### A-3-3 Evaluation of February 1982 Flood

A flood hydrograph has been observed by JICA study team during the second phase field investigation (Nov. 6, 1981-Mar. 31, 1982) at the point of Mulayyinah (S-1) situated about 500 m upstream from the point S-2. Figure-3 present the observed data of rainfalls and discharges which are computed by use of Manning's formula on the basis of measured river stages and cross-section. Under the assumption that n = 0.038, which is commonly adopted in PAWR's estimates, the peak flood discharge has been estimated at 770 cu.m/sec.

#### The following are reviewed from observations:

- No precedent rain has been observed in the drainage basin before 13th February. Accordingly, rains during Feb. 13 and up to 5 a.m of Feb. 14 are considered to be initial losses of rains, which are consumed to saturate surface layers of drainage basin.
- After once soil layers have been saturated with rain water, sudden floodings, characterized by a rapid rise, sharp peak and rapid recession, have been observed. Two peaks of the flood are likely caused by series of rainfalls, whose peaks were observed at 8 a.m and 2 to 3 p.m.
- Although large variations in rain intensity (and hence runoff rate) during a storm are reflected in the shape of the resulting hydrograph, the time scale of intensity variations depends mainly on basin size. Based on the conceptions of a unit hydrograph, <u>flood concentration time</u>

of 3 hours as read on Figure-3 could be applicable to any other given intensity of storm rainfalls.

If the former of two sharp peak hydrographs is extracted from observations so as to represent the typical shape of flood hydrograph within the drainage basin under consideration, a ratio of peak discharge to the volume of flood is given as:

A Company of the Company

Ratio = 
$$\frac{\text{Peak Discharge}}{\text{Flood Volume}} = \frac{770 \text{ cu.m/sec.}}{7.7158 \text{ MCM}} = 9.975 \times 10^{-5} (\text{sec}^{-1})$$

In connection with the shape of a typical flood hydrograph in Oman, the PAWR's report on "Surface Water Gauging Station Network of the Sultanate of Oman (PAWR 83-2 Report)" mentions as follows:

"A flood hydrograph is the record of a flood event from the beginning to the end; in a desert hydrograph, the discharge rises sharply from no flow to some peak value and then drops off again, returning somewhat more slowly to zero flow. The shape of the hydrograph will vary from basin to basin and from storm to storm, but within an area that is geographically similar, hydrographs will be similar in shape."

- The coefficient of peak flood runoff, fp, is also calculated as:

$$fp = \frac{770 \text{ cu.m/sec x 3.6}}{654 \text{ sq.km x 10.5 mm/hr}} = 0.40$$

Where, concentration time of peak food is taken at 3 hours after inspection of observed pattern of rainfall and hydrograph, and rainfall intensity during this period in terms of areal average is seen as 10.5 mm/hr against the second peak of rain and discharge, after the surface layers of drainage basin have once been saturated by rain water.

The same procedure was employed to evaluate the flood observed at the river mouth (S-4) for the same duration of February 1982. The findings are then summarized as below:

- Peak Flood Discharge: Qmax = 163 cu.m/sec
- Flood Volume : V = 4.03 MCM
- Ratio of Peak Discharge to Flood Volume

$$= \frac{163 \text{ cu.m/sec}}{4.03 \text{ MCM}} = 4.045 \times 10^{-5} \text{ (sec}^{-1}\text{)}$$

## A-3-4 Maximum Probable Flood

### 1. Maximum Probable Rainfall

Rainfall measurements are obtainable from six raingauge stations installed within and the vicinity of the drainage basin of the proposed dam for the period of 11 years from 1974 up to 1984. Areal rainfalls computed by using Thiessen method are tabulated in Tables A-3-4 and A-3-5. Since the observation period is too short, these data are unusable for the foundation of statistical assessment. Therefore, the statistical evaluation of rainfall records should be to a great extent based on the data collected at Muscat, only where a long-term rain records are available since the late 1800's.

Table A-3-8 summarizes the rain records collected from Muscat. A long-term fluctuation of annual rains is visualized as shown in Figure A-3-7. From this, the recent years after 1974, during which rain data are available within the project boundary, are not always abundant in rain with exception of the year 1976.

In areas along the Northeast coast of Oman involving the project area and Muscat, most of annual rainfall is concentrated during winter season from November to April. In addition, in many years, more than 80% of annual rain concentrates in a particular month of the year. Table A-3-9 presents characteristics of maximum rainfalls observed in the Northeast coast area.

As a long-term rain record, only monthly rain data are obtainable from Muscat for a relatively long period of about 90 years. To characterize the maximum probable rainfalls of a relatively short duration, reation of monthly maximum and daily maximum rains to annual rains are extracted from all of available

data and plotted as given in Figure A-3-8. Envelope curves so drawn to enclose all the points would show the maximum probable monthly or daily rainfalls when a reasonable estimate of annual rainfall is made.

Probability study on annual amount of rainfall at Muscat gives 255 mm of annual rainfall with a recurrence interval of 50 years.

With regard to regional difference of annual rains, the dainage basin of the proposed dam has been receiving about 1.4 times of annual rainfall reported at Muscat.

## Maximum Probable Rainfall

- Annual rainfall of 50-years return period = 255 mm/yr.\*/
- Maximum probable annual rainfall at Muscat = 255/06.6 = 425 mm/yr. = 430 mm/yr.

(Note: A ratio of provisional estimate of the maxmimum probable flood to the 50-years return period flood, 1:0.6, is referred to in the "Report on Water Resources Study, Phase II, and Technical Proposal for Construction of Water Recharge Projects for the Government of Oman: Ministry of Agriculture and Fisheries.")

