CHAPTER 4 ANALYSIS OF RAINFALL

4.1 Rainfall Mechanism

4.1.1 Characteristics of Rainfall

Monthly rainfall and rain-days observed at the rain gauges installed by the project (Table B-1-3, Fig. B-1-1) during June, 1983 to July, 1985 are shown in Table B-4-1. Daily rainfall is shown in Fig. B-4-1. The data of the rain gauges installed after 1984, are excluded from the figure. The rain gauges in the upper part on this figure -- from RA3 to RK6 -- are in the mountain area, in the middle part -- from RA2 to RM4 -- are on and around the boundary between the mountain and the marginal wadi plain area, and in the lower part -- from RA1 to RM2 -- are in the marginal wadi plain and the gravel plain area. Monthly mean rainfall and rain-days in the mountain and in the gravel plain are shown in Fig. B-4-2.

The characteristics of rainfall in the project area, suggested from those figures, are described below.

- There was no rain in the gravel plain in summer (June to September) except one rainfall on August 10 - 11, 1983.

- Rainfall was observed at a few rain gauges on the same days, in summer except one rainfall on August 10 - 11, 1983

Although rain-days are not many in winter, rainfall was observed at many rain gauges on same days.

Three representative rainfalls (on August 10 - 11, 1983, on August 25, 1984, and during December 28, 1984 to January 1, 1985) are analyzed and the results are described below.

1) The rainfall on August 10 - 11, 1983

The distribution of the rainfall on the weather chart is shown in Fig. B-4-3, and the distribution in comparison with NOAA data is shown in Supporting Report (F). Time series of hourly rainfall at each gauge is shown in Fig. B-4-4.

The tropical cyclones, originating in the Arabian Sea and migrating westerly to Mashirah Island in southeast of Oman, brought rainfall over the project area. The rainfall moved from east to west. It brought heavy rainfall to the mountains and light rainfall to the gravel plain; no rainfall was observed at Saham on the coast of Wadi Ahin. The rainfall continued intermittently during more than 24 hours. The heaviest rainfall was observed at RF1, 88.5 mm.

2) The rainfall on August 25, 1984

The distibution of the rainfall on the weather chart is shown in Fig. B-4-5 and time series of hourly rainfall in shown in Fig. B-4-6. Though the low pressure spreads to Northwest India and Arabian Peninsula, this is the typical weather chart in summer. Even though rainfall continued less than an hour in a very small area, the intensity was very high.

3) The rainfall during December 28, 1984 to January 1, 1985

The distribution of rainfall on the weather chart is shown in Fig. B-4-7, and the distribution in comparison with NOAA data is shown in Supporting Report (F).

The cold front, stretching from the cyclone, moved east over Oman and brought the rainfall. Time series of hourly rainfall on December 29, 1985 when the front passed, are shown in Fig. B-4-8. Although much rain came on the coastal area, the rainfall occured over the whole project area except for a part of the mountain. As the cold front was passing easterly, the rainfall was also moving from west to east.

The maximum rainfall intensity observed at each rain gauge is shown in Table B-4-2. Although all the rainfall in the mountain on the figure is in summer, almost of all rainfall in the gravel plain and the marginal wadi plain (RAI, RG1, RG2, RF4, RM1, RM2) are in winter. The strong rainfall for long periods in the mountain were caused by the tropical cyclone on August 10 - 11, 1983. On the other hand, the strong rainfall during a short period in the mountain was recorded on each of the days.

As the result of the analysis, the rainfalls in the project area can be divided into three types: rainfall caused by tropical cyclones, rainfall in summer caused by other phenomena, and rainfalls caused by synoptic disturbances in winter.

The second type of rainfall have strong intensity. Seventy percent of the rainfalls lasted less than one hour, and almost all rainfall came during one hour even when that rainfall continued more than one hour.

The correlation of the summer daily rainfalls between RF1 and RF2, which are the nearest rain gauges in the mountain, is shown in Fig. B-4-9. Although those are only 4 km apart, both of them experienced only six rainfalls out of twenty-four rainfalls which were either observed. Consequently, the rainfall area is narrow. These characteristics (narrow are a, intense and short rainfalls) suggest that rainfalls are caused by cumulus and cumulonimbus clouds formed under strong ascending air current.

The distribution of occurence time of the summer rainfall is shown in Fig. B-4-10. It shows that 85% of the rainfalls comes during one to six in the afternoon, when the sea breeze circulation grows strongly. Consequently, the cause of the rainfall can be considered as the sea breeze circulation.

In summary, there are three clear causes of the rainfalls in the project area: tropical cyclones, sea breeze circulations, synoptic disturbances.

Table B-4-1	Monthly Rainfall and R	ain-day (June 1983-July 1985)

				1983				<u>.</u>			•••••••	1007			~ <u>~</u>	· · · · · ·	÷	a						Lt: Nu	umber	, ณ
Station	JUN		F	r	(<u> </u>	•. 1 •		:		r 	·····	1984	· · · · · ·			:. 1	.				·		1985			.
RA 1	(0)	<u>JUL</u> (0)	AUG (0)	SEP (0)	0CT (0)	<u>NOB</u> (0)	DEC (0)	JAN (0)	FEB (0)	<u>MAR</u> 0)	APR (0)	<u>MAY</u> (0)	JUN (0)	JUL (0)	AUG (0)	SEP (0)	0CT (0)	(0)	DEC (1) 26.0	JAN (0)	FEB (0)	MAR (1) 0.5	APR (0)	MAY (0)		טנ נס)
RA 2	(x)	(0)	(2)	(0)	(0)	(0)	(0)	(1)	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(1)	(0)		(1) 3.5	(0)	(0)	I	(0)	(0)	(0)	(0)
RA 3	(x)	(0)	(3) 64.0	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(1) 2.5	(0)	(3) 18.0	(1) 2.5	(4) 11.0	(0) 0.0	(0)		(0)	(0)	(1) 0.5			(0)	1.0
RA 4	(x)	(x)	(x)	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(1) 0.5	(0)	(0)	(2) 8.5	(0)	(2) 4.5	(0)	(0)	(2) 1.0	(0)	(0)	(0)		(0)	(0)	
RA S	(x) ()	(x) (-)	(3) <u>12.0</u> (-)	(1) <u>6.0</u>	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(1) 0.5	(0)	(4) 36.5	(0)	(0)	(2) 3.5	(0)	(0)	(1) 2.5		(0)	(0)	(2 17
RA 6	(0)	(0)	(1)	(~) (0)	(-) (0)	(-) (0)	() (0)	() (1)	(-) (0)	(-) (0)	() (0)	(-) (0)	(-) (0)	(-)	(-)	(-) ((0)	(-) (0)	() (0)	(3) <u>18,5</u> (1)	(0) (0)	(0)	. (0) - (0)	، سمی میں	(0) (0)	(0)	(0) (0)
RG 1	(0)	(0)	(1) 1.5 (1)	(0)	(0)	(0)	(0)	(1) 4.5 (0)	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(1) 28.0 (1)	(1)	(0)	(1)		(0)	(0)	(0
RG 2 RG 3	(0)	(0)	7.0 (x)	(1)	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(1)	(0)	(0)	(0)	(0)	18.5	<u>0.5</u> (0)	(0)	$\frac{1.5}{(1)}$	(2)	(0)	(0)	(0)
RG 4	(0)	(1)	(3)	16.5 (1)	(0)	(0)	(1)	(1)	(0)	(0)	(2)	(0)	(1)	2.5	(x)	(x)	(x)	(0)	3.5	(0)	(0)	0.5	5.0 (0)	(0) [:]	(0)	(0)
RG 5	(0)	<u>13.5</u> (1)	<u>47.5</u> (3)	<u>1.0</u> (0)	0.0	0,0 :(0)	<u>3.0</u> (0)	<u>1.0</u> (0)	(0)	(0)	11,0 (2)	(0)	2,5 (1)	<u>13.5</u> (4)	+ (2)	+ (1) 3.5	+ (0)		9.0 (2) 1.0	(0)	(0)	5.0 (2) 3.0	(0)	(0)	(2) 27.0	(0)
RF 1	(0)	7.5 (1) 1.0	41.0 (3) 93.5	(1) 7.0	(0)	(0)	(0)	(0)	(0)	(0)	2,5 (2) 24,5	(0)	<u>1.0</u> (0)	5.0 (3) 26.0	9.5 (1) 0.5	(2)	(0)	(0)		(1) 1.5	(0)	(2)	(0)	(1) 0,5	(1) 9.5	(0)
RF 2	(0)	(2)	(4) 59.5	(3)	(0)	(0)	(0)	(0)	(0)	(1)	(2)	(0)	(0)	(5)	(4) 27.0	(5) 41.5	(0)	(0)	(0)	(1) 3.0	(0)	(2) 17.5	(0)	(2) 10.5	(1) 26.5	(2) 15
RF 3	(0)	(2) 18.0*	(3) 29.0*	(0)	(0)	(0)	(0)	(0)	(0)	(1) 3.0	(2) 26,5	(1) 1,0	(x) x	(x) x	(2) 3.0	(3) 8.0	(0)	(0)		(2) 7.0	(0)	(3) 15.5	(0)	(1) 5.0	(1) 0.5	(4) 11
RF 4	(0)	(0)	(1)	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(0)	° (0)	(0)	(0)	(0)	15.5	(0)	(0)	(0)	(0)	(0)	(0)	(0)
MF 1	(x)	(x)	(1) 6.5	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(0)	9.0	(1) 0.5	(0)	(1) 1.0	(0)	(0)	.(0) 	(0)
MF 2	(0)	(0)	(8) <u>45.0</u>	(3) <u>13,6</u>	(0)	(0) (0)	(2) 0.4 (0)	(2) 0.4 (0)	(1) <u>0.2</u> (0)	(x) (0)	(x) (0)	(0)	(0) (0)	(0)	(0) (0)	(0)	(0)	(0) (0)	(3) <u>13,5</u> (1)	(0) (0)	(0) (0)	(1) 3.0 (2)	(0) (0)	(0) (0)	(0) (0)	() (0
RK 1	(0) (0)	(0)	(1) 7.5 (3)	(1) <u>2.0</u> (1)	(0) (0)	(0)	(0)	(0)	(0)	(0)	(1)	(0)	(0)	(1)	(3)	20.5	(0)	(0)	8.0	(0)	(0)	4.5		(1)	(1)	(1
RK 2	(0)	(0)	10,5 (3)	3.5	(0)	(0)	(0)	(0)	(0)	(x)	10.5 (x)	(1)	(0)		30,5 (5)	18,0 (6)	(0)	(0)	6.0 (0)	(1)	(0)	4.0	(1)	(1)	$\frac{11.5}{(2)}$	9
RK 3	(0)	(1)	<u>49.0</u> (8)	2.5_ (0)	(0)	(0)	(0)	(0)	(0)	(1)	(2)	2.5	(0)	(3)	(5)	<u>16.5</u> (3)	(0)	(0)	(0)	4.5	(0)	(2)			$\frac{10.0}{(1)}$ 18.0	(3
RK 4	(0)	<u>11.0</u> (0)	51.5 (4)	(2)	(0)	(0)	(0)	(0)	(0)	3.0 (0)	(2)	(0)	(0)	(2)	<u>32.5</u> (0)	(1)	(0)	(0)	(x) ++	2.0 (x) ++	(x) ++	10.5 (1) 1.5	(0)	(1)	(1) 0.5	(0
RK 5 RK 6	(0)	(5)	19.5 (4)	<u>35.5</u> (2)	(0)	(0)	(0)	(0)	(0)	(1)	2.0 (2)	(0)	(0)	5.5 (4)	(5) 22,0	3.0 (5) 23,0	(0)	(0)		(0)	(0)	(2) 13.5	(2)	(2) 10.0	(0)	(5 52
	· ()	<u>31,0</u> (-)	76.5 ()	<u>16.0</u> (-)	ı(-)	·.(-)	(-) ¹	(-)	()	(-)	<u>14.0</u> (-)	(-)	(-)	(-)	(-)	(-)	(-)	(-)		.(-)	(-)	1	(0)	(0)	(0)	(0
RK 7 RM 1	(0)	(0)	(1)	(0)	(0)	(0)	(0)	(1) 0,5	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(1) 38.5	(0)	(0)	(1)	(0)	(0)	(0)	(0
RM 1 RM 2	(0)	(0)	4,0	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(1) 7.0	(0)	(0)	(0)	(0)	(0)	(0)	(0
RM 2	(0)	(0)	(1)	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(0)		(1) 11.5	(0)	(0)	4.0	(1) 1.0	(0)		((
RM 4	(0)	(x)	<u>11.5</u> (x)	(x)	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(0)	(1). 0.5	(0)	(0)	(2) 1.0	(1) 1.0	(0)	(2) 23,5	(0)	(0)	(0)	10

Note: -: Before installed of rain gauge x: No record *: Recorded partly and total rainfall in July and August is 60.0mm +: Total rainfall in August September and October is 3.0mm ++: Total rainfall in December 1984 and January, February 1985 is 3.0mm

Table B-4-2 Maximum Rainfall Intensity (June 1983-July 1985)

					***	· .	····		
					D	uration			
Ra	in gauge	15 min.	30 min.	60 min.	120 min.	6 hr.	12 hr.	24 hr.	48 hr.
	RAI	8.5 (29/12/84)			ł . ł	26.0 (29/12/84)			
LN	RA2	3.5 (10/8/83)	4.0 (10/8/83)	4.0 (10/8/83)	7.5 (10/8/83)	11.5 (10/8/83)	18.5 (10,11/8/83)	20.0 (10,11/8/83)	
I AHIN	RA3	9.0 (29/8/83)	12.5 (29/8/83)	13.0 (29/8/83)	15.0 (10/8/83)	29.5 (10/8/83)	41.5 (10,11/8/83)	48.0 (10,11/8/83)	49.5 (10,11/8/83)
NADI	RA4	5.0 (15/7/84)	5.0 (15/7/84)	7,5 (15/7/84)	8.0 (15/7/84)				
	RA5	10.0 (5/9/84)	19.0 (5/9/84)	20.5 (5/9/84)					
Ř	RC1	19.5 (29/12/84)	19.5 (29/12/84)	20.5 (29/12/84)	24.0 (29/12/84)	26.0 (29/12/84)	26.0 (29/12/84)	28.0 (29/12/84)	
GHAFIR	RG2	12.5 (29/12/84)	12.5 (29/12/84)	16.0 (29/12/84)	16.0 (29/12/84)	18.5 (29/12/84)			
BANT G	RG3	10.0 (15/9/83)	14.0 (15/9/83)	16.0 (15/9/83)	16.5 (15/9/83)		-		
	RG4	13.0 (28/7/83)	13,5 (28/7/83)	13.5 (28/7/83)	13.5 (28/7/83)	28.0 (28/7/83)	37.0 (10/8/83)		
WADI	RG5	10.0 (10/8/83)	16.5 (10/8/83)	17.5 (10/8/83)	18.0 (10/8/83)	30.0 (10/8/83)	32.5 (10/8/83)	37.0 (10/8/83)	
	RF1	13.5	14.5 (10/7/84)	23.0 (10/8/83)	34.5 (10/8/83)	61.5 (10/8/83)	88.5 (10/8/83)		
	RF2	19.5	26.0 (22/6/85)	26.5 (22/6/85)	39.0 (10/8/83)	53.0 (10/8/83)	54.0 (10/8/83)		
AL-FARA'	RF3	12.5 (10/7/83)	14.0 (11/4/84)	14.0 (11/4/84)	14.0 (11/4/84)	15.0 (10/8/83)	18.5 (10/8/83)	19.5 (10,11/8/83)	
	RF4	12.0 (29/12/84)	12.5 (29/12/84)	12.5 (29/12/84)	12.5 (29/12/84)	15.5 (29/12/84)			
WADI	MF1	3.0 (29/12/84)	5.0 (29/12/84)	7.0 (29/12/84)	9:0 (29/12/84)				
	MF2	6,4 (28/8/83)	11.0 (28/8/83)	17.6 (28/8/83)	19.0 (28/8/83)	19.0 (28/8/83)	20.4 (10/8/83)	21.0 (10/8/83)	21.2 (10,11/8/83
	RKI	15.0	20.0 (8/9/84)	20.5 (8/9/84)				·	
KHARUS	RK2	14.0 (24/8/84)	18.0 (24/8/84)	19.0 (24/8/84)		÷.,			
	RK3	18.0 (25/8/84)	35.0 (25/8/84)	43.0 (25/8/84)	· · ·				r 9
BANI	RK4	15.0 (15/6/85)	18.0 (15/6/85)	19.5 (10/8/83)	20.0 (10/8/83)	25.0 (10/8/83)	25.5 (10/8/83)		
WADI	RK5	16.0 (15/9/83)	18.5 (15/9/83)	19.5 (15/9/83)	н. 1				
3	RK6	15.5 (3/9/84)	15.5 (3/9/84)	23.0 (25/7/83)	23.0 (25/7/83)	33.5 (10/8/83)	39.5 (10/8/83)	72.0 (10,11/8/83)	
Н	RMI	27.0 (29/12/84)	29.5	29.5 (29/12/84)	37.5 (29/12/84)	38.5 (29/12/84)			
AL-MA'AWIL	RM2	4.5 (29/12/84)	4.5 (29/12/84)	4.5 (29/12/84)	5.0 (29/12/84)	7.0 (29/12/84)			· .
AL-M	RM3	2.5	2,5 (10/8/83)	4.5 (10/8/83)	5.5 (10/8/83)	10.0 (10/8/83)	11.5 (10/8/83)		
NADI	RM4	18.5 (18/4/85)	18.5 (18/4/85)	18.5 (18/4/85)	18.5 (18/4/85)	18.5 (18/4/85)	18.5 (18/4/85)	18.5 (18/4/85)	23.5 (18,19/4/8

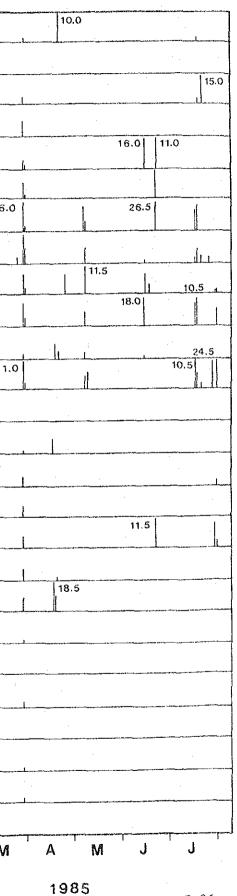
lower row: (date)

