

### 2-2-3 Measurement of Radon Concentration

A hole of 60 cm depth is drilled with a hand-auger and an etch cup (made of Terradex) is buried bottoms up in the hole and covered (Fig. II-2-10).

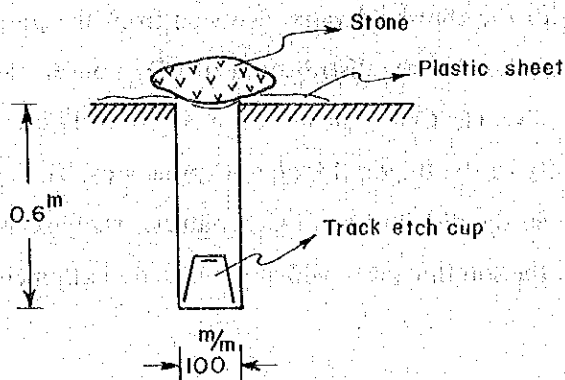


Fig. II-2-10 Planting of Track Etch Cup

After about 3 weeks, the cup is retrieved and sent to Terradex for track etch count. The corrected values (Table II-2-1) correspond to the number of tracks in 30 days for a  $1 \text{ mm}^2$  area. Fig. II-2-11 shows the frequency distribution of the logarithm of radon concentration (in tracks per  $\text{mm}^2$  per 30 days). The distribution is very regular and the probability curve (Fig. II-2-12) gives a threshold value of 1.6 corresponding to 39.8 tracks per 30 days per  $1 \text{ mm}^2$ .

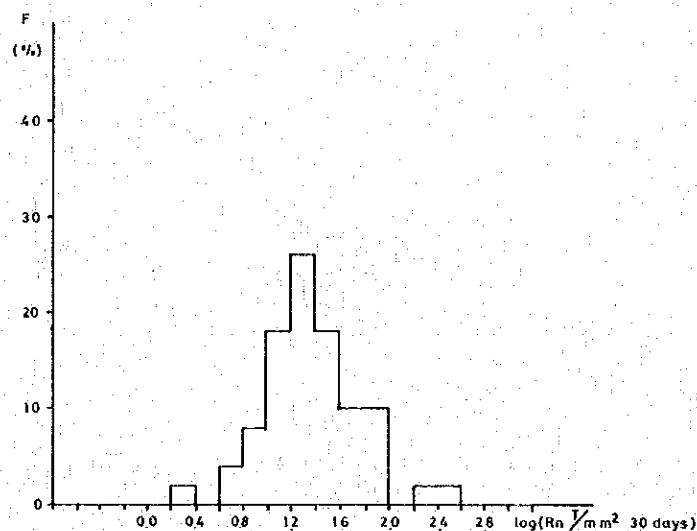


Fig. II-2-11 Frequency of Diagram of Rn Gas Concentration

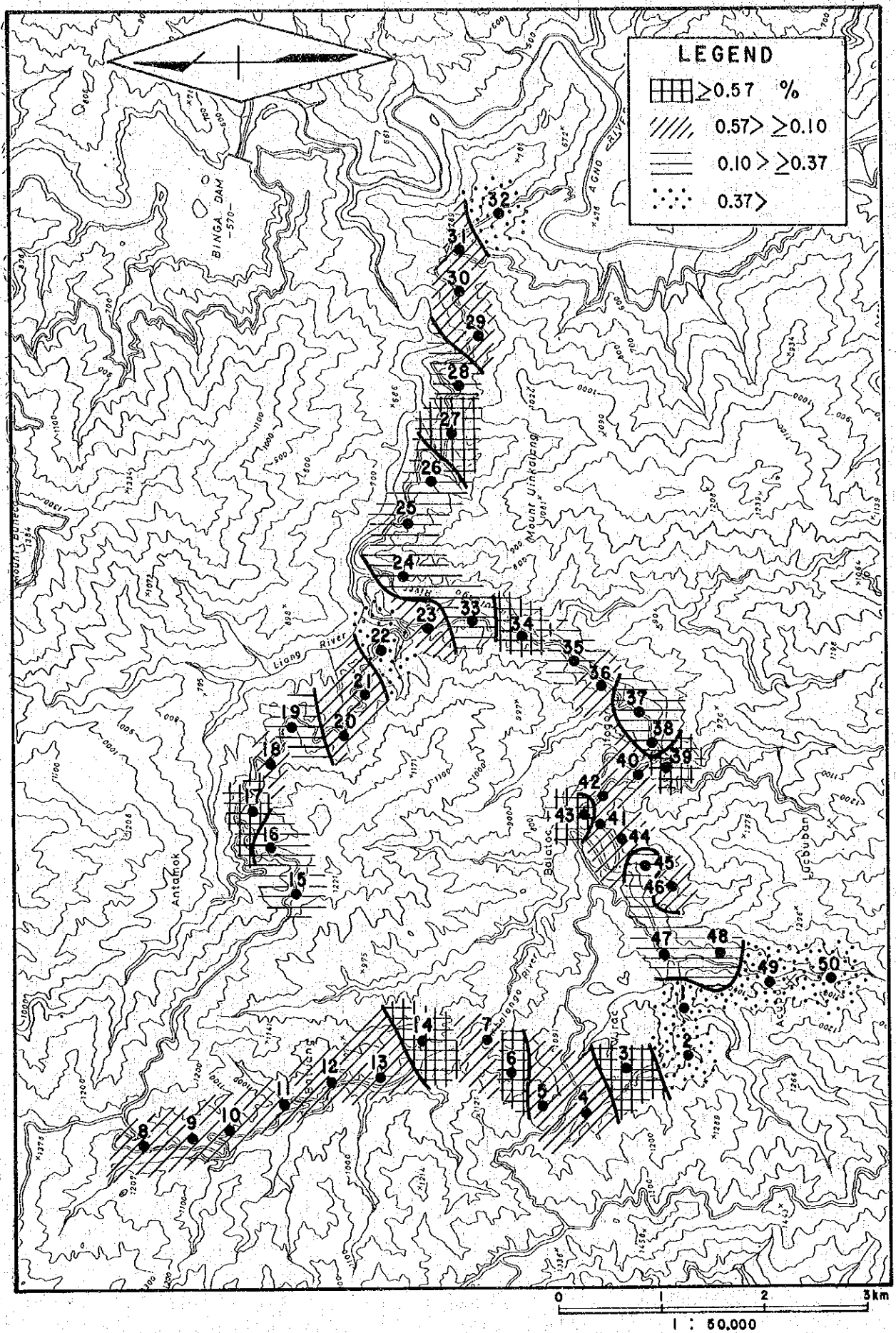


Fig. II-2-9

Distribution Map of CO<sub>2</sub> Gas Concentration



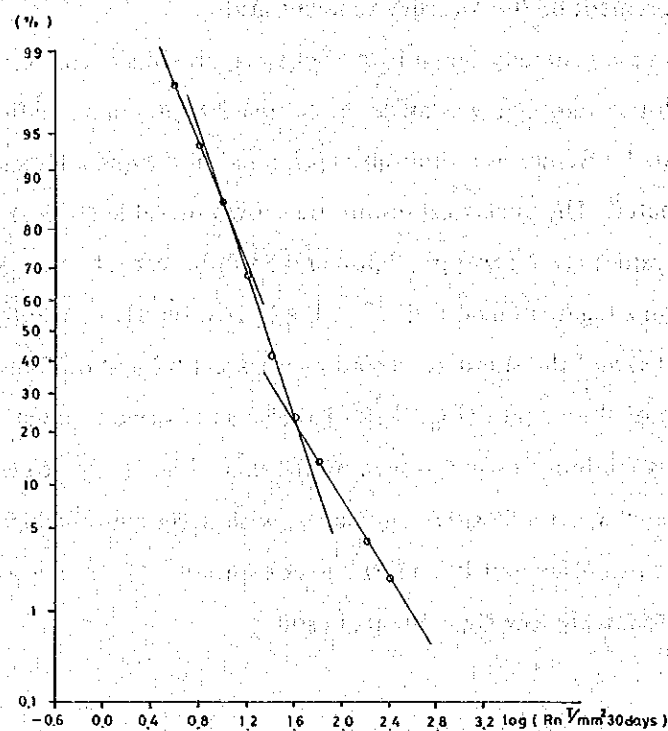


Fig. II-2-12 Probability Plot of Rn Gas Concentration

The mean value ( $\bar{x}$ ) and standard deviation ( $v$ ) obtained from the frequency diagram (Fig. II-2-11) are 1.37 and 0.38 respectively. The map of radon concentration (Fig. II-2-13) has been drawn using  $\bar{x}$ ,  $\bar{x} \pm v$  and  $\bar{x} - 2v$  as limits. It shows a large domain of relatively high radon concentration from the Itogon Bridge hot spring to Acupan Mine. But the most striking feature of the radon distribution is a large positive anomaly in the vicinity of Batuang which corresponds geologically to the presence of andesite and granodiorite-porphyry. The outlined anomaly thus reflects this geological difference rather than the presence of geothermal fluids at depth. On the other hand, a large negative anomaly appears in the northeastern part of the field which corresponds geologically to intermediate and basic eruptive rocks.

#### 2-2-4 Measurement of the Mercury Concentration

Soil samples are generally taken from the  $A_1$  layer. But as the soils of the surveyed area are poorly developed sampling was made in the first 20–30 cm excluding the  $A_0$  layer. The analysis was made by flameless atomic absorption method using a Hiranuma HG-1 type mercury concentration meter. The analytical results are shown in Table II-2-1. The maximum, minimum and mean values are 4,385 ppb, 9 ppb and 83 ppb, respectively. As shown by the frequency diagram drawn on a logarithmical scale (Fig. II-2-14), the Hg concentrations fall in two groups. The mean value ( $\bar{x}$ ) and the standard deviation ( $v$ ) are 1.92 and 0.65 respectively on the logarithmic scale. The probability plot (Fig. II-2-15) also shows two populations of Hg values. The distribution of Hg content in soil is drawn on the map (Fig. II-2-16) with the limit concentrations of  $\bar{x}$ ,  $\bar{x} \pm v$  and  $\bar{x} - 2v$ . Positive anomalies, with a Hg content higher than  $\bar{x} + v$  (2.57), occur around Acupan Mine and Itogon Bridge hot springs. The northeastern part of the area however, shows relatively low Hg contents in soil.