<sup>\*/</sup> The rainfall of 255 mm per year was decided from a curve which plotted annual rainfall records observed at Muscat from 1894 to 1983 on a logarithmic probability paper. (The rainfall record refer to Table A-3-8 "Monthly Rainfall at Muscat".)

(see Figure A-3-8) =  $430 \times 0.66$ 

= 284 mm/mo.

- Maximum probable daily rainfall at Muscat (see Figure A-3-8) =  $430 \times 0.40$ 

= 172/mm/day

- Maximum probable monthly rainfall within the proposed basin = 284 x 1.4 = 398 mm/mo.
- Maximum probable daily rainfall within the proposed basin = 172 x 1.4 = 241 mm/day

### 2. Maximum Probable Flood

Many empirical formulas have been presented for estimating flood peaks from relatively small basins. For the most part, these formulas are misleading and unsound and they can be discussed only in a general way, because all of them contain coefficients or constants which can be evaluated only where sufficient volume of storm-flood observations are available.

The most widely used design equation for basins in known as the rational formula. The equation states that the rate of runoff equals to the rate of supply (excess rainfall) if the rain lasts long enough to permit the entire area to contribute. As is indicated that the formula is ideally applicable for catchments of 5 to 10 sq. km. and is commonly used for surface drainage design in urban areas, the above assumption is literally correct for an area of very small area and may not involve serious error for basins up to a few square kilometers in area when the time of concentration is taken correct.

For larger areas, the coeeficient of run-off not only expresses the proportion of the total rainfall which runs off but

also the effect of overland flow and channel storage on the peak. The coefficient is defined as the ratio of the theoretical absolute peak (100% runoff and no storage) to the actual peak, and fortunately this value can be obtained from the actual observation of flood hydrograph collected at Mulayyinah in February 1982. Moreover, according to the advanced studies made by Kadoya and others of Kyoto University in Japan based on actual measurements of hydrologic data collected from rivers of 0.13 to 740 sq.km catchments, it is varified that the rational formula is still applicable to rivers of more than several hundred square kilometers catchments when the strict application of the concentration time of peak flood discharge is accompanied. With an assumption that the application of this formula can still be extended to the range of 812 sp.km.: -

- a) The coefficient of peak flood runoff has been calculated as fp = 0.4
- b) If the maximum probable rainfall of 241 mm/day is distributed within a short duration of the concentration time of 3.0 hours, the average intensity of rainfall is then given as:

$$i = \frac{241}{3.0} = 80.3 \text{ mm/hr}$$

c) The maximum probable flood discharge is thus computed as:

$$Qp = 1/3.6 \times fp \times 1 \times 812 \text{ sq.km}$$
  
= 1/3.6 x 0.40 x 80.3 x 812  
= 7,245 cu.m/sec

# Peak-Volume Relation

A ratio of peak flood discharge to volume of flood runoff, as derived from actual measurement of flood hydrograph is February 1982, was applied to evaluate the maximum probable flood which would be caused by the maximum probable rainfall.

- Flood Volume =  $812 \text{ sq.km} \times 241 \text{ mm} \times 40\%$ = 78.277 MCM

## - Maximum Probable Flood

 $Q_p = \text{Flood Volume } \times \text{Peak Discharge Flood Volume}$ = 78.277 x 10<sup>6</sup> x 9.9795 x <sup>-5</sup>

= 7,828 cu.m/sec

Areal Rainfall at Dam Site (Above Dam: CA = 812 sq.Km)

		: -						-	:	***	*** AREAL RATIO **	*** AR	
127.3	3.5	3.1	8	1.7	10.8	5.4	3.2	5.5	13.1	22.7	47.5	8.8	MEAN
1399.9	39.0	33.8	19.9	19.1	118.6	59.8	35.6	9.09	144.6	249.5	522.0	97.2	TOTAL
11.8	2.5	0.0	0.0	3,3	0.1	6.0	0.0	0.0	0.0	0.0	0.0	0.0	1984
148.6	0.0	1.7	0.0	0.0	38.3	0.4	0.0	0.0	32.3	29.8	42.5	3.6	1983
209.3	8.0	7.9	0.8	0.8	1.4	0.0	0.0	5.4	0.1	41.2	150.9	0.0	1982
47.2	0.0	0.0	0.0	0.0	0.0	12.5	2.4	17.6	10.1	7.7	0.0	3.5	1981
70.5	o. 0.	0.0	0.0	0.0	0.1	10.7	0.0	1.5	0.0	21.8	20.5	6.1	1980
669.0	19.5	1.7	5.6	3.7	0.0	0.9	4.6	0.5	4.3	1.3	1.2	21.6	1979
8.76	0.0	0.0	0.0	2.1	30.0	16.4	1.9	0.0	3.0	4.6	39.9	0.0	1978
160.7	0.0	8	3.9	0.0	0.0	0.0	22.9	33.8	39.0	0.8	25.9	25.5	1977
420.8	6.3	13.5	0.0	4.1	29.5	4.4	1.7	0.0	53.6	142.5	138.4	26.8	1976
94.5	0.0	0.0	0.0	3.0	15.3	1.5	1.9	2.0	0.0	0.7	63.0	7.2	1975
68.6	0.0	0.0	9.6	2.1	4.0	2.0	0.3	0.0	2.3	5.8	39.7	2.9	1974
AINDAL	DEC:	NOV.	OCT.	SEP.	AUG.	ነባቦ፣	JUNE	MAX	APK.	MAR	FEB.	JAN.	YEAR

0.0 FARFAR = 0.291 SOHAR HAYL(WJ) = 0.204 HAYL(WH) = 0.279DAQIQ = 0.095 KITNAH = 0.131

sq.Km)