	RA3	43.5 14.5	
	RA4		- -
	RA 5	20.5 11.6	
ea	RG4	13.5 37.0	
n Area	RG5	36.5	
Mountaín	RF1	88.5 24.0 14.5 14.0 1	
	RF 2	13.5 53.0 15.0 21.0 21.0 11.5 11.5	16.0
Major	RF 3	1.25 19.0 12.5 14.0	
Σ	RK3	37.5	
	RK4	11.0 25.5	· · · · ·
	RK5	17.5 19.5 16.0	
	RK6	23.0 40.0 32.0 11.5	11.0
	RA2	15.0	
n/ reas	RG3	16.5	
Marginal Wadi Plain/ Frontal Mountain Areas	MF 2	21.0 19.0	· · · ·
adi unta	RK1	20.5	
al W l Mo	RK2	10.5 19.0 10.5	
rgin onta	RM 3	11.5	
Ma Fr	RM4		
	RA 1	26.0	
	RG 1	28.0	· · · · · · · · · · · · · · · · · · ·
Area	RG 2	18.5	· · · · · ·
Plain Area	RF4	15.5	
	MF 1		
Gravel		38.5	
0	BM1 RM2		
	, g ; ==	JASONDJFMAMJJASOND	JFM

Fig. B-4-1 Daily Rainfall Time Series

1983

1984



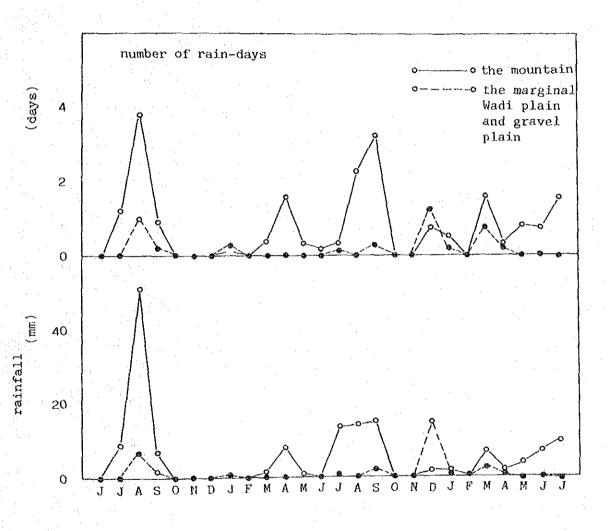
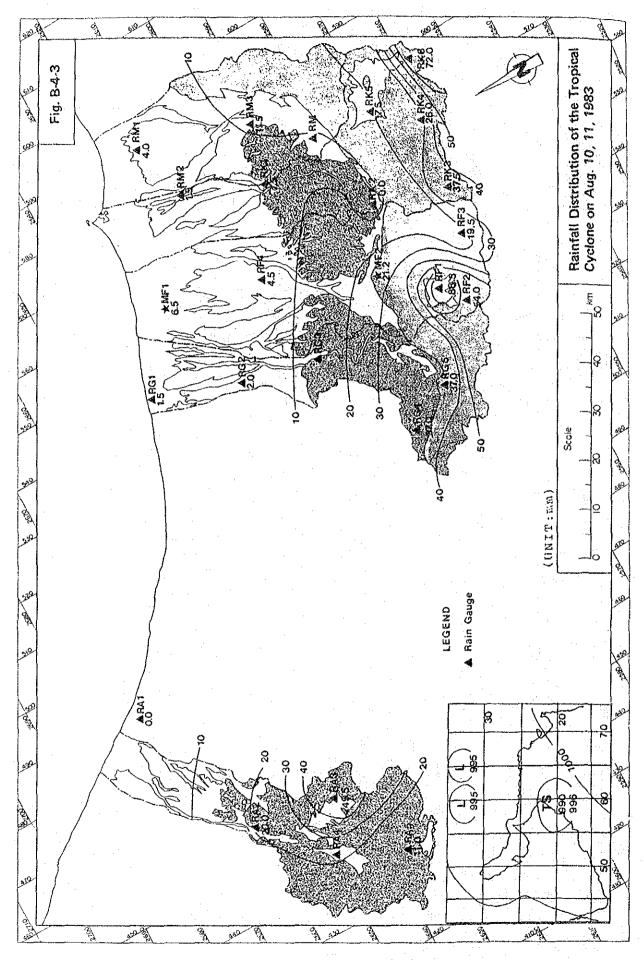
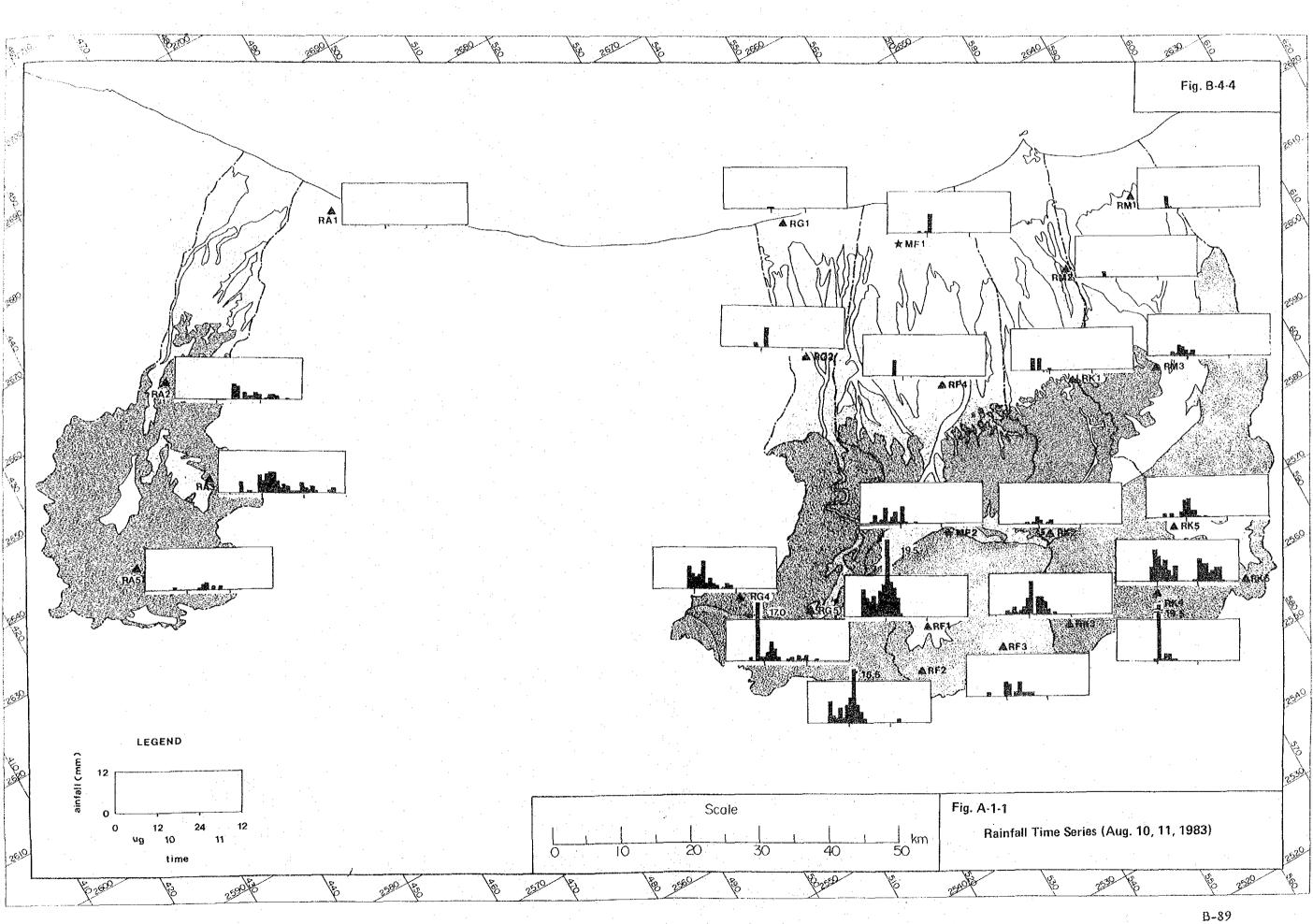
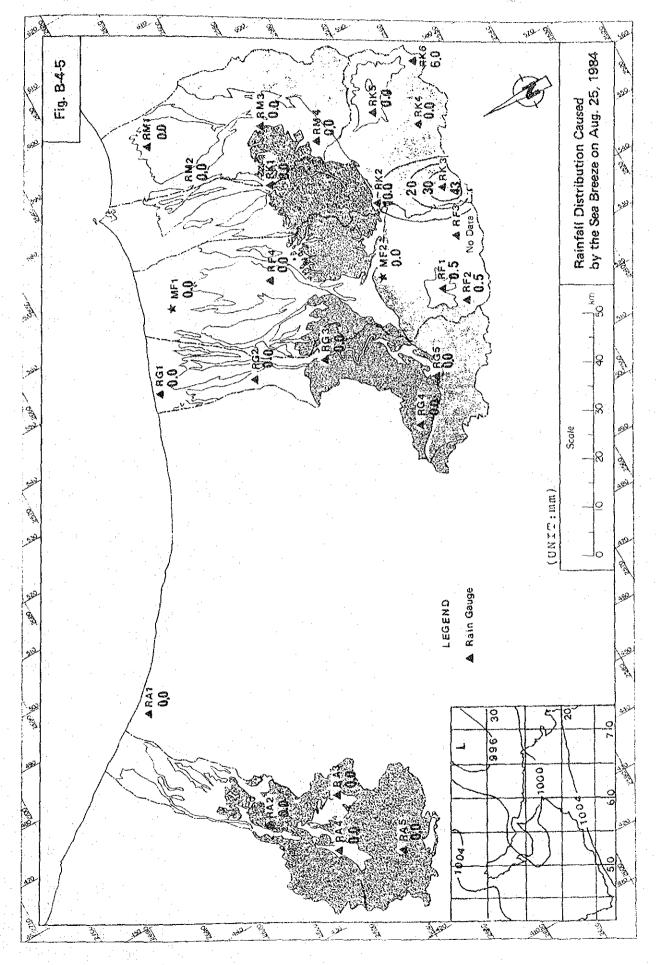


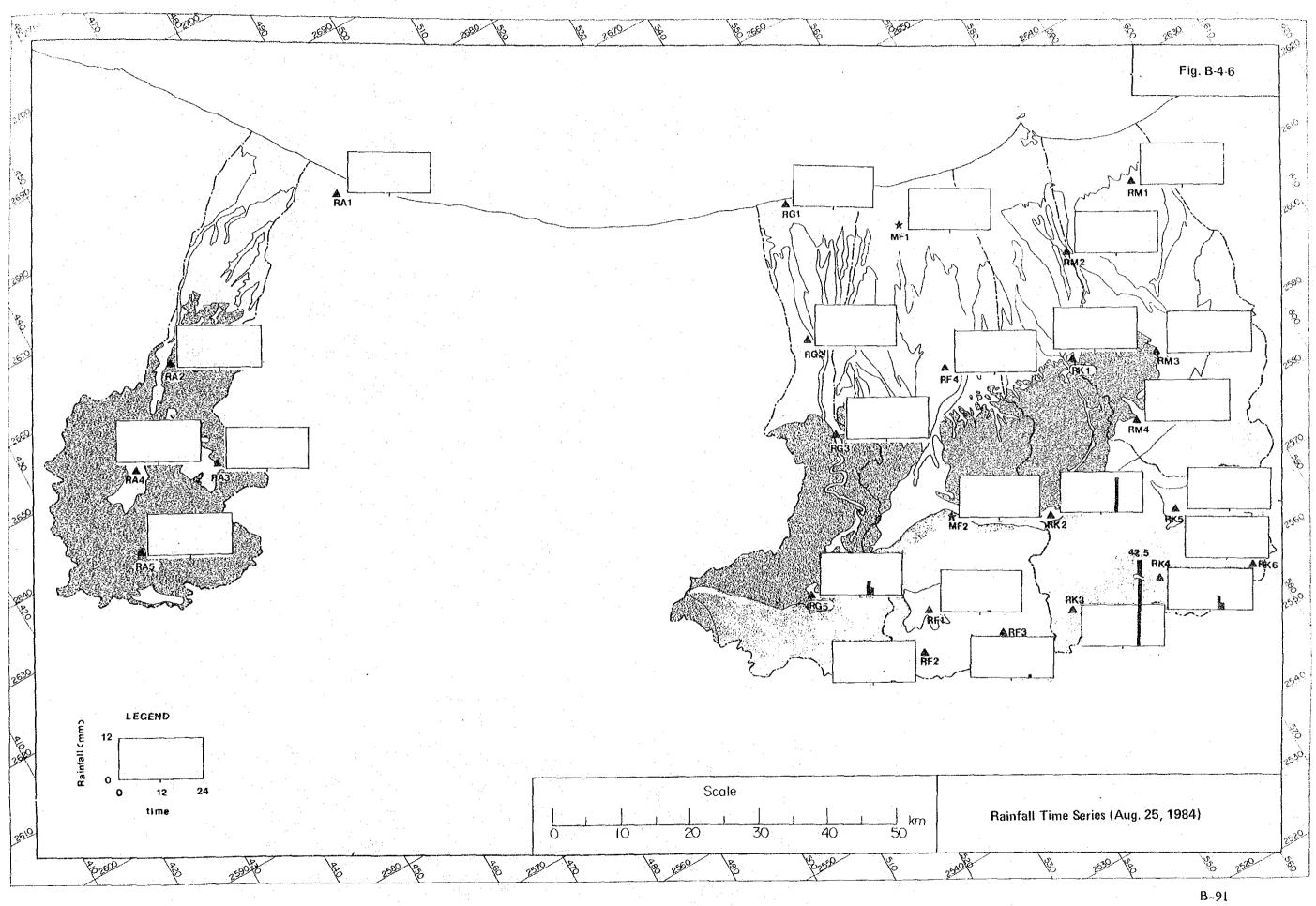
Fig. B-4-2 Comparison of Regional Rainfall in Mountain and in Marginal Wadi Plain and Gravel Plain



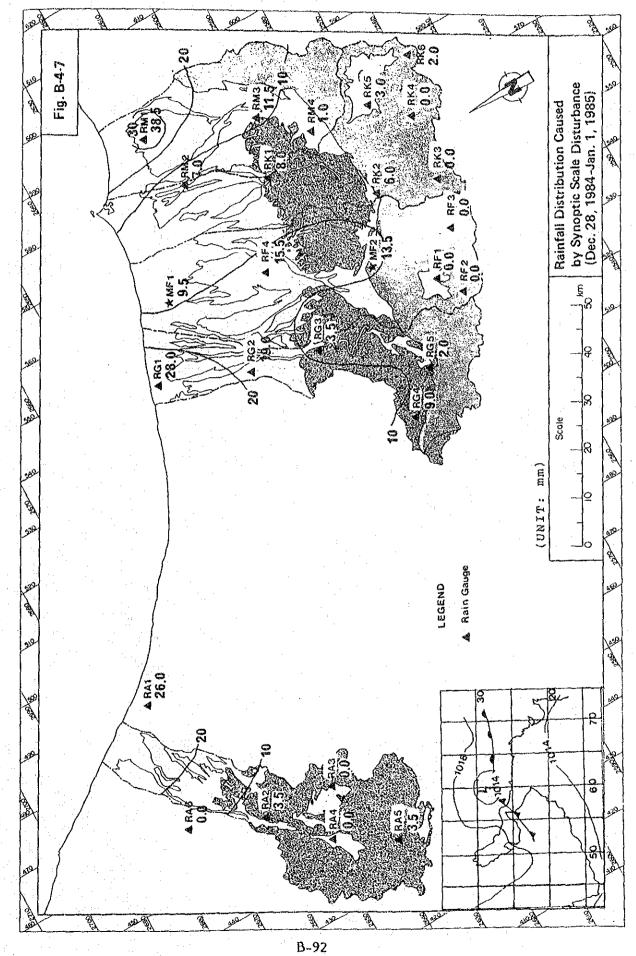


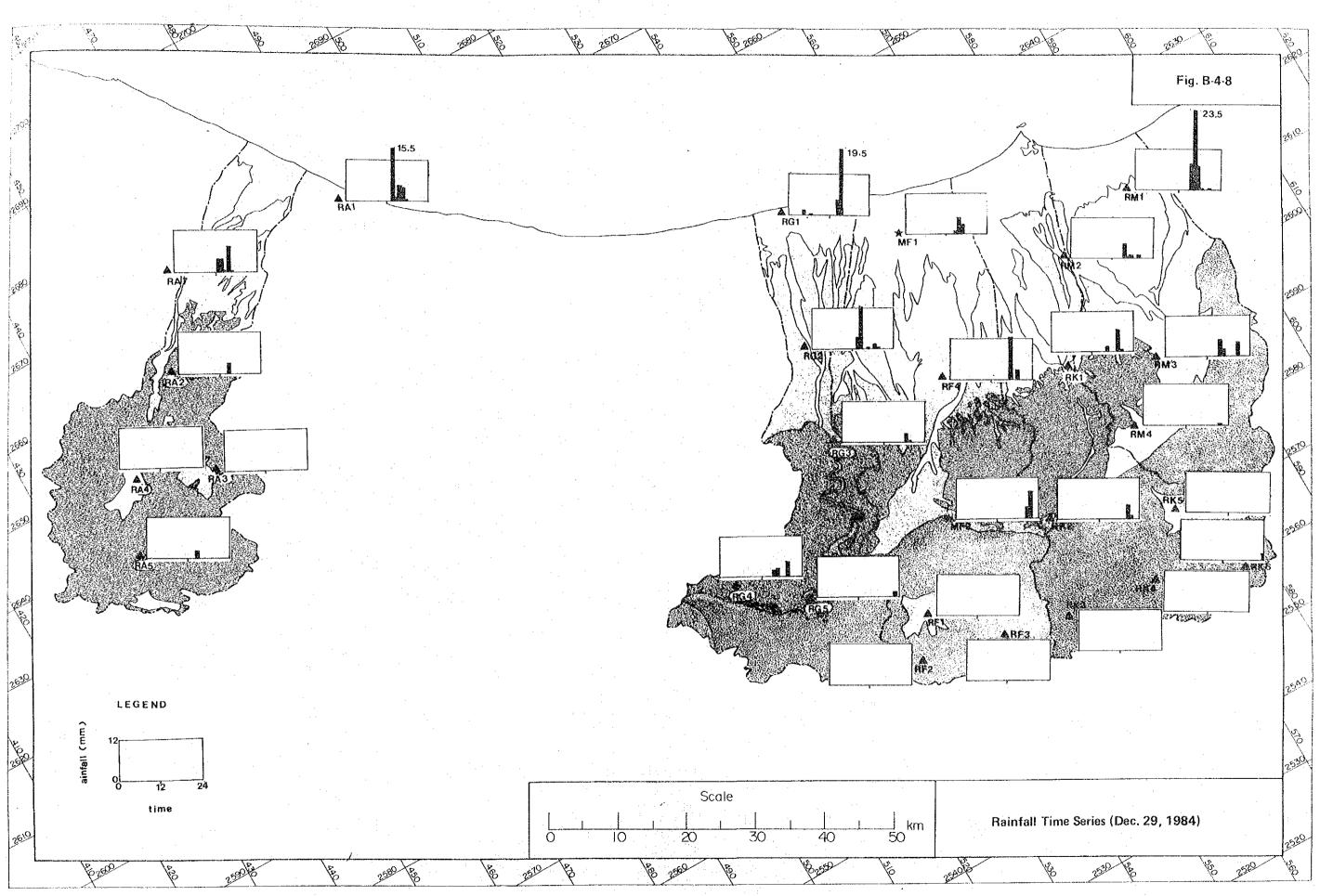






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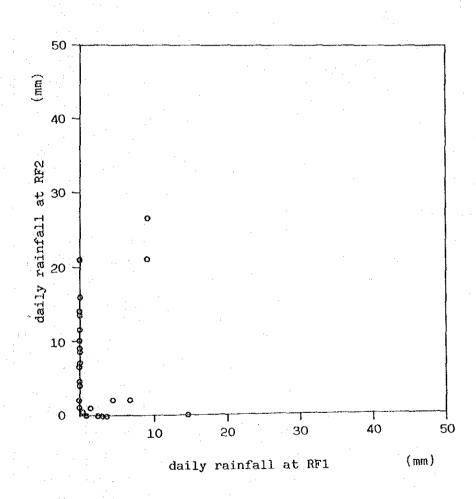




