the relationships between geomorphological conditions and Lahar and Banjir in order to estimate the endangered area to Lahar and Banjir.

There is no clear dividing line between Lahar and Banjir; however it is generally understood that Lahar includes more sand and gravel than Banjir.

The endangered area is mainly on the slope of the active volcano, G. Merapi. K. Krasak (K. Bebeng) and K. Putih were the most endangered tributaries in recent years. People in villages along these rivers experienced the heavy damage from G. Merapi's explosion in 1969 which plugged the upper reach of K. Blongkeng with volcanic deposits and changed its channel to K. Putih. A similar alternation of river course happened to K. Batang with the plugging of K. Bebeng. When Lahar and Banjir are regarded as a geomorphic stage, geomorphological analysis can be used to determine the probable endangered areas to Lahar or Banjir in the near future.

The geomorphological land condition map in this report was prepared with using aerial photo-interpretation (scales: map 1/25,000 and photos 1/40,000). According to the geomorphological study, there has been quite a close relationship between landforms and Lahar/ Banjir occurences. Figs. 4 and 5 show that the rivers generally overflow at the places where their channels turn sharply or where riverbed gradients change and become more gentle. These topographical changes correspond to the geological and lithological differences in the slope of the volcano. The landforms of the volcano's slope are the result of volcanic activity: lava flow, nuée ardente d'avalanche, pumice-flow, Lahar, Banjir, etc. Usually Lahar/Banjir flow along former river cources, small valleys, shallow valleys and so on. Today's landforms suggest the area which will probably be covered by the next Lahar/Bajir occurances. The isoclinal slope map, drainage system map and flow direction map are prepared as supplemental information.

3.1.2 The relationship between landform and Lahar/Banjir

According to the survey of the Lahar and Banjir in 1975 and 1976, the relationships are found between them as follows:

Points of overflow (breaking point)

- 1) The point where the stream gradient becomes more gentle.
- 2) Crossing points with former river courses.
- 3) A riverbed zone which is raised up.
- 4) River terraces.

- 5) Zones where the capacity of the valley is decreasing as compared with its upper reach.
- 6) At a sharp river bend course of the river.

The flooding areas

- 1) The boundary of a flood area is ill-defined except by small landforms such as small undulations, stream, roads, etc.
- 2) The flood spread into small rivers or former river courses.
- 3) If the small river or the former river course does not have enough capacity to pass all the volcanic material, the flood will overflow again.
- 4) A lot of the sand transported by the flood remains: along small channels, in former river courses or on roads over which the main flood current has passed.
- 5) The damage to houses and loss of life are mainly distributed in the sand deposit areas.

Location of previous Lahar damaged areas and the breaking zones are shown in Fig. 5.

3.1.3 Estimation of the endangered area

Using the relationships between the landform and Lahar/Banjir, the endangered area is estimated.

The degree of probable flooding is divided into five categories based on the height of the bank.

Degree of probable flooding

Height of bank	Degree of probable flooding
0 - 5m	very high
5 - 10m	hígh
10 - 20m	middle
20 - 30m	1ow
30 - m	very low

The area was determined by using the geomorphological landcondition map and other information. In the cases where areas overlapped, the higher degree of probable flooding was adopted. The most endangered areas such as the point of overflowing, and former river courses and terraces are shown in the map.

Flood types are also categorized depending on the volume of sand contained as follows:

Types of floods

- 1) Flood contains a lot of sand and gravel
- 2) Flood contains less sand
- 3) Flood contains little sand and the inundation period is rather long

3.2 Damage by Lahar and Banjir in Recent Years

By means of a questionnaire the areas damaged and flowded in 1975 and 1976 were surveyed. The survey was conducted along the K. Putih and K. Bebeng. The items questioned were as follows:

Table 1 Question Card

(1)	Did you experience Lahar or Banjir? yes, no
(2)	When did it happon?
(3)	At what time did it attack here?
(4)	The time of the highest flooding?
(5)	How deep was the flood?m
	Sand deposited?m
(6)	What was the main direction of Lahar flow? (Show on the map)
(7)	Where was the boundary of the sand deposition? (Show on the map)
(8)	Where was the end or boundary of the water flow? (Show on the map)
(9)	How many houses were damaged?
	How many people were injured or killed?

Data of Lahar/Banjir in 1975, 1976

The results of the questionnaire and field survey are summarized in Table 2 and shown in the map, Fig. 7. In the map, areas of sand deposit have a thickness of more than 30cm.

Table 2 Questionaire Survey Results

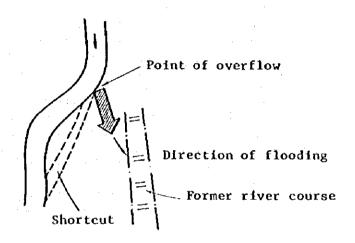
item No.*	River *	Year of Flooding	Places Flooded	Area Flooded (m ²)	Area of Sand Deposition	Approx.	Buildings Damaged	Deaths
					(m ²)	(m ²)		
1	K. Putih	1976	Salamsari Ngablak	132,000	34,000	26,000	2	
2		1975	Gemblongan	135,000	90,000	68,000	2	
3	K. Bebeng	1976	Brojonalan Pakel	596,000	227,000	86,000	5	-
4		1975	Gembokan	184,000	46,000	28,000	40	
5	K. Krasak	1976	Tempel	3,000	3,000	2,000	2	1
6		1976	Kricakan*	2,501,000	1,500,000	1,180,000	67	26

^{*}Also including Sumokaton, Madesan, Mriyan, Guling, Pokok, Kerisan, Banjarejo

^{**}Further details explaining the conditions at each site of flooding are given in subsequent paragraphs by item no.

- Item 1: This area is located on the junction of present and former river courses. At first the food went into the old river course and then overflowed its banks. The flood water rejoined K. Gremeng and K. Putih again at the end.

 On the left bank at Ngablak (near the bridge) was attacked by sandy flood waters and a house and a school were damaged. Sand deposited only along K. Putih and the smaller river, K. Gremeng. Other places were covered only by water.
- Item 2: This point is at a sharp bend in the river course. A shortcut has now been constructed to avoid the same overflowing again. Near the present river the flood spread a lot of sand over the area and went into the former river course.



- Item 3: K. Bebeng overflowed near its confluence with K. Krasak. This point is located at the upper side of the sand pocket which was not completed at that time.
 The level of the riverbed had been raised by sand pocket construction. The relative height between the volcanic slope and the riverbed is presently less than 2 meters.
- Item 4: This point is just at the bend in the river course. The flood split into two courses: one along the small valley and the other along the road. The flood along the road went back to the K. Krasak 600m lower downstream. The 40 houses damaged at that time were along this road.
- Item 5: At a bridge crossing the national road. K. Krasak narrows and changed the direction of flow suddenly to the right at an angle wider than 90 degree. The flood attacked the left bank. A man in a car was killed by the flood. The flooded area was very small.
- Item 6: The overflow occured at four points in this area.
 - (1) The biggest flood attacked the right bank at the Kricakan village. Most of the flood waters flowed into a small channel passing through villages such as Sumokaton,

Madesan, Mriyan and Guling. Along this small channel most of the damage resulted from the flow of sandy water,

The heaviest damage was seen in this area as shown below.

	the state of the s	the state of the s
village name	houses damaged	deaths
Madesan	48	9
Mriyan	8	. 1
Guling	4	16
Total	60	26

- (2) On the right bank 1.3km below the first point. The flood spread from the channel to the fields, and then poured into K. Petel. The sand deposition in this area was small.
- (3) On the right bank 3.5km below the first point. The flood attacked the channel of K. Petel near Banaran. K. Petel has the very big channel as compared with its present discharge. It might be old river course of K. Krasak. As main current of the flood passed through the K. Petel channel, the channel was filled by a lot of sand. Along the channel, 7 houses were destroyed by the secondary overflow from K. Petel channel.
- (4) On the left bank just opposite the third point. The flood poured into the small valley and was contained; it did not overflow again.

3.3 Simulation Study of the Endangered Area

Using the changes of the capacity for each section along the river course, the endangered area can be computed. Data such as a morphometeorological, hydrological and experiences of flooding were collected in order to compare the capacity of each section along the river valley. The following is a summary of the factors considered and computational formulas used in the computer simulated model.

- (1) Cross-Section
 - The cross-section was measured using photogrametry techniques and aerial photographs taken in 1976.
 - 2) The cross-section was taken each 100 meters as a rule.
 - 3) Plotting scale: 1/1,000 horizontal and 1/400 vertical.
- (2) The River Valleys Studied
 - 1) K. Putih

measuring line number 202300 -- 500800

. 2) K. Krasak and K. Bebeng

measuring line number 402150 -- 500800

(3) Formulas for Valley Flood Capacity Calculation

The capacity of each valley section to pass the flood discharge was calculated one by one. Average velocity without sand was calculated using Manning's equation, and velocity with sand was calculated with Wang's formula respectively as follows:

	Capacity for water only Manning's equation	Capacity for water and sand Wang's equation
Flood velocity :	$V = 1/n \cdot R^2/3 \cdot I^{1/2}$	$V_1 = \frac{\gamma W}{\gamma W + \alpha (d - \gamma W)}$
Quantity of runoff:	Q = Av	$Q_1 = Av_1$
where	V: average velocity of V1: velocity of flow with a : Manning's roughness R: hydraulic mean radiu I: slope of riverbed Q: quantity of runoff of Q1: quantity of runoff IA: cross-sectional area γW: specific weight of γ d: specific weight of γ d: ratio of sediment to	h sand coefficient (see below) is (R = A/P) of only water including sand tater ater

: wetted perimeter

Roughness coefficient (n)

The manning's roughness coefficient should be based on data from each cross-section. Nowever, since data about roughness could not be obtained from the gauged run-off, the Manning Strickler equation was used to estimate it as an expendient as follows:

$$n = 0.0417 \, dm^{1/6}$$

n : roughness coefficient where

dm: mean diameter of grain

(4) Parameters for the above equations

- Area of cross-section, hydraulic mean radius and slope of the riverbed were obtained from the photogrametrical measurement data.