Areal Rainfall below Dam (CA = 471

							-					TIND	= MM/MONTH
YEAR	JAN		MAR.	APR.	MAY	JUNE	JULY	AUG.	SEP.	OCT.	NOV.	1	ANNUAL
1974	1.4	l	3.5	0.8	0.0	0.0	0.0	1.1	1.0	5.0	0.0		77.2
1975	4.2		0.0	0.0	0.7	0.0	0.0	10.4	0.0	0.0	0.0		67.4
1976	9.5		77.9	58.6	0.0	0.0	0.5	12.1	1.4	5.7	₽.		307.4
1977	56.1		1.8	41.9	22.0	10.1	0.0	0.0	0.0	6.T	16.9		177.1
1978	0.0		5.3	3.0	0.0	0.9	4.6	13.8	0.0	0.0	0.0		65.1
1979	20.3		1.6	3.8	0.0	0.0	0.0	0.0	0.0	28.3	7.5		8.66
1980	2.4		8.7	8.0	0.9	0.0	3.0	0.0	0.0	0.0	9.0		27.8
1981	4.3		1.0	16.5	14.0	0.0	7.6	0.0	0.0	0.0	0		43.4
1982	0.0		52.3	1.4	6.0	0.0	0.0	0.0	0.0	7.0	14.4		217.0
1983	0.8		34.4	26.1	0.0	0.0	0.0	13.3	0.0	0.0	0.8		129.0
1984	0.3		0.0	0.0	0.0	0.0	0.3	0.0	4.7	0.0	0.0		12.9
OTAL	99.3	502.4	186.6	152.9	38.5	10.9	16.0	50.7	7.0	41.3	40.4	78.1	1224.1
AAN	σ	ı	0 / [	13.9	3.5	O L	1.5	4.6	0.6	3.8	3.7	•	111.3

Areal Rainfall at River-Mouth (CA = 1,283 sq.Km)

	15	277	ţ			***		i c	E	77.074	0.00	ANTHITAT
크	EB.	MAR.	APR.	⋖I	JUNE	JULY	AUG.	SEP.	OCT.	NOV.		ANNUAL
7	3.2	5.0	1.7	0.0	0.2	щ СЭ	2.9	1.7	7.9	0.0	9.0	71.8
Ŋ	σ	7.0	0.0	•		•	13.5		0.0	0.0		84.5
57	7	118.7	55.4				23.1		2.1	10.8		379.1
7	Q	1.2	40.1				0.0		3.2	11.9		166.8
(1)	Ò	4.9	3.0				24.1		0.0	0.0		85.8
	0.7	1.4	4.1			•	0.0		13.9	1.7		80.9
,	15.4	17.0	0.3				T-0		0.0	0.2		54.8
	•	1.1	12.4				0.0		0.0	0.0		45.8
$\ddot{\rightarrow}$	47.1	45.2	9.0	3.7	0.0	0.0	6.0	0.5	0.7	10.3		212.1
7	46.6	31.5	30.0				29.1		0.0	1,3		
	•	0.0	0.0			3.9	0.0	٥	0.0	0.0	4.4	12.2
5	14.	226.4	147.6	52.5	26.6	43.7		14.6	27.8	•	53.4	1335.3
	46.8	20.6	13.4	7.8	2.4	4.0	8.5	1,3	2.5	3.3	6.4	121.4
- T-1	AREAL RATIO	* * * O										
H	0.060	KITNAH	= 0.083	HAYL (WJ)	אָנוֹ) - = 0	129	HAYL (WH)	= 0.224	4 FARFAR	0	.249 SOI	SOEAR = 0.255
		α 1. α 1. α	12 1 2 TO	7. 1.								
											TNT.	HINOW/ MM =
i											.	
	I POI	MAR.	APR.	MAY		-	AUG.		OCT.	NOV.	DEC.	ANNUAL
	71.6	2.7	9.0	0.0		•	0.0		0.0	0.0	2.2	77.7
		0	0.0	0.0			6.2		0.0	0.0	0.0	53.3
		51.9	55.2	0.0		•	0.0		8.2	0.0	17.3	253.1
	26.3	2.6	38.7	10.9		•	0.0		0.0	17.6	0.2	164.4
	4	6.5	2.0	0.0			0.0		0.0	0.0	0.3	46.6
	. •	1.5	2.3	0.0		•	0.0		36.6	0.7	58.4	120.9
	2.2	3.9	1.2	0.0	0.0	0.0	0.0	0.0	0.0	8.0	0.0	8.6
		1.5	18.2	12.6			0.0		0.0	0.0	0.0	37.1
		56.3	2.0	0.0		•	0.0		0.0	16.6	10.4	222.5
	ö	37.9	23.7	0.0		•	0.0		0.0	0.0	0.0	121.8
		0.0	0.0	0.0		•	0.0		0.0	0.0	9.7	15.0
	0.467	164.8		23.5	0.0	3.6	6.2	•	44.8	35.7	98.5	• • •
!	6.44	15.0	13.1	2.1	0.0	•	9.0	0.4	4.1	3.2	9.0	102.0