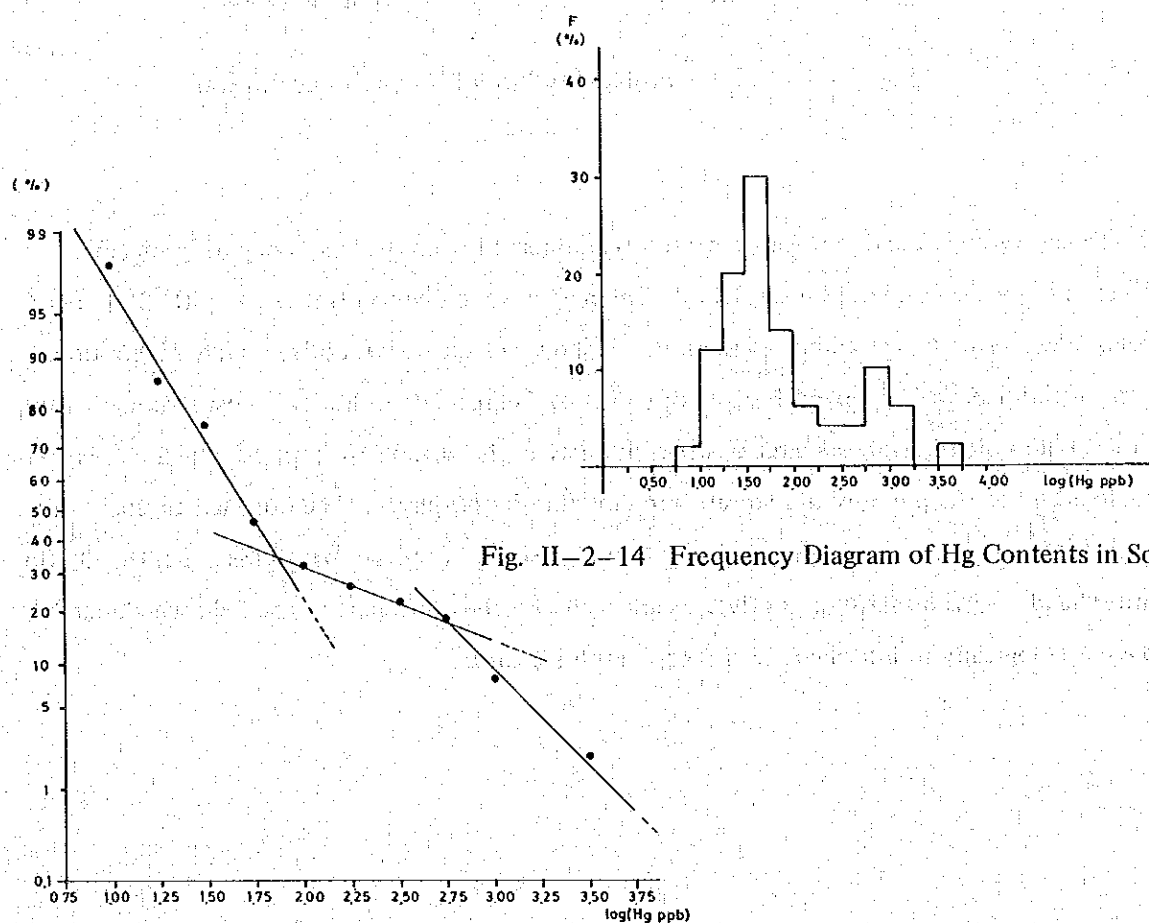


Fig. II-2-14 Frequency Diagram of Hg Contents in Soil

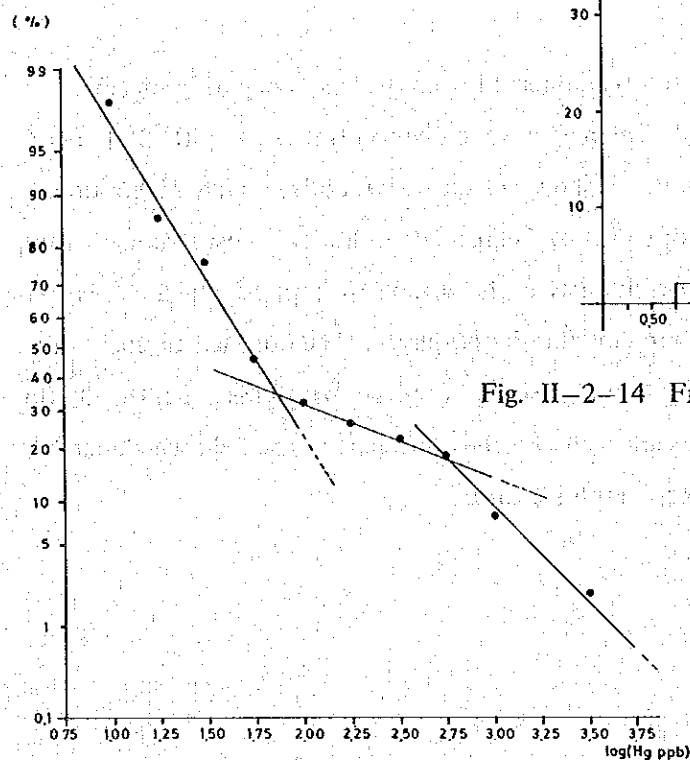


Fig. II-2-15 Probability Plot of Hg Contents in Soil

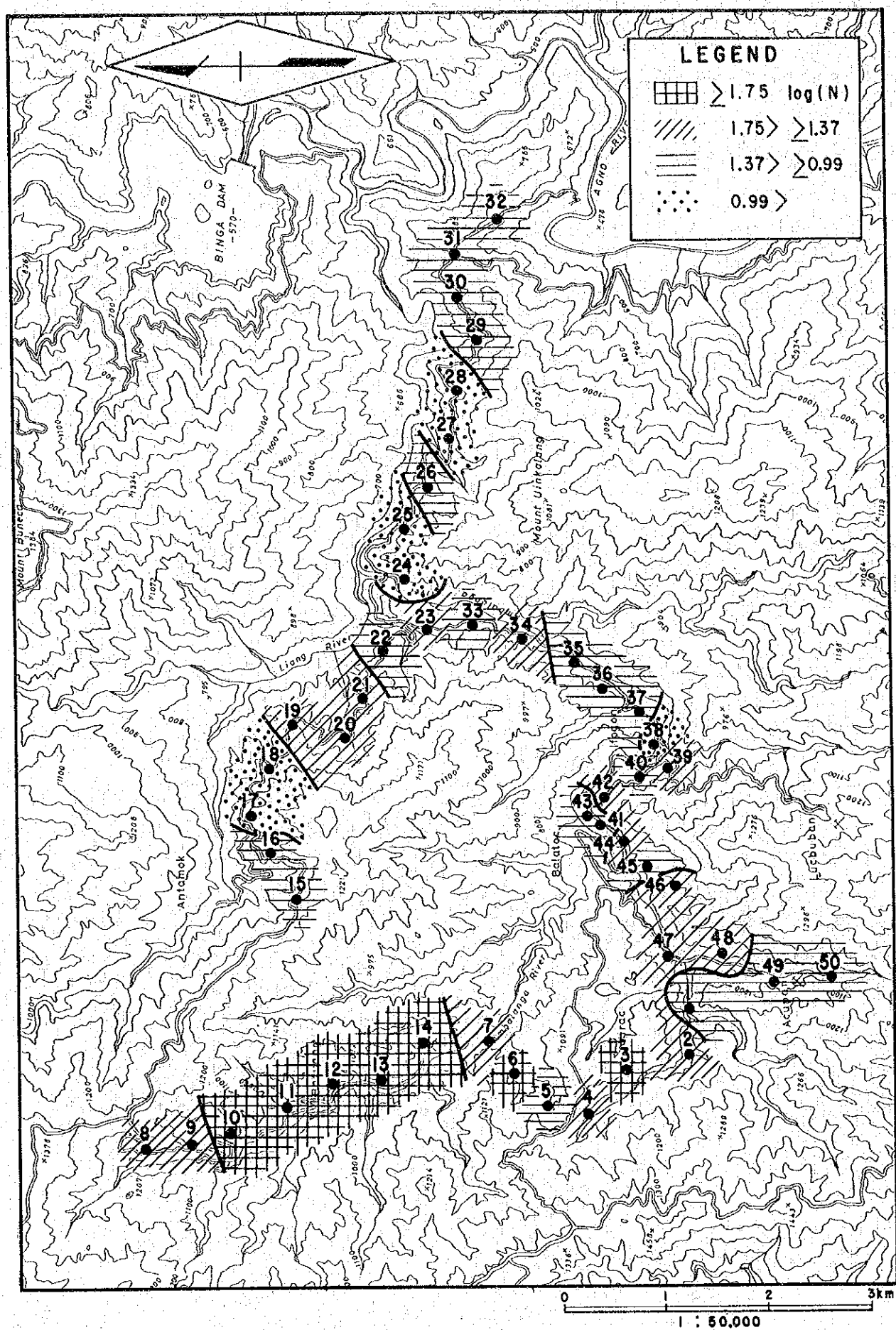


Fig. II-2-13

Distribution Map of Rn Gas Concentration



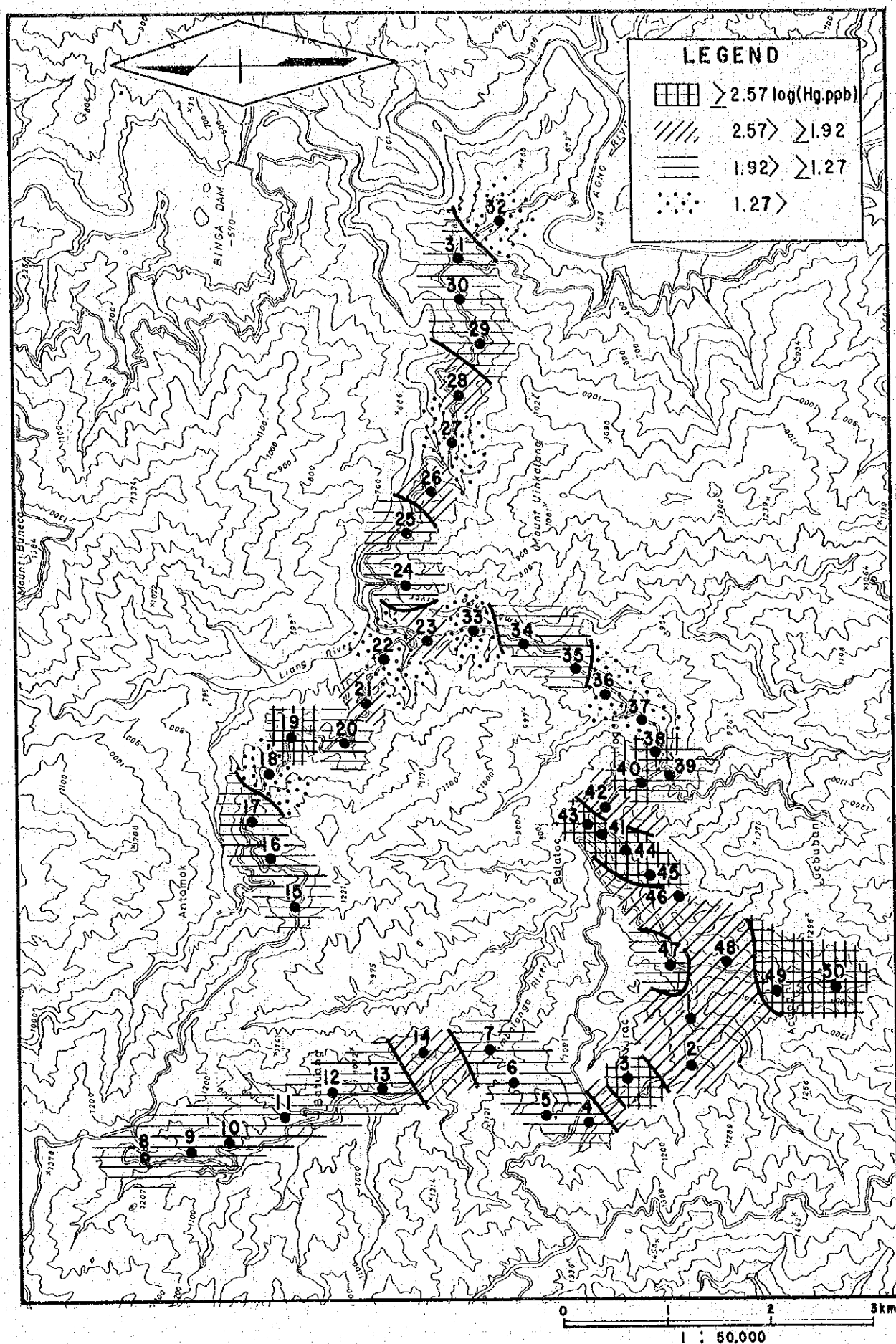


Fig. II-2-16 Distribution Map of Hg Contents in Soil



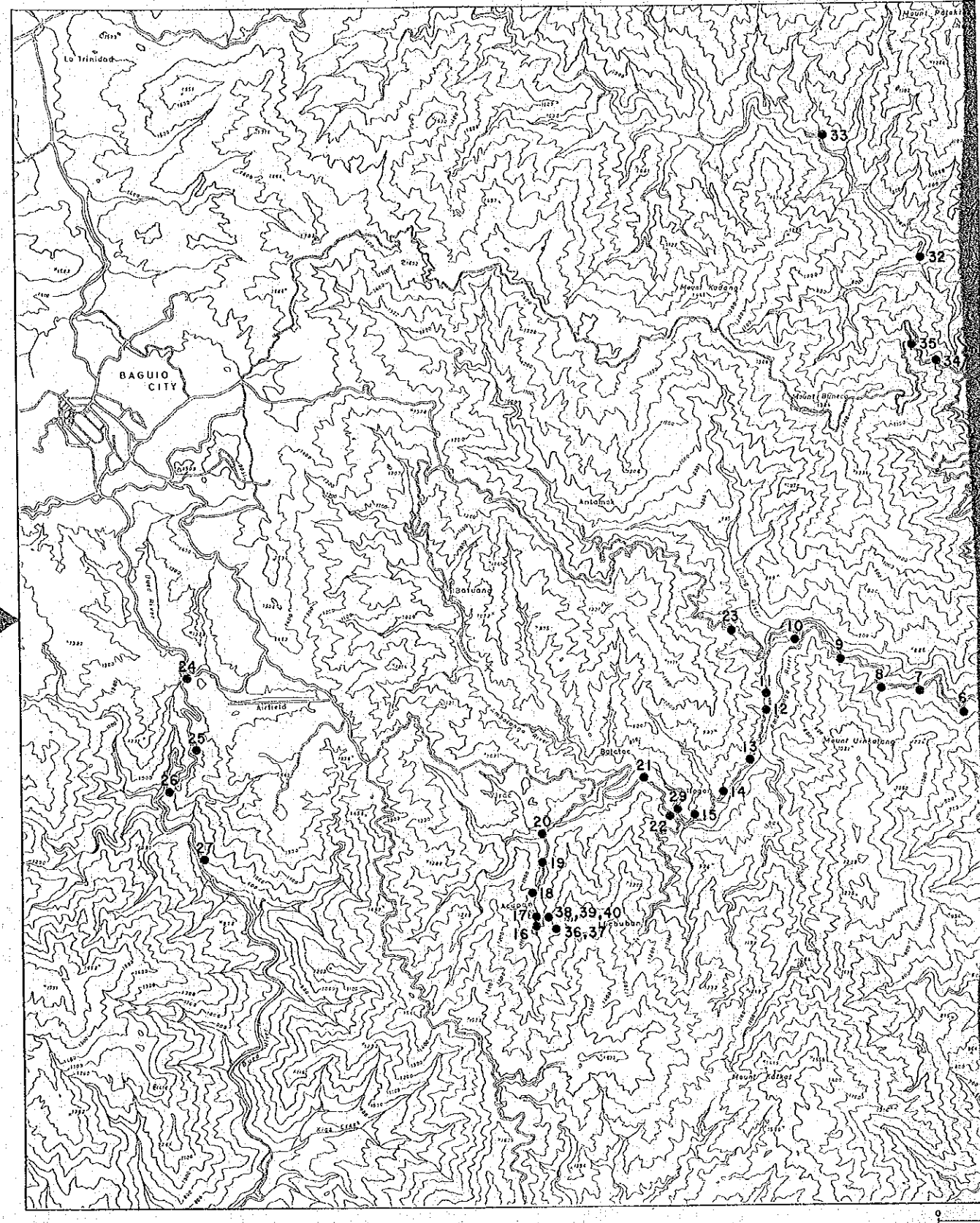
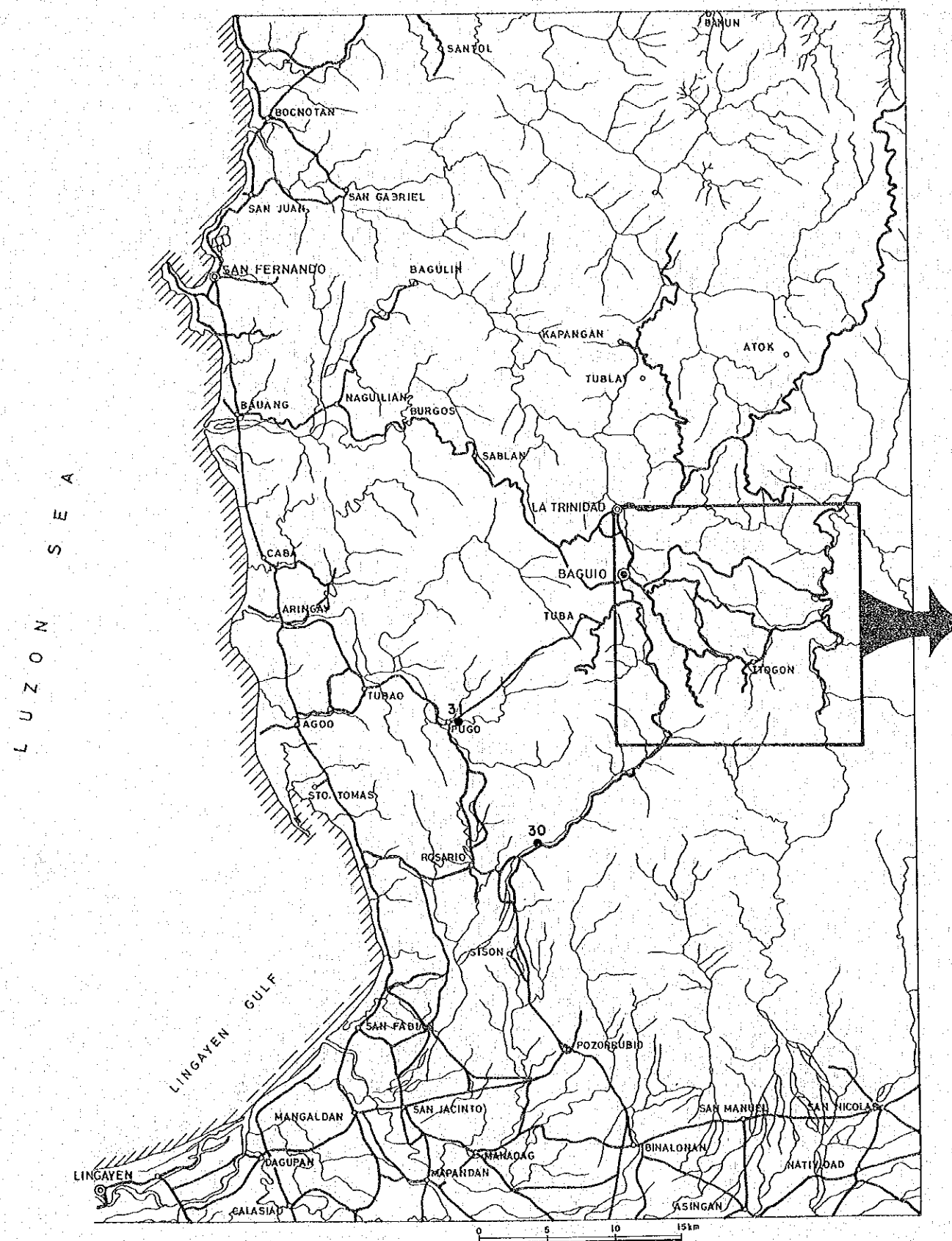


Fig. II-2-17 Location Map of Rock Samples for Hg Analy





Table II-2-2 shows whether a given sample analyzed for mercury is altered or fresh. Sampling locations are shown on Fig. II-2-17. With reference to Table II-2-2, the following are the average Hg content of host-rocks of hot springs;

Dalupirip	:	25 ppb (mean of no. 1, 2 and 3)
Itogon	:	1,078 ppb (no. 29)
Klondyke	:	19 ppb (no. 30)
Pugo	:	269 ppb (no. 31)
Laboy	:	21 ppb (mean of no. 32 and 33)
Acupan	:	156 (mean of no. 36, 37, 38, 39 & 40).

It can be noted from the above data that host rocks from the Itogon Bridge and Pugo hot springs and from Acupan Mine show particularly high mercury contents. The frequency

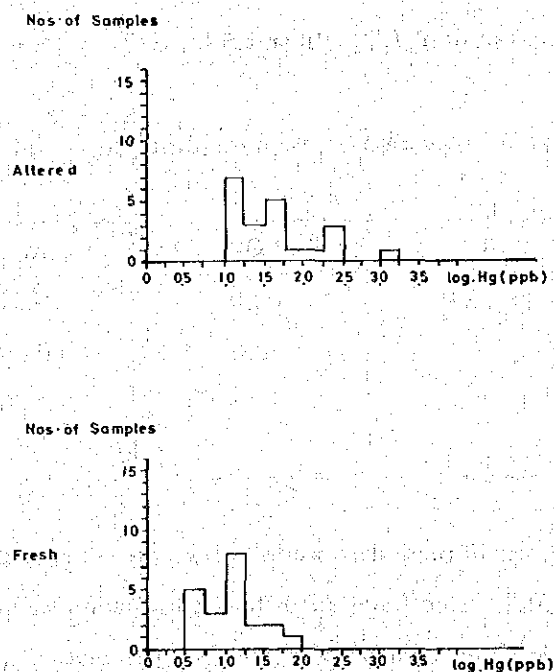


Fig. II-2-18 Histograms for Hg Contents in Rocks

diagrams of mercury contents of altered samples on one and of unaltered ones on the other, are shown on a logarithmic scale in Fig. II-2-18. The minimum duplication appears on both diagrams between 1.00 and 1.25, but in most cases the Hg content is higher in altered samples than in fresh ones. It is not sure whether this alteration is linked with the hydrothermal alteration associated with the gold mineralization in the area, or whether it is connected with the hot spring alteration. For the time being, we can consider it to be related to the post-mineralization hot spring activity. As a matter of fact, relatively high Hg content (43 and 34 ppb) in no. 21 and 22 are found at Itogon Bridge even in altered rocks that are not the hot spring host rocks, whereas, in the Acupan Mine, rocks that underwent strong mineralization-alteration (shown by abundant pyrite and chalcopyrite) have relatively low Hg contents (12–24 ppb, no. 16 to 20). This is also the case for sample no. 14 from Ambalanga River. On the other hand, the host rocks of some hot springs, for example the Laboy River hot spring, show low Hg contents. This reveals that the geothermal fluids perhaps did not always cause an enrichment in mercury.

#### 2-2-5 Factor Analysis of T, CO<sub>2</sub>, Rn and Hg in Soil

The coefficients of correlation between the different parameters are as follows:

	T	CO <sub>2</sub>	Rn	Hg
T	1.00	-0.20	0.24	0.19
CO <sub>2</sub>		1.00	0.29	-0.18
Rn			1.00	0.07
Hg				1.00

The number of sets of these data seems to be quite small to allow statistical considerations. The correlation coefficients between the different parameters are characteristically very small, too. However, the degree of connection of the four parameters may be tested using, for convenience sake, the following criteria:

0.7 to 1.0 (-0.7 to -1.0): rather strong connection

0.4 to 0.7 (-0.4 to -0.7): good connection

0.2 to 0.4 (-0.2 to -0.4): slight connection

0.0 to 0.2 (-0.0 to -0.2): really no connection

As the radon (Rn) content shows correlation of 0.24 and 0.29 with temperature (T) and carbon dioxide (CO<sub>2</sub>) content respectively, there are slight connections between Rn and T and