- Roughness "n"

The relationship of the mean diameter of the riverbed material to the slope of the riverbed was estimated using the data shown in Table 3 and the formula below.

The angle of slope was decided for each section from its longitudinal profile,

Relationship of the slope and mean diameter (D)

$$D = 95.52 e^{I} - 92.77$$

where

D : mean diameter

I : slope of riverbed

: base of natural logarithms with a value

of 2.71828

Correlation coefficient of D and I is 0.96. (r=0.96)

Table 3 The Mean Diameter of the Riverbed Material and the Slope of the Riverbed

Location No.	Measurement line No.	Slope (I)	Riverbed mean diameter (D) in mm
1 2 3 4	401090 401150 401270 401370	0.045 0.042 0.037	7.43 6.92
6	401740 402610	0.033 0.018	5.31 4.78

(5) Calculation of the Capacity

K. Putih and K. Krasak river courses were each divided into four longitudinal zones according to slope gradient. The slope, diameter and roughness coefficient are given for each section as shown in Table 4 below.

Table 4 Topographical Parameters per Zone

River	Zone	Slope (I)	Riverbed diameter (D)um	Roughness (n)
K. Putih	No. 202300 201830	0.024	5.07	0.055
	No. 201820 201500	0.033	5.95	0.056
	No. 201490 201160	0.042	6.85	0.057
	No. 201150 500800	0.057	8.35	0.059
K. Krasak	No. 402150 401920	0.020	4.68	0.054
	No. 401910 401310	0.038	6.45	0.057
	No. 401300 500980	0.047	7.35	0.058
	No. 500970 500800	0.063	8.96	0.060

Wang's equation is used to obtain the velocity when the flood contain a lot of sand. Each parameter was defined as follows:

(K. Putih)

 $yw = 1.0 \text{ t/m}^3$

 $d = 2.8 t/m^3$

 $\alpha = 0.41$ (Data from Salam Jan. 9, '78, on K. Krasak used here)

(K. Krasak)

 $yw = 1.0 \text{ t/m}^3$

 $d = 2.8 t/m^3$

 $\alpha = 0.11$ (Zones 402150 - 401920 gaged on Jan. 9, '78: Blanburan)

 $\alpha = 0.41$ (Zones 401910 - 500800 gaged on Jan. 9, '78: Salam)

(6) The Estimation of Probable Endangered Area

The capacity of each river section to pass the flood volume through its banks was studied by computer to determine the following characteristics of probable following:

- (1) The runnoff volume
- (2) Location of the river section with insufficient capacity to pass the total flood volume received from the section just prior to it.

The method used for calculation is based on three major assumptions as follows:

- (1) That the starting point of the runoff would be zone 500800 for both K. Putih and K. Krasak.
- (2) That four volumes for the runoff would be considered:

Case a:	runoff equal	to 1/4	the height of	the rive	erbank
Case b:		1/2	11	18	
Case c:	11	3/4		1f	
Case d:	11	the	full hèight	и	

A summary of the data resulting from the first two assumptions is shown table below.

Name of river	Flooding start point of assumed runoff			Volume of (b) 1/2 height of bank	(c) 3/4	noff m3/sec.) (d) full height of bank
K. Putih		12	393	1,598	4,750	8,577
K. Krasak		20	634	4,968	12,535	27,616

(3) That the maximum volume of a river section would be passed on to the next section regardless of whether overflow occurred or not. If overflow did occur, the volume of the overflow (i.e., the difference between the received runoff and the capacity of the section) would be assumed to be permanently deducted from the river system (i.e., from all the subsequent river sections).

For example, assume that 393 m3/sec is the starting runoff, and that the lower 20 sections of K. Putih have enough
capacity to pass through this volume, but that the 21st
section (No. 201060) has only 365 m3/sec as capacity. This
means an overflow of 28 m3/sec will probablly occur in the
21st section. On the lower reach, the new runoff amount of
365 m3/sec will be discharged. If a subsequent section has
less capacity than 365 m3/sec, it will be labelled as an
overflow point in a similar fashion, and so on.

The results of a computer simulation based on the above assumptions are shown in Tables 5 and 6 for each section of K. Putih and K. Krasak respectively. A summary of the probable endangered flooding areas are shown on the maps in Figs. 9 and 10 for the K. Putih and K. Krasak respectively.

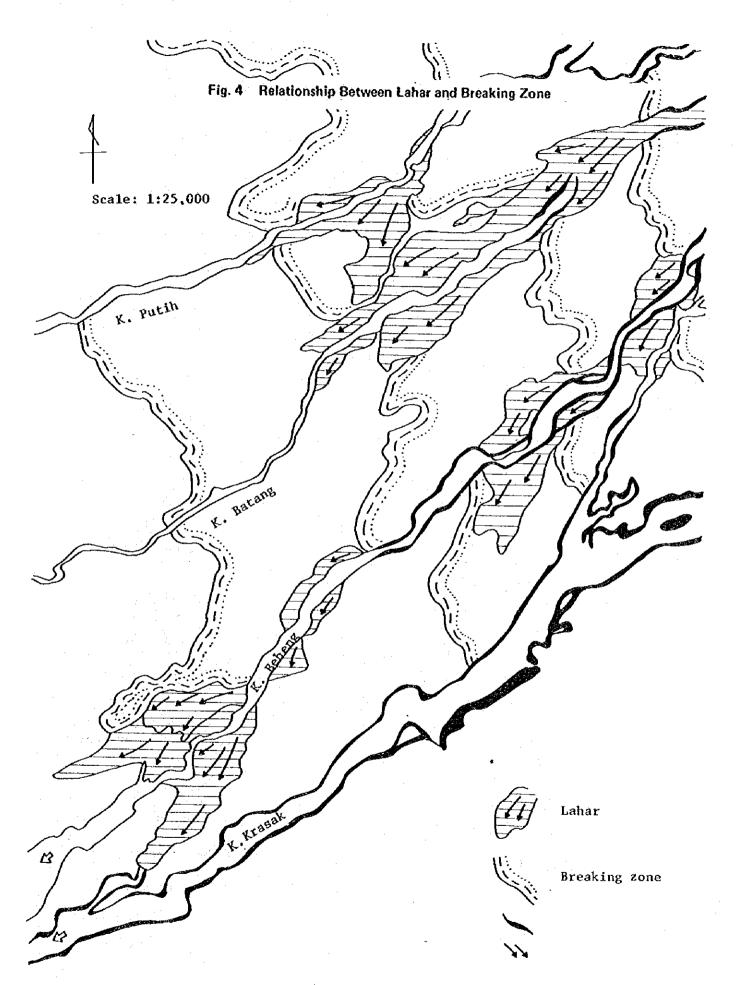
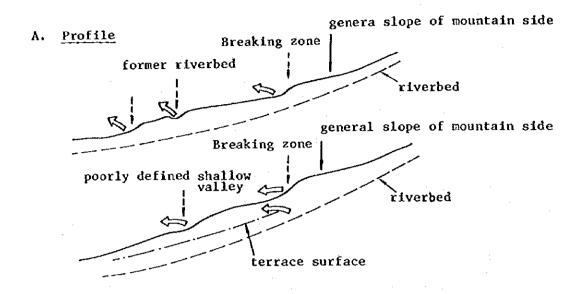