Table A-3-6 Flood Runoff at Proposed Dam Site

													٠							÷ .						
	mm)	Annual	2.35	:		5.65	٠.			34.70			8.10	3.25	0.50	1.10	08.0	1		. 1	18.06			7.11	i	81.62
	(Unit: mm)	Dec.	1			1						٠	1	ı	0.20	i	-1		·	i	1			ı.	1	
		Nov.	ı			1			٠	0.35			0.05	ı	ŀ	1	ı				i			ŧ	ı	
		Oct.	í			ſ				ı			ı	1			ı				ı			, I		
te   		Sep.	ı			ı				ı	٠		ì	ι	1	ı	ı		٠		1			ł	1	
Dam Site	-	Aug.	1			0.20		(0.40)	(0.20)	09.0			į	0.35	ì	1	1				i	•		1.74	i	
at Proposed Dam		Jul.	J			J				l			1	ì	. 1	. 1	0.15						•	ı	ı	
	·	Jun.	i			1			r	1	(0.20)	(0.15)	0.35		ı	I					ı			ι	t	
Flood Runoff		May		•		;				ı.			2.55	ì	ı	1	0.65				1			1	1	
F1.0		Apr.	ı			ı		(4.00)	(0.15)	4.15			3,30	ì	ı	ı	1				Į	(0.30)	(0.77)	1.07	1	
		Mar.	ı			i	(1.05)	(2.90)	(9.85)	13.80				i							1.02			0.38		
		Feb.	2.35	(0.35)	(5.10)	5.45	(3.35)	(2.55)	(8.60)	14.50			1.10	2.90	1	0.20	i	(0110)	(10.50)	(6.38)	17.04			3.92		
		Jan.			•	ŀ		:		1.30			0.75	Ī	0.30	ı	1				1			ı	ı	
		Year	1974			1975				1976			1977	1978	1979	1980	1861			. •	1982			1983	1984	Total

Table A-3-7 Flood Runoff at River-Mouth

Annual 2.704 3.51 1.13 0.50 (Unit: mm) 0.38 Dec. Nov. Flood Runoff at River-mouth 0.53 0.53 Jun. 0.83 0.83 -(0.10) (0.30) 0.40 3.19 1.33 1.46 Apr. (0.26) (3.27) 3.53 (0.22) 0.24 0.25 5.07 (3.14) (1.68) 4.82 (0.04) (1.59) 1.63 (2.38) (2.28) 4.66 0.65 0.40 0.12 1.22 Jan. Total Year 1974 1975 1976 1978 1979 1980 1981 1982 1983 Mean 1977

Table A-3-8 Monthly Rainfall at Muscat

													(Unit:	mm)
Year	Jan.	Feb.	Mar.	Apr.		June	July	Aug.	Sep.	Oct.	Nov.	Dec.	Annual	Remark
1893		39	*	*	*	*	*	*	*	*	1	13	*	British Emb.
94	11	57	7	5	_	-	8			**	-	31	119	18
95	106	16	65	_	-			· <del>-</del>		-	18	-	205	. u
96	63	4	39		-	_	· _	_		4.	77	-	187	11
97	13	16	1	-	_	_	-0.0	_	· <u>-</u>	. ~			30	11
98	2	3	36	_	-	64		<i>→</i>		· -	14	3,	122	. •
99	_	.7	40	_	-	-	~			_		1	48	1 11
1900	64	34	17	<u></u>					_		23	63	201	11
01	_	12	28	_			-	_	<b>-</b> .	_		13	53	21
02	_	7	-	7	-	_	_			25	_	13	52	H
03	10			11	-	1	· •	_	-	-	1.	3	26	п
04	-	3	. 3	_	_	_	_	·	-		18	1	25	0
05	31	46	56	_	_	_	_	`	_	_	5	2	140	
06	15	33	37	_	_	6	_	1	_	-	-	40	132	et et
07	6	79	_	22	_	_	_	_	_	_	5	4	116	11
80	6	_	10	3	_	_	_		_	-	-	5	24	18
09	115	_	_	_	_	_			-	_	_	54	169	. 37
1910	24	_	11	-	_	-	_		-	-		38	73	47
11	67	3	7		_			_	_		18	6	101	11
12	60	12	0	97	_	_	_	7	٠		5	25	206	ti .
13	-	99	22	_		_	_	_			_	14	135	11
14	3	42	2	_	_	9	3	1	_	14	45	22	141	H
15	7	1	3	32	_	_		-			_	7	50	и
16	98	30	5	98	_	_	_	15	-	20	_	_	266	. "
17	60	19	_	2	_	_	_	_	_			24	105	н
18	4	-	10	8	-			_	_	_	_	39	61	11
19	22	22	20		4	-		_	_	_	_	_	68	н
1920	6	14	1	_			_	_	_	· _	4		25	<b>n</b>
21.	4						_				25	15	44	11
22	6	6	_	<u>.</u> .	_		_	_	_		_	_	12	IT
23	7	8	_	36		_	_	_	_	_	20	39	110	o i
24	3	6	_	_		_	<u> -</u>	-	_	_	_	19	28	22
25	7	2	5	_	_	_	_	••	_	44	1	_	59	17
26	25	_	9	6	_	_	_	_	_	-	_	32	72	n
27	-	17		10	_		4	_	-	_	10	9	50	п
28	47	56	_	-		_	-	_		_		19	175	11
29	8	3	_	_	_	_	_	_	_	_	34	116	161	11
1930	142		1	6	_	_				<del>.</del>	_	110	150	Ministry of
31	142	1	<u>+</u>			Recor	d Arai	labla					170	Defence
32					NO	Kecor	u Avas	Table			*			Defeuce
							11							•
33 34							11							
							 H							
35	140	_	00					٠					100	nutrice to Test
36	143	6	28	-	-	-	-	-	-	-	12	-	189	British Emb.
37	29	45	-		-	-	-	-	~	-	-	27	101	
38	-	-	-	-	-		-		-	10	_	20	30	"
39		75	-	-		-	_			-	-	23	98	
1940	21		10		-	-				<del>-</del>	_	55	86	

Table A-3-8 Monthly Rainfall at Muscat (Contined)