Table II-2-2 Analytical Data Sheet of Rock Samples for Hg Analysis

Sample No.	Location	Geological formation	Rock name	Alteration	Hg contents (ppb)	Remarks
1	Daluping	Pugo	andesitic meta-volcanics fine grained, schistose	deep green, altered hydrothermally, Py. diss.	37	host rock for discharging hot spring
2	"	"	"	silicified strongly, quartz vein rich.	27	"
3	"	"	"	altered hydrothermally, calcite vein rich.	12	"
4	"	"	andesitic vol. breccia	quartz and calcite vein rich.	5	
5	"	Aguno batholith	hornblende diorite	fresh, although hb. is altered to chlo. partially.	10	
6	along S. Liang	Pugo	meta-sediments of coarse basic pyroclastics	fresh, dark gray in colour.	7	
7	"	"	meta-sediments of fine basic pyroclastics	fresh, deep gray in colour.	40	
8	"	Basement	green schist	fresh.	5	
9	"	"	meta-sediments of basic pyroclastics	fresh, although Py. imp. imp.	12	
10	"	"	actinolite schist	silicified strongly.	45	
11	along S. Ambalanga	Aguno batholith	hornblende diorite	fresh.	29	
12	"	"	"	fresh, although hb. is altered to chlo. and Py. diss. weakly	14	
13	"	"	"	fresh, no Py. and hb. is fresh.	12	
14	"	"	"	altered by ore mineralization. hb. is altered to chlo. & bio.	10	
15	"	"	"	fresh, although hb. is altered to chlo. partially. no Py.	10	
16	Acupan	Baliboc plug	quartz diorite	altered by mineralization. calcite, Py. Cp. rich. hb. is fresh.	12	
17	"	Aguno batholith	quartz diorite	altered. Py. & Cp. rich. hb. is altered to chlo. & bio. partially.	24	
18	"	"	hornblende quartz diorite	altered. Py. diss. & stringer rich. hb. is altered to actinolite.	12	
19	"	"	hornblende diorite	altered. Py. stringer rich. hb. is altered to chlo.	14	
20	"	"	"	altered. much Py. stringer & diss. hb. is altered to chlo.	14	
21	Itogon	Zig. Em.	andesitic lapilli tuff	altered. argillified.	45	
22	"	"	andesite dyke	altered. chloritized & quartz veined. much Py. stringer & diss.	34	
23	Antamok	Aguno batholith	hornblende diorite	fresh. Py. diss. weakly and bearing malachite.	7	
24	Kannon Rd.	Klondyke	auto-brecciated andesite	fresh. grayish green in colour. chlo. & calcite rich.	5	
25	"	"	andesitic tuff breccia	fresh. deep green in colour. chlo. & calcite rich.	24	
26	"	"	diorite	fresh. hb. is altered to chlo. partially.	7	
27	"	Klondyke	andesite coarse tuff	fresh. calcite rich.	5	
28	along S. Liang	Pugo	chlorite schist	fresh. calcite veinlet network and bearing malachite.	5	host rock for discharging hot spring
29	Itogon	Baliboc plug	dacite breccia	altered. Py. diss. with calcite veinlet network.	19	"
30	Klondyke	Klondyke	tuffaceous sandstone	fresh. dark gray in colour.	269	"
31	Pugo hot-spring	Rosario	"	altered. chloritized and much Py. diss.	17	"
32	Laboy	Aguno batholith	hornblende quartz diorite	altered to white. calcite veinlet rich.	24	"
33	"	"	"	altered. hb. is altered to chlo. & epi. calcite & Py. rich.	12	
34	"	"	hornblende diorite	fresh. greenish gray. hb. is black in colour. no Py.	19	
35	"	"	micro hornblende diorite	fresh. grayish green. hb. is fresh. no Py.	161	
36	Acupan	"	diorite	silicified strongly. hb. is altered to chlo. Py. diss.	82	
37	"	"	andesite	silicified strongly.	190	host rock for discharging hot spring
38	"	Baliboc plug	dacite breccia (fine)	altered. pale green. chlo. & calcite rich.	48	"
39	"	"	" (coarse)	altered. pale green chlo. & calcite rich	298	"
40	"	"	" (older)	altered. pale green. chlo. & calcite rich		"