Fig. 5 Idealized Pattern of Lahar Occurrence



B. Morphological aspect

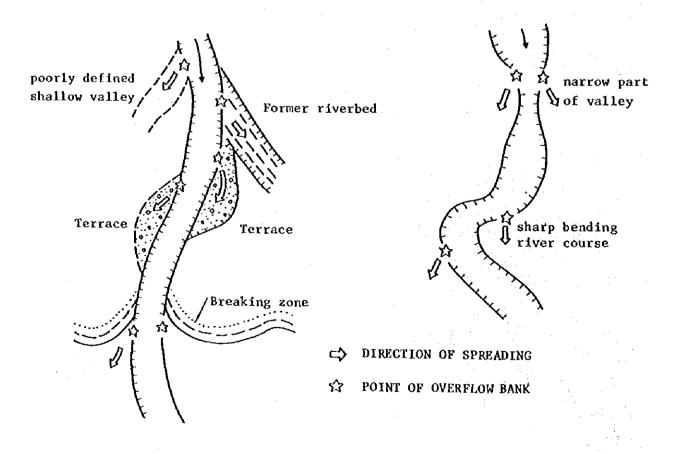
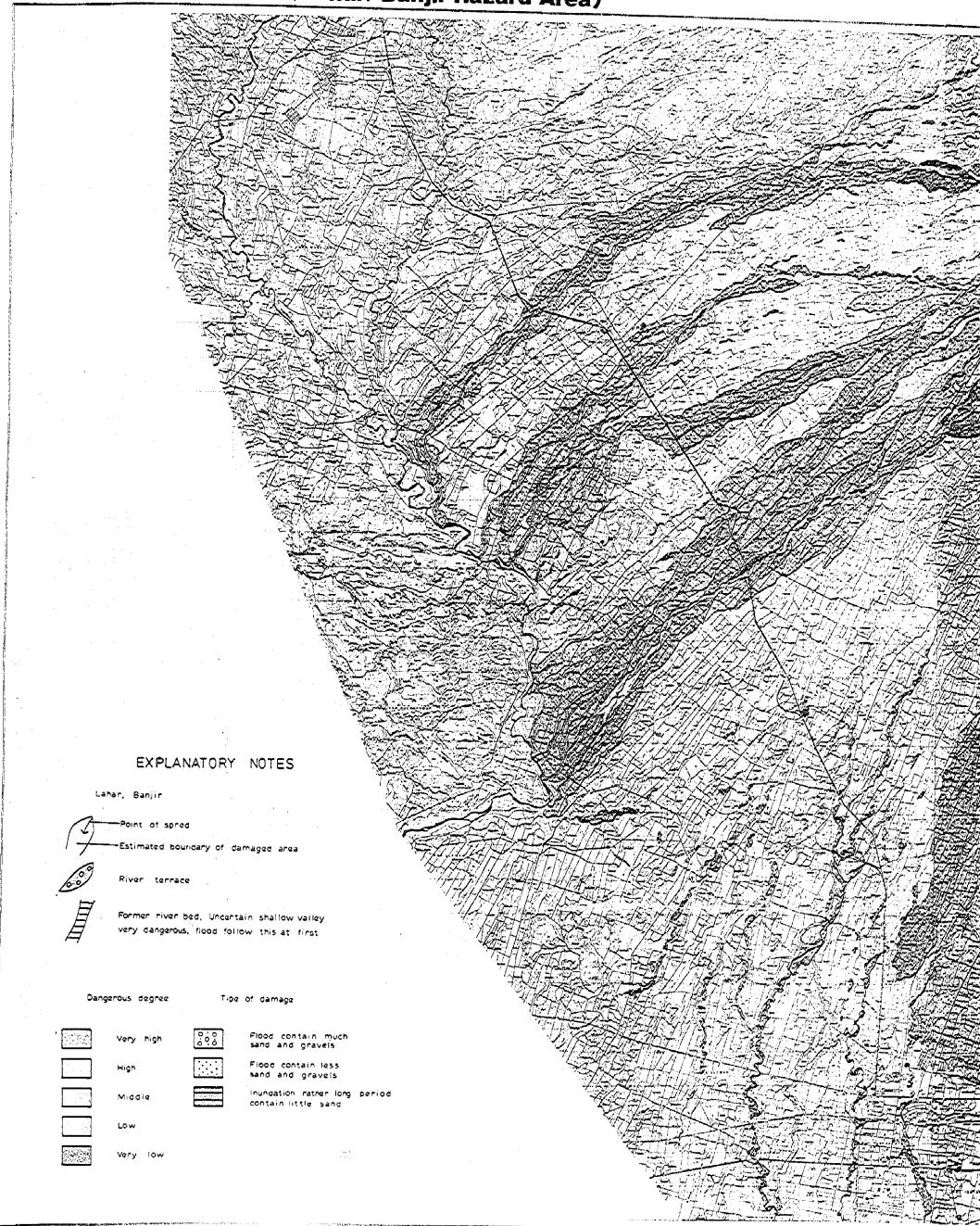


Fig 6 PROBABLE ENDANGERED AREA MAP (Lahar/Banjir Hazard Area)



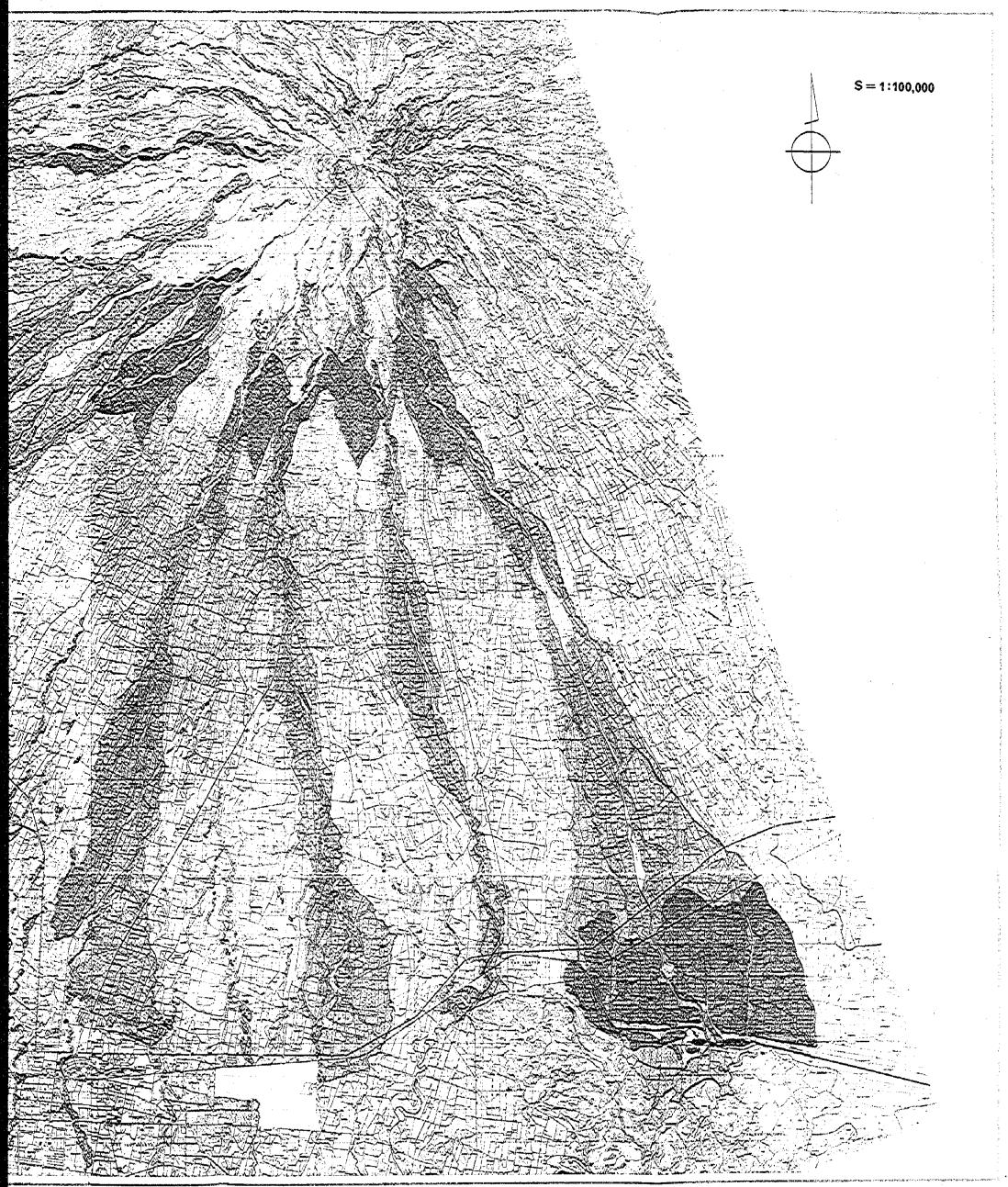
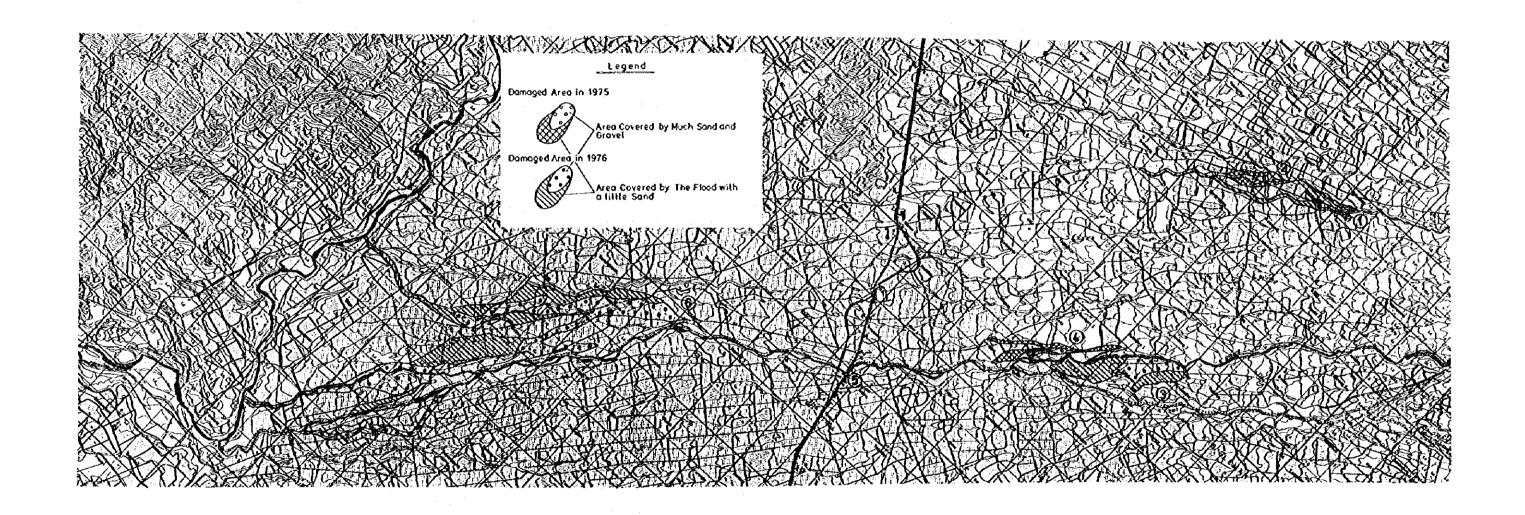