(Unit: mm) Apr. May June July Aug. Sep. Oct. Nov. Dec. Annual Remark Jan. Feb. Mar. Year 0.5 13.5 20 34 British Emb. 41 8.5 41.5 29 42 4 \* 87 43 166 44 \* 45 1 14 46 2.5 14 47 44.5 25 2.5 29 48 49 \* 2.5 0.5 1950 16.5 62 1 6 69 51 16 69 1 52 52 15 49 53 2 27 1 10 1 2 54 11 14 188 97 7 70 55 0.5 37.1 171 234 12.2 13.2 56 36 225 62 57 109.0 9 5 16 53.0 58 68.6 12.7 115.5 2.5 59 10.9 21.8 16 108.9 13.7 18.3 36.8 24.1 1960 12.2 0.5 1.8 10 34.6 2.3 14.5 2.3 61 20.3 119.1 72.1 -6.9 62 19.8 11.4 141.0 2.3 24.9 94 63 5.1 26.9 64 10.4 11.4 83.1 2.0 107.7 22.6 65 96.6 P.D.O. 66 88 1 : 7 6.2 0.1 5.9 22.6 67 0.6 29.2 144.8 90.3 68 22.8 2.5 74.2 69 56.4 5.6 12.2 141.7 1970 31.7 37.5 44.9 97.6 71 265.4 95.5 47.7 18.3 72 103.9 96.8 73 96.8 74 0.3 20.0 23.3 DAR SITE 3.0 83.6 1.0 75 2.7 79.9 2.0 228.6 76 51.0 56.0 66.3 43.3 10.0 56.9 187.9 6.8 77 64.5 22.0 5.7 32.0 0.6 66.8 78 12.4 39.4 13.4 1.0 32.3 39.2 79 6.9 3.7 1980 0.3 123.7 0.5 81 4.2 16.0 103 11 132.5 4.5 29.5 82 2.9 59.6 35.1 0.9 3.6 80.3

Source: Water Resources Department

46.7

Notes: 1. - no rainfall

25.6

2. \* no data

4.5

0.9

Characteristics of Maximum Rainfall

			ļ		Annual Ra	Rainfall				•	
Rain Gauge	River	Altitude (m)	Record	Maximum (1)	Minimum	Average	(Rate)	Monthly Maximum $(1)$ $(2)/(1)$	Maximum $(2)/(1)$	Daily Maximum (3) (3) (1)	(1)
Wadi Jizzi Basın											
Daqiq	Al Jizzi	800	1974-84	630.8	17.5	182.1	(1.88)	234.8	(0.37)	74.3 (0.12)	(2)
Kitnah	=	655	1974-84	353.1	8.8	108.7	(1.12)	140.8	(0.40)	78.0 (0.22)	(7)
Hayl (W.J)	=	200	1974-84	328.5	19.0	104.6	(1.08)	155.7	(0.47)	71.0 (0.22)	(2)
Hayl (W.H)	Ξ	430	1974-84	0.494	0.0	116.2	(1.20)	190.9	(0.41)	64.9 (0.14)	(4)
Farfar		260	1974-84	406.1	0.0	144.1	(1.49)	155.0	(0.38)	67.0 (0.16)	(9)
Sohar	E	15	1974-84	253.1	8.6	102.0	(1.05)	137.2	(0.54)	91.0 (0.36)	(98
Basin Average at	Dam		1974-84	420.8	11.8	127.8	(1.32)	150.9	(0.36)	54.4 (0.13)	(E)
Other Basin											
Ghozaifah	Ahin	780	1974-83	281.1	51.6	115.1	(1.19)	125.3	(0.45)	48.6 (0.17)	[7)
Qufays	<b>\$</b>	009	1974-83	345.7	31.2	145.8	(1.50)	201.5	(0.58)	59.4 (0.17)	17)
Haybi	=	570	1974-83	268.6	28.3	148.9	(1.54)	163.0	(0.61)	59.0 (0.22)	22)
Hawqayn			1975-83	314.0	39.0	160.2	(1.65)	131.0	(0.42)	98.0 (0.31)	31)
Rustaq		350	1974-83	377.6	78.2	172.8	(1.78)	123.8	(0.33)	81.5 (0.22)	22)
Nakhl			1974-83	267.7	16.0	139.4	(1.44)	145.0	(0.54)	70.0 (0.26)	76)
Muscat	Aday		1893-83	266.0	3.7	101.5	(1.05)	171.0	(0.64)	1	
-op-	Ė		1974-83	228.6	3.7	97.0	(1.00)	103.0	(0.45)	. 1	
Quryat	Miglas		1975-84	339.5	35.0	145.7	(1.50)	140.0	(0.41)	140.0 (0.41)	41)

Table A-3-10 Calculation of Recharge Rate Based on Hydrological Analysis

After Project Calculation for Recharge Rate Based on Hydrological Analysis Before Project 6.63 3.47 Expected Flood Runoff at River Mouth (MCM) Analyzed Flood Runoff at River Mouth (MCM) Total Recharge below Dam (MCM): (4) - (5) (6) Flood Runoff below Dam (MCM): (4) - (3) (3) Analyzed Flood Runoff at Dam (CMC) (2) Arcal Rainfall at River Mouth (mm) Recharge Rate  $(%): (7)/(4) \times 100$ (1) Areal Rainfall at Dam Site (mm) Description 3 (4) (2) 8

Remark: 1/ (3) x Areal Ratio x Rainfall Ratio = 6.63 x 1,283/812 x 121/128

## A-4 Outlfow Capacity of Outlet Conduit

A recharge potential at the downstream Wadi course of the dam was estimated to get the optimal recharge rate for the development plan.

Figure A-4-1 will provide a theoretical curve to represent relationship between peak flood discharges at the dam site and river mouth when the equivalent amount of areal rainfalls are expected in both the drainage basin.