 $\frac{1}{2} \left(\frac{1}{2} + \frac{1}{2} \right) = \frac{1}{2} \left(\frac{1}{2} + \frac{1}{2} + \frac{1}{2} \right) \left(\frac{1}{2} + \frac{1}{2} + \frac{1}{2} + \frac{1}{2} \right)$







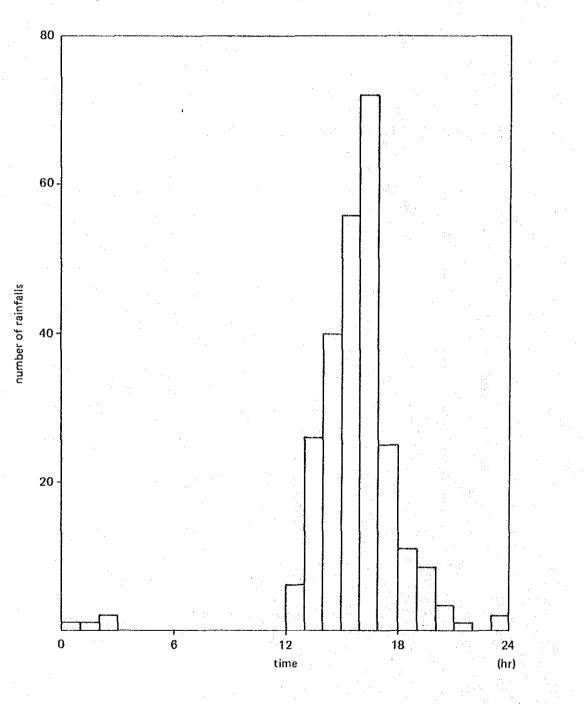


Fig. B-4-10 Occurrence Time Distribution of Summer Rainfall

4.1.2 Rainfalls Caused by Tropical Cyclone

The largest rainfall observed during the project, was brought by the tropical cyclone. However, according to D.E. Pedgley who summarized the record of the cyclones in the Arabian Sea (Cyclones along the Arabian Coast, Weather, London, Volume 24, 1969, pp.456-468), tropical cyclones that bring more than 10 mm rainfall on Mashirah Island or at Salalah come once in more than five years. As Mashirah Island and Salalah are on the Arabian Coast, they are considered under stronger influence of tropical cyclones than Muscat.

According to Renardet Sauti Ice (Water Resources Survey in North-East Oman, Annex A Climatology, March 1975) tropical cyclones bringing more than 10 mm rainfall at Muscat come once in more than 50 years. Renardet Sauti Ice also reported more than 300 mm rain came on June 5, 1890 from a tropical cyclone. This amount of rain is very rare.

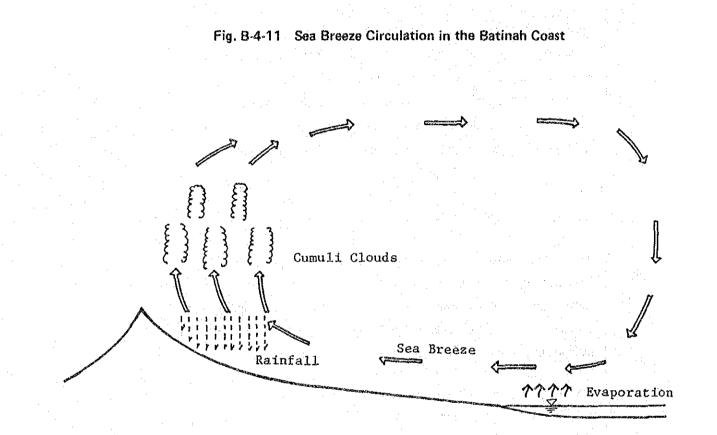
4.1.3 Rainfall Caused by Sea Breeze

The mechanism of rain formation by sea breeze is shown in Fig. B-4-11. The sea breeze transports humid air masses from Gulf of Oman to Hajar Mountain. The air mass is lifted by the sea breeze circulation joined with valley wind circulation. Then it is transformed to cumulus or cummulonimbus clouds where rainfall is formed.

The relation between elevation and rain-days and rainfalls caused by the sea breeze circulation is shown in Fig. B-4-12. All of the data was observed during June to September of 1983, 1984 and in June and July of 1985. As elevation increases, the rainfall and the rain-days increase. No rain was observed from the rain gauge below 150 m a.s.l.

The total number of the rainfalls during the period was 202. The frequency distribution of the rainfalls on normal probability paper is shown in Fig. B-4-13. On this figure, the X-axis is daily rainfall using a cube root scale, and the Y-axis is cumulative frequency. The cube root of rainfalls generally form a normal distribution. The distribution of the rainfalls caused by sea breeze circulation fits the normal curve well. The characteristics of rainfall caused by the sea breeze circulation are summarized below.

- The rainfall increases, as elevation increases.
- The rainfall probably does not occur at lower elevation.
- The rainfall occurs mainly in the afternoon.
- The rainfall continues for a short time and almost all of it stops in less than one hour.
- The rainfall area is narrow.



Hajar Mountains

Plain

Gulf of Oman

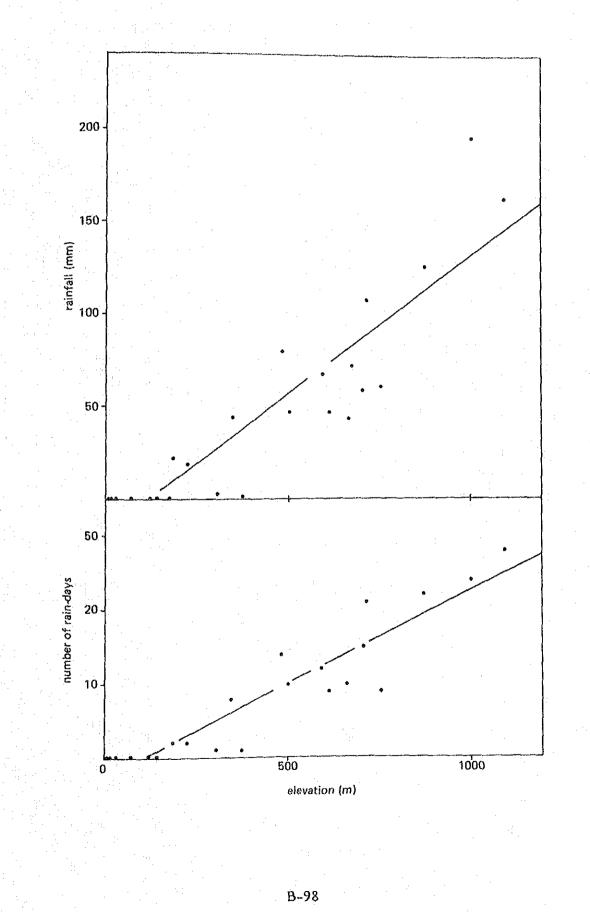


Fig. B-4-12 Relation between Elevation and Rainfall and Rain-days Caused by the Sea Breeze Circulation

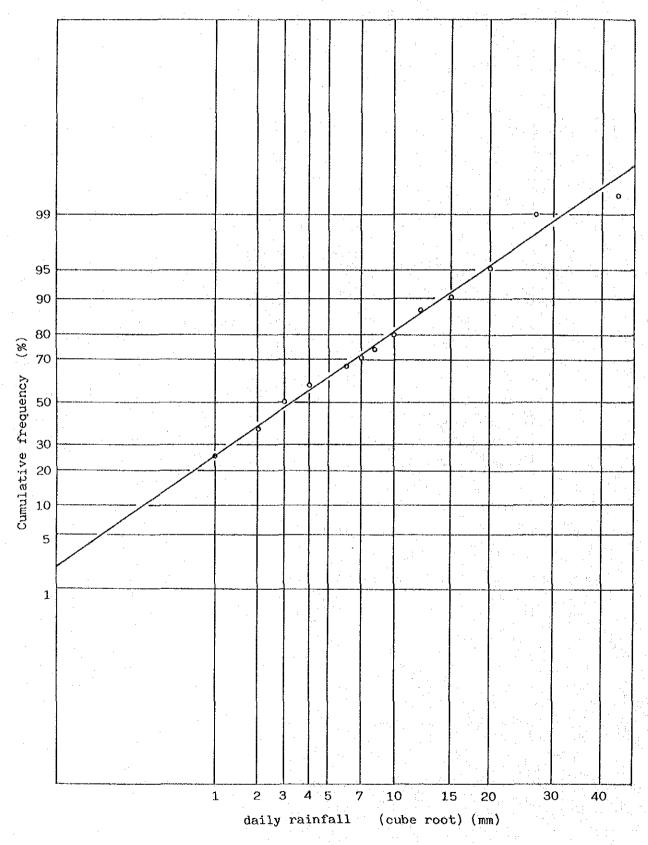


Fig. B-4-13 Cumulative Frequency of Summer Rainfall (Normal Curve)

4.1.4 Rainfall Caused by Synoptic Disturbance

Although the rainfall caused by synoptic disturbance occures mainly in winter, only a few rainfalls were observed in winter of 1983-84 and 1984-85. Therefore, the characteristics are summarized below from past data.

- The rainfall occurs mainly in winter.

- The rainfall occurs over a broad area.

4.2 Probability of Occurence

Muscat alone could supply rainfall data for a long period in and around the project area. The data is for seventy-three years intermitently from 1872. They are plotted on probability paper (Fig. B-4-14). They fit double exponential type distribution well. It shows that annual rainfall more than 100 mm occurrs once in two years; more than 200 mm once in ten years; and more than 300 mm once in a hundred years. The numbers on the figure indicate rainfall after the year 1977. In years where there was no data at Muscat, the rainfall at Darsite or Ruwi (MAF) neighbouring Muscat are indicated.

The annual rainfall is 2.5 mm at Ruwi in 1984 and 3.5 mm in 1985. Those are less than any other data observed at Muscat. The total rainfall in two years, after the observation network was installed by the project, is only 22 mm; consequently, the duration was very dry. Although the rainfall probability at Muscat can apply to the project area, even though the rainfall caused by sea breeze circulation may not occur, it is clear that during the observation period it was very dry.

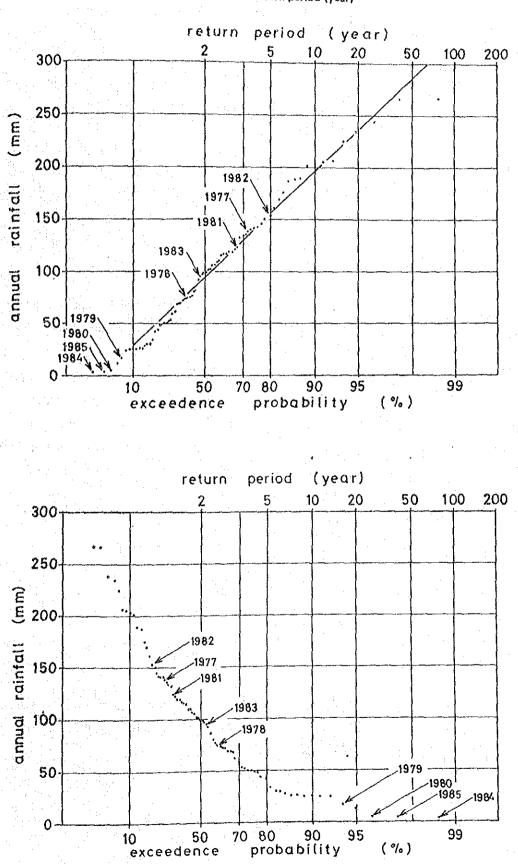


Fig. B-4-14 Annual Bainfall Frequency at Muscat return period (year)

4.3 Estimation of Rainfall over Each Basin

4.3.1 Outline of Estimation Method

The long duration rainfalls are estimated by the use of the rainfall data of MAF network. Examination of rainfalls lasting for a long period are necessary in order to evaluate the water resources, but this project could get data only during dry years.

The rainfalls should be estimated daily, because daily rainfall are necessary in the calculation process of water balance, which are calculated in chapter five. The period for the estimation is after June of 1976, because daily rainfall observation started on that day at all of the MAF network shown in Table B-1-3.

Interpolation or extrapolation of rainfall observed by MAF cannot apply to the estimation of winter rainfall. However, it can not apply to the estimation of summer rainfall, because summer rainfall covers a smaller rain area than winter rainfall. Accordingly, different methods are applied to rainfall of each season. Since there were no summer rains anywhere in the project area, it was assumed that sea breeze circulation brought all of summer rainfall during June to September. It was also assumed that synoptic disturbance brought all of winter rainfall during October to May.

After the estimation of rainfall at each rain gauge, rainfall over each basin was calculated by the use of the Thiessen method.

4.3.2 Rainfall in Summer

Rainfall caused by sea breeze is estimated by the probability model created in the project.

The dependence of the rain days on elevation is made clear in Fig. B-4-12. Together with the rainfall data observed at Al-Saiq (2000 m a.s.l), the relation between elevation (a) and rain-days in summer (n) is expressed as follows:

n = 0.02190a - 2.87.

Accordingly, the relation between rain-days at Al-Saiq in summer (ns) and the rain-days at a m a.s.l. (na) is expressed as follows:

na = (a-131).ns/(2000-131).

This expression can be used to estimate the rain-days at a m a.s.l. with the use of the rain-days at Al-Saiq (ns).

According to daily rainfall, the cube roots are distributed on a normal curve (Fig. B-4-13). The probability model which produces the rainfall along the normal curve was created and the daily rainfall at each rain gauge was estimated using the model. Annual summer rainfall, the sum of daily rainfall, is shown in Table B-4-3. The values after 1983 on the table are not estimated ones, but rather observed data.

					·····	. : 						
		1976	1977	1978	1979	1980	1981	1982	1983	1984	1985 (June-July)	Mean (1976-1984
RA	1	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0
	2	1.1	22.2	5.5	11.0	3.7	0.5	10.3	20.0	3.0	0.0	8.6
-	3	17.9	12.9	18.1	9.1	22.6	3.6	16.7	64.0	31.5	1.0	21.8
	4	44.9	13.4	62.2	32.1	19.9	26.4	11.2	8.9	13.0	0.0	25.8
	5	52.9	20.1	56.6	20.4	28.7	19.8	29.9	18.0	37.0	17.0	31.5
RG	1	0.0	0.0	0.0	0.0	0.0	0.0	0.0	1.5	0.0	0.0	0.2
	2	0.0	0.0	0.0	0.0	0.0	0.0	0.0	7.0	0.0	0.0	0.8
	3	8.9	7.4	0.5	1.2	10.2	0.0	0.5	16.5	2.5	0.0	5.3
	4	65.2	11.0	54.9	20.5	13.2	5.5	5.4	62.0	19.0	0.0	28.5
	5	46.0	13.9	64.7	26.0	32.8	16.9	25.3	48.5	19.0	27.0	32.6
Œ	1	27:4	32.1	49.2	7.7	21.4	1.5	12.8	101.5	34.0	9.5	32.0
	2	41.6	41.5	74.9	36.4	73.8	38.3	76.6	94.5	113.5	42.0	65.7
	3	45.3	11.7	25.2	12.4	30.6	19.9	8.6	60.0	20.5	11.5	26.1
	4	0.0	0.0	0.0	0.0	0.0	0.0	0.0	4.5	0.0	0.0	0.5
ſF	1	0.0	0.0	0.0	0.0	0.0	0.0	0.0	6.5	0.0	0.0	0.7
	2	42.1	3.8	16.4	15.4	17.7	0.5	3.0	58.6	5.0	1.5	18.1
K	1	13.3	0.0	9.4	0.0	0.0	0.0	0.0	9.5	20.5	0.0	5.9
	2	29,9	6.9	20.8	11.2	13.7	22.9	24.5	14.0	51.0	21.0	21.7
	3	57.0	36.9	28.0	22.0	62.8	11.9	14.7	51.5.	82.5	11.0	40.8
	4	48.9	71.9	75.6	20.3	47.5	59.1	11.0	62.5	48.0	42.0	49.4
	5	29.7	31.3	17.0	31.4	5.6	49.6	18.4	55.0	8.5	0.5	27.4
	6	38.9	100.8	69.7	56.8	29.7	21.9	16.4	123.5	59.5	52.0	57.5
M	1	0.0	0.0	0.0	0.0	0.0	0.0	0.0	4.0	0.0	0.0	0.4
	2	0.0	0.0	0.0	0.0	0.0	0,0	0.0	1.5	0.0	0.0	0.2
	3	0.0	0.0	4.4	0.0	0.0	0.0	0.0	11.5	0.0	0.0	1.8
		4	12.2	1.6	5.1	1.1	1.0	12.3	13.0	7.2	0.0	0.0
											$e^{-i\omega t}$ (1)	

Table B-4-3 Estimated Summer Rainfall

* The values before 1982 are estimated, and the values after 1983 are observed.

4.3.3 Rainfall in Winter

The rainfall in winter can be estimated directly from the daily data observed by MAF, because such rainfall covers a broad area.

The general method for estimating rainfall at non-observed points was used. As shown in Fig. B-4-15, non-observed points were plotted at origin and observed points A, B, C were plotted against the X-Y-axes. If the observed rainfall rA, rB, rC are plotted in Z-direction, one plane including rA, rB, rC is determined. The intersection of the plane and Z-axis r_0 is assumed to be the estimated rainfall. This method can be used, even if the intersection is outside the triangle rA rB rC.

This method can be expressed as follows:

	rA	rB	r _C	1	1	1
r _o =	XA	$\mathbf{x}_{\mathbf{B}}$	x _C	XA	ХB	XC
	YA	YB	r _C X _C Y _C	YA	YB	l X _C Y _C

Because it is clear that the rainfall in winter also depends on elevation (Fig. B-2-2), one of the coordinates chosen should be elevation and the other longitude. Latitude was taken into consideration when A,B,C were chosen from the rain gauges of MAF network.