Fig. 7 Map of The Area Damaged by The Floods (RAHAR/BANJIR) in 1975, 1976



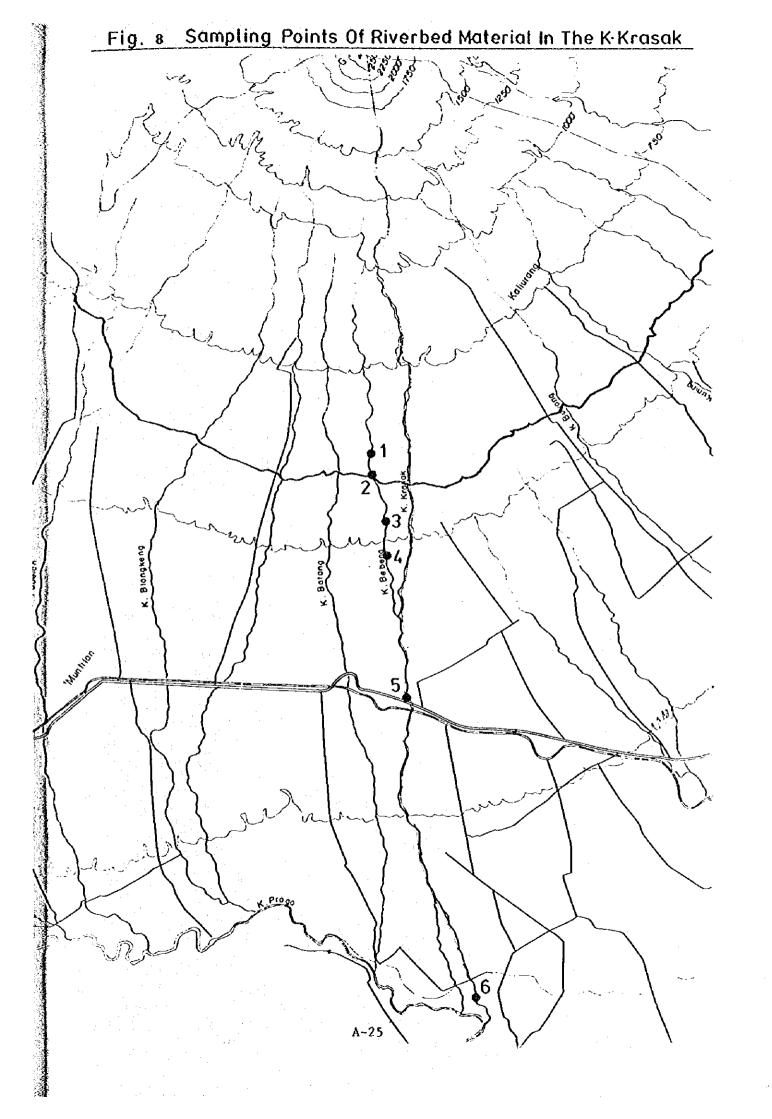


Fig. 9 Probable Endangered Areas (K. KRASAK)

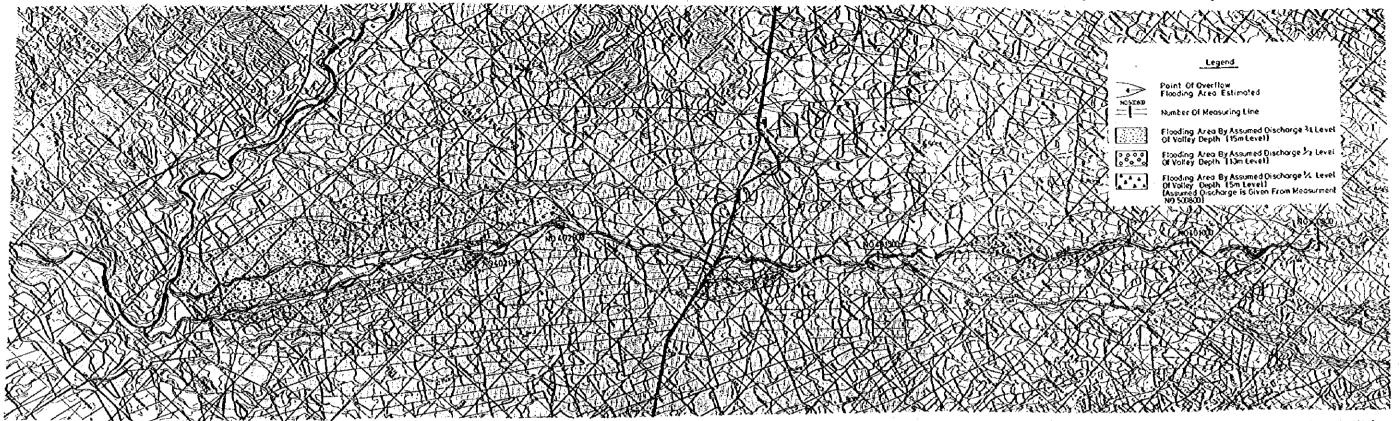
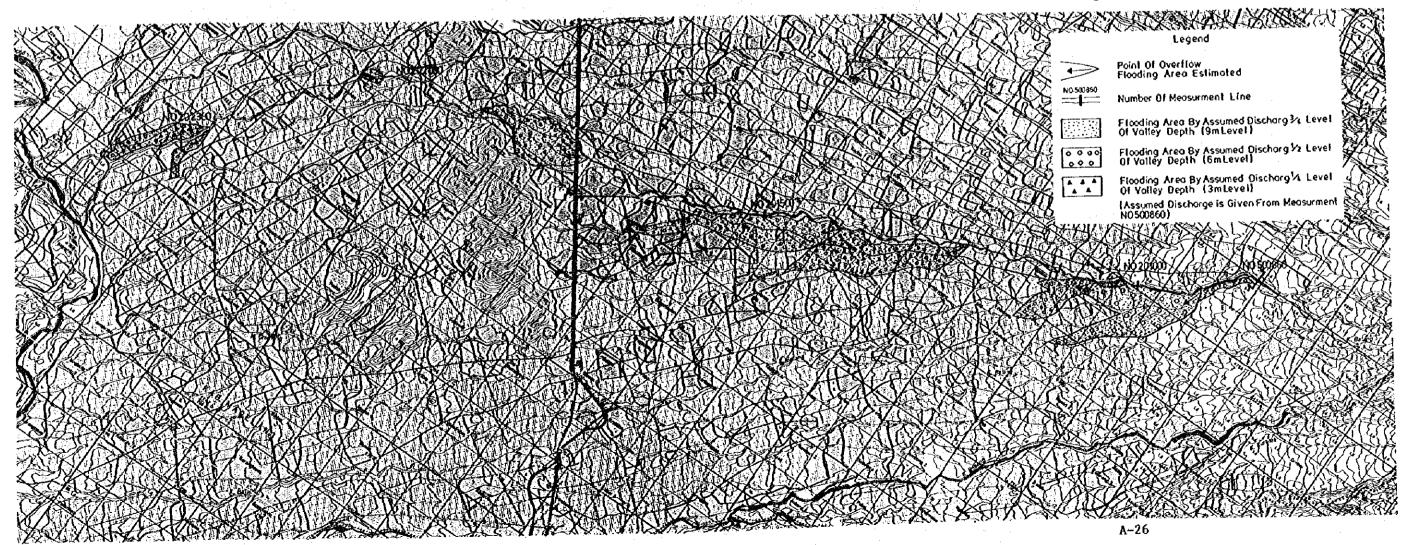


Fig. 10 Probable Endangered Areas (K. PUTIH)