## Theoretical Curve for Peak Flood Discharges

	Proposed	Dam-Site		Ri	ver - Mou	ith
Areal	Specific	Runoff	Peak	Specific	Runoff	Peak
Rainfall	Runoff	Volume	Discharge	Runoff	Volume	Discharge
(mm)	(mm)	(MCM)	(m³/sec)	(mm)	(MCM)	(m <sup>3</sup> /sec)
11.5	0.21	0.1705	17.0	0	0	0
13.0	0.30	0.2436	24.3	0.09	0.1155	4.7
15.0	0.42	0.3410	34.0	0.20	0.2566	10.4
20.0	0.83	0.6740	67.3	0.45	0.5774	23.4
25.0	1.35	1.0962	109.4	0.68	0.8724	35.3
30.0	2.00	1.6240	162.1	0.90	1.1547	46.7
35.0	2.73	2.2168	221.2	1.11	1.4241	57.6
40.0	3.50	2.8420	283.6	1.31	1.6807	68.0
45.0	4.30	3.4916	348.5	1.51	1.9373	78.4
50.0	5.20	4.2224	421.4	1.70	2.1811	88.2
60.0	7.33	5.9520	594.0	2.09	2.6815	108.5
70.0	9.85	7.9982	798.2	2.47	3,1690	128.2
80.0	12.60	10.2312	1021.0	2.82	3.6181	146.4
		The second second		The second secon	and the second second	

Note: Peak discharges are converted from runoff volume by use of ratios of peak discharge to flood volume.

- (1) At dam site ratio =  $9.9795 \times 10^{-5}$
- (2) At river mouth ratio =  $4.045 \times 10^{-5}$

However, analyzed flood runoffs are not always plotted on the theoretical curve, but plotted elsewhere around the curve, mainly due to difference in catchment's rainfall. These relations are also shown in Figure A-4-1, meaning that (1) when the entire Wadi Jizzi basin receives the same amount and distribution of areal rains the dam can discharge 17.0 cu.m/sec of water without wasting any portion of surface water into the sea, and (2) when downstream area receives a greater amount of rains than upstream the dam can release a smaller amount of discharge not so as to waste water

into the sea. In consideration of such condition, an envelope curve was so drawn, as shown in Figure A-4-1, as to enclose all the points distributed above the theoretical curve, for the purpose of finding optimal release rate of stored water in the dam.

Consequently, 13.0 cu.m/sec was obtained from the study as the optimal rate of release.

An outlet conduit is to be embedded beneath the dam body in consideration of the existing condition of the water route, effective recharging of groundwater in the downstream river deposits, and necessity to flush-out the fine sediments around the entrance of the outlet conduit.

A circular shape steel pipe conduit with the inner diameter of 1,500 is planned for the outlet facilities of the detention dam. At the entrance of outlet conduit, trush-rake is equipped in order to prevent the objectionable materials flowing into the conduit.

On outlet discharge of this conduit at the varying reservoir water levels can be obtained as a pipeline flow by the following equation.

$$Q = \frac{\sqrt{2g} \cdot A}{\sqrt{fv + fe + fr + fb}} \cdot \sqrt{H}$$

$$A = 0.7854 D^2$$

where, Q; outlet discharge of conduit as a pipeline flow

g; gravitational acceleration

A; flor area =  $0.7854 \text{ D}^2 = 1.767$ 

D; inner diameter of outlet conduit pipe, adopted by 1.50m

- fv; coefficient of changing velocity loss, adopted by  $1.0\,$
- fe; coefficient of the entrance loss adopted by 0.5
- fr; coefficient of friction loss, fr =  $124.5 \text{xn}^2/\text{D}^{4/3} \text{xL}$ = 2.056
- fb; coefficient of bend loss, fb = 0.99
- n; coefficient of roughness, adopted by 0.015
- L; length of conduit, adopted by 126.0 m
- H; total head, measured from top surface at the end of conduit

Based on the results of the above-mentioned computation, the relationship between an outlet discharge and a reservoir water level can be estimated as follows:

$$Q = 3.669 \sqrt{H}$$

$$Q_{\text{max}} = 3.669 \sqrt{163.9 \text{ m} - 151.9 \text{ m}} = 12.7 \text{ m}^3/\text{sec}$$

ingle Event Flood Runoffs at Dam Site and River Mouth

								Unit: MCM and (cu.m/sec)	(cu.m/sec)	
YEAR	JAN	FEBRUARY	MARCH	APRIL	MAY	JUNE	JULY	AUGUST	NOV.	DECEMBER
1974		(190.4) 1,908 (71,1) 1,758					2			
1975		(28.3) (413.3) 0,284 4.141 (2.1) (82.5) 0.051 2.040						(16.2) 0.162 (-)		
1976		(271.4) (206.7) (696.9) 2.720 2.071 6.983 (123.5) (118.3) 3.054 2.925	(85.1) (235.0) (798.2) 0.853 2.355 7.998 (13.5) (169.7) 0.334 4.195	(324.1) (12.2) 3.248 0.122 (75.8) (-) 1.873 (-)				(32.4) (16.2) 0.325 0.162 (-) (-)	(28.3)	
1977	(60.8) 0.609 (36.3) 0.898	(89.1) 0.893 (33.7) 0.834		(267.5) 2.680 (69.0) 1.706	(206.7) 2.071 (43.1) 1.065	(16.2) (12.2) 0.162 0.122 (-) (-)		. · · · . · . · · · · · · · · · · · · ·	(4.1) 0.041 (-)	
1978		(225.0) 2.355 58.7 1.450	·					(28.3) 0.284 ( - )		
1979	(24.3) 0.244 (6.2) 0.154									(16.2) (-) 0.162 - (-) (19.7) (-) 0.488
1980		(16.2) 0.162 (-)	(73.0) 0.731 (12.5) 0.308				:			
1981				·	(52.7) 0.528 (12.9) 0.321		(12.2) 0.122 (-)			
1982		(13.0) (850.9) (517.0) 0.130 8.526 5.181 (-) (163.0) (87.2) - (4.029 2.155	(19.5) (63.2) 0.195 0.633 (11.4) (24.4) 0.282 0.603							
1983		(317.6) 3.183 (80.9) 2.001	(30.8) 0.309 (18.7) 0.462	(24.3) (62.4) 0.244 0.625 (5.2) (15.6) 0.128 0.385				(141.0) 1.413 (27.5) 0.680		

Remark: The upper row shows single event flood volume in MCK and pask discharge (parenthesized) in cum/sec at the proposed dam site, while the lower row presents those figures at the Wad: First tives mouth.