In case A,B,C are located nearly on a line or fan from the origin, this method produces inadequate values. Since estimation of the rainfall in Wadi Ahin is this type of case, latitude is included in the expression as follows:

	rA	rB	rC	۲D	1	1 -	1	1	
r -	XA	X_{B}	Х _С	X _D	XA	ХB	Х _С	XD	
10 =	YA	YB	Y _C	YD	YA	ΥB	YC	YD	
r _o =	ZA	ZB	ZC	ZD	z _A	z _B	X _C Y _C Z _C	ZD	

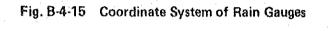
Annual winter rainfall and the sum of daily rainfall are shown in Table B-4-4. The values after October 1984 on the table are not estimated ones, but actually observed ones.

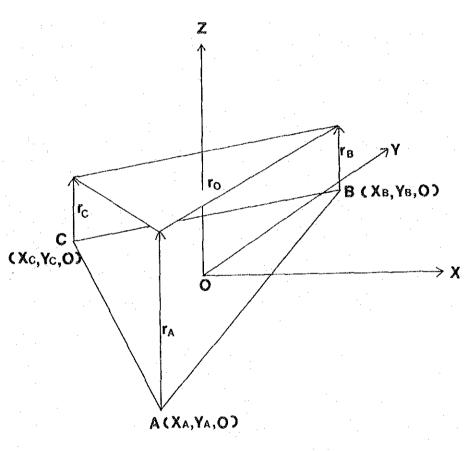
Table B-4-4 Estimated Winter Rainfall

Rai gau		1976 (OctDec.)	1977	1978	1979	1980	1981	1982	1983	1984	1985 (JanMay)	Mean
RA	1	15.2	120.2	59.9	116.8	1.2	49.4	117.1	76.9	26.0	0.5	64.8
	2	25.9	104.4	43.1	89.5	34.4	63.1	152.4	142.0	4.0	0.0	73.2
	3	47.2	146.3	71.3	71.4	28.3	75.8	236.0	102.5	4.5	10.5	88.2
	4	31.2	114.2	46.6	67.5	25.0	31.2	193.8	128.6	1.5	0.0	71.1
	5	58.9	172.7	75.6	110.0	37.1	63.9	281.2	163.1	3.5	2.5	107.6
RG	1	9.8	94.2	37.8	82.3	44	62.2	146.4	86.3	32.5	0.0	61.8
	2	41.2	168.7	30.4	107.3	64.1	68.5	230.4	102.9	18.5	2.0	92.7
	3	71.1	240.1	22.0	130.0	122.0	76.0	314.0	120.0	3.5	5.5	122.7
	4	55.2	179.6	73.6	90.3	68.1	95.6	345.4	110.2	21.0	5.0	116.0
	5	56.8	182.4	72.4	94.0	72.5	96.7	357.7	103.1	3.5	3.0	115.8
RF	1	58.5	191.3	73.6	98.1	82.7	102.2	385.7	106.3	24.5	8.5	125.7
	2	53.7	203.7	85.4	87.9	89.1	113.8	399.7	145.0	7.5	31.0	135.2
	3	60.5	193.8	71.9	102.6	87.2	103.0	398.6	100.2	30.5	27.5	130.6
	4	30.8	122.3	42.2	87.8	33.9	74.4	241.2	76.8	15.5	0.0	80.5
MF	1	13.6	94.3	33.9	75.8	11.8	66.8	172.8	85,8	9.0	1.5	62.8
MF	2	65.2	178.8	59.1	112.5	78.2	89.5	377.6	59.0	14.1	3.0	115.2
RK	1	8.6	198.4	93.2	94.5	17.2	101.2	289.8	118.8	8.0	4.5	103.8
	2	47.2	188.9	72.4	99.2			364.9	1 - E		4.5	117.6
	3	55.8	199.2	75.5	100.7	and the second	24 - 24 - C	403.3	10 T (10 - 10 - 10 - 10 - 10 - 10 - 10 - 10	1997 - 1997	and the second second	130.4
	4	37.1	210.5	89.8	86.7			391.7	for a		17.5	132.8
	5	25.1	199.3	87.4	83.6			344.6	÷	5.0	4.0	117.0
	6	24.7	233.5	106.4	79.6	1.1	1.1	402.8		이 사람이	30.0	145.4
RM	1	6.2	82.5	25.3	55,3	8.7	1 I.I.	180.8	1.815		1.5	63.2
	2	7.3	115.9	46.3	69.2	10,1	1	200.1	99.0		0.0	70.4
	-3	6.5	180.0	82.5	79.5		:	267.7		1 A 4	5.0	95.7
	4	21.1	187.7	81.7	85.1			312.3	1	1.0	24.5	108.1

* The values before May 1983 are estimated, and the values after Oct. 1983 are observed.

(mm/season)





4.3.4 Calculation of Rainfall over Each Basin.

Seasonal and annual mean of estimated rainfalls at each rain gauge are shown in Table B-4-5. The rainfalls for each basin were calculated by the use of Thiessen method and the estimated rainfalls at each rain gauge. The results, mean rainfalls over each basin, are shown in Table B-4-6. Comparing the rainfall on each basin, rainfall on Wadi Bani Kharus is the greatest with Wadi Bani Ghafir, Wadi Al-Fara, Wadi Ahin and Wadi Al-Ma'awil following Wadi Bani Kharus in turn.

and a second	20 and 40 and 10 and	(197	6 ∿1985)
Rain Gauge Station	Winter	Summer	Annua 1
RA 1	64.8	0.0	64.8
2	73.2	8.6	81.8
3	88.2	21.8	110.0
4	71.1	25.8	96.9
5	107.6	31.5	139.1
RG 1	61.8	0.2	61.8
2	92.7	0.8	93.5
3	122.7	5.3	128.0
4	116.0	28.5	144.5
5	115.8	32.6	148,4
RF I	125.7	32.0	157.7
2	135.2	65.7	200.9
3	130.6	26.1	156.7
4	80.5	0.5	81.0
MF 1	62.8	0.7	63.5
2	115.2	18,1	133.3
RK 1	103.8	5.9	109.7
2	117.6	21.7	139.3
. 3 .	130.4	40.8	171.2
4	132.8	49.4	182.2
5	117.0	27.4	144.4
6	145.4	57.5	202.9
RM 1	63.2	0.4	63.6
2	70.4	0.2	70.6
3	95.7	1.8	97.5
4	108.1	5.9	114.0

Table B-4-5 Annual Mean Rainfall

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έ.

Table B-4-6 Mean Rainfall of Wadi Basin (Seasonal & Annual)

UNIT : NM

	л Вајск	7.0 150.5 157.5	66.0 70.2	2.5 74.9 77.4	20.2 20.8 20.8	827.5 87.6 9.7	3.4 254.0 257.4	11.5 122.0 122.0 2.1 2.1	11
	LOWER	4.1 127.3 131.5	84 N 86 N 86 N	0.0 89.1 89.1	0.0 12.0 12.0	85.4 85.4	0.0 221.2 221.2	· · · ·	11 20 r- VI
	W.AL UPPER	10.8 181.8 192.6	7.0 83.1 90.0	5.78 83.78 83.73 83.73	1.4 31.3 32.6	13.1 102.8 115.9	в. 0 298. 0 305. 0	1111 1319 1319 1319 1319 1319 1319 1319	1113. 121.
	-¥	29. 2¥ 179. 0¥ 208. 2¥	24. 4 * 77. 2 * 101. 6 * * *	101.0**		20.9* 103.1* 123.9*	10.0** 322.2** 333.1*	мим мим ж ж ж ж ж ж ж ж ж ж ж ж ж ж ж ж ж ж ж	141.1*
, ,	. KHARU LOWER	140.u 140.u 140.u	61.2 64.2 7	0.0 79. W	0 m m 0 m m	ວິດ ເດີຍ ເດີຍ ເດີຍ ເດີຍ ເດີຍ ເດີຍ ເດີຍ ເດີ	0.0 230.6 230.6	01 200 00 00 00 00 00 00 00 00 00	ពីសាល
	UPPER UPPER	41.6 197.6 239.2	10 0 10 0 10 0 10 0 10 0	22.5 90.1 112.6	25. 0 65. 4 90. 4	30.8 111.5 142.3	15.0 365.1 382.1	700 TN7 000 TN7 700	mmu
	 BAISN	10.4¥ 157.1¥ 167.4¥	• ສຸສຸສຸສຸສຸສຸສຸສຸສຸສຸສຸສຸສຸສຸສຸສຸສຸສຸສຸ	100-7-24 100-44 103-44	1000 1000 1000 1000 1000 1000 1000 100	ນ 4 00 ສຸສຸສຸສຸສຸສຸສຸສຸສຸສຸສຸສຸສຸສຸສຸສຸສຸສຸສ	* * * * *	* * * * * * * *	14.0* 14.0* 121.0*
	AL-FARA LOWER	2.3 119.3 121.6	0.0 37.2 37.2	88.4 0.0 86.4	31. 1 31. 4	0.0 70.7 70.7	215.4 215.4 215.4	0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	8 8 3 7 7 8 8 7 7 8 8 7 7 8 8 7 7 8 8 7 7 8 8 7 8 8 8 7 8
	UPPER	16.8 187.5 204.4	28.0 64.0 91.9	13.0 104.5 117.5	24.7 82.3 106.9	10.5 96.1 106.7	15.0 369.7 384.7	8000 900 900 900 900 900 900 900 900 900	127.0 151.3
	¥20109	11.0* 177.0* 188.5*	* * * * * ການ ຍາຍ • • • • •	***** *****	* * * * * * * * * * * * * * * * * * *	0.00 	12.0 8 9 10.0 8 10.0 8 10.0 8 9 10.0 8 9 10.0 10.0 10.0 10.0 10.0 10.0 10.0 10	NUC 004	10 M N
	GHAF IR LOWER	136.6 136.6	33.1 33.1 33.1	ဝ ထ စ ဂ ဂ ဂ ဂ ဂ ဂ	38.9 38.9 39.9	6 6 6 6 6 7 6 7 6 6 7 6 6 7 6 7 6 7 6 7	0.0 195.4 195.5	880 844 770 1000 770	្រ ហ ហ រា ហ ហ រា យ យ
	UPPER	16.4 194.5 210.9	47.3 61.7 109.0	19. 2 101. 2 120. 4	25. 3 84. 2 109. 5	11. 2 93. 1 104. 3	11 140.0 000.0 000.0	2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2	125. 151.
	* 	10.0¥ 100.0¥ 107.7¥	* * * * * * O 0 0 N 0 0 N 0 0	16.0* 01.0* 107.0*	547.04 54.04 54.14 54.14 55.0	00.7 × ×	15.0* 15.0* 201.3* 216.4*		**************************************
		11.6 111.8 123.4	2.9 54.0 54.0	5.7 102.4 108.2	1.9 18.5 20.4	ចលក ជួល លក	133.4 135.4	1121. 38 21. 38	==== 5.0 80.1 80.1
	UPPER LOWER	16.8 142.1 155.0	44.0 62.4 106.4	20.8 87.1 107.8	22.5 31.4 53.5		19.5 231.8 251.3	1 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2	24. 4 24. 4 118. 4
	SEAGON	SUMMER WINTER MNNUML	SUMMER WINTER ANNUAL	SUMMER WINTER ANNUAL	SUNNER WINTER RINUAL	SUMMER WINTER ANNUAL	SUPPIER WINTER NNNUAL	SUMMER WINTER ANNUAL SUMMER WINTER ANNUAL	SUPPER UINTER PINUAL
	YEAR	17	78	62	80	ដ	83	0 0 N 4	1f 11 14 19 19 11
				1				·····	11 - C - C - C - C - C - C - C - C - C -

CHAPTER 5 ANALYSIS OF INFILTRATION CAPACITY AND EVAPOTRANSPIRATION

5.1 Experiment on Infiltration Capacity

5.1.1 Experimental Sites

A series of experiments for infiltration capacity of surface soil were carried out at five sites within the project area. The sites and their surface soil characteristics are as follows:

- 1) Sand———— Al-Muladdah

- 4) Semi-consolidated sand and gravel ----- North of Jamma
- 5) Semi-consolidated sand and gravel----- West of Al-Rustag

No rainfall was recorded at any of the sites six monthes before the experiments.

5.1.2 Method and procedure of experiments

The method and procedure of the experiments are as follows;

- 1) Set up a block of area $(1 \times 1 \text{ m}^2)$ at flat ground without vegetation, and surround it with a small mound of soil and gravel.
- 2) Sprinkling water uniformly on the ground by watering device at a constant sprinkling intensity (mm/h) determined beforehand.
- 3) Keep sprinkling water and watching whether all the water infiltrates or not into soil without any remaining on the ground surface.
- 4) Check time and cumulative water amount when 30%, 50% and 70% of the block (1 m^2) are covered with water remaining on the surface.
- 5) Take soil samples at different depths with soil sampler, right after the end of water sprinkling to determine the water content of the soil.
- 6) Repeat taking soil samples after a few days interval.
- 7) Measure the water content of all the soil samples.

The procedure above was carried out for all sites. In addition, the following measurement was performed for some of the sites.

- 8) Observe the features of lowering wet front during water pouring.
- 9) Measure the grain size of soil samples.
- 10) Measure the density of dehydrated soil samples.

The dehydration of major soil samples and weight measurement of dehydrated samples were carried out at Rumeis Laboratory, MAF.

Some samples were brought to Japan and measurement was done there.

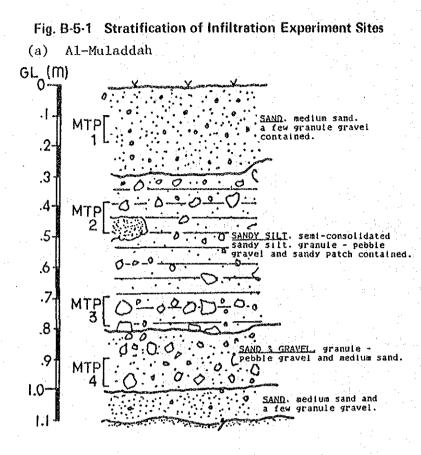
5.1.3 Experiment results on infiltration capacity

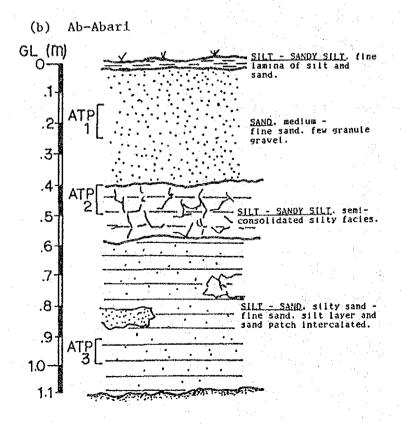
Experiment conditions and results are summarized in Table B-5-1. The soil stratification of the two sites is shown in Fig. B-5-1.

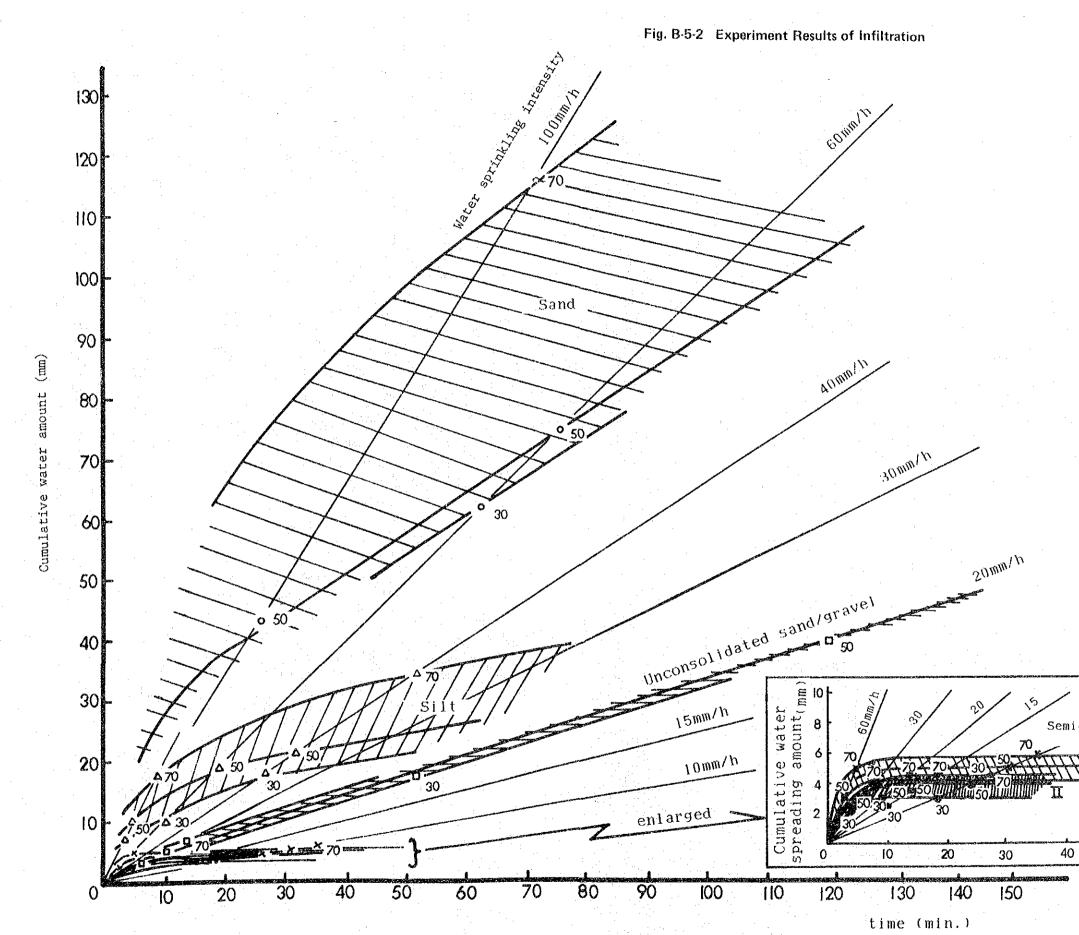
The relationships between cumulative water amount and infiltration capacity, are shown in Fig. B-5-2. A great difference in capacity is revealed between surface soil materials. Sand has the highest capacity. Semi-consolidated sand and gravel has the lowest capacity.