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		Runoff	973.0	641.6 424.6 6.426.0	186.3	1958.8	612.7	533.7	766.1	687.7	1302.7	1001.3	201.3 201.3	923.9	901.5	192.6	763.5	306.2	1306.7	1084.1	1365.0	1033.2	836.8 1266.3	894.4 2938.4	981.5 3786.5
		Velocity with sedi-	3.8	anc Jul	113	4.4	in-	in.	y cy cy 4 m 4	8,4 6,4	54.4 54.6	န လူ မ ၁ စာ ဝ	n o e	i co c	e Gen	355	4 0. L'A	4.4.0 4.4.0	i en ei o or ei	988 174		4 4 4 4 4 <i>c</i>	4 K K C	w4.	รพ.ศ. วังวัง
		Runoff	1691.0	738.5	323.8	3404.4	1064.9	927.5 1145.0	1331.5	1195.2	2264.2	1740.3	1893.6	1605.8	1566.8	334.8	4346.5 1327.0	532.2	2271.1	1159.2	1740.6	1795.8	1454.4	1554.4 5107.0	1705.9
, , , ,	Putih)	Velocity	9.9	6 29 4 4 70 00	(), (),4	000	, 9°	24.4	s s s s	, v, v	14.0	, t, o	o vi v	145. V	n n n	ယူ လ ဆို ဆို	6.1 6.12	044 Juin	16.0 0.00 0.00 0.00 0.00 0.00 0.00 0.00		4.0 4.0	- L- L		₩. ₩.4.c	108 108
į	Flood (K. 1	Hydraulic Radius (R)	3,02353	2.71042	2.38983	4.07692	2.07059	3.92308	3,13924	3,30769	4.13513	3,43421	2.83871	2.69792	2.71845	1,31818	5.05000 3.51613	2.08197	2.84034	3.49412	3.64000	4.29630 4.18841	4,21739	2.62963 4.11905 2.44526	1.70476
	for The Fl	Wetted Perimeter		\$ 4 %	. 65 . 65	<u>5</u> 8		8,5	6.4	\$5.	4.	19.59	28.5	388	103	140.	180		119	885 5885	75 89 89	. 8 . 4 . 4 . 6 . 6 . 6 . 6 . 6 . 6 . 6 . 6 . 6 . 6	85. 23.	108. 168.	105. 162.
	Valley fo	Slope of River Course (1)	ő	\$ #	voo	30.5		2 2 3 3 3 3 3 3	19.	191	325	366	;;; <u>;</u>	32.	325	<u>ಕ್ಷಕ್ಷ</u>	15.	90,4 90,4	**	223	444	4% 4% 4	25	ង្កង់ដ	12,21
	The	River Course Area (A)	257.	133	141.	424	176.	153.	25.5 8,5 8,5 8,5 8,5	215.	300	761. 261.	2669. 2044.	288 288 208 208 208 208 208 208 208 208	580 530 730 730 730 730 730 730 730 730 730 7	87. 457.	218. 218.	127.	333	297. 193.	273. 387.	232 802,	342	284 922.	284. 776.
	Capacity of	Width of Riverbed	67.2	36.6 36.6 6	53.4 69.2	90.6	6.00 8.00 8.00	30.69 4.0	8.40	63.0		0.85 0.85 0.45 0.45 0.45 0.45 0.45 0.45 0.45 0.4	108.6	83.0 72.8	9.68 6.69	62.0	111.8 64.8 8.8	1.00%	1122	24.8 8.8 8.8	82.6 0.45 0.65	044 4.69 8.80 8.80	34.0 65.40	93.8 2.4.2 2.8.2.4	95.8 146.2
		Height of Riverbed	10.2	,40 40	8 5 4 0	10.3	N 00	Ģ	7.2	7.7	∞ ∞	20.5	84	46.6 4.0	8.0 0.0	55.23 1.65.23	0.00	o, ∧o, ∧o o, oo, ⊲	4.8 0.1	8.0 8.0	v.v.c	~ C- 00 V 00 C	89.00 89.00	2000 500	800 800
	Table >	Elevation of Riverbed	219.4	225.3 228.8	230.9.	232.6	243.2 244.0	245.2 247.6	250.4	254.4	257.5	263.1	768.4 268.4 268.4	272.2	277.5 280.3	283.6 286.4	290.0 290.0 200.0	293.3 293.3 294.0	298.4 300.7	303.2 305.2	307.5 309.6	314.7 316.8	319.5	3237 326.6 326.6 328.6	333. 331. 33. 33. 33. 33.
		Altinde	229.6	232.8 235.7	239.3 241.7	242.9 245.3	248.8 252.8	252.4 254.8	257.6 259.0	261.0 262.8	265.8	269.8	274.2	278.6	282.5 285.8	285 292.6	296.0 296.0 296.0	299.3 299.3 200.3	302.4 305.8	309.0	6.60 6.40 6.80	322.5 322.5 325.0	3263 3308 3308	332.8 332.8 335.28	337.3
		Point No.	202300	202280 202270	202260	202240	202220	202200 202190	202180 202170	202160	202140	202120	202100	202080	202060 202050	202040 202030	202010	201991 201990 201990	201980	201960 201950	201940 201930 201930	201910 201910 201900	201890 201880	201870 201860 201880	201840 201830

+ Flooding by assumed runoff equal to full of valley depth (depth of water more than 12 meter) O:8,577 m³/sec ++ Flooding by assumed runoff equal to 3/4 of valley depth (depth of water more than 9 meter) O:1,598 m³/sec +++ Flooding by assumed runoff equal to 1/2 of valley depth (depth of water more than 6 meter) O:1,598 m³/sec +++ Flooding by assumed runoff equal to 1/4 of valley depth (depth of water more than 3 meter) O: 393 m³/sec

Overflow point

Runoff (O1)	2008	
Velocity with sedi- ment (VI)	ĸĸąĸĸąĸĸąĸĸĸĸĸĸĸĸĸĸĸĸĸĸĸĸĸĸĸĸĸĸĸĸĸĸĸĸĸ	
Runoff (Q)	80887 80	
Velocity (V)	マッコペネとNコピースペルスペンでは、Composer Ather States Composer States Comp	8,577 m ³ /sec 4,750 m ³ /sec 1,598 m ³ /sec 393 m ³ /sec
Hydraulic Radius (R)	2.25.25.25.25.25.25.25.25.25.25.25.25.25	12 meter) Q: 8,6 9 meter) Q: 4,7 6 meter) Q: 1,3 3 meter) Q: 1,3
Wetted Perimeter (P)	<u>ชัชสีพีรธุรชุรชุรชุรชุรชุรธุร</u> รุรธุรชุรชุรธุรธุรฐรรฐรรรฐรรรฐรรฐรฐรฐรฐรฐรฐรฐรฐรฐร	more than more than more than
Stope of River Course (1)	ਫ਼	of water of water of water of water
River Course Area (A)	2014444444 2014444444 2014444444 2014444444 20144444444 20144444444 2014444444444	depth (depth depth (depth depth (depth depth (depth
Wicth of Riverbod	82282202222222222222222222222222222222	to 3/4 of valley to 3/4 of valley to 1/2 of valley to 1/4 of valley
Height of Riverbed	๐๑๒๙๙๛๛๛๛๚๛๛๛๛๛๛๚๚๚๛๚๚๚๛๛๛๚๛๛๛๚๚๛๛๚๚๚๚๚๚๚๚	runoff equal runoff equal runoff equal
Ekwation of Riverbed	######################################	y assumed y assumed y assumed y assumed
Aftitude	%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%%	+ Flooding b ++ Flooding b ++ Flooding b
Point No.	201820 201820 2017790	Overflow point

‡	ŧ	‡	11	ŧŧ		ŧ	ŧ	1:		
Runoff (O1) (O2) (O2) (O3) (O3) (O3) (O3) (O3) (O3) (O3) (O3	279.3 405.5	377.4 997.9 2312.6 2172.6 1070.6	1241.4 365.3 550.7 7.020	873.0 873.0 6520.7	2501.5 2501.5 2638.5 1627.7 2032.0 1131.6	2329.0	1216.7	1904.4	8577.3 13876.8 14149.9 22756.5 17948.6	;
V (1)	11.K 16.4	प्रकल्प क्षेत्रकृत	4 w u u L'Oùx	4 A Q	0000000 04-01-00	16.00 16.00 16.00	\$ 30 ·	14.0	. oo oo daa aa oo daa aa oo daa	2
Runoff (Q) (Q) (1732.3 1761.7 1761.7 1761.7 1872.3 1872.7	485.4 704.7	655.9 1734.4 4019.3 3776.0 1860.7	2157.5 634.8 957.2 4.99.1	1260.5	2875.0 4347.7 2828.8 3531.6 1965.1	1663.4 4047.8	2114.5	3309.9	24117.8 24592.6 39550.8 31195.0 20416.3	
> 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	120	& C & & C &	∞vv.« ⊣d∺v	0.880 0.70 0.70	01121 01122 0114336 0114336	11.3	10.0	11.1	22355 2355 2455 2455 2455 2455 2455 2455	33/86 33/86 33/86 33/86
Hydraulic (R)	2.51724	1,61728 2,75000 3,57480 2,53846 2,53846	2.65000 1.38636 1.67257 2.10345	2.63333 3.18182 9.00000	4.69643 5.09091 5.44118 5.50588 4.17500 3.49152	3.63043 4.85135 4.74627	4.30612	4.38235	6.38461 6.47325 6.87156 9.82979 8.22857	0 : 8,577 0 : 4,750 0 : 1,598
Wetted Primeter Co. 1 (2) (2) (3) (4) (4) (4) (4) (4) (4) (4) (4) (4) (4	87. 56.	8821 127,240 14,7,240	5 5 5 5 5 5 7		<u> </u>	. 4. C	9.0		222 222 243 243 244 245 245 245 245 245 245 245 245 245	` ១៩៩ ៩
<u>ა</u>	4,4	% \$0,80,80 \$0 \$0,80 \$0 \$0,80 \$0 \$0,80 \$0 \$0,80 \$0 \$0 \$0 \$0 \$0 \$0 \$0 \$0 \$	304. 304.	861. 84.	ద్ది దే ద్వర్ణ దేది దేది.	25.25.25 2.45.25.25	i oʻx	188	88.52.8.88.88 8.52.8.88.88	\$ \$ \$ \$ \$
Area Course (A) 125 (A	22.5	1848 1984 1984 1984 1984 1984 1984 1984	2021 1822 1892 1893 1893	158. 175. 666.	9889888888888888888888888888888888888	359.	211.	25.4 28.8 8.8	296 1525 1286 1286 1286 1286 1286 1286	ភ្ទុំភូទូ
Width	77.0 48.8	846 846 846 846 846 846 846 846 846 846	% % % % % % % % % % % % % % % % % % %	4 4 8 2 2 4 4 6 6 6 6 6 6 6 6 6 6 6 6 6 6 6 6	4004000 9004000 90004990	4.00 4.00 4.00 4.00 5.00	88.4 88.4 1215	200 24.07	2001112 200112 200440 200842 200842	full of valley d 3/4 of valley d 1/2 of valley d 1/4 of valley d
H 978 18.00 84 84 44 44 44 88 88 88 88 86 14 14 18 18 18 18 18 18 18 18 18 18 18 18 18	4 8.0 1.0.	๛๛๛๛ ๔๗๚๕๗๔	৸৻৻৻৻৻৻ ড়৻৻৻৻৻৻৻৻৻৻৻৻৻৻৻৻৻৻৻৻৻৻৻৻৻৻৻৻৻৻	4 8 8 6 2 8	<i>ᲑᲡᲡᲡ</i> ᲑᲐ Ბ Ქ ᲠᲥॐ	8.7.9	10°C	% 6.6 1.0	######################################	qual to qual to qual to
Elevation of the first of the f	\$99.8 600.6	609 609 614 619 619 619 619 619	26.00.00 4.00.00 4.40.00 5.4.4.4.	647.4 653.6 658.6	4488 4488 4488 4488 4488 4488 4488 448	705.5	721.2	732.2	44.87.77.77.77.77.77.77.77.77.77.77.77.77.	by assumed runoff e by assumed runoff e by assumed runoff e by assumed runoff e
Attitude	888 448 844	00000000000000000000000000000000000000	635.0 641.7 646.8	652.3 658.8 665.2	64717 64747 68818 68818 6978 6060	711.3 717.1 722.3	728.1	738.8	0,445 0,445 0,445 0,448 0,548 0,548	퓛렃퓛뀵
Point No. 2012/2013/00/2012/2012/2012/2012/2012/20	201120	201115 201110 201000 201080	201060 201060 201050	201030 201020 201010	201,000 \$00990 \$00980 \$00970 \$00960 \$00950	500930 500920 500910	\$00800 \$00890	\$00800 \$00870	500850 500880 500830 500820 50080 50080	.š