between Rn and CO<sub>2</sub>. There is also a slight negative correlation between T and CO<sub>2</sub> content. However, there is really no correlation between Rn content and Hg content. Using the four variables: T, CO<sub>2</sub> and Rn content in soil air, and Hg content in soil, a principal factor analysis can be made although the number of parameters and sets of data are both small. Nevertheless, the factor analysis, within the range covered by the data, allows examination of the advantages and problems of the different methods used in the geochemical exploration.

The Pearson integrated correlation factors are obtained from the raw data, and the following results are derived after the Varimax rotation for factor loading (See also the geographical representation on Fig. II-2-19).

	first factor	second factor
T	0.530	0.120
CO <sub>2</sub>	-0.349	0.491
Rn	0.202	0.572
Hg	0.400	-0.049

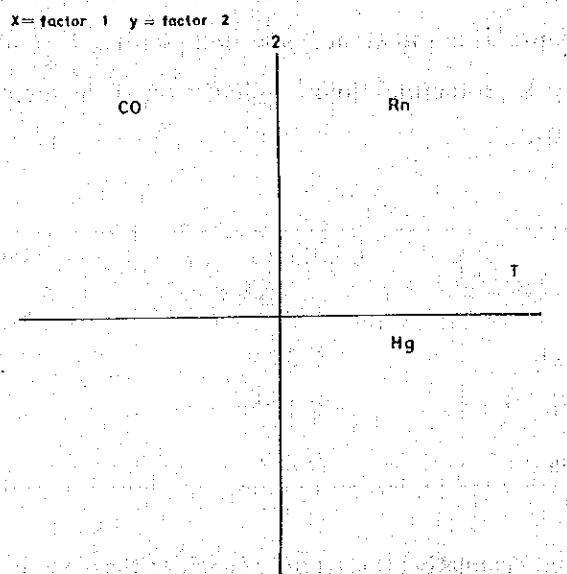


Fig. II-2-19 Factor Loading for Variables

This shows the T and the Hg content as first factors, and the CO<sub>2</sub> and Rn content in soil air as second. The contribution of the factors are similar (0.604 for the first factor and 0.585



for the second), and this may be explained by the general dilution of the values taken by each variable. However, analysis of each variable yields the following communalities:

T	0.296
CO <sub>2</sub>	0.363
Rn	0.368
Hg	0.162

All these values are small, especially those concerning the T and the Hg content. This means that there is a strong influence of some peculiarity outside the first and second factors, as previously seen in the examination of the degree of connection of one variable with respect to the others. In conclusion, judging from the factor analysis of the 50 sets of four variables used in the survey, there may be some doubts about the efficacy of some of these factors as exploration indices. Conversely speaking, the positive or negative correlation between the different variables probably become inexplicit because the geothermal manifestations at the surface are not very strong.

As regards the significance of this factor analysis, although many problems remain unsolved, the T and the Hg content apparently act as first factors and, as the distribution maps of these two parameters (Fig. II-2-5 and II-2-16) shows a good correlation with hot springs, we may say that these reflect the movements of geothermal fluids at depth. On the other hand, since the CO<sub>2</sub> and Rn contents in soil air act as second factors, these parameters do not seem to be related to the activity of geothermal fluids. Calculation of the factor weights of each parameter yields the following results:

	first factor	second factor
T	0.387	0.091
CO <sub>2</sub>	-0.274	0.376
Rn	0.167	0.443
Hg	0.271	-0.031

Using these data, and calculating the scores of each of these factors for every measurement point, the distribution maps of Fig. II-2-20 and II-2-21 showing factor scores are obtained.

Without forgetting the limitations already pointed out earlier, it may be inferred that the first factor corresponds to the activity of geothermal fluids, and the second one to the presence around Batuang of andesites and granodioritic porphyries.

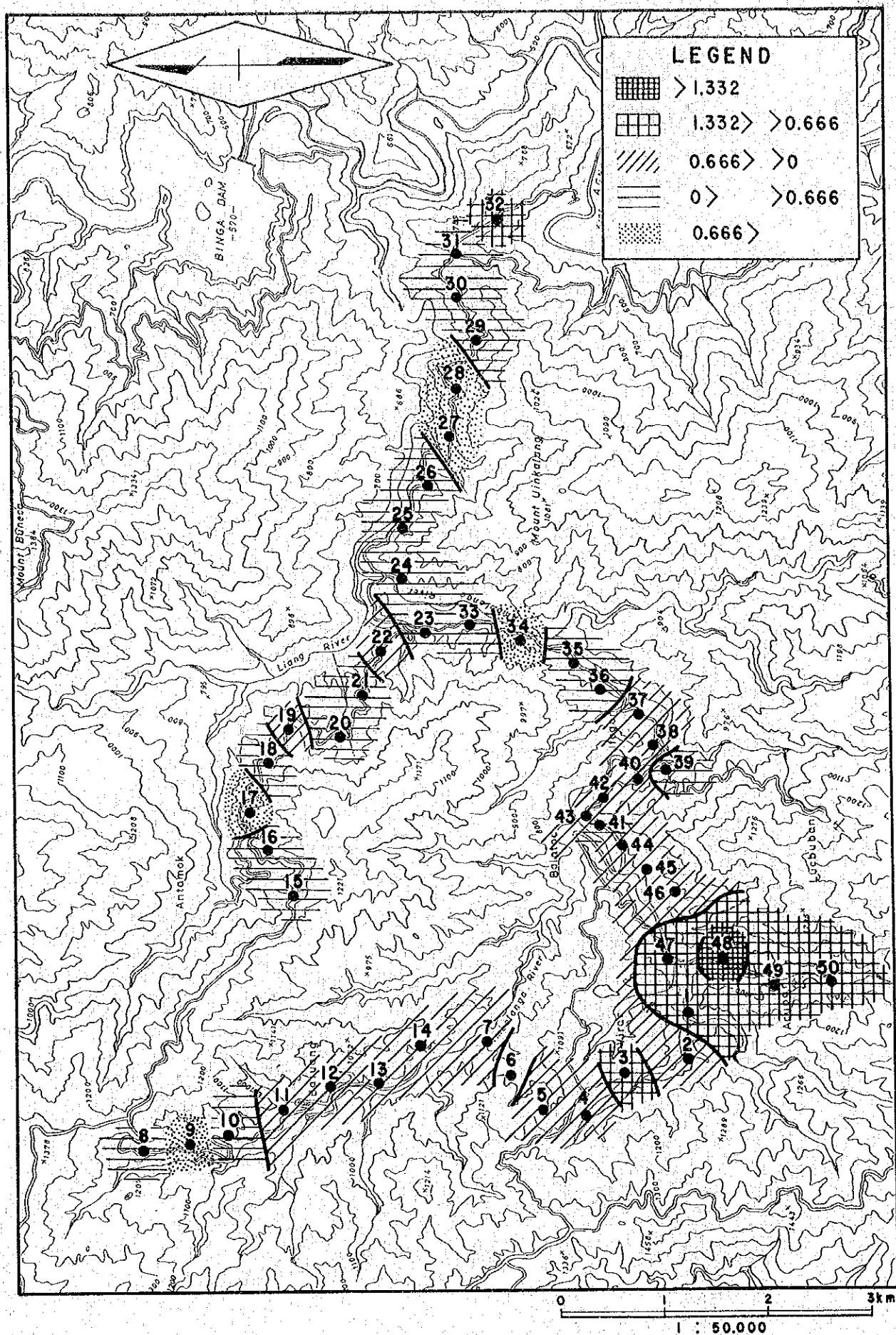


Fig. II-2-20 Distribution Map of Factor Score (Factor-1)



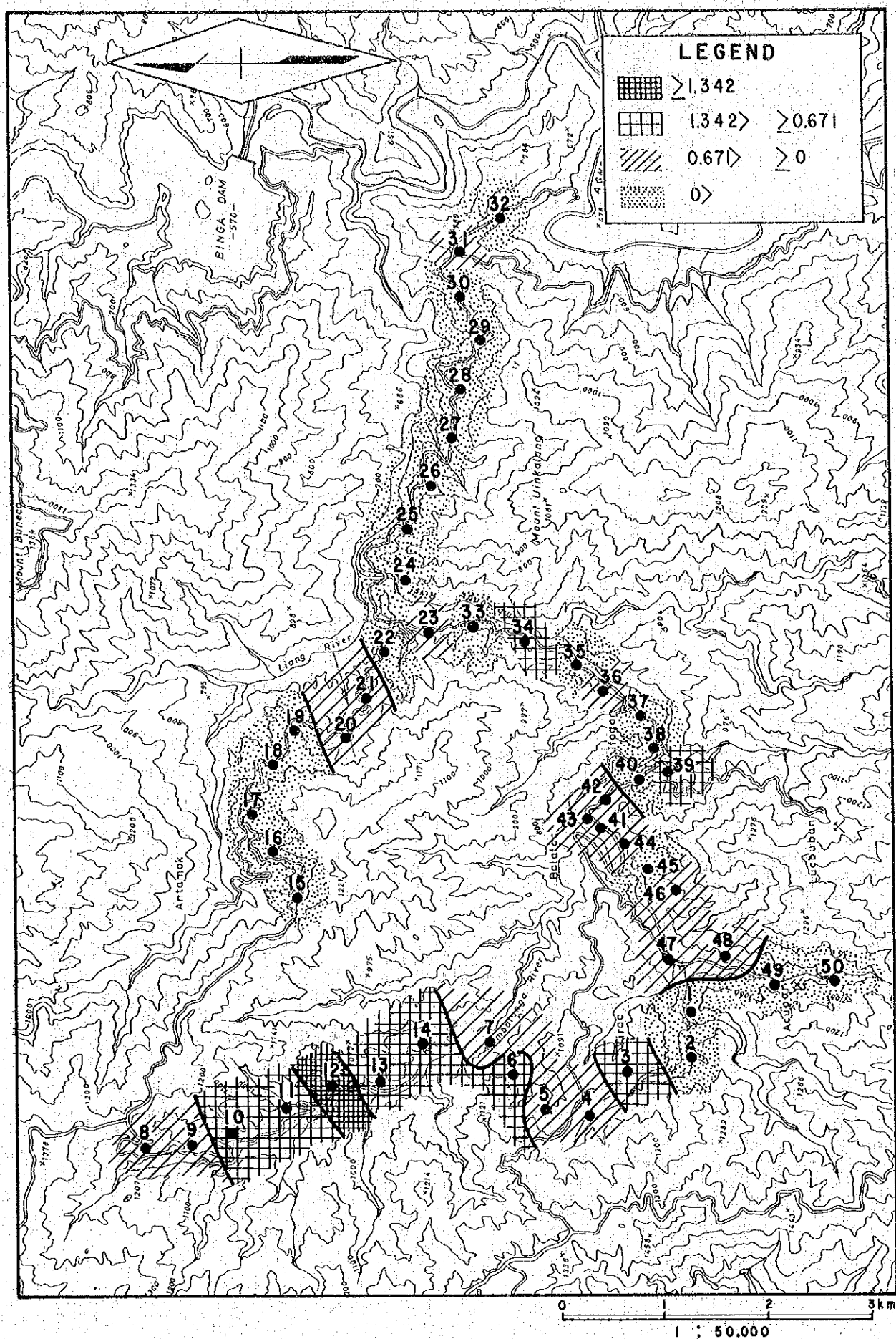


Fig. II-2-21

Distribution Map of Factor Score (Factor-2)



#### 2-2-6 Measurement of T, CO<sub>2</sub>, Rn and Hg in Soil

Of the four parameters considered in the geochemical exploration at one meter depth, the temperature and the mercury content in soil reflect the activity of geothermal fluids, whereas the other two variables which do not seem to be directly linked with the activity, reflect geological differences. Concerning the connection between mercury content and mineralization, the relation is not very clear between gold-silver mineralization and mercury content. However, within the limits of the sampling method, may be disregarded any relation between mercury content and copper mineralization. It is generally observed that for copper and gold-silver mineralizations occurring in a given area at a given period of time, the gold-silver stage, in many cases, postdated the copper mineralization. The mercury enrichment, likewise probably occur in the later stage of a given series of geothermal fluid activity. The reason for this may be the evolution of the chemical characteristics of the associated intrusive and extrusive rocks. In any case the distribution of mercury contents reflect the whole activity of geothermal waters accumulated at depth over a long period of time.