Regarding the depth of wet front in case of 120 mm/h on sand, its change with time as shown in Table B-5-2 and Fig. B-5-3 is uniform. The descending speed of the front is speed almost constant, 9.2 mm/min., as shown in Fig. B-5-4. The average volume water content of the transmission zone was 22%. Table B-5-1 Experiment Results of Infiltration

						· · ·									
Location of experiment site	te	Al-Muladdah	addah	Ab-	Ab-Abari		South	of	Tarif	North o	of Jamma	West	of	Al-Rustaq	staq
Surface soil		Sand		S S	Silt		Un-co sand	Un-consolidated sand and gravel	lated avel	Semi-cor sand a	Semi-consolidated sand and gravel	Sen	Semi-consolidated sand and gravel	isolid nd gr	ated avel
Date of exper	experiment	13/7/ '85	15/7/ '85	15/7	15/7/*85		18/	18/7/ 185		17/7/*85	7 *85		21/	21/7/ 985	
Experiment site No.	te	1 - M	M-2	A-1	A-2	A-3	T - 1	T-2 T	T-3	<u>й-1</u>	J2	ы-1 К-1	R-2	氏 で、 の	R-4
Water sprinkling intensity (mm/h)	ling n/h)	60	120	40	60	120	20	30	60	10	60	01	15	20	30
<u>р.</u>		62.5	22.5	26	10	œ	21	S	1	25	2.5	18	10	ത	4
water illm	cumulative water amount (mm)	62.5	11.0	17.5	10	16	17	ო	,	4.2	2.5	3.5	2.5	3.0	2.0
50% surface covered with	time (min)	75.0	45	31	- 19	10	120	10	:	30	£	24	12	12	cc
water film	cumulative water amount (mm)	75.0	22.5	21	19	20	40	ິນ		5.0	I	4.0	3.5	4.0	3.0
70% surface covered with	time (min)	1	117.5	50	L	18	3	S 1	ى. ت	35	a	27	18	13.5	ŝ
water film	cumulative water amount (mm)	1	59.0	34	ł	36	I	6.5	ы	5.8 8	5.0	4. U	4.5	44 10	4.0
Cumulative wate at the end of sprinkling	water amount of water	92.5	125	54	25.5	42	40	51 1	6.5	6.7	7.5	4.5 5	5.0	2.0	4.5
Depth of wet front at the water sprinkl	of wet at the end of sprinkling(cm)	32	1	28	Ø	11	1	0	0	cu	G	1.3		1.0	1.0







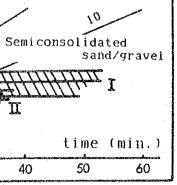


Table B-5-2 Depth of Wet Front in Sand in Case of Water Sprinkling Intensity 120mm/h

Al-Muladdah (M-2)

Time after starting water sprinkling Cumulative sprinkling water amount Α:

B :

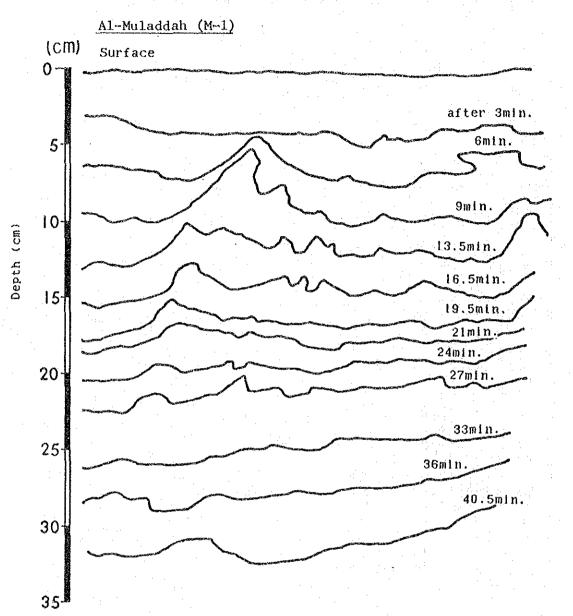
Ċ: Depth of wet front

· . · ·			
A (min)	B (mm)	C (cm)	
$\begin{array}{c} 0\\ 1.5\\ 2.8\\ 4.0\\ 5.3\\ 6.2\\ 7.4\\ 8.8\\ 10.0\\ 11.1\\ 12.5\\ 13.7\\ 14.7\\ 16.0\\ 17.4\\ 18.5\\ 20.0\\ 21.1\\ 25.4\\ 27.5\\ 29.0\\ 30.0\\ 31.2\\ 32.5 \end{array}$	$\begin{array}{c} 0\\ 3.0\\ 5.6\\ 8.0\\ 10.6\\ 12.4\\ 14.8\\ 17.6\\ 20.0\\ 22.2\\ 25.0\\ 27.4\\ 29.4\\ 32.0\\ 34.8\\ 37.0\\ 40.0\\ 42.2\\ 50.8\\ 55.0\\ 58.0\\ 60.0\\ 62.4\\ 65.0\\ \end{array}$	$\begin{array}{c} 0\\ 1\\ 3\\ 5\\ 8\\ 9\\ 9.5\\ 10\\ 11\\ 12.5\\ 14\\ 16\\ 16.5\\ 17\\ 18\\ 18.5\\ 19\\ 20.5\\ 23\\ 24\\ 25\\ 26\\ -28, 5\end{array}$	

A (min)	B (mm)	C (cm)
$\begin{array}{c} 33.8\\ 35.0\\ 36.2\\ 37.5\\ 39.0\\ 4.04\\ 42.0\\ 43.0\\ 44.4\\ 45.5\\ 46.8\\ 48.2\\ 49.5\\ 50.0\\ 51.1\\ 52.2\\ 54.5\\ 55.6\\ 57.0\\ 58.0\\ 59.5\\ 60.0\\ 63.0\\ 63.0\\ \end{array}$	67.6 70.0 72.4 75.0 78.0 80.8 84.0 86.0 88.8 91.0 93.6 96.4 99.0 100.0 102.2 104.4 109.0 111.2 114.0 116.0 120.0 124.0 126.0	29.5 32 32.5 35 38 - - - - - - - - - - - - - - - - - -

۰.

Fig. B-5-3 Depth and Shape of Wet Front in Sand in Case of Water Sprinkling Intensity, 120mm/h



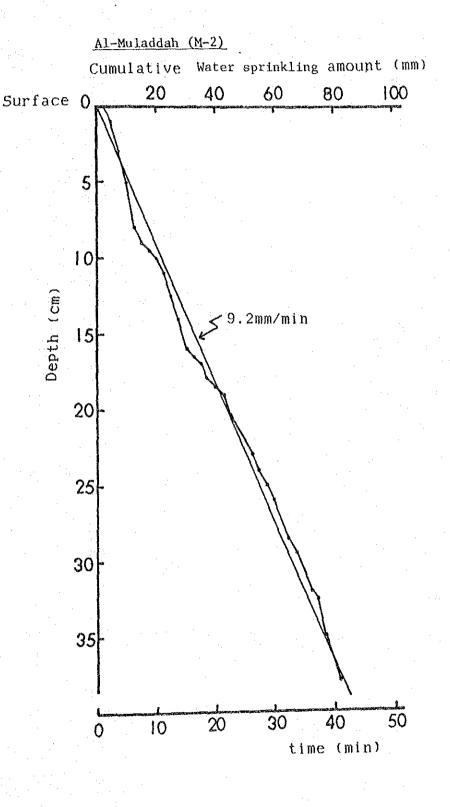


Fig. B-5-4 Descending Speed of Wet Front in Sand in Case of Water Sprinkling Intensity, 120mm/h

8-119

5.1.4 Experiment Results on Soll Water Content

All the measured results of the experiment are summarized in Table B-5-3.

The sand and silt which were taken with the soil sampler display a similar tendency in the change of water content with time, as shown in Fig. B-5-5 and Fig. B-5-6.

Regarding the other sites, quick evaporation within 24 hours was confirmed.

Table B-5-3(1) Soil Water Contents after Water Sprinkling (1/4)

Site: Experimental site No. (cf. tableB-5-1) Date: Sampling date Time: Sampling time V: Volume of soil sample V: Volume of soil sample W⁵: Weight of dry soil sample W⁵: Weight of soil water w: Water content (W_VW_)x100 6: Volume water content⁵ (W_VV_)x108 P: Density of dry soil sample⁵ (W_SV_5)

			and the second						
Site	Date	Time	Depth	V _s	Ws	Ww	w	8	ያ
			CM	cm ³	g	a,	%	%	g/cm ³
M-1	Jul.13	10pm	0~5	ł	118.3	21.1	17.8	-	-
			1~6	100	138.9	35.0	25.2	35.0	1.39
			23~28	100	156.0	26.5	17.0	26.5	1.56
			32~		H		dry	dry	
	Jul.14	5am	2~7	100	151.8	14.5	9,6	14.5	1.52
			2~7	100	153.1	15.4	10.0	15.4	1.53
			11~16	100	153.2	16.0	10.4	16.0	1.53
			19~24	100	154.2	20.7	13.4	20.7	1.54
	Ju1.15	5am	2~7	100	143.2	10.5	7.3	10.5	1.43
			8~13	100	153.8	13.0	8.4	13.0	1.54
			10~15	100	160.1	13.6	8.5	13.6	1.60
			17~22	100	149 0	12.5	8.5	12.5	1.49
	an an ann an Anna Ann an Anna Anna Anna Anna		22~27	100	152.0	13.1	8.6	13.1	1.52
	Jul.16	6am	0~5	100	145.5	6.0	4.2	6.0	1.46
n ang tao sa Tao sa sa			7~12	100	156.2	11.8	7.6	11.8	1.56
			13.5~18.5	100	158.9	13.5	8.5	13.5	1.59
			21.5~26.5	100	153.8	12.0	7.8	12.0	1.54
			29~34	100	158.3	12.7	8.0	12.7	1.58
			35~40	100	147.2	10.6	7.2	10.6	1.47
			43~48	100	134.9	12.0	8.9	.12.0	1.35
· ' '					ļ	.L	<u> </u>		

Site	Date	Time	Depth	V _s	Ws	Ww	w	9	p
			CM	cm ³	g	g	%	%	a\cm3
M-1	Jul.17	6am	0~1		145.3	1.0	0.7	_	
			1~6	100	144.8	6.7	4.6	6.7	1.45
			8.5~13.5	100	141.0	9.9	7.0	9.9	1.41
			15~20	100	155.5	10.4	7.4	10.4	1.56
	Jul.18	6am	0~5	100	134.3	3.8	2.8	3.8	1.34
			6~11	100	146.1	8.7	6.0	8.7	1.46
			13.5~18.5	100	150.4	9.3	6.2	9.3	1.50
			22~27	100	150.2	9.1	6.0	9.1	1.50
			28.5~33.5	100	146.3	8.4	5.7	8.4	1.46
			35.5~40.5	100	139.6	8.3	6.0	8.3	1.40
			42.5~48.5	100	129.4	9.6	7.4	9.6	1.29
			49~54	100	116.3	11.1	9.6	11.1	1.16
			55~58		62.2	4.8	7.7		-
	Jul.21	8am	0~5	100	146.9	4.0	2.7	4.0	1.47
	-		5.5~10.5	100	146.0	7.9	5.4	7.9	1.46
	-		11~16	100	164.3	9.9	6.0	9.9	1.64
			23~28	100	140.5	7.7	5.5	7.7	1.41
			30~35	100	147.9	7.5	5.1	7.5	1.48
			37~42	100	131.5	12.9	9.8	12.9	1.32
			44~49	100	134.6	9.1	6.7	9.1	1.35
	Jul.24	6pm	0~5	100	141.4	1.2	0.8	1.2	1.41
			8~13	100	154.4	7.2	4.7	7.2	1.54
			13.5~18.5	100	162.7	8.3	5.1	8.3	1.63
			19~24	100	153.4	7.6	5.1	7.6	1.53
			30.5~35.5	100	157.3	6.5	4.1	6.5	1.57
			37.5~42.5	100	146.0	7.1	4.8	7.1	1.46

Table B-5-3(2) Soil Water Contents after Water Sprinkling (2/4)

Table B-5-3 (3) Soil Water Contents after Water Sprinkling (3/4)

		Í.m.			}	I		· .	
Site	Date	Time	Depth	V _s	Ws	Ww	w	θ	ß
			Cm	cm ³	g	à	~ %	%	g∕cm3
M-1	Aug.1	10am	7~12	100	150.9	5.6	3.7	5.6	1.51
			13~18	100	152.8	6.3	4.2	6,3	1.53
			19~24	100	144.4	6.4	4.4	6.4	1.44
			25~30	100	150.6	5.9	3.9	5.9	1.51
			30~34	-	128.6	4.3	3.4		
			36~40	-	156.1	8.8	5.6		
	Aug.11	10am	(0~10)	e	(201.3)	(1.34)	(0.7)	-	
			10~18		182.8	6.3	3.4	-	
			40~48	-	157.9	15.0	9.5		
			70~78		190.7	8.8	4.6	-	-1
			90~98	~	171.9	11.6	6.7		
A-1	Jul.20	9pm	1.5~6.5	100	153.8	25.7	16.7	25.7	1.54
			7.5~12.5	100	152.4	21.4	14.0	21.4	1.52
			14~19	100	153.1	21.9	14.3	21.9	1.53
			21~26	100	153.6	18.5	12.0	18.5	1.54
			28~			••	dry	dry	_
A-2	Jul.20	6pm	1~6	100	141.2	22.2	15.8	22.2	
			8~				dry	dry	-
	Ju1.22	1 pm	0~5	100	127.9	3.8	3.0	3.8	1.28
		-	6~11	100	146.0	7.7	5.2	7.7	1.46
		-	15.5~20.5	100	142.2	7.5	5.2	7.5	1.42
			25~				dry	dry	
	Ju1.24	2pm	0~5.5	-	152.9	3.2	2.1	•-1	-
			7~12	100	148.4	5.0	3.4	5.0	1.48
			15~20	100	148.2	8.2	5.5	8.2	1.48
			22~27	100	146.6	8.4	5.7	8.4	1.47
			30~	_	 	_	dry	dry	
						<u> </u>			<u> </u>

Table B-5-3 (4) Soil Water Contents after Water Sprinkling (4/4)

· · ·		· · ·			1999 - 1999 - 1999 - 1999 - 1999 - 1999 - 1999 - 1999 - 1999 - 1999 - 1999 - 1999 - 1999 - 1999 - 1999 - 1999 -				
Site	Date	Time	Depth	۷ _s	₩ _s	Ww	w.	θ	P
			Cm	cm ³	g	a	%	%	g/cm ³
A-2	Jul.29	2pm	0~4		151.4	1.7	1.1		***
			4.5~9.5	100	148.3	4.1	2.7	4.1	1.48
. '			11.5~16		139.4	5.2	3.8		-
· .			17~22	100	149.7	5.2	3.5	5.2	1.50
·			23~28	100	126.5	7.1	5.6	7.1	1.27
			29~30	-	47.6	2.0	4.1		
	Aug.1	7am	0~3		187.0	1.8	1.0	<u></u>	-
-			3~6	-	112.8	1.8	1.6		
۰.			8~12	-	102.2	3.0	3.0		
		:	14~19	100	141.0	4.4	3.1	4.4	1.41
			20~25	100	139.5	5.8	4.1	5.8	1.40
·			26~30		137.3	8.6	6.3	-	-
	Aug.11	llam	(0~8)	-	(186.1)	(1.9)	(1.0)	-	-
			15~23	-	162.6	4.6	2.8		
			40~48	-	151.2	9.6	6.4		
			90~98		141.8	11.6	8.2		*3
A-3	Jul.20	6pm	0.5~5.5	100	136,5	23.4	17.2	23.4	1.37
			11~	-	-		dry	dry	
T-3	Jul.18	9pm	0~2		1429.8	41.4	2.9	-	**
			10~	-		-	dry	dry	
J-1	Jul.17	7pm	0~3		318.4	16.7	5.2	<u>-</u>	
			0~5	-	773.6	54.1	7.0	-	_
			5~	~		-	dry	dry	
	Jul.20	2pm	0~	-	-		dry	dry	-
J-2	Jul.17	7pm	2~5		684.1	56.4	8.2		
			6~				dry	dry	
	Jul.20	2pm	0~	-			dry	dry	-

(): sample taken outside the water spreading sites

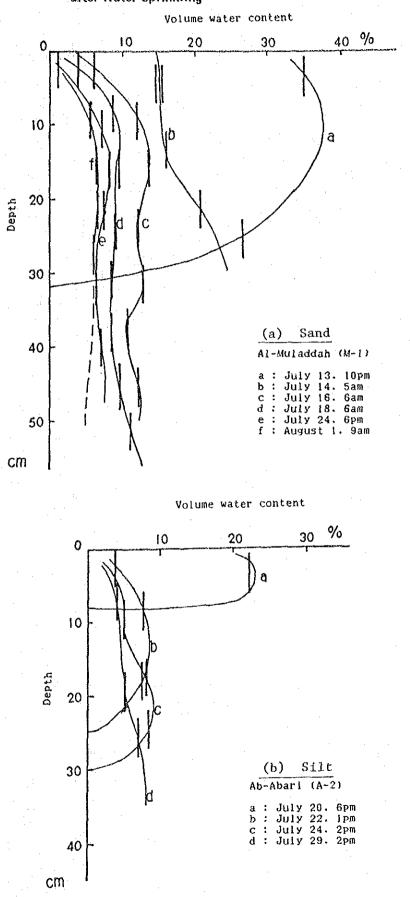


Fig. B-5-5 Change of Volume Water Content in Sand and Silt with Time after Water Sprinkling