		Ī		Ī					Ī																			
	Runoff (O1)	137.6	3523 3523	359.5 359.5	375.1	1090.0	954.7	1641.4	179.5	1133.2	804.1	434.6	1356.7	2176.2	2166.2	2111.3	2540.2 5317.2	7539.9	6565.5	6874.1	3233.7	4102.9	\$403.1	5944.1 5944.1	1194.5	2824 29292 5 5 5 5	3519.6 2555.1 3122.7	
	Velocity with sedi-	1.7	, ci -i . ci -i	6.6.6 6.6.6	 	6.5 6.4	iw ব বৃত্ত	্ৰ ল গ্ৰাক	11 € 20 €	9.5°	25	io.	9 4 4 4 6 5	i wi	4 4 0 J AJ 0	2.4 2.7	4 m Ów	6.6 8.0	2,2 2,2	χ, λ, α α		4.0	9 v9 v	4.00 4.00	3.67	0 -1 0 -1	ભ 4 બ 4 4	
	Runoff	164.8	617.4	216.2 430.7 638.1	1872.8	1305.8	1143,7	1966,4	215.1	1357.6	963.3	520.7	1625.3	3782.3	3764.8	5402.1	4414,9 9241.3	13104.3	11410,8 15962,1	11947.2 8464.8	\$620.2 \$033.6	7130.9	9390.6	10330.9	2076.1	9099.0 10.5	6117.0 4440.8 5427.3	
Krasak)	Velocity (V)	77. 0.2	inni inni	16.4 19.4		2.9 9.0	5.7	8,4 E.F.	ų ų x	4 4 C. S.	6.59 6.49	en e l∞a	κ. « Αν 4	w. Go	- 1- A	0 0 0 00 0	. 6 . 7	11,3	9.0 11.3	10.1 10.7	& C & O	8 9	1000	0 II 7 4 4		4.16379	27.6	7,616 m ³ /sec
Flood (K. X	Hydraulic Radius (R)	0.67213	1.15306	1.47561	1.25000	1,82514	2.03650	2.60563	0.68243	2,25000	3.08721	1.83784	2,70909	4.58993	3.53676	5.18750	6.26250	6.36364	5.40171 7.00000	6.22632 5.77622	4,03145	3.98617	5.84337	5.10112	3.30303	116.	4,61111 3,81579 4,39264	meter) Q : 27,61
the	Wetted Perimeter (P)	122. 168.	196. 208.	102. 101.	108. 158.	183. 211.	137. 140.	142.	\$ 50 \$ 50 \$ 50 \$ 50 \$ 50 \$ 50 \$ 50 \$ 50	128.	140. 172.	7 .00	173.	139.	136.	128	160.	708. 209.	202 202.	15.00 15.00 15.00	159. 161.	217.	166.	178	132	8	152. 163.	than 20
Valley for	Slope of River Course (1)	50. 20.	3.5. 5.0.	15.5	23.4		22.5	25.33 15.433	ដូនូវ	5 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1	នុង	6,4	ន្តន	15. 28.	37.	145	77.	ද්දේ	33.73	တို့ ကို လူ	5.23	38. 0.	3.5	4 .6	2.5		ઌ૽ૢઌૢ૽ઌૢ ૹ૽ઌ૽ૢૹ૽	of water more
the	River Course Area (A)	822. 183.	226. 191. 95.	121.	135. 375.	223 425 425	409	370.	101. 236.	354 354 54	230. 531.	204.	298. 815.	638. 696.	539.	466	1002	1330.	1414.	28.5 82.6 9.5 9.5 9.5 9.5 9.5 9.5 9.5 9.5 9.5 9.5	641. 716.	865. 496.	970.	900	436.	101.4	580. 716.	(depth
apacity of	Width of Riverbed	161.8	200.0 203.6 31.2	77.8 92.2	103,6	162.2	130.0	135.0	171 170 140 140	116.8	102.8	8 4 33.2	100.8 159.4	125.6	125.8	109.4	143.0	1895	182.4	124 244 244	46.8	196.6 201.2	154.6	139.8	120.8	00 f	135.8 135.6 144.8	full of valley depth
6 The Ca	Height of Riverbed	E.9.	ичи 00	22.2	น ผู้ผู้จ	20,00	14. 20:	ત્રુલ જુ⊄(in.	14.4 1.60	င့်သင့် လ ဂါလ	44. VO	0, 0, 4, 4,	7.1	8.2 6.0	3.7 5.0	000	200	\ 0\ 0	n es	9 85 5 80 6 80 6 80 8 80 8 80 8 80 8 80 8 80 8	2. Q 8. 4.	8.2 10.2	9.5 9.6	6.2 9.1	346.0	, %, ₇ ,	runoff equal to fi
Table (Elevation of Riverbed	234.6 236.6	2223 2403 2403 2403 2403	244.6	251.0	255.0 255.0	258.8	263,4 263,4	167.8	272.2	276.2	7,877 780.0	281.0 284.6	286.7 287.6	292.2 295.0	297.2 299.8	301.0	308.7	0.00 0.00 0.00 0.00 0.00 0.00 0.00 0.0	319.5 219.5 27.5	323.4	327.6	329.8 332.8	336.4 324.4	342.8 344.2	346.0	350.0	sumed
	Altitude	2335.6 238.5 25.5 25.5	2242 242:2 243:2	246 446 80 80	0.425 8.44 8.25 8.35 8.35	256.9 260.0	263.3	266.8 266.8 267.6	271.0	276.8	10,5	284.0	284.0 291.0	295.2 295.2	297.4 301.0	305.8 307.0	311.0	318.4	100 100 100 100 100 100 100 100 100 100	25.00 25.00	329.2	334.0	338.0 343.0	346.0 348.0	359.0	354.2	358.0 359.0	+ Flooding by a
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(depth of water more than 20 meter) Q: 27,616 m3/sec (depth of water more than 15 meter) Q: 12,535 m3/sec (depth of water more than 10 meter) Q: 4,968 m3/sec (depth of water more than 5 meter) Q: 634 m3/sec Flooding by assumed runoff equal to full of valley depth Flooding by assumed runoff equal to 3/4 of valley depth Flooding by assumed runoff equal to 1/2 of valley depth Flooding by assumed runoff equal to 1/2 of valley depth *‡‡

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Hydraulic Radius (R)	3.59091 4.15584 5.16552 4.11765 3.72951	3.57143 2.55207 2.53843 2.58853 2.58859 2.64314 3.44776	2.78289 2.88989 2.88989 2.88989 2.889889 2.889889 2.89889	1.76623 3.57143 3.08965 1.60060 748461 748461	3.73 3.60 3.60 3.60 3.90 3.90 3.90 5.00 5.00 5.00 5.00 5.00 5.00 5.00 5	6,000 6,000	4.76554 5.30025 6.11290 5.15262 5.66738 3.88400	1.20 meter) Q: 27,616 1.15 meter) Q: 12,535 1.10 meter) Q: 4,968 1.5 meter) Q: 634
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River Course Area (A)	4444 8860 8860 888 888 888 888	000 000 000 000 000 000 000 000 000 00	7077 7077 7085 7085 7087 7087 7087 7087	22.24 22.24 22.24 23.24 24.24 25.24 25.24 26.24	227 227 227 227 227 227 227 227 227 227	2002 1986 1986 1761 2093 77	1687, 20136, 1648, 1940, 1940,	depth (depth depth (depth depth (depth depth (depth
Width of Riverbed	0.02211111 0.022111111 0.032191101	90111144 901080444 90108041444 90108040444	22 22 22 24 25 25 25 25 25 25 25 25 25 25 25 25 25	24 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1	25.00 20.00	28.14.14.14.1 98.44.14.14.1 140.44.14.88.04.0	7114411111 782442011 7684664001	full of valley 3/4 of valley 1/2 of valley 1/4 of valley
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omt No.	401680 401680 401675 401670 401665	401659 401658 401655 401656 401640 401640	401633 401623 401615 401610 401603 401600 401595	401585 401580 401575 401570 401565 401555	401550 401545 401535 401535 401524 401524 401526	401510 401505 401495 401495 401480 401485 401475	401450 401465 401450 401450 401440 401440	erflow point