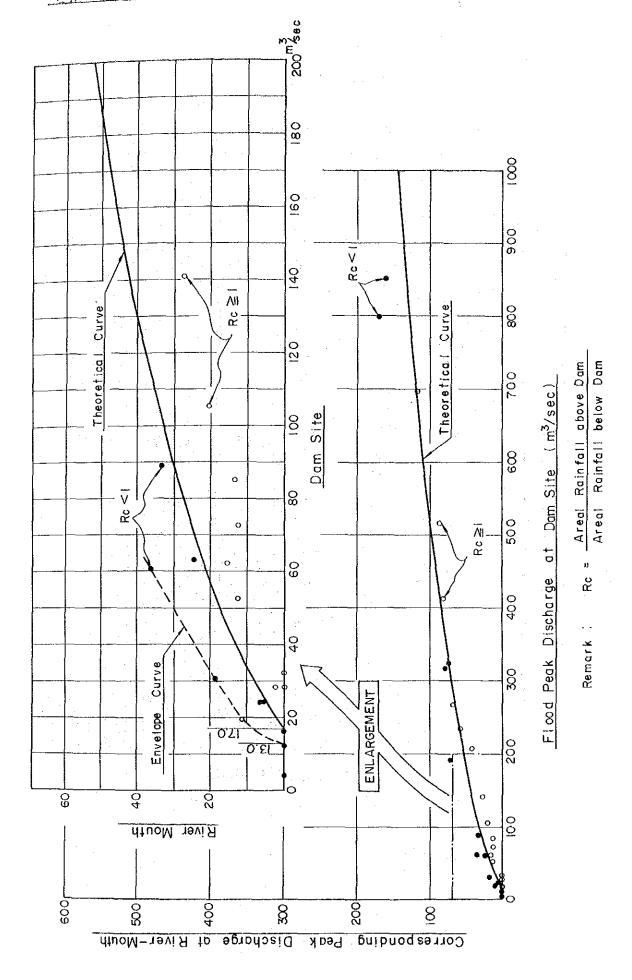
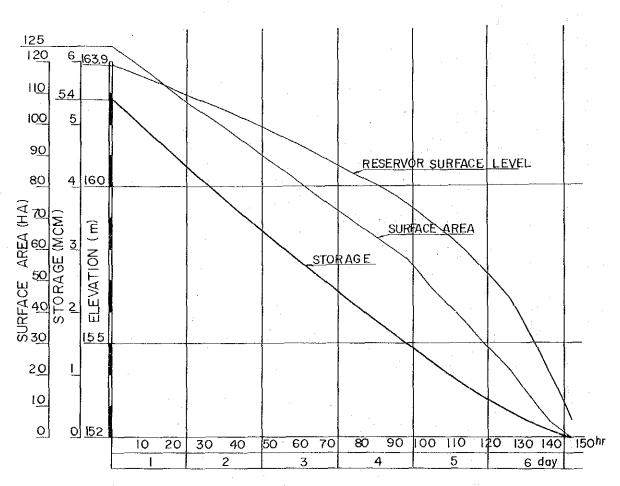
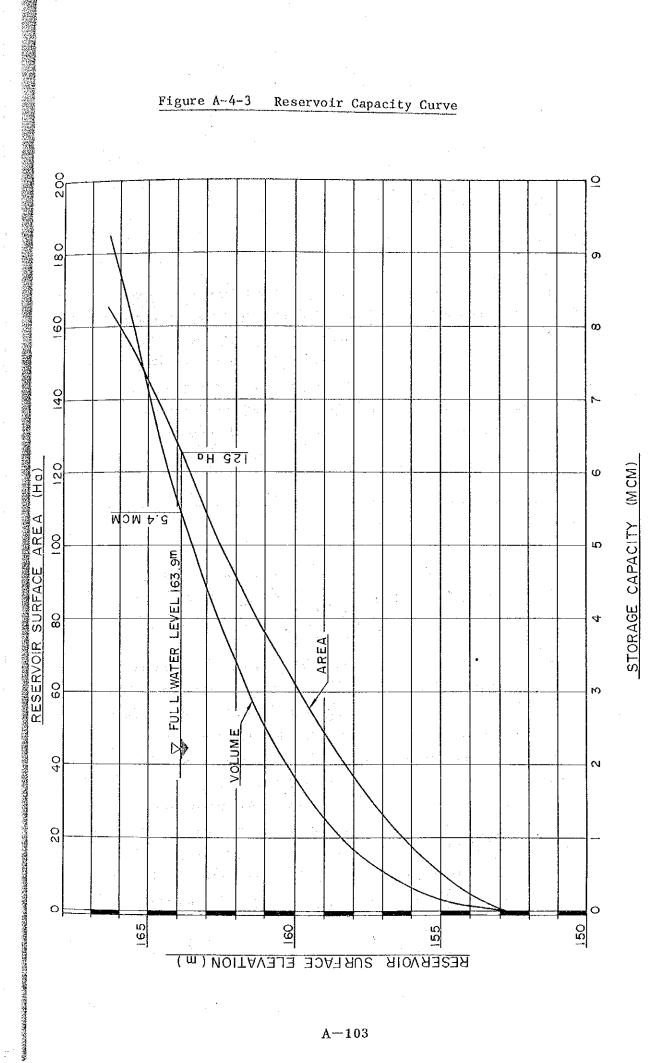


Figure A-4-2 Reservoir Emptying Time



TIME (DAYS & HOURS)



## A-5 Irrigation Water Requirement

## Present Irrigation Conditions

#### 1. Existing Irrigation System

The irrigation system in the area is predominantly basin irrigation. Each basin is either circular one meter to three meters in diameter and 10 centimeters to 60 centimeters deep holding one tree, or rectangular five square meters to 30 square meters in size and a few centimeters lower than the conveyance canal holding one or two trees and occasionally undergrown by feed crops. The later may almost be called the border-strip irrigation.