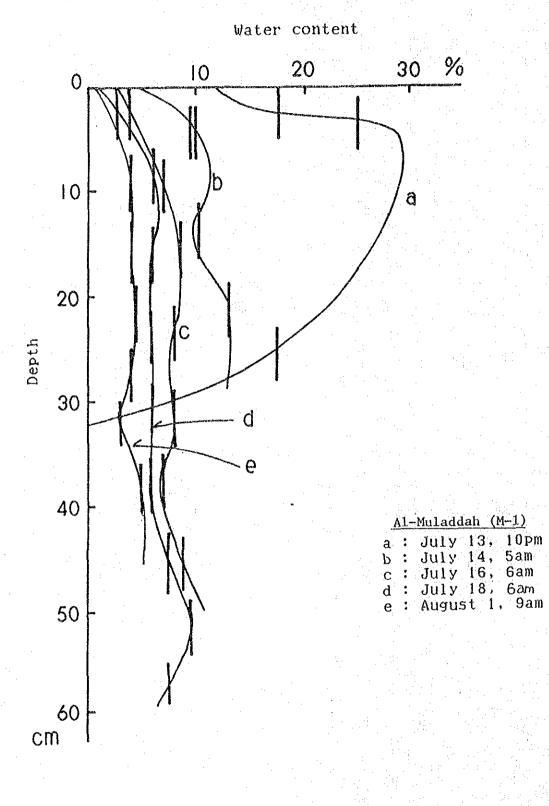
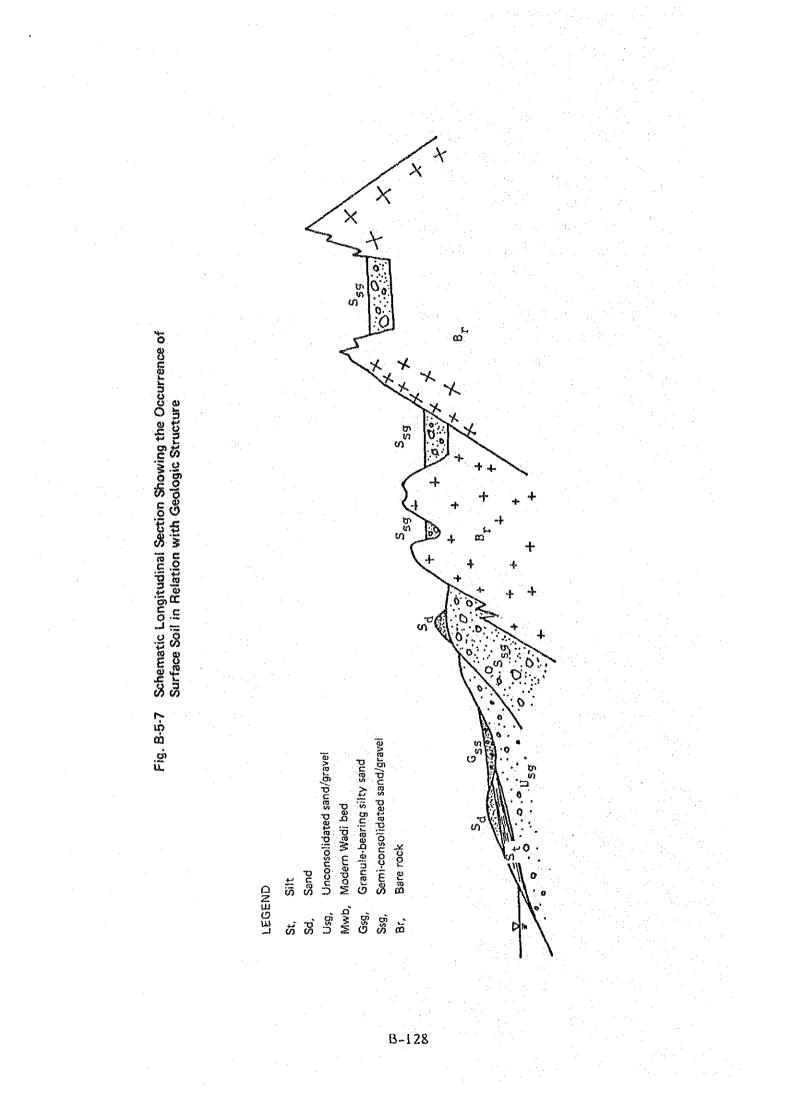


Fig. B-5-6 Change of Water Content in Sand with Time after Water Sprinkling

5.2 Areal Distribution of Infiltration Capacity

As shown in the schematic longitudinal section in Fig. B-5-7, silt, sand, unconsolidated sand/gravel and semi-consolidated sand/gravel distribute in the Batinah plain roughly in that order from the coast. This section is based on the topographic and geologic maps and also on the satellite imagies of LANDSAT and NOAA. A surface soil distribution map was drawn and is shown in Fig. B-5-8.

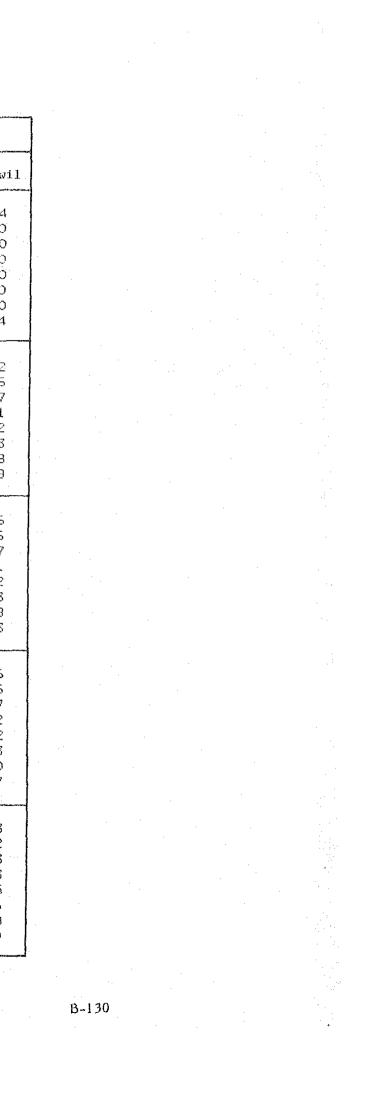
The foregoing experiment sites were selected to represent each type of soil surface in this map. The map exhibits the distribution of infiltration capacity. The area of each surface soil is compiled in Table B-5-4.





			Area	l extent	(km ²)	S & Conference of the second		۸rea	l percenta	age (%)	
P. Ca	ri adi	Ahin	Bani Ghafir	Al-Fara'	Bani Kharus	Al-Ma ^t awil	Ahin	Bani Ghafir	Al-Fara'	Bani Kharus	Al-Ma'awil
tt)	St Sd Usg Mwb	5,511 0,000 0,000 0,000	17.694 0.000 0.000 0.000	13.900 0.000 0.000 0.000	34,399 0,000 0,000 0,000	17,343 0,000 0,000 0,000	. 585 0.000 0.000 0.000	1.859 6.000 6.000 0.000	1.215 0.000 0.000 0.000	2,662 0,000 0,000 0,000	1,684 0,000 0,000 0,000
Coastal (A)	Gss Ssg Br Total	0,000 0,000 0,000 5,511	0,000 0,000 0,000 17,694	0.000 0.000 0.000 18.800	0,000 0,000 0,000 34,399	0,000 0,000 0,000 17,343	0.000 0.000 0.000 .585	0.000 0.000 0.000 1.859	0.000 0.000 0.000 1.215	0.000 0.000 0.000 2.652	0.000 0.000 0.000 1.684
Sand/gravel plain (B)	St Sd Usg Mwb Gss Ssg Br Total	18.216 8.741 34.890 11.973 5.950 140.368 47.744 267.882	$\begin{array}{c} 35.\ 474\\ 20.\ 443\\ 0.\ 000\\ 19.\ 053\\ 35.\ 474\\ 5.\ 533\\ 0.\ 000\\ 115.\ 042 \end{array}$	35,891 40,608 86,959 44,983 28,576 228,609 0,000 465,626	31.231 59.444 66.987 31.080 45.714 40.208 24.441 293.105	17.73626.32319.53231.9365.276207.329112.533420.664	1.516.7753.0941.062.52812.4504.23523.759	3.727 2.148 0.000 2.003 3.727 .587 0.000 12.191	2.320 2.525 5.522 2.908 1.847 14.779 0.000 30.103	$\begin{array}{c} 2.\ 417\\ 4.\ 600\\ 5.\ 184\\ 2.\ 405\\ 3.\ 537\\ 3.\ 111\\ 1.\ 391\\ 23.\ 145\end{array}$	1.722 2.555 1.897 3.101 .512 20.133 10.928 40.849
(A)+(B)	St Sd Usg Mwb Gss Ssg Br Total	$\begin{array}{r} 24.827\\ 8.741\\ 34.890\\ 11.973\\ 5.950\\ 140.363\\ 47.744\\ 274.493\end{array}$	0,000	0.000	$\begin{array}{c} 65.\ 630\\ 59.\ 444\\ 66.\ 987\\ 31.\ 080\\ 45.\ 714\\ 40.\ 208\\ 24.\ 441\\ 333.\ 504 \end{array}$	$\begin{array}{c} 35.\ 079\\ 26.\ 323\\ 19.\ 532\\ 31.\ 936\\ 5.\ 276\\ 207.\ 329\\ 112.\ 533\\ 438.\ 007 \end{array}$	$\begin{array}{r} 2.202 \\ .775 \\ 3.094 \\ 1.062 \\ .528 \\ 12.450 \\ 4.235 \\ 24.345 \end{array}$	5, 585 2, 148 0, 000 2, 003 3, 727 , 587 0, 000 14, 049	$\begin{array}{c} 3.536 \\ 2.625 \\ 5.522 \\ 2.903 \\ 1.847 \\ 14.779 \\ 0.000 \\ 31.318 \end{array}$	5.079 4.800 5.184 2.405 3.537 3.111 1.891 25.807	3.406 2.556 1.897 3.101 .512 20.133 10.928 42.533
Mountain area (C)	St Sd Usg Mwb Gss Ssg Br Total	24.827 8.741 34.890 11.973 5.950 207.578 559.049 853.007	535.230	448.467	65,630 59,444 66,987 31,080 45,714 114,814 575,127 958,796	35.079 26.323 19.532 33.283 5.276 237.301 235.000 591.793	2.202 .775 3.094 1.062 .528 18.410 49.583 75.655	5.585 2.310 0.000 3.501 4.250 14.077 56.227 85.951	3.536 2.625 5.622 3.456 1.847 22.602 28.993 68.682	5.079 4.600 5.184 2.405 3.537 8.884 44.504 74.193	3.406 2.555 1.897 3.232 .512 23.043 22.820 57.467
(A)+(B)+(C)	St Sd Usg Mwb Gss Ssg Br Total	49.654 17.492 69.780 23.946 11.899 347.946 606.793 1127.500	42.431 0.000 52.395 75.930 139.577 535.230	81,216 173,918 98,444 57,152 578,221 448,467	131.259 118.888 133.975 62.160 91.429 155.022 599.568 1292.300	70.158 52.646 39.064 55.219 10.552 444.630 347.532 029.800	4.404 1.550 6.189 2.124 1.055 30.860 53.818 100.000	11.171 4.458 0.000 5.504 7.977 14.663 56.227 100.000	7.072 5.251 11.244 6.364 3.695 37.382 28.993 100.000	10.157 9.200 10.367 4.810 7.075 11.995 45.395 100.000	6.813 5.112 3.793 6.333 1.025 43.176 33.748 100.000
	Mountain area(A)+(B)Sand/gravelCoastal area2a/a(C)(C)(A)(C)(A)	St Sd Usg Mwb Gss Ssg Br Total St Sd Usg Mwb Gss Ssg Br Total St Sd Usg Mwb Gss Ssg Br Total St Sd Usg Mwb Gss Ssg Br Total St Sd Usg Mwb Gss Ssg Br Total (9)+(4) (0) (1)+(4) (2) St St Sd Usg Mwb Gss Ssg Br Total (1) (2) (1) (2) (1) (2) (1) (2) (2) (2) (1) (2) (1) (2) (2) (1) (2) (2) (2) (1) (2) (2) (2) (2) (2) (2) (2) (2) (2) (2	St St<	Pres Pres Ahin Bani Ghafir St 5.611 17.694 Sd Sd 0.000 0.000 0.000 Usg 0.000 0.000 0.000 Wab 0.000 0.000 0.000 Sd 0.000 0.000 0.000 St 18.216 35.474 Sd 8.741 20.443 Dotal 5.511 17.694 Mab 11.973 19.063 Br 47.744 0.000 Total 257.832 116.042 St 24.827 53.168 Sd 8.741 20.443 Sg 140.363 5.533 Br 47.744 0.000 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 Table B-5-4
 Areal Classification of Surface Geologic Coverage



5.3 Infiltration Capacity and Surface Runoff

Some rainfall infiltrates into soil, and the rest flows on surface as surface runoff. For a single rainfall,

p = m + f

where p, m and f are the amounts of rainfall, infiltration and surface runoff, respectively.

Based on the experiment results in Fig. B-5-2, the interpolation and extrapolation of measured 50% values were calculated and the cumulative infiltration amount verses time is shown in Fig. B-5-9. The infiltrationrate with time derived from the figure above was drawn and is shown in Fig. B-5-10. The initial and final infiltrabilities of each soil thus obtained are tabulated in Table B-5-5, where initial infiltration is defined as the value at 0 hour of the extrapolated straight line.

The actual relationship between p, m and f applied to the rainfall on Aug. 10, 1983 is exhibited in Fig. B-5-11.

and the second				·			
	St	Sd	Gss	Usq	Ssq	Br	Mwb
Initial infiltration loss (mm)	15	25	15*	3	5	2*	
Final infiltration rate (mm/h)	10	40	25*	20	0	0*	

Table B-5-5 Initial Infiltration Loss and Final Infiltration Rate

*Estimation value

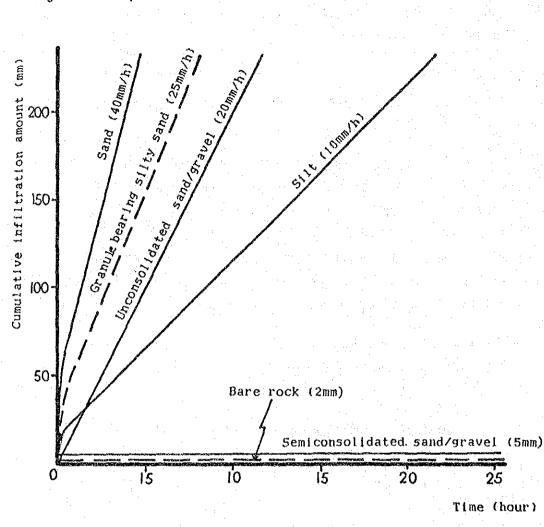
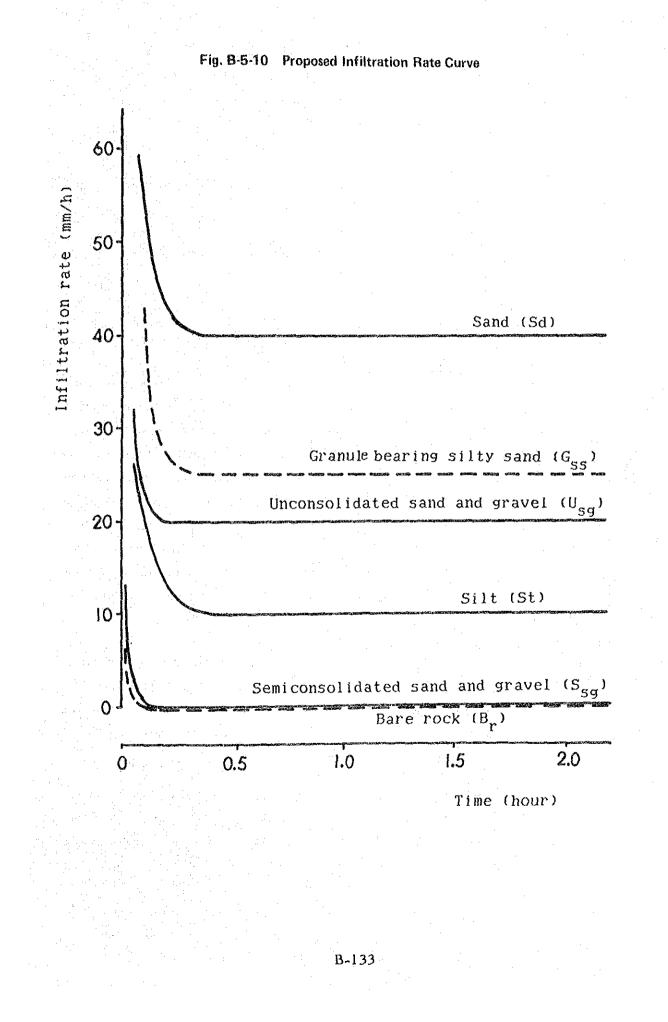


Fig. B-5-9 Proposed Curve of Cumulative Intiltration Amount with Time



5.4 Infiltration Capacity and Evapotranspiration

The majority of water which infiltrates into soil is later lost due to evaportranspiration. The balance percolates down to the groundwater zone. For a single rainfall,

 $m = e^* + r$

where m, e^* and r are the amount of infiltration, evapotranspiration and ground water recharge due to percolation.

In the region of semi-consolidated sand/gravel and bare rock, the final infiltration rate is zero (cf. Table B-5-5), then

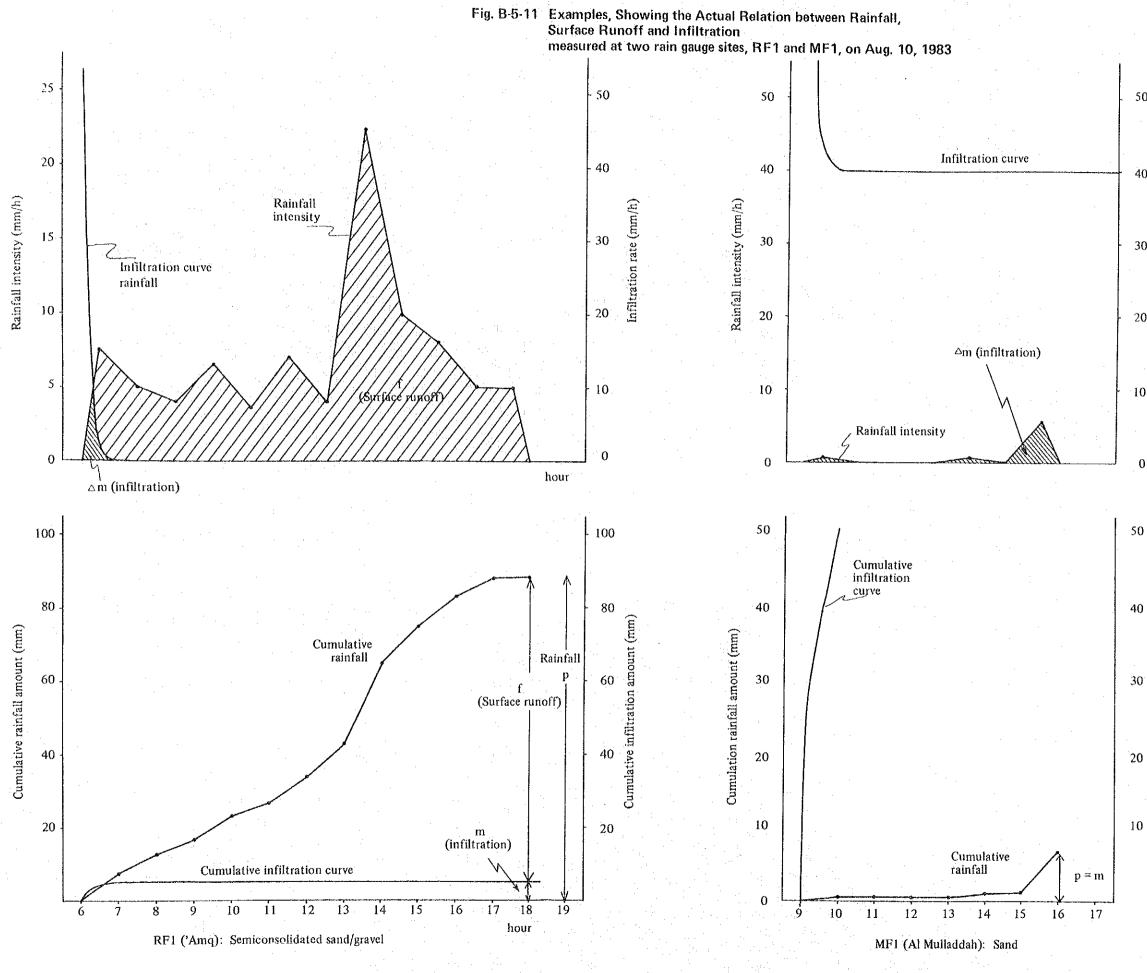
m = e*

The region of high final infiltration-rate extends along the coastline where rainfall is rare, and then only in small amounts.