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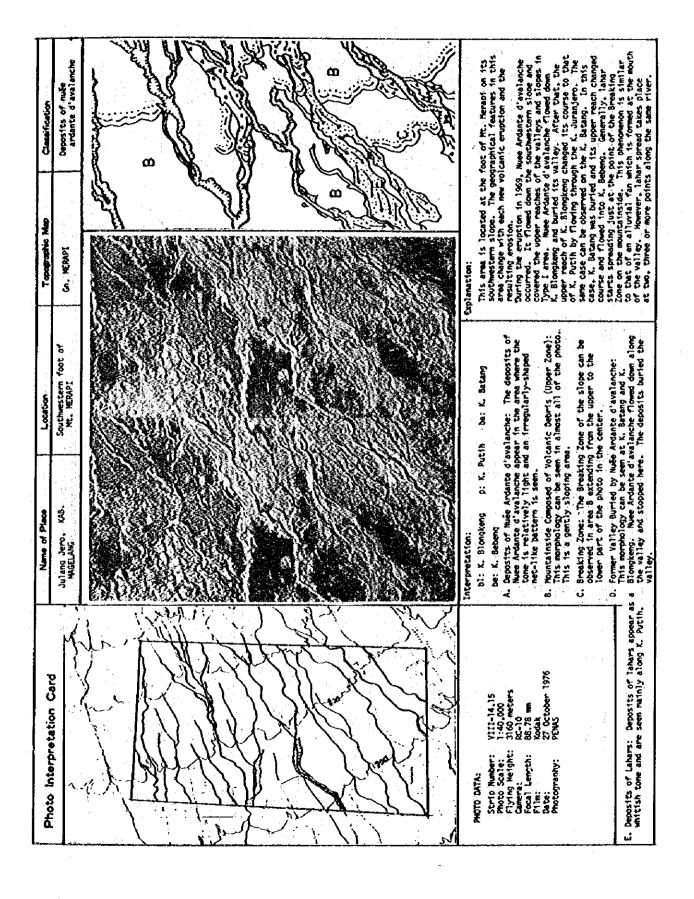
		‡‡	‡ .	
Runoff (Q1)	27983.9 28530.7 21104.7	21480.1 9081.4 9645.7 21826.5	17529.2 32106.6 50020.1 23013.4	
Velocity with sedi- ment (VI)	6140 6446 6446	31188 9 125 357 8 6 7 5	15.0.15.	
Runoff (Q)	48636.0 49586.3 36680.1	373324 373324 167643 379344 420938	304065.8 55801.3 86934.9 39997.2	•
Velocity (Y)	23.2 21.3.2 17.8	225 225 275 275 275 275 275 275 275 275	26.1 26.1 26.1 23.7	ć
Hydraulic Radius (R)	12.48809 11.48768 8.91342	8.46875 7.68345 8.30400 12.01399	10.68243 12.90361 12.69349 10.89032	
Wetted Perimeter (P)	283. 231. 531.	44.85.44.55 44.85.44.55	148. 166. 261. 155.	
Slope of River Course (1)	6699	89484444	<u> </u>	
River Course Area (A)	2098. 2332. 2059.	1897. 1068. 1038. 1718.	1581. 2142. 3313. 1688.	
Width of Riverbed	126.0 162.5 177.5	1828 82.0 82.0 82.0 82.0	92.0 123.5 176.0 115.5	
Height of Riverbed	18. 2002 5000 5000	711811 71811 719 719 719 719 719 719 719 719 719 7	12241 10400	:
Elevation of Riverbed	690.4 697.8 703.0	72007 72007 73007 7007 7007 7007 7007 7007 7007 7007 7007 7007 7007 7007 7007 7007 7	788.0 784.6 764.2 7.1.7	
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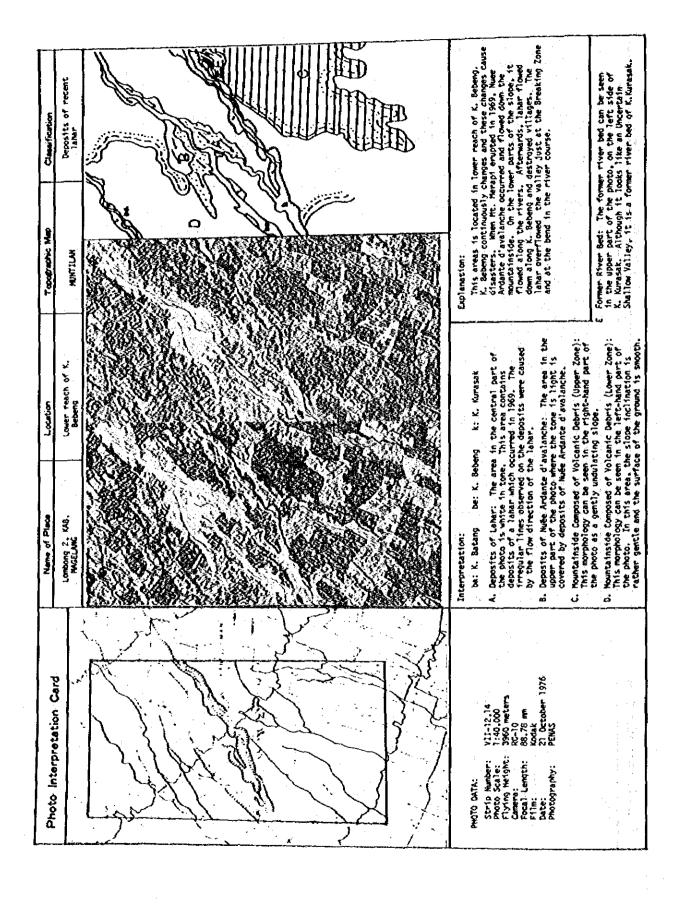
Point No.

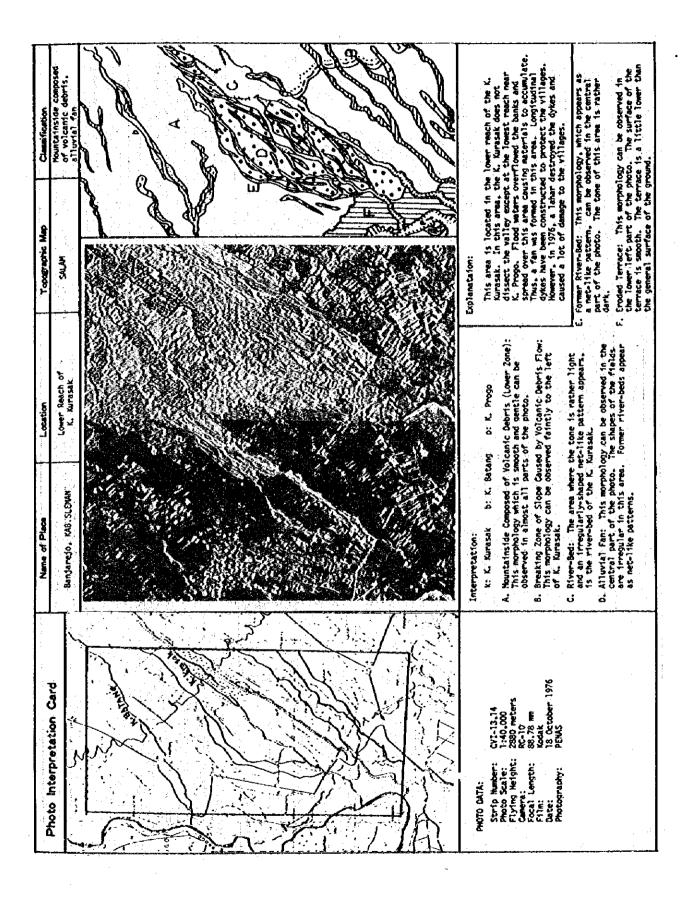
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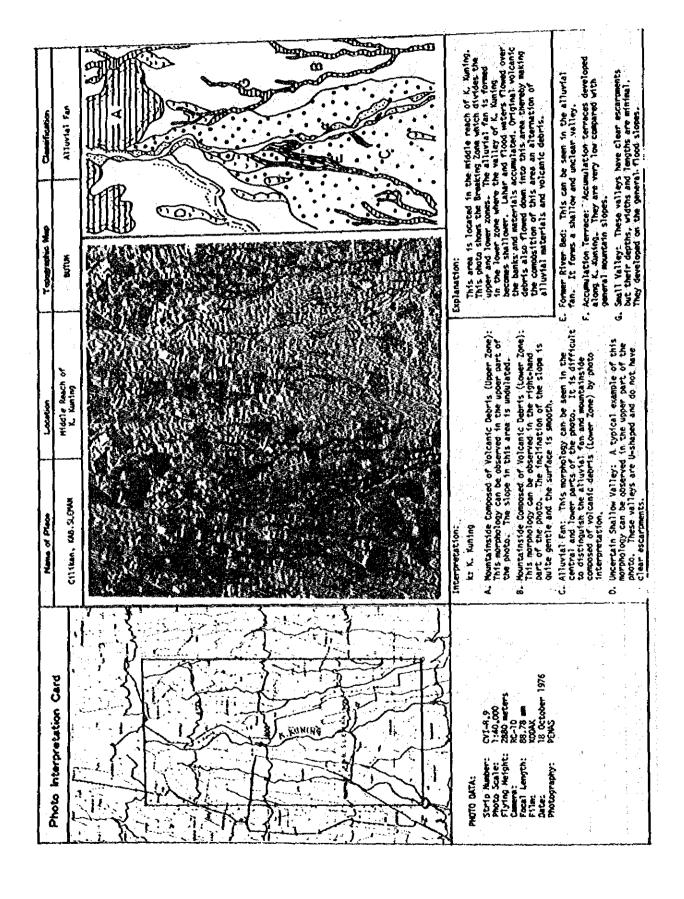
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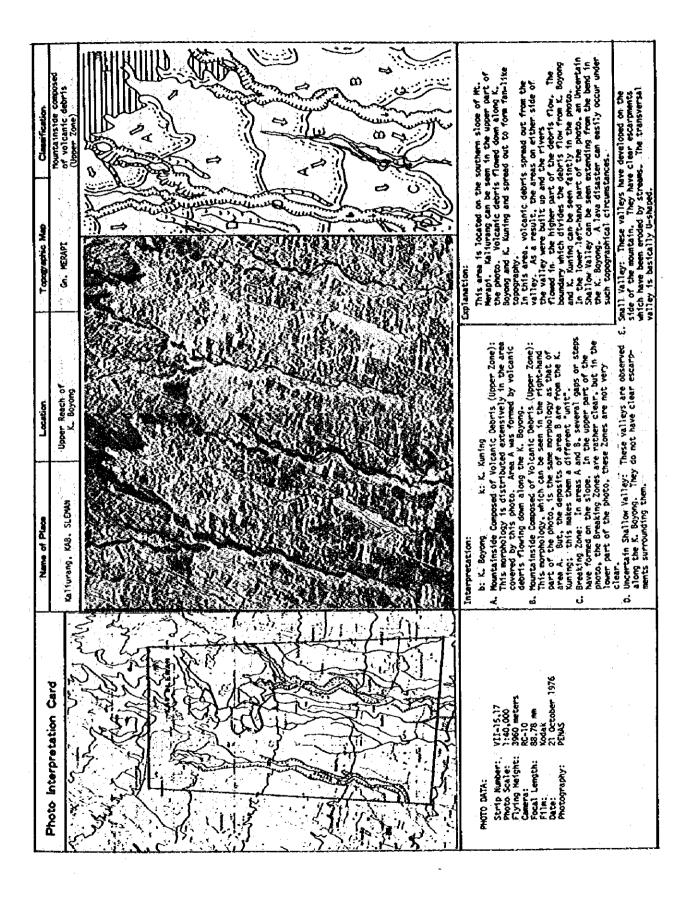
Photo Interpretation Cards