Carbon dioxide and radon contents, which were measured at intervals of about 500 m reflect geological differences rather than geothermal fluid activity. It is possible, however, that other correlations will appear if the intervals between sampling points will be reduced and the superficies of the surveyed area be reflected.

Unfortunately, the activity of underground geothermal fluids does not act very strongly on the four parameters measured at the surface. This causes limitations on the effectiveness of these parameters with regards to the understanding of underground processes. For interpretative investigation of movements of hydrothermal fluids at depth, measurements not only of the first factor, namely the temperature and the mercury content, but also the carbon dioxide and radon concentrations should be made. However, concerning the measurement of carbon dioxide content, as the soils of the surveyed area are poorly developed, it is difficult to obtain at each point the same condition for measurement of the gas concentration, and it is perhaps sufficient, for this geochemical exploration at one meter depth, to measure the temperature, the mercury content in soil and the radon concentration.

#### 2-3 Geochemical Survey of the Hot Spring Water

During this survey, 10 hot spring water samples were taken from seven (7) locations: Dalupirip, Itogon, Klondyke, Asin, Pugo, Laboy and Acupan mines. Gas samples were also taken from exhalations. Chemical and isotopic analyses were done on the water samples. The locations of the hot spring are shown in Figure II-2-22; and summary is given in Table II-2-3.

### 2-3-1 The Source Regions of Hot Springs

The Dalupirip hot spring ascends from within the boulders below the Agno River. The source area measures approximately 6m x 7m where three (3) horizontal pipes were constructed within cubicles for drainage of the hot spring. The sampled water was taken from the drainage pipe having the highest temperature (42.5°C). The total discharge was observed to be 150 l/min., temperature ranges from 37.2°C to 42.5°C and field pH of 5.5. Found under the sedimentary pebbles along the Agno River are green andesitic altered pyroclastic rocks, which suffered strong pyritization and silicification but no bleaching was observed. Many quartz veins and calcite veinlets were also observed as well as white crystals of gypsum and green crystals of iron sulfate.

Itogon hot spring is located at 150 m to 300 m above Itogon bridge. The host rock could be the Balatoc plug or the andesites that intrude it. There are three main springs in the area where temperature ranges from 41.3°C to 89.5°C, pH 6.5 to 7.5, and discharge at approximately 10 l/min. Two samples were taken from the springs having the highest temperatures. The host rock, where spring water discharges along its joint surfaces, suffered grey green alteration with calcite veinlets. Besides gypsum and iron sulfates, there are also large amounts of calcite sinters. Hot spring gas bubbles were observed at the spring with the highest temperature. Of the 87 ml of the gas collected, only 0.05 ml contained H<sub>2</sub>S while 86.95 ml contained CO<sub>2</sub>.

Klondyke hot spring is located near downstream of Bued River. The host rocks consist of andesitic pyroclastic and conglomeratic formations. Numerous calcite veins were also observed. The hot waters probably discharges along a NE-SW trending fault zone and along the associated joint surfaces. Temperature ranges from 44°C to 49°C and pH is 7.5. White to yellow precipitates of acicular sulfur crystals are locally observed. Collection of gas bubbles were done using gas detection tube and showed that CO<sub>2</sub> concentration is less than 0.05% and H<sub>2</sub>S is below detection limit.

Asin hot spring is located at approximately 10 km from the city of Baguio. The locality was reconstructed so that the thermal discharge flow to a 4m x 8m area of the fluvial sediments and accumulates into a pond-like reservoir, thus no discharge rate was taken. Measured temperature is 73.8°C and pH is 8.0. Gas bubbles in the pond were collected and analyzed using gas detector and showed that CO<sub>2</sub> and H<sub>2</sub>S concentration are below detection limit, i.e. less than 0.05% and 0.005%, respectively. No precipitates of sulfur were observed.

Pugo hot spring is located at the left bank of the Pugo River. The fluid discharges at two locations from within sedimentary fluvial sand. Temperature is from 35.3°C to 36.3°C while





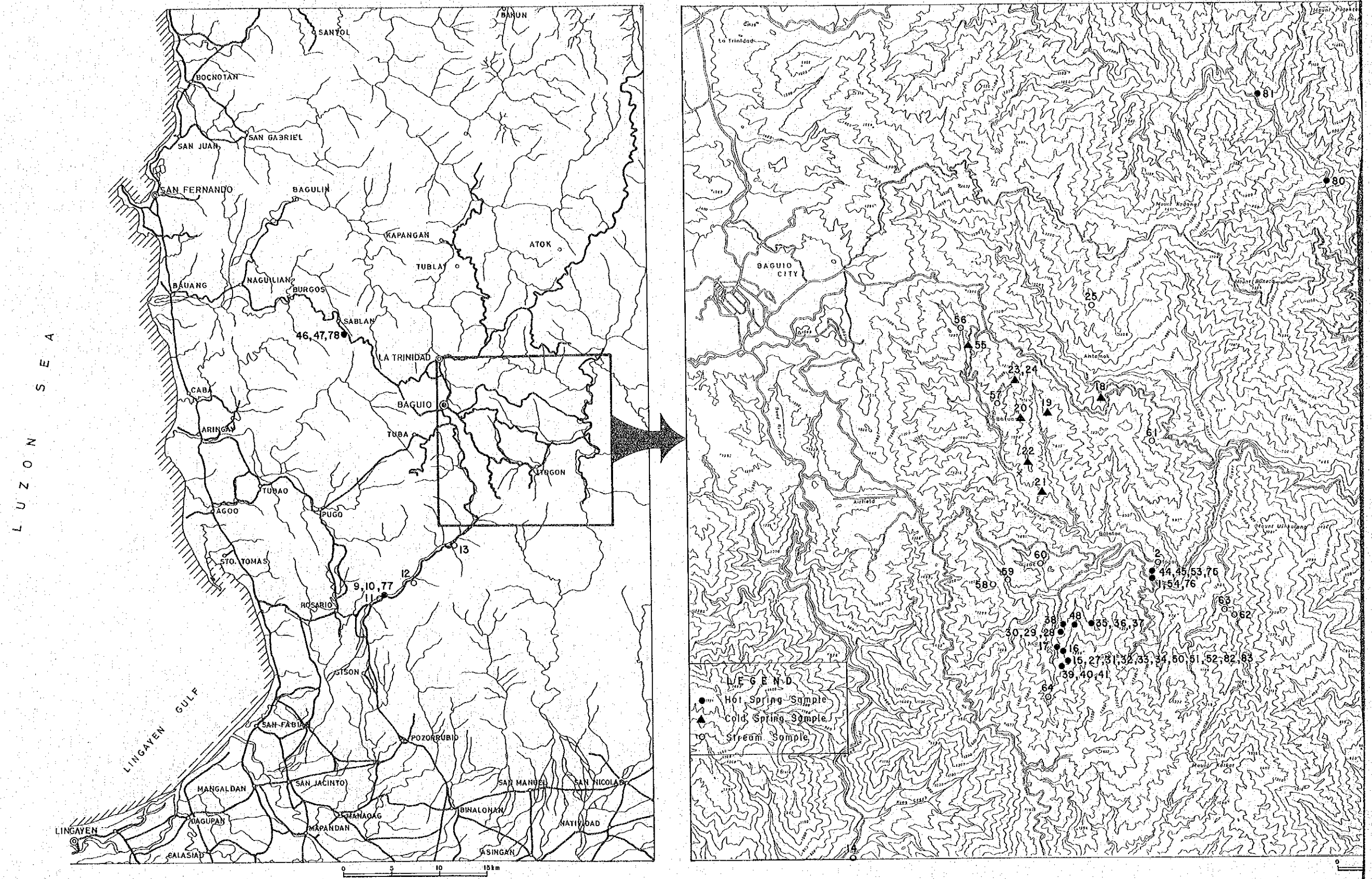


Fig. II -2-22 Location Map of Spring Water Samples





Table II-2-3 List of Spring Water Samples

SAMPLE CODE	HOT SPRING LOCATION	ELEVATION M. a.s.l.	TEMPERATURE (°C)	pH	DISCHARGE VOL. (L/min.)	%V GAS		DEPOSIT	COUNTRY ROCK/GEOLOGY
						CO <sub>2</sub>	H <sub>2</sub> S		
DA	Dalupin	420	42.5	5.5	> 50	—	—	calcite & quartz vein, CaSO <sub>4</sub> & FeSO <sub>4</sub> films	gravels on Pugo formation
IT-1	Itogon	680	89.5	7.5	5	99.94	0.06	calcite veinlet, calcite sinter, gypsum-like crystals, FeSO <sub>4</sub>	Balatoc Plug along the contact between lower Balatoc Plug and overlying detritic
IT-2			62.2	7.5	5	—	—		
KL	Klondyke	240	49.5	7.5	10	<0.05	<0.005	calcite vein	andesitic tuffaceous sandstone (Rosario Formation)
AS	Asin	260	73.8	8.0	not sure	<0.05	<0.005	none	gravels
PU	Pugo	80	36.3	7.0	5-6	<0.05	<0.005	none	gravels on tuffaceous sandstone (Rosario formation)
LA-1	Laboy	660	47.5	7.0	10	—	—	calcite thin veinlet along jointing	hb. qtz diorite
LA-2		700	47.5	7.0	15	—	—	calcite epidote veinlet;	granite
BA-1	Balatoc (Acupan Mine)	Acupan mine 3300 level (519m)	81.0	7.5	25	—	—	gypsum and calcite vein	diorite
BA-2		3150 level (565m)	62.1	8.0	60-80	?	?		andesite













No.	23	24	25	26	27	28	29	30	31	32	33
Location	Batuang	Batuang	Antemok	Antemok	Acupan	Acupan	Acupan	Acupan	Acupan	Acupan	Acupan
Code	E24	E25	E26	Antemok	Benquet-1	Benquet-2	Benquet-3	Benquet-4	Benquet-5	Benquet-6	Benquet-7
Sampling Date	7, June, '82	7, June, '82	9, June, '82		19, Aug., '77	18, Aug., '77	19, Aug., '77	19, Aug., '77	18, Aug., '77	19, Aug., '77	19, Aug., '77
T (°C)	19.0	18.0	17.5	37.0	79.5	99	80	72.5	64.5	55	73
pH	5.51	5.30	6.79	3.35	7.84	8.36	-	8.12	8.05	7.80	7.25
Conductivity (µl/cm)	518	647	258	4.653							
Discharge (l/min)	1	0.5	20		40		50	2~3		20	60
Chem. Ana.	mg/l	mg/l	mg/l	mg/l	mg/l	mg/l	mg/l	mg/l	mg/l	mg/l	mg/l
Na <sup>+</sup>	12.6	0.548	5.10	15.659	866	37.669	1.077	43.062	856	37.234	41
K <sup>+</sup>	0.61	0.016	0.60	0.570	117	2.992	158	3.530	122	3.120	4.7
Ca <sup>2+</sup>	40.0	1.996	32.0	28.942	121	6.038	57	67	210	438	408
Mg <sup>2+</sup>	28.8	2.370	7.75	94.0	5.4	0.444	2.0	2.2	8.3	11.2	8.5
Cl <sup>-</sup>	7.09	0.200	2.36	613	1.202	33.904	1.405	35.202	1.260	35.540	12
SO <sub>4</sub> <sup>2-</sup>	230	4.789	52	45.786	392	8.162	540	10.306	410	29.983	29.566
HCO <sub>3</sub> <sup>-</sup>	16.5	0.270	83.9	1.375	157	2.573	-	198	336	87	85
CO <sub>3</sub> <sup>2-</sup>					2	0.067	-	4	5	1	0
H <sub>2</sub> CO <sub>3</sub>					6	0.097	-	4	8	3	10
Li				2.40	4.2	7.0	5.9	5.6	4.6	0.05	0.07
As											
B				10.60	47	4.348	55	5.088	47	<0.05	0.96
SiO <sub>2</sub>	34.1	0.568	21.8	47.6	177	2.946	236	3.928	167	26	28
NH <sub>4</sub> <sup>+</sup>											
T.S.M.(mg/l)	369.7	415.5	205.5	3,926.5	3,092.4	4,154.3	3,530.0	3,409.2	3,429.3	2,054.2	2,018.2
Alkalinity											
Acidity											
Remarks	cold spring	cold spring	stream water	hot spring	hot spring	hot spring	hot spring	hot spring	hot spring	hot spring	hot spring
Reference	B.E.D. (9/Aug/82)	do	do	do	KRTA (Sep/77)	do	do	do	do	do	do



No.	34	35	36	37	38	39	40	41	42	43	44											
Location	Acupan	Acupan	Acupan	Acupan	Acupan	Acupan	Acupan	Acupan	Acupan	Acupan	Irogon											
Code	Benquet-8	Benquet-9	Benquet-10	Benquet-11	Benquet-12	Benquet-13	Benquet-14	Benquet-15	Benquet-16	Benquet-17	Benquet-18											
Sampling Date	19, Aug., '77	20, Aug., '77	20, Aug., '77	20, Aug., '77	20, Aug., '77	20, Aug., '77	20, Aug., '77	20, Aug., '77	20, Aug., '77	20, Aug., '77	19, Aug., '77											
T (°C)	57	65	52	97	63	51.5	—	45	24	43	94											
pH	7.25	8.23	7.89	—	7.57	7.04	7.61	7.75	—	—	7.47											
Conductivity (μS/cm)																						
Discharge(l/min)	5	15	10	<1	8	8		50	200	20,000	5											
Chemt. Ana.	mg/l	m.eq./l	mg/l	m.eq./l	mg/l	m.eq./l	mg/l	m.eq./l	mg/l	m.eq./l	m.eq./l											
Na <sup>+</sup>	23	1.000	845	36.755	942	40.974	1,041	45.281	85	3.697	32	1.392	42	1.827	44	1.914	18	0.783	124	5.394	570	24.793
K <sup>+</sup>	3.3	0.084	118	3.018	124	3.172	139	3.555	7.2	0.184	2.2	0.056	2.8	0.072	3.0	0.077	1.7	0.043	15	0.384	81	2.072
Ca <sup>2+</sup>	425	21.208	50	2.495	98	4.890	119	5.938	312	15.569	505	25.200	592	29.541	550	27.445	274	13.673	445	22.206	86	4.291
Mg <sup>2+</sup>	12.0	0.987	1.7	0.140	7.0	0.576	1.3	0.107	2.9	0.239	11.0	0.905	3.4	0.280	3.8	0.313	12.9	1.062	15.9	1.308	4.0	0.329
Cl <sup>-</sup>	8	0.226	935	26.373	1,142	32.212	1,360	38.361	12	0.338	6	0.169	7	0.197	10	0.282	<5	0.141	151	4.259	397	11.198
SO <sub>4</sub> <sup>2-</sup>	1,400	29.150	740	15.408	622	12.951	488	10.161	1,000	20.821	1,520	31.648	1,500	31.232	1,475	30.711	770	16.032	1,185	24.673	655	13.638
HCO <sub>3</sub> <sup>-</sup>	173	2.835	188	3.081	266	4.360	—	72	1.180	259	4.245	195	3.196	192	3.147	—	—	—	—	—	473	7.752
CO <sub>3</sub> <sup>2-</sup>	0	0.000	5	0.167	3	0.100	—	1	0.033	0	0.000	2	0.067	2	0.067	—	—	—	—	—	2	0.067
H <sub>2</sub> CO <sub>3</sub>	21	0.339	3	0.048	8	0.129	—	4	0.064	58	0.935	10	0.161	6	0.097	—	—	—	—	—	38	0.613
Li	0.05		4.4		4.7		5.2		0.10		0.07		0.05		0.07		0.05		0.61		3.0	
As																						
B	<0.05	0.005	31	2.868	45	4.163	55	5.088	0.05	0.005	<0.05	0.005	0.32	0.030	<0.05	0.005	<0.05	0.005	7.8	0.722	24	2.220
SiO <sub>2</sub>	26	0.433	140	2.330	153	2.546	177	2.946	46	0.766	36	0.599	34	0.566	32	0.533	30	0.499	32	0.533	137	2.280
NH <sub>4</sub> <sup>+</sup>																						
T.S.M.(mg/l)	2,091.4	3,056.7	3,410.0	3,380.3	1,542.2	2,429.3	2,388.5	2,317.9	1,111.7	1,975.7	2,467.0											
Alkalinity																						
Acidity																						
Remarks	hot spring	hot spring	hot spring	hot spring	hot spring	hot spring	hot spring	hot spring	hot spring	hot spring	hot spring	mine drainage	mine drainage	mine drainage	mine drainage	mine drainage	mine drainage	mine drainage	mine drainage	mine drainage	mine drainage	hot spring
Reference	KRTA (Sep./77)	do	do	do	do	do	do	do	do	do	do	do	do	do	do	do	do	do	do	do	do	do



No.	45	46	47	48	49	50	51	52	53	54	55
Location	Iogon	Asin	Asin	Acupan	Acupan	Acupan	Acupan	Acupan	Iogon	Iogon	Batuang
Code	Benquet-19	Benquet-20	Benquet-21	BED-1	BED-2	BED-3	BED-4	BED-5	BED-6	BED-7	BED-8
Sampling Date	19, Aug., '77	21, Aug., '77	21, Aug., '77	25, Oct., '78	25, Oct., '78	25, Oct., '78	25, Oct., '78	25, Oct., '78	25, Oct., '78	25, Oct., '78	26, Oct., '78
T (°C)	85	70	41	96	87	61	59	79	86	75	20
pH	7.37	8.78	7.43	8.5	6.8	8.8	7.8	7.7	6.3	6.0	2.7
Conductivity (μm/cm)				4.115	-	4.478	1.757	4.267	1.470	2.373	638
Discharge (l/min)			2.5	60	<6	30	very low	40	20	10	30
Chem. Ana.	mg/l	mg/l	mg/l	mg/l	mg/l	mg/l	mg/l	mg/l	mg/l	mg/l	mg/l
Na <sup>+</sup>	355	523	122	783	-	775	8.134	32.623	183	376	8.67
K <sup>+</sup>	56	11	1.7	125	-	88.2	0.473	88.2	36.1	62.2	0.57
Ca <sup>2+</sup>	152	222	64	12.2	-	234	11.277	85.7	103	54.3	47.4
Mg <sup>2+</sup>	16.0	N.D.	5.0	1.5	-	10	0.823	7.31	17.2	7.81	16.6
Cl <sup>-</sup>	233	835	200	984	-	1,440	5.754	31.309	417	191	24.2
SO <sub>4</sub> <sup>2-</sup>	670	475	108	407	-	449	12.160	5.663	560	481	447
HCO <sub>3</sub> <sup>-</sup>	347	26	170	382	-	292	2.065	168	244	387	-
CO <sub>3</sub> <sup>2-</sup>	0	0.067	0								
H <sub>2</sub> CO <sub>3</sub>	34	0.548	21								
Li	1.8	0.2	0.1	3.36	-	4.11	0.48	4.55	1.28	2.50	-
As				2.25	-	2.50	-	4.0	0.4	1.0	-
B	9.5	0.879	3.3	38.9	-	40.9	5.26	41.8	3.57	13.7	-
SiO <sub>2</sub>	103	32	32	267	-	210	1.964	244	172	195	81.3
NH <sub>4</sub> <sup>+</sup>				0.05	0.38	0.046	0.035	0.031	0.042	0.036	0.040
T.S.M (mg/l)	1,975.5	2,135.5	727.0	3,000.6		3,539.1	1,471.4	2,767.0	1,735.9	1,768.0	625.7
Alkalinity											
Acidity											
Remarks	hot spring	hot spring	hot spring	hot spring	hot spring	hot spring	hot spring	hot spring	hot spring	hot spring	cold spring
Reference	KRTA (Sep./77)	do	do	BED-ELC (Feb./79)	do	do	do	do	do	do	do



No.	56	57	58	59	60	61	62	63	64	65	66
Location	Batuang	Batuang	Virac	Virac	Virac	Balinguay	Itogon	Itogon	Acupan	Acupan	Acupan
Code	BED-9	BED-10	BED-11	BED-12	BED-13	BED-14	BED-15	BED-16	BED-17	76-B-3	76-B-4
Sampling Date	26, Oct., '78	26, Oct., '78	26, Oct., '78	26, Oct., '78	26, Oct., '78	30, Oct., '78	1, Nov., '78	1, Nov., '78	31, Oct., '78		
T (°C)	20	19.5	20	20	20	20	20	20	20	63.7	46.1
pH	3.9	4.2	7.3	8.2	8.0	8.0	7.4	7.4	7.2		
Conductivity (µS/cm)	532	247	275	224	183	279	104	134	870		
Discharge (Q/min)	6,000	6,000	600	6,000	6,000	300	400	400	500		
Chem. Anal.	mg/l	mg/l	mg/l	mg/l	mg/l	mg/l	mg/l	mg/l	mg/l	mg/l	mg/l
Na <sup>+</sup>	10.0	0.435	7.33	0.319	4.0	0.174	4.67	0.203	6.33	0.275	5.67
K <sup>+</sup>	0.86	0.022	0.43	0.011	0.29	0.007	0.43	0.011	0.29	0.007	0.43
Ca <sup>2+</sup>	12.7	0.634	13.0	0.649	42.6	2.126	30.4	1.517	20.0	0.998	47.8
Mg <sup>2+</sup>	8.75	0.720	9.31	0.766	5.88	0.484	9.69	0.797	7.63	0.628	8.13
Cl <sup>-</sup>	24.2	0.683	27.8	0.784	3.55	0.100	4.43	0.125	3.55	0.100	4.43
SO <sub>4</sub> <sup>2-</sup>	284	5.913	115	2.394	86.6	1.803	50.9	1.060	33.3	0.693	25.3
HCO <sub>3</sub> <sup>-</sup>	-	-	-	-	71.6	1.173	80.8	1.324	77.1	1.264	145
CO <sub>3</sub> <sup>2-</sup>	-	-	-	-	-	-	-	-	-	-	-
H <sub>2</sub> CO <sub>3</sub>	-	-	-	-	-	-	-	-	-	-	-
Li	-	-	-	-	-	-	-	-	-	-	-
As	-	-	-	-	-	-	-	-	-	-	-
B	-	-	-	-	-	-	-	-	-	-	-
SiO <sub>2</sub>	74	1.232	59	0.982	24.2	0.403	27.4	0.456	42.1	0.701	49.6
NH <sub>4</sub> <sup>+</sup>	0.040	0.050	0.015	0.140	0.015	0.015	0.015	0.015	0.015	0.015	0.015
T.S.M (mg/l)	414.5	231.7	238.7	208.7	190.3	285.5	149.7	809.0	877.6	4,478.2	1,726.2
Alkalinity											
Acidity											
Remarks	stream water	stream water	stream water	stream water	stream water	stream water	stream water	stream water	stream water	hot spring	hot spring
Reference	BED-ELC (Feb./79)	do	do	do	do	do	do	do	do	Sawkins F.J. et al., (1979)	do





No.	67	68	69	70	71	72	73	74	75	76	77
Location	Acupan	Irogon	Antamok	Antamok	Antamok	Acupan	Acupan	Dalupinip	Irogon	Irogon	Klondyke
Code	76-B-5	76-B-11	76-B-12	76-B-14	76-B-15	76-B-26	76-B-27	DA	IT-1	IT-2	KL
Sampling Date								24, Sep., '82	24, Sep., '82	29, Sep., '82	25, Sep., '82
T (°C)	46.1	87.9	51.7	40.6	46.1	61.1	Cold	42.5	89.5	62.2	49.5
pH								7.90	7.68	7.79	8.36
Conductivity ( $\mu\text{S}/\text{cm}$ )											
Discharge (l/min)											
Chem. Ana.	mg/l	mg/l	mg/l	mg/l	mg/l	mg/l	mg/l	mg/l	mg/l	mg/l	mg/l
Na <sup>+</sup>	30	1.305	26.968	1.620	70.465	167	7.264	120	5.220	920	40.017
K <sup>+</sup>	4.3	0.110	8.6	0.220	107	2.737	6.1	0.136	4	0.102	110
Ca <sup>2+</sup>	509	25.399	13.2	0.659	565	28.194	1.000	49.900	753	37.575	77
Mg <sup>2+</sup>	4.25	0.350	8.3	0.683	92.0	7.570	20.5	1.687	19.1	1.572	14.1
Cl <sup>-</sup>	4.4	0.124	516	14.554	1.890	53.310	828	23.355	302	8.518	1,270
SO <sub>4</sub> <sup>2-</sup>	1,370	28.525	480	9.994	34.563	1,360	28.317	1,240	25.818	87	1,811
HCO <sub>3</sub> <sup>-</sup>	162	2.655	522	8.555	886	14.521	590	9.670	798	13.079	470
CO <sub>3</sub> <sup>2-</sup>											
H <sub>2</sub> CO <sub>3</sub>											
Li	<0.01	2.80									
As											
B	0.12	0.011	2.35	0.217	4.83	0.447	0.64	0.059	0.52	0.046	3.95
SiO <sub>2</sub>	42	0.699	191	3.179	143	2.380	82.3	1.370	69.5	1.157	254
NH <sub>4</sub> <sup>+</sup>											
T.S.M. (mg/l)	2,126.1	2,361.5	6,967.8	4,054.5	3,306.1	3,206.1	407.2	345	1,436	1,686	1,548
Alkalinity											
Acidity											
Remarks	hot spring 2,600L	hot spring	hot spring 1,850L	hot spring 1,850L	hot spring 1,550L	hot spring 3,150L	cold spring 1,500L	hot spring	hot spring	hot spring	hot spring
Reference	Sawkins F.J. et al (1979)	do	do	do	do	do	do	BED-JICA (1982)	do	do	do







pH is from 5.5 to 7.0. Total discharge is estimated at about 10 to 15 l/min. The water samples was taken from a well-like reservoir where hot spring gas bubbles were observed. The detection tube shows that this gas contains less than 0.05% CO<sub>2</sub> and less than 0.005% H<sub>2</sub>S. The rocks outcropping in the bottom of the river at 6 to 7 meters beneath the hot spring are middle grained pyroclastic sandstones, which are chloritized and pyritized.

There are three occurrences of hot springs in the Laboy River. One at upstream, one at downstream, and one 100 m above the downstream. Near the upstream and downstream, the water discharges along the joint surfaces of an amphibole-bearing quartz-diorite which suffered a general silicification with transformation of amphibole into epidote or chlorite and development of numerous calcite veinlets. The pyritization is particularly strong in the host rocks of the upstream hot spring. For each of these two hot springs, the discharge is around 15 l/min, temperature range is 47.5 to 48.6°C while pH is 7.0. No gas bubbling was observed in the area. However, the rocks around are covered by a relatively thick deposit of white crystals of gypsum. At about 100 m upstream from the downstream spring, hot water discharges from the pebbles of the bottom of the river and mix with the river water. Temperature is 38.5°C, pH is 7.0 and discharges at a few l/min.

Hot waters are emitted at several places in the galleries of Acupan Mine. Water sampling was done at levels 3,300 and 3,150. For sample at 3,300 level, the discharge was running down a concrete channel after its springing from a drilling hole. Temperature is 81.0°C, pH is 7.5 and discharge rate is 25 l/min. Ambient temperature was 66°C. Geologically, the area consists of dioritic rocks intruded by andesitic dykes. The pale green amphibole is altered to chlorite; the rocks are strongly pyritized and silicified. On the ventilation pipes, which are in contact with the gallery walls, a large amount of white crystal resembling gypsum has deposited. At the 3,150 level, the fluid also discharges from a drilling hole but flows directly to the bottom of the gallery, where it was sampled. The andesite is particularly silicified in this area. Discharge rate is 60 to 80 l/min, temperature is 62.1°C, and pH is 8.0. Ambient temperature was 62.0°C. At this level, gas is also emitted together with the hot water but sampling was impossible due to extreme heat which made breathing difficult.

### 2-3-2 Chemical Composition of the Hot Spring Waters

The results obtained on the samples collected during this survey are shown on Table II-2-4 together with those reported by B.E.D. (Aug. 9, 1982), KRTA (Sept. 1977), B.E.D.-E.L.C. (Feb. 1979) and F. J. Sawkins et. al. (1979). The sampling locations of the present survey are shown on Fig. II-2-22.

Key-diagrams of the chemical compositions of the hot spring waters are shown on Fig. II-2-23 for Acupan and Asin hot springs, on Fig. II-2-24 for Itogon, Antamok, Dalupirip, Laboy, Klondyke and Pugo hot springs and on Fig. II-2-25 for stream waters and cold waters. On the basis of these key diagrams, the following characteristics were obtained for the hot springs of the different areas:

Acupan Mine: Two types of water composition can be distinguished: the  $\text{Cl}^-$  type, which is rich in chloride ions, and the  $\text{SO}_4^{2-}$  type, rich in sulfate ions. This difference in anionic composition affects the cation proportions of the waters such that the  $\text{Cl}^-$  type is rich in  $\text{Na}^+$  and  $\text{K}^+$ , and the  $\text{SO}_4^{2-}$  type is rich in  $\text{Ca}^{2+}$ . A close relationship between this chemical difference and the space distribution of the springs was found, i.e. the  $\text{Cl}^-$  type waters are restricted to the Balatoc plug and the  $\text{SO}_4^{2-}$  type waters are distributed outside the plug (Fig. II-2-26).

Itogon: As regards the proportion of anions, the  $\text{SO}_4^{2-}/\text{HCO}_3^- + \text{CO}_3^{2-}$  ratio fluctuates around 3/2 and the  $\text{Cl}^-$  from 0 to 50%. For the cations,  $\text{Na}^+ + \text{K}^+$  accounts for 60% to 100%, the rest being  $\text{Ca}^{2+}$  and very small concentrations of  $\text{Mg}^{2+}$ . The anionic composition is similar to that of the  $\text{SO}_4^{2-}$  type of Acupan Mine but since the  $\text{Na}^+ + \text{K}^+$  concentrations are high, the waters are also similar to the  $\text{Cl}^-$  type. The Itogon hot spring waters may thus be considered as waters of the  $\text{Cl}^-$  type of Acupan Mine correlatively enriched in  $\text{SO}_4^{2-}$  and  $\text{HCO}_3^- + \text{CO}_3^{2-}$  (Fig. II-2-24).

Antamok: Most of the waters have  $\text{SO}_4^{2-}$  as dominant anion and  $\text{Ca}^{2+}$  as dominant cation, thus resembling the  $\text{SO}_4^{2-}$  type of Acupan Mine. Sample no. 69 (76-B-12), however, is rich in  $\text{Cl}^-$  and  $\text{Na}^+ + \text{K}^+$  and thus similar to the  $\text{Cl}^-$  type waters of Acupan Mine. This sample was taken from the 1,980 level.

Laboy: Only two samples were analyzed, but in both cases,  $\text{SO}_4^{2-}$  is the dominant anion and  $\text{Ca}^{2+}$  is the dominant cation;  $\text{Na}^+$  and  $\text{K}^+$  are absent. Both samples clearly belong to the  $\text{SO}_4^{2-}$  type of Acupan Mine (Fig. II-2-24).

Dalupirip: The principal anion is  $\text{SO}_4^{2-}$ , with small amounts of  $\text{HCO}_3^- + \text{CO}_3^{2-}$ , while the principal cations are  $\text{Na}^+ + \text{K}^+$ . This composition is similar to that of the Itogon spring waters (Fig. II-2-24).

Asin, Klondyke and Pugo: All these samples have  $\text{Cl}^-$  rich and  $\text{Na}^+ + \text{K}^+$  rich compositions.

As shown in Fig. II-2-25, the stream waters and cold waters have  $\text{Cl}^-$  as anion and  $\text{Ca}^{2+} + \text{Mg}^{2+}$  as cations.

The chemical compositions of hot spring waters may thus be classified according to the relative proportions of the anions  $\text{SO}_4^{2-}$ ,  $\text{Cl}^-$  and  $\text{CO}_3^{2-} + \text{HCO}_3^-$  on one hand and the cations  $\text{Na}^+ + \text{K}^+$ ,  $\text{Mg}^{2+}$  and  $\text{Ca}^{2+}$  on the other.

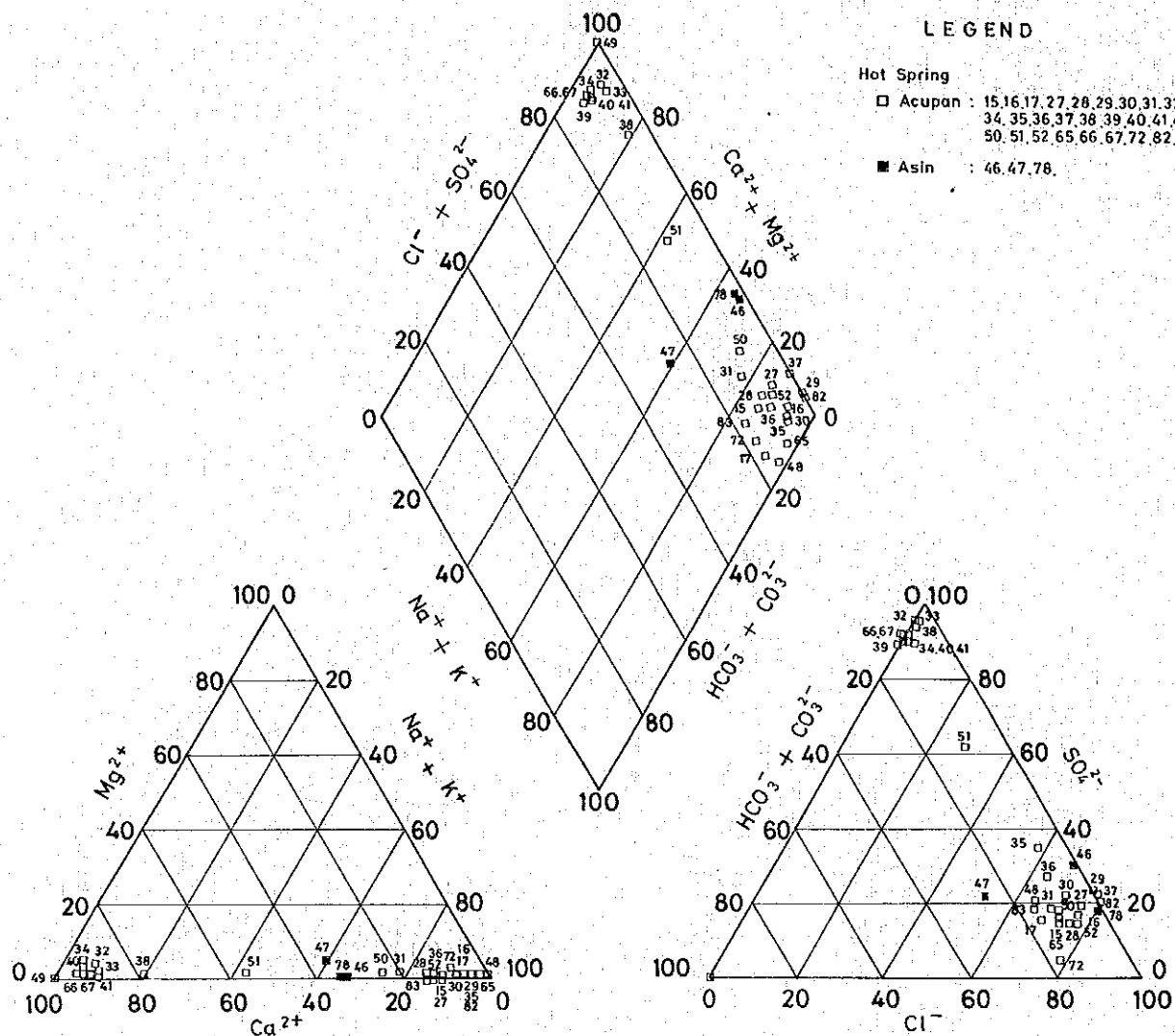


Fig. II-2-23

Key-Diagram for Acupan and Asin Hot Spring





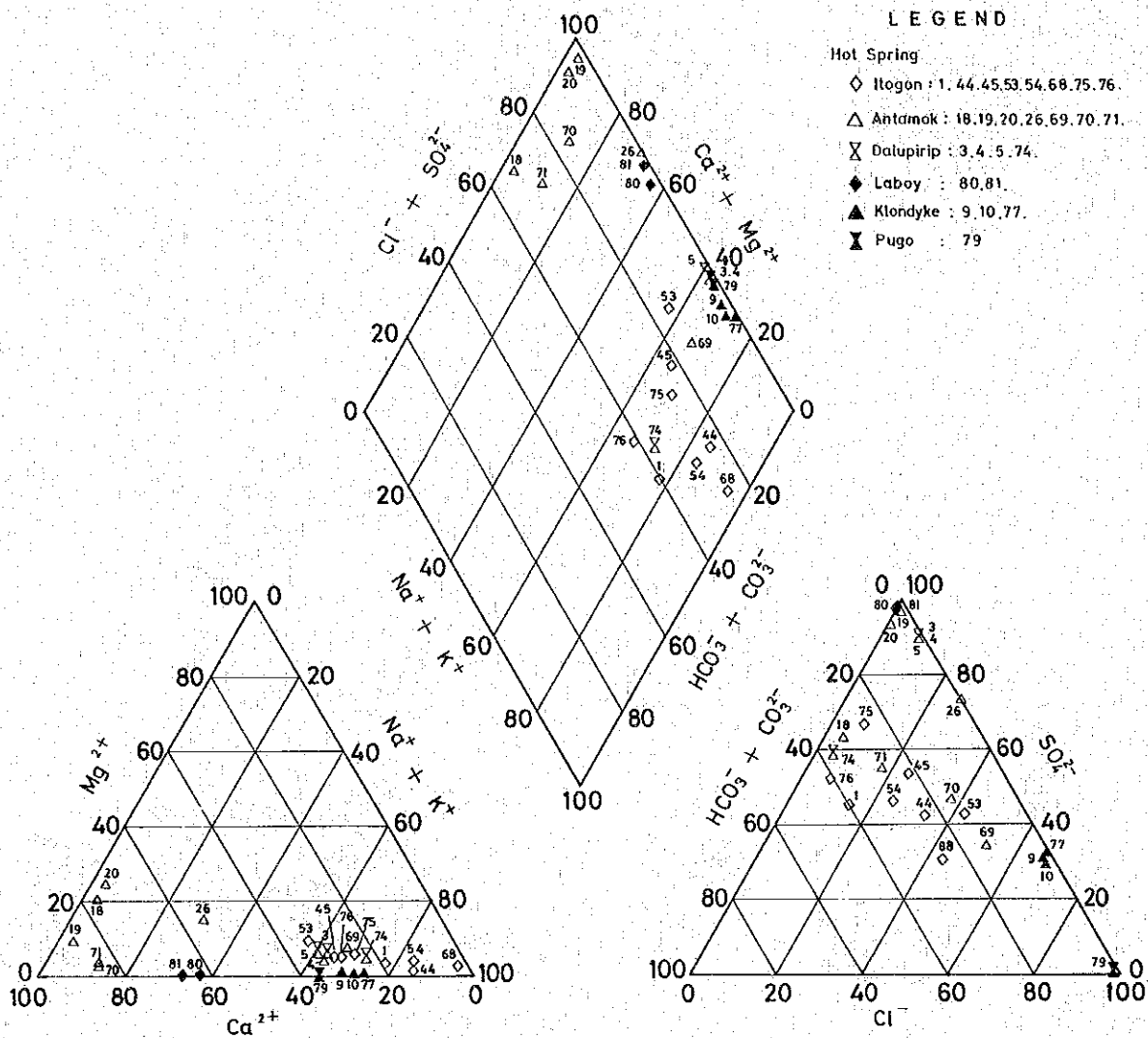


Fig. II-2-24

Key-Diagram for Itogon, Antamok, Dalupirip, Laoby, Klondyke and Pugo Hot Spring



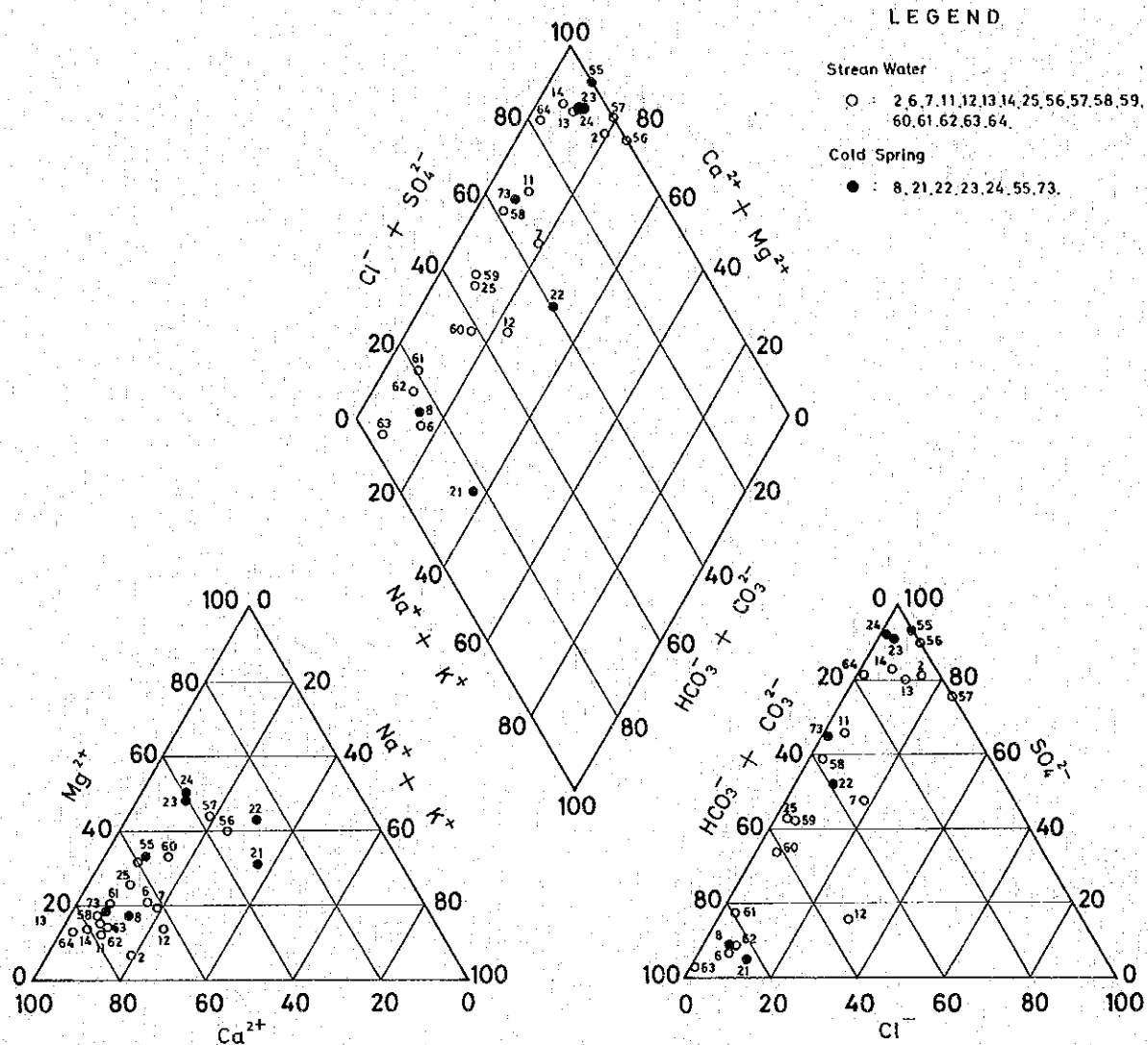


Fig. II-2-25

Key-Diagram for Stream Water and Cold Spring



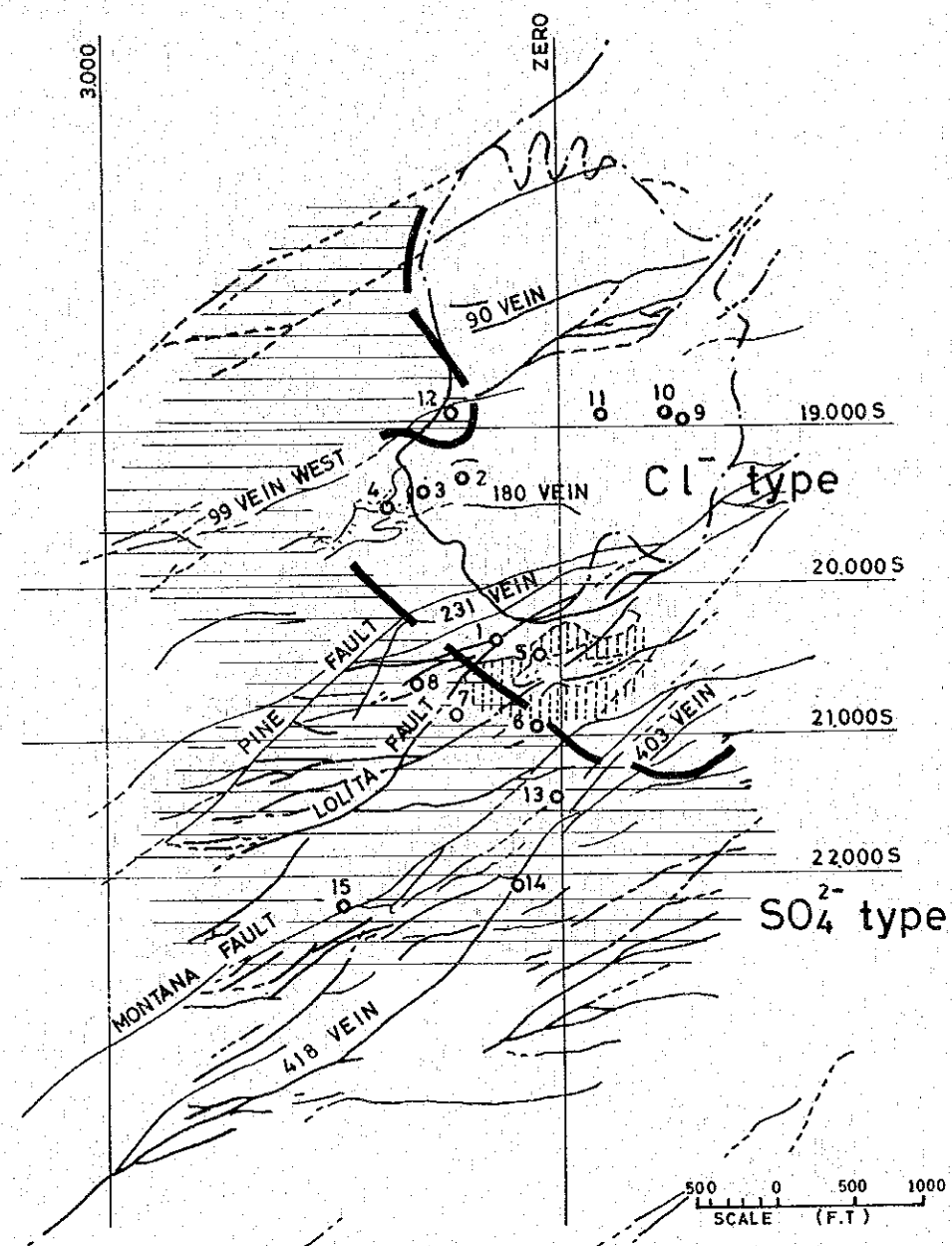


Fig. II-2-26

Chemical Zoning of Spring Water Samples at Acupan Mine