When rainfall is small, m is also small and the resultant r becomes very small. Fig. B-5-12 shows the depth of infiltration observed at the Al-Muladdah Agro-meteorological station for the next day after rainfall of 6.5 mm. The average infiltration depth is 8.7 cm, and the average volume water content in wet layer is 7.4%. It is routhly equal to the field capacity of 6% approx. obtained through the experiment (cf. Fig. B-5-4). This fact implies that tens of mm of rainfall do not percolate down to groundwater zone but stay in the surface soil layer within a few meters. A similar situation may hold true for the other soils. The percolation down to groundwater zone should be negligible for all soils, that is, r = 0.

The bare rock and semi-consolidated sand/gravel contain the smallest volume of water among all soils and it may evaporate within 24 hours after rainfall. The other soils take several days or weeks until they return to the initial soil moisture conditions before rainfall. Evapotranspiration takes place during these days.

What is mentioned above corresponds with the analytical results of NOAA images. The surface soil distribution in the coastal plain corresponds especially well with the final NOAA images obtained for the aerial change on soil water content after rainfall (cf. Supporting Report (F)).



50 40 Infiltration rate (mm/h) 30 20 10 0 50 Cumulative infiltration amount (mm) LEGEND P : Rainfall

f : Surface Runoff

m: Infiltration

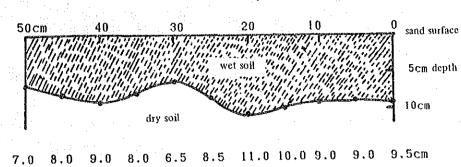


Fig. B-5-12 Infiltration Depth of Rain at Al-Muladdah on Dec. 29, 1984, Measured at 14:30 on Dec. 30, 1984

Estimation of Actual Evaporation After Rainfall

The equilibrium equation for energy at soil surface level can be expressed:

LeE + H + Qn + G = O,

where Le is latent heat of evaporation, E is rate of actual evaporation, H is specific flux of sensible heat into atmosphere, Qn is specific flux of net radiation, and G is specific flux of heat conducted into the earth. The Bowen ratio Bo, which is the ratio between latent heat and sensible heat, is defined by:

$$Bo = H/LeE$$

5.5

Derived from two equations above is the actual evaporation E, namely:

$$E = \frac{-Qn - G}{(1 + Bo)Le}$$

Bo, Bowen ratio, can be determined by:

Bo =
$$\frac{Cp(T_1 - T_2)}{Le(q_1 - q_2)}$$

where o_p is specific heat under constant pressure, T_1 and T_2 are air temperature at two different heights, and q_1 and q_2 are specific humidity at two different heights.

A items except E observed at Al-Muladda, Agrometeorological Station.

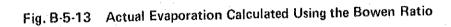
By using the equations above, the daily amount of actual evaporation E, was calculated for the period before and after the rainfall on August 10, 1983, the results shown in Fig. B-5-13.

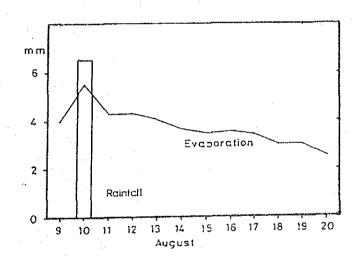
A clear tendency can be seen on daily change of calculated amount. However the calculated actual evaporation exceeds the total rainfall. One of the reasons of the excess is unsuitable observation height of the thermometer. Therefore, the height of the thermometer were adjusted in November 1983. The data observed after that time are considered to be more appropriate for estimation of actual evaporation.

5.6 Estimation of Dew

The same method which is used in estimation of actual evaporation normally can be applied to the estimation of dew. However, the daily value of dew is less than the actual daily evaporation after rainfall and the method has the same problem mentioned in paragraph 5.3 above. Accordingly, the estimated dew cannot be evaluated quantatively. However, the results of estimation suggest dew will probably occur in summer, in August and September especially. This coincides with personal experience.

Dew on tree leaves is greater than that on the ground because the temperature of leaves is lower than that of the ground at night.





CHAPTER 6 STUDY OF SURFACE RUNNOFF

6.1 Observed Surface Runoff

During the survey period, surface runoffs were observed eight times at the six wadi gauges installed by the Project. However, large floods which reach the sea did not occur in the project period as expected. The survey period was an extremely dry one. Further investigation will be required for studying the relationship between the precipitation and topographic structures. Eight flood dischanges recorded during the survey period are summarized in Table B-6-1, B-6-2(1) - (8) and FiG. B-6-1(1) - (8).

Manning's equation was used to estimate outflow from the water level data, given the geometry of wadi channel found by surveying. The roughness coefficient was assumed as 0.040 according to the wadi situation and the previous studies. The ground surveys such as longitudinal profile and cross-section leveling surveys were executed at all the gauge sites during the survey period. The results of survey are shown in Fig. B-1-7 (1) (15).

Wadi Basin	Discharge Site	Date	Duration	Outflow (m ³)	Peak Discharge (m ³ /sec)
Wadi Al-Fara'	WF1 (Mazahit)	Sep.15,'83	15:50- 00:02	26,040	5.8
Wadi Bani Ghafir	WG1 (Al-Houqain)	Sep.15'83	13:46- 20:57	1,263	0.9
Wadi Al-Fara'	WF4 (Tabaqah)	Apr.12,'84	14:55- 16:15	815	0.7
Wadi Al-Fara'	WFl (Mazahit)	Jul.30'84	18:00- 22:01	65,017	13.0
Wadi Ahin	WAI (Al-Heil)	Sep.15,'84	17:00- 18:08	2,036	1.6
Wadi Ahin	WA1 (Al-Heil)	Sep.16'84	19:25- 22:52	8,376	1.7
Wadi Al-Ma'awil	WM2 (Afi)	Apr.18,'85	16:10 - 20:07	41,891	19.1
Wadi Bani Kharus	WK I (Al-Abiyad)	Apr.19,'85	16:20 06:51	127,658	44.1

Table B-6-1 Observed Flood during Survey Period

Fig. B-6-1(1) Observed Runoff Hydrograph WF1, Sep. 15, 1983

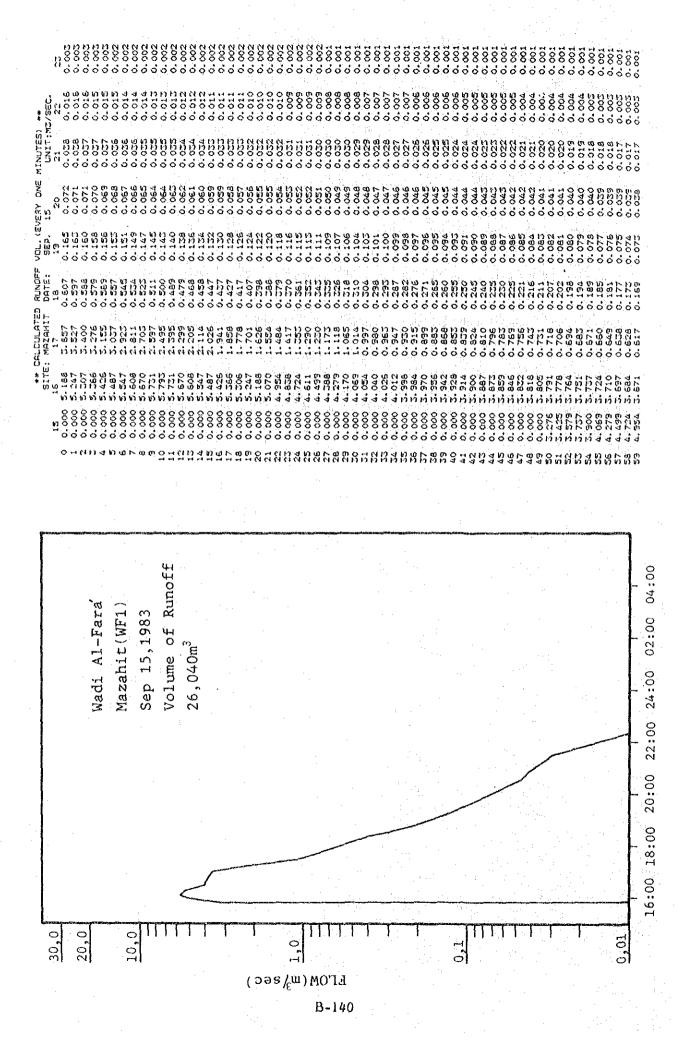
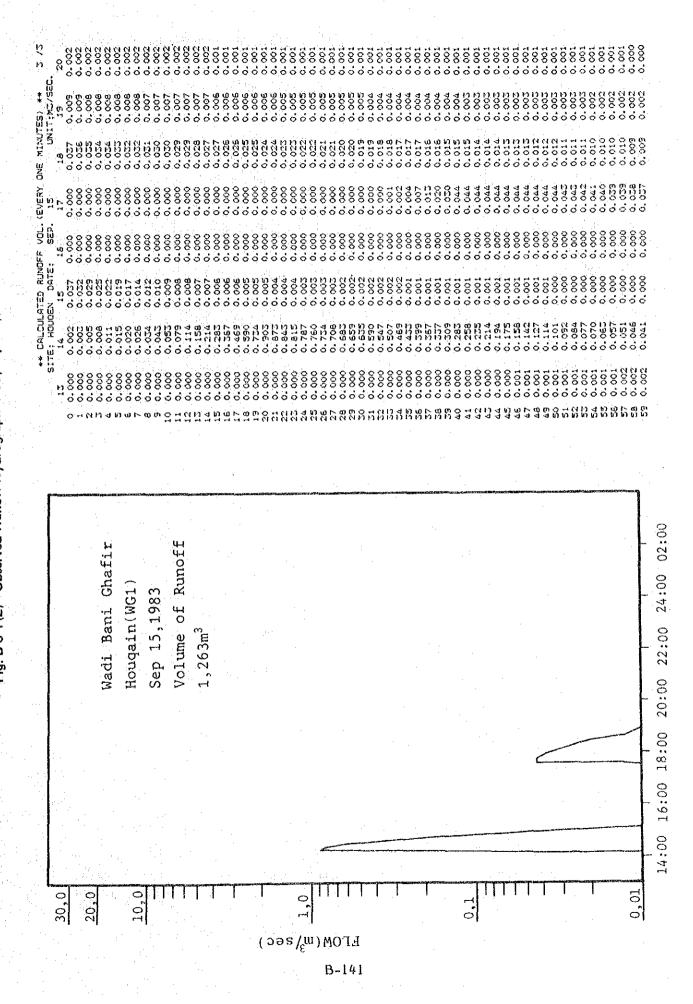