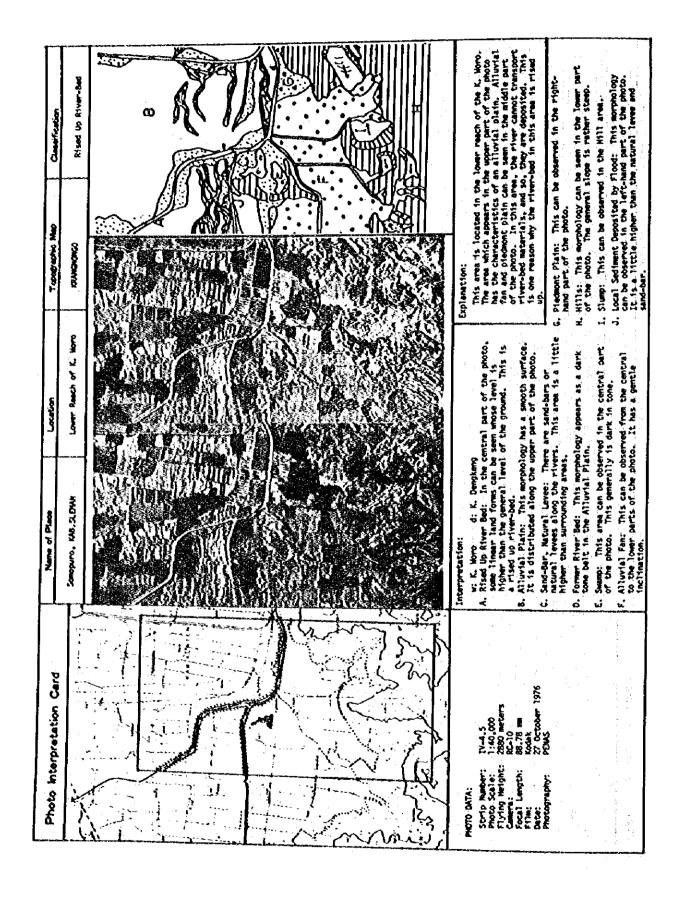


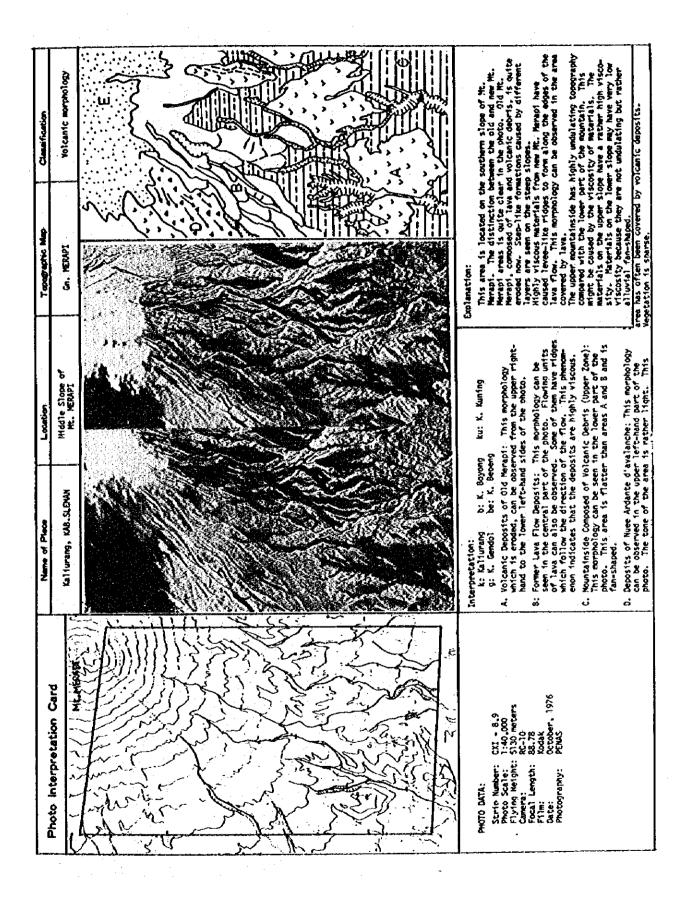












SUPPORTING REPORT B GEOLOGY AND LAHAR DEPOSITS

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1. GEOLOGICAL BACKGROUND OF G. MERAPI

1.1 Stratigraphy

The geology in the area of G. Merapi can be classified into the following five groups which are also presented in Table 1.

(1) Recent Volcanic Products from Eruptions of G. Merapi

This group consists mainly of lava flows and pyroclastic fragments from eruptions since 1888, the exact dates of which are known, and lahar that has occurred since 1930.

(2) Terrace and Other Unconsolidated Deposits

This group consists mainly of present riverbed deposits, terrace deposits, alluvial fan deposits in the lower stream area, and other alluvial deposits.

(3) Young Merapi Volcanic Products

This group includes labar deposits and lava-flows before 1930 (exact dates are not known).

(4) Old Merapi Volcanic Products

This group consists of Old Merapi lava flow, intrusive and pyroclastic rocks.

(5) Base Rocks

This group includes volcanic rocks and deposits of the tertiary period which form the foundation of G. Merapi.

1.2 History of Volcanic Activity of G. Merapi

G. Merapi has had a long recorded history of eruptions. There is a record of damage caused by a major lahar flow in 1006 on a stone monument with Sanskrit inscriptions of old Javan stories. There is also a record of some three thousand deaths caused by a major eruption in 1672 which produced glowing clouds and volcanic mud flows. A record of the mountain's volcanic activity after 1800 was kept by R.W. Bemmelen. This record is summarized in Table 2 along with the records for recent years.

The summary suggests a cycle of 1-7 years of activity followed by a dormant period of 1-6 years. Except for recent years, the exact scale of damage caused by such activity is unknown.

Records indicating the direction of lava flows in recent years is summarized in Fig. 1. This summary shows that there has been a shift in the direction from NW-W-SW-N-NW-W-SW. Since the old Merapi volcanic wall in the southeast remains, it can be expected to control lava overflows in that direction. The principal nuée ardente danger zone should continue to be in the direction of SW-W-NW.

1.3 Terminology of Volcanic Products

In this "Study for Making Master Plan" the terminology of volcanic products used by the Japanese Study Team is based on R.W. van Bemmelen, I. Suryo and field surveys by the team as shown Fig. 2.

(1) Nuée Ardente: Avalanche Type

During an active stage of G. Merapi, viscous lava may flow-over the crater rim and form a lava tongue. Parts of the unstable lava mass may callapse and red hot lava blocks will roll down the slope, breaking into numerous lava fragments. Some of the fragments will turn into glowing sand and ash, and form a suspension of glowing particles in hot air, which will descend as a dense cloud.

N.W. Van Bezzelen is of the opinion that lava blocks measuring about 10023 do not cause nuée ardente, but only common avalanches, and that the development of smoke and dust from such avalanches is small. If thousands of cubic meters of lava break off, the development of heat and pulverized lava is so great, that the avalanche appears to expand rapidly. This incandescent flow of volcanic debris is called "ladu" in Indonesian.

Such nuce ardente of the avalanche type at G. Merapi usually move 1 - 3km from the crater, but big avalanches may reach a distance of 7km.

(2) Nuée Ardente: Explosion Type

Poring volcanic explosions, lava masses may be emitted from the eruption vent. The lava is thrown upwards and a part of it flows over the crater rim and rushes down the slope. This flow of incandescent debris is accompanied by the nuce ardente.

The difference between nuée ardente of type 1 and 2 is that the explosion type is always accompanyed by a rising califlower cloud of volcanic ash.

(3) "Lahar" Mud Flows

The Indonesian term "lahar" is used to describe the volcanic mud flows on the Merapi and Kelut volcanoes. Lahar of Merapi are classified as rain lahar.

During or after eruption, a huge amount of volcanic material is deposited on the slope of the volcano. Rain water mixed with this unconsolidated hot volcanic material turns it into a high density mass, which descends with increasing velocity and forms lahars in the middle and lower reaches. Such lahar is called hot rain lahar until they cool.

The concept of lahar is very similar to the Japanese "Dosekiryu", in respect to massive straight movement and the inclusion of big blocks and thick and poorly sorted sediments. The characteristics of sediment products and discharges in Indonesia are very similar to those in Japan. (See Appendix-1; 8-48)