There are three types of water conveyance systems to lead water to these basins.

- i) direct diversion from the main and lateral canals.
- ii) a series of basins used as a conveyance canal.
- iii) combination of i) and ii).

No. i) is the system where the water is directly diverted from the conveyance canal to the basin. A few farms adopt this system. However, seldom is the main canal straight or wellkept for efficient conveyance of water, causing excessive seepage and evaporation losses from the canal. Yet this system has the potentiality to be most efficient of the three, as far as the water saving is concerned, because there is no intermediate conveyance losses before the cropping area. No. ii) is the system where water is fed directly into the first basin. After filling the first basin the water flows into the second one and then the third to the last basin. This is the most inefficient method of

the three as the percolation loss as the beginning of the system is enormous and application rate towards the end of the system is insufficient. There are varying degrees of combination of the above two. If sensibly laid out this No. iii) system proves to be the most economical of the three as this saves the expensive construction of the long main and lateral canals and also the spaces occupied by them. Most of the farm lands employ the last system. But the basin-to-basin conveyance distance is usually too much (30 meters to 60 meters) resulting in the deficiency described above for No. ii).

Apart from the basin irrigation on the border-strip and furrow irrigation seem to gain popularity especially among the newly developed large scale farms along the highway. The conveyance canals are generally straight and well kept. But there is still a lot of room for improvement in canal routing and quality as well as for introduction of other irrigation facilities.

### 2. Present Irrigation Practices

The procedure of irrigation in the survey area varies little from place to place, be it of different efficiency, i.e., pumped water is let to the crop area by gravity through ditches, furrows and basins. All the irrigation water in the area is presently supplied by well water including in the village of Falaj Al Awhi on the north-western corner where the irrigation water was formerly supplied by the Falaj. The pumps are usually installed singly or in a pair in a pit about five meters deep. Suction pipes extend into the adjoining well of 80 centimeters to 120 centimeters diameter. The depth to the water surface was observed to be about seven to 10 meters below ground surface. The pump sizes are mainly 3" x 3" and sometimes 4" x 3".

Water is first poured into a rectangular concrete basin of the size 1.2 m x 1.8 m to 2 m x 3 m and 20 centimeters to 50 centimeters deep. The top rim of the basin is 60 centimeters to 70 centimeters higher than the ground. Water is then let out through an elevated concrete ditch 15 centimeters to 20 centimeters wide and 10 centimeters to 15 centimeters deep. The length of this concrete ditch is usually less than 10 meters and from there on a crude earth ditch continues. The earth ditch is usually winding and not of a uniform shape. Water is diverted by breaking a side of the ditch and clogging the immediate downstream by filling mud into the ditch. This daily operation helped in deforming and deteriorating the ditch, such a ditch will cause more evaporation losses due to low speed of flow and much seepage due to constantly disturbed and softened ditch surfaces.

Farmers seem to know approximately the right way of irrigating each crop after many years or generations of experience. Basin irrigation is adopted for tree crops, border-strip for fodder crops and furrow irrigation for various vegetables. However, it was almost always the case that the conveyance route was arranged in a rather complicated and wasteful way with a lot of detours which could be simplified and rationalized. These detours are possibly the consequences of the efforts to achieve best results in uniform water distribution. By the time the purpose was achieved the conveyance route was apparently grossly elongated. It is therefore important to study the topography carefully when rectifying the canal route.

#### 3. Irrigation Efficiency

The classification according to the crop condition described in previous paragraph is reconsidered here in terms of irrigation intensity.

In No. 1 (good condition) area it was observed that water was abundantly applied both in quantity and frequency. In some cases water was applied every day or every other day, where the quantity was estimated to be from 100 percent up to even 500 percent of the consumptive use. On these locations plant growth generally looked good. Possible reasons for it is ample application of fertilizers and good leaching of hazardous salts coupled with sufficient supply of water and favourable soil conditions. Further study on harmful effect of over-irrigation should be made.

In No. 2 (moderate condition) area water application was estimated at between 50 percent to 100 percent of the consumptive use of water. On these farms the irrigation method and facilities were usually poor to fair. Only the beginning part of the irrigation system received ample to excessive supply of water and the end part of it suffered from water deficiency.

No. 3 (poor condition) are obtained only 20 percent to 50 percent of water compared to the consumptive use of the plants. In these areas the farmers attended the farm only occasionally or only part of the farm due to lack of sufficient available water.

The summary of water consumption of each of these areas is given below:

Classification	Area Percentage	Water Application Percentage of Consumptive Use	Average
	(%)	(%)	(%)
Good area	20	100	$100^{\frac{1}{2}}$
Moderate area	49	50 - 100	75
Poor area	31	20 - 50	35

<sup>1/</sup> Some inconsistencies involved in the Feasibility Study Report were revised. The maximum water application percentage of consumptive use was considered to be 100%.

Under these conditions the overall irrigation efficiencies are considered to be 50 percent. Most of the losses occur through deep percolation and evaporation caused by inefficient and deficient conveyance canals and partly due to over-irrigation.

## 4. Present Irrigation Water Requirement

From the hydrogeological study, about 17 kilometers long coastal plain extending from the Wadi Khadaq to the Village of Majis