Fig. B-6-1 (2) Observed Runoff Hydrograph WG1, Sep. 15, 1983



Fig, 8-6-1 (3) Observed Runoff Hydrograph WF4, Apr. 12, 1983

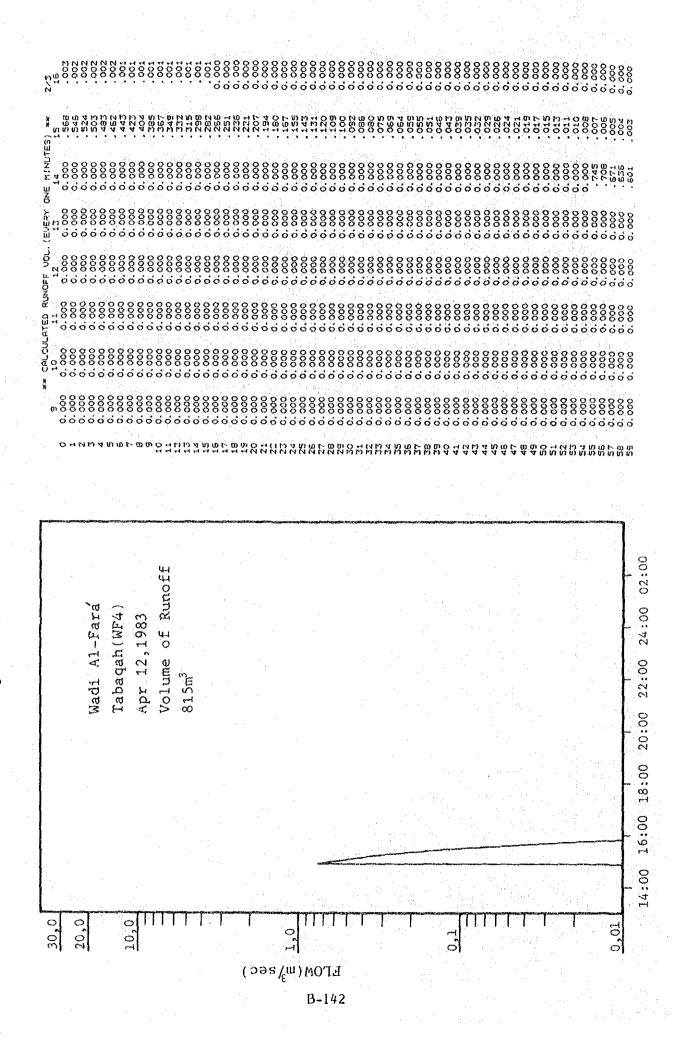


Fig. B-6-1 (4) Observed Runoff Hydrograph WF1, Jul. 30, 1984

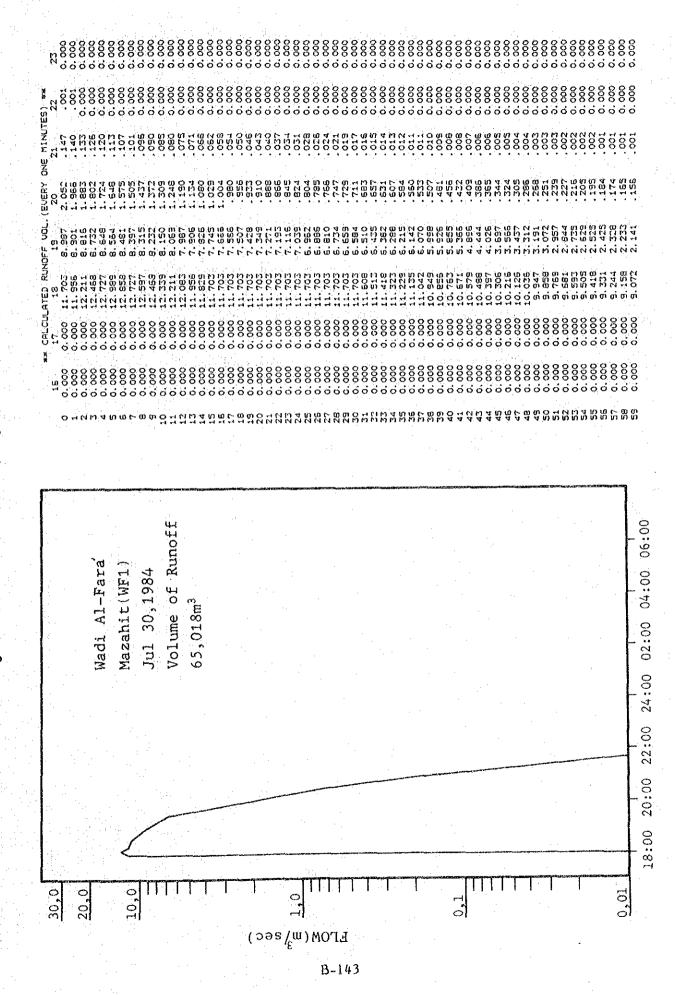


Fig. B-6-1 (5) Observed Runoff Hydrograph WA1, Sep. 15, 1984

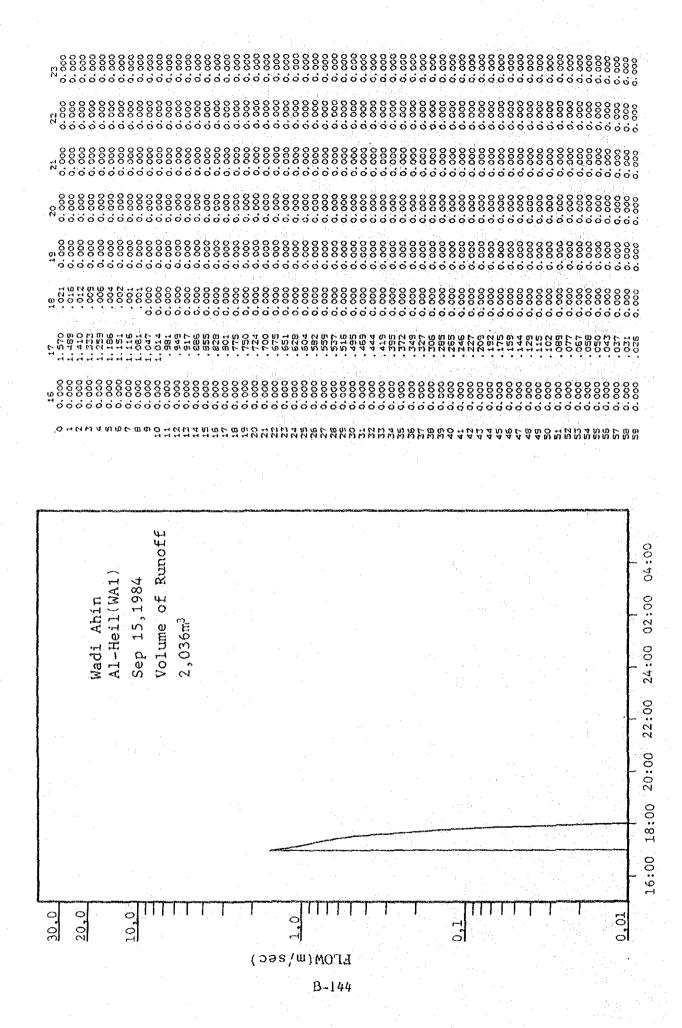


Fig. B-6-1 (6) Observed Runoff Hydrograph WA1, Sep. 16, 1984

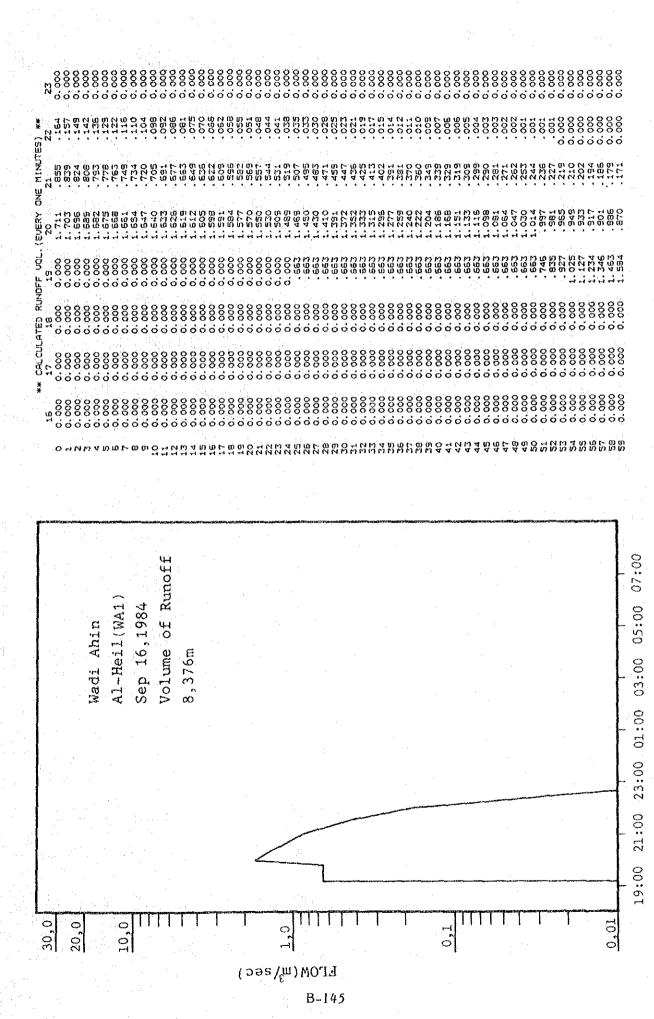
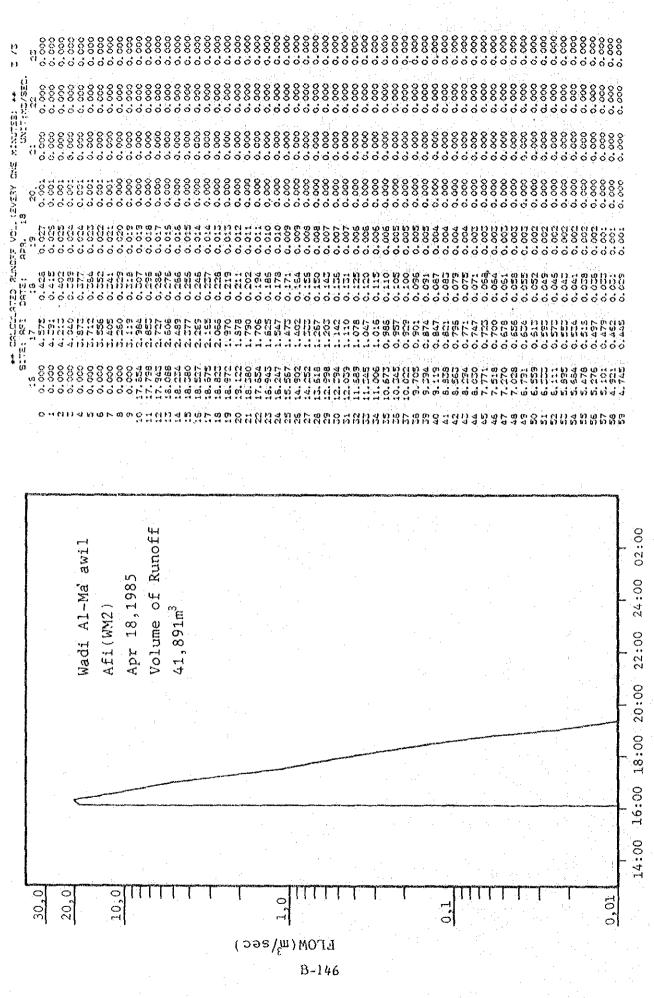


Fig. B-6-1 (7) Observed Runoff Hydrograph WM2, Apr. 18, 1985



Observed Runoff Hydrograph WK1, Apr. 19, 1985 (1/2) Fig. B-6-1(8)

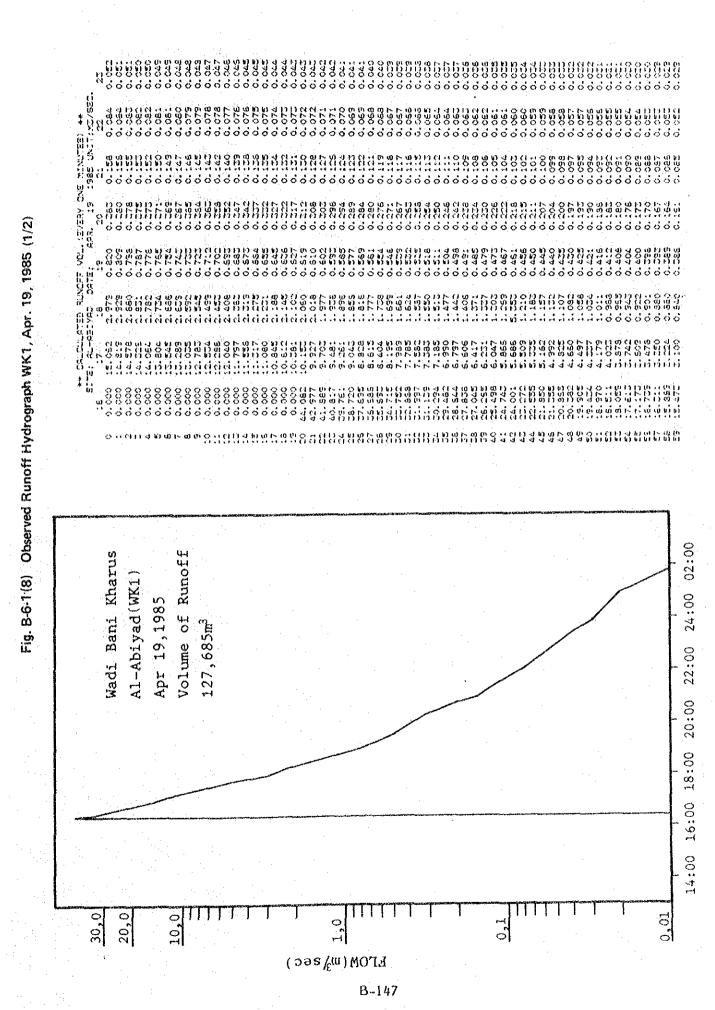


Fig. B-6-1 (8) Observed Runoff Hydrograph, WK1, Apr. 19, 1985 (2/2)

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6.2 Surface Runoff Features

It was observed that the rainfalls which caused the runoff observed at the wadi gauges, were not recorded at the rain-gauges in located in the basins of the wadi gauges, because of their small coverage. The density of the rain gauges might be reviewed after several years observation for water resources development purposes. Although only a small amount of data was obtained during the Project, the general relationship between rainfall and runoff are described based on this observation data and previous studies. (ILACO and GIBB, 1977).

1. Short duration of runoff

The minimum duration of runoff was about 1 hour and the maximum was about 24 hours in both of mountain and coastal area. Hence, runoff is usually of very short duration.

2. Rapid runoff in the mountain area

The Project area belongs to the semi-arid zone and the major mountain area, which are covered with steep bare rocks and sparce vegetation so that the flood reaches the wadi bed rapidly. Wadi gradients in the mountain area are as steep as 1/75 - 1/145. These basin conditions are the cause of rapid runoffs in the mountain areas.

3. Great infiltration capacity of the wadi-bed

Fewer observed runoff events comparing the number of rainfall imply that the wadi-beds have a large infiltration capacity. There are many cases of the rainfall which were recorded at the rain-gauges in the mountain area, however no discharges were recorded at the wadi gauges downstream.

Floods of the coastal area occur mainly in winter

The number of the floods in winter is greater than the ones in summer in the coastal area. This tendency suggests that there is a relationship between the rainfall intensity and the infiltration capacity of the gravel plain.

As the area which has infiltration capacity exist only along wadi beds in the gravel plain (about 18%), the other area (about 82%) is classified as non-permiable.

6.3 Flood Discharge to the Sea

Flood discharge to the sea was not observed during the observation period (Sep. '83 to Aug. '84). Flood discharges to the sea which had been observed in previos studies are summarized in Table B-6-2. Those 7 floods occurred in the winters from 1974 to 1976. The year of 1976 had heavy rainfall according to the rainfall data of Muscat (See Table B-6-3). From the previous studies and our observation, flood discharge to the sea is rare for the following reasons:

- 1. The coastal area which is composed mainly of sand and silt seems to have a large storage infiltration capacity. Consequently, most of the rainfalls was captured along the wadi bed.
- 2. The Gravel plain area spreads out from about 10 km upstream of the seaside for about 35 km to the interior. The surface of the Maginal wadi plain area is cemented hard which will easily convey surface runoff to the wadi bed rapidly. However, the wadi beds have a large infiltration capacity, which will accelerate the depletion of the discharge into the ground. Thus, the average annual rainfall would be about 100 mm/year and each rainfall would be small scale.

3. Mountain area and runoff

Mountain area, is mainly covered with bare rocks and sparce vegetation. Surface runoffs occur easily, but the wadi-beds are covered with gravels where the surface flow can easily infiltrate. Accordingly, surface runoff to the sea occurs only under the following conditions;

a) Heavy rainfall over the whole water-shed.

b) Rainfall with strog rainfall intensity in the coastal area.

Regarding the water resources in the Batinah area, Horn (1979) and Credew (1980) reported that the average runoff to the sea is about 20-23 MCM/year (Table B-6-4). The estimations by Horn and Credew are very similar except for the discharge from Wadi Bani Kharus.

Table B-6-2	Enumeration of Observed Floods	

	· · ·	1. A. A. A.							
Wadi Basin	Wadi Gau- ge	Topolo- gical Location	Observa- tion Date	Season	Observation time	Duration	Peak Disch m ³ /s	Runoff arge Volume m ³	Öbser- ver
W. Ahin	WAI	Mountain	Sept.15,'84	Summer	17:00 - 18:08	I.I hrs	1.6	2,036	JICA
•	WA1	Mountain	Sept.16,'84	Summer	19:25 - 22:52	2.5 hrs	1.7	8,376	JICA
W.B.Ghafir	N7	Mountain	Feb.12, 75	Winter	18:00 - 18:00	24 hrs	90.9	738,360	Gibb
	N7	Mountain	May 13, 75	Winter	15:20 - 17:30	2.2 hrs	16.9	26,520	Gibb
	WG1	Mountain	Sept.15,'83	Summer	13:46 - 20:57	7.2 hrs	0.9	1,263	JICA
W. A.I. F	1.1			· ·			1.5	9,630	Gibb
W. Al-Farð	N9	Mountain	Jul.7,'74	Summer	16:00 - 18:06	2.0 hrs 8.6 hrs		24,237	Gibb
· · ·	N9	Mountain	Feb.2,'76	Winter Winter	18:25 - 03:00	20.8 hrs	3.3 2.9	24,237 31,470	Gibb
	N9	Mountain	Feb.6,'76		06:00 - 02:50 13:00 - 23:00	20.8 hrs 10.0 hrs	59.0	177,810	Gibb
	N9	Mountain	Jul.29, 76	Summer Summer	15:50 - 00:02	8.2 hrs	5.8	26,040	JICA
	WF1	Mountain	Sept.15,'83	Summer	19:00 - 22:01	4.0 hrs	13.0	65,017	JICA
	WF1	Mountain	Jul.30,'84	Winter	14:55 - 16:15	1.3 hrs	0.7	815	JICA
:	WF4	Mountain	Apr.12,'84					1	
W.B.Kharus	N3	Mountain	Mar.25,'76	Winter	17:30 - 20:30	3.0 hrs	7.3	41,925	Gibb
*.	N4	Coastal	May 14,'75	Winter	17:45 - 02:00	8.3 hrs	3.4	74,506	Gibb
	N4 -	Coastal	Feb.3,'76	Winter	00:00 - 05:00	5.0 hrs	4.6	29,715	Gibb
	N4	Coastal	Mar.26,'76	Winter	02:00 - 03:30	1.5 hrs	1.2	6,720	Gibb
н н	N4	Coastal	Mar.30,'76	Winter	23:00 - 07:00	8.0 hrs	1.9	23,640	Gibb
-	N4	Coastal	Apr.9,'76	Winter	02:30 - 06:00	3.5 hrs	6.9	51,900	Gibb Gibb
	N5	1	Oct.,2'74	Winter	15:00 - 22:45	7.8 hrs	5.9	37,800	
	N5	Mountain	Oct.4,174	Winter	15:00 - 18:12	3.2 hrs	77.7	244,080	Gibb Gibb
	N5	Mountain	Oct.6,'74	Winter	16:00 - 17:36	1.5 hrs	2.0	5,040	Gibb
	N5	Mountain	Dec.27,'74	Winter	20:00 - 23:20	3.3 hrs	8.2	38,160 68,040	Gibb
	N5	Mountain	May 12,75	Winter	15:20 - 16:40	1.3 hrs	40.3		Gibb
	N5	Mountain	Aug. 16, 75	Summer	18:00 - 13:50	19.8 hrs	71.2 69.5	295,459 635,463	Gibb
	N5	Mountain	Aug. 17, 75	Summer	14:40 - 04:40	14 hrs 7.7 hrs	5.1	47,175	Gibb
	N5	Mountain	Mar.30,'76	Winter	17:00 - 00:40	10.8 hrs	85.6	672,270	Gibb
	N5	Mountain	Apr.8'76	Winter	19:00 - 05:50 16:15 - 09:30	10.8 hrs	107.3	781,065	Gibb
	N5	Mountain	Sept.15,'76	Summer	21:50 - 05:20	7.5 hrs	1.5	20,565	Gibb
	N6	Mountain	May 22,'74	Winter	19:00 - 10:55	16.0 hrs	7.9	43,305	Gibb
	N6	Mountain	Oct.1,'74	Winter	19:00 - 10:33	6.3 hrs	9.6	60,930	Gibb
	N6	Mountain	Oct.2,174	Winter Winter	16:20 - 06:51	13.5 hrs	44.1	127,658	JICA
	WK1	Mountain	Apr.19,'85		1 a a	· · ·			
W. Al-Maawi	1 W M 2	Mountain	Apr.18,'85	Winter	16:10 - 20:07	4.0 hrs	19.1	41,891	ЛСА

Table B.6.3	Observed Flood Discharge to	the Sea during Previous	Survey Period
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Wadi Basin	Name of Place (Name of gauge)	Runoff date	Runoff duration	Peak flow m ³ /sec.	Amount of Runoff m ³
Wadi *1) Ahin	Khishdah	Feb.15,'74	No data	1.2	12,000
		Feb.15,'74	No data	1.4	20,000
Wadi *2) Bani Kharus	Al-Sawadi (N4)	May 14,'75	17:45 - 02:00	3.4	74,506
		Feb.02, '76	00:00 - 05:00	4.6	29,715
		Mar.25,'76	02:00 - 03:30	1.2	6,720
		Mar.30,'76	23:00 - 07:00	1.9	23,640
		Apr.09,'76	02:30 - 06:00	6.9	51,900

*1) ILACO Jul. 1975

"WATER RESOURCES DEVELOMENT PROJECT NORTHERN OMAN"

*2) SIR ALEXANDER GIBB AND PARTNERS JUN. 1976 "WATER RESOURCES SURVEY OF NORTHERN OMAN"

Table B-6-4 Annual Rainfall at Muscat during Previous and JICA Survey Period

Year of Survey Period	Annual Rainfall (mm/year)				
1974	3.7				
1975	80.0				
1976	203.3				
1983	123.0				
1984	18.5				
Mean *	105.0				

*) Mean of rainfall data from 1951 to 1984

(unit = mcm/year)							
Reorters	Wadi Ahin	Wadi B. Ghafir		Wadi B. Kharus	Wadi Al-Ma'awil	Total	
ILACO *1	2,58	مىنىدىرىيىتى بىرى بىرى بىرى بىرى بىرى بىرى بىرى ب		La Charles and C		2.58	
GIBB *2	1999 - 1999 - 1999 - 1999 - 1999 - 1999 - 1999 - 1999 - 1999 - 1999 - 1999 - 1999 - 1999 - 1999 - 1999 - 1999 -	21.2	13.9	27.7	7.0	69.8	
HORN *3	5.5	4.5	4.1	5.4	0.4	19.9	
CARDEW *4	4.8	3.8	3.6	10.9	0.3	23.4	

Table B-6-5 Estimated Flood Discharge to the Sea for the Median Year

*1 ILACO JUL. 1975 "WATER RESOURCES DEVELOPMENT PROJECT NORTHERN OMAN"

*2 SIR ALEXANDER GIBB AND PARTINERS JUNE 1976 "WATER RESOURCES SURVEY OF NORTHERN OMAN"

*3 P.M. HORN-F.A.O. FEB. 1979 "WATER RESOURCES OF THE BATINAH"

*4 PRECCE CARDEW AND RIDER/SIR M MACDONALD AND PARTNERS SEP. 1980 "POWER AND URBAN WATER SUPPLY STUDY: PHASE II, WATER DEVELOPMENT PROGRAME"

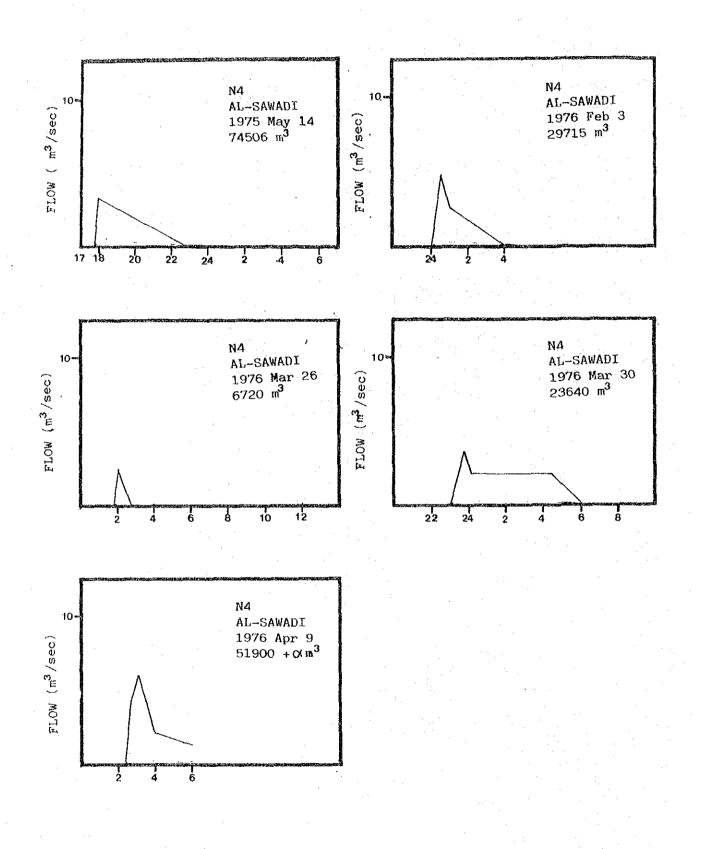


Fig. B-6-2 Observed Runoff to the Sea during Previous Survey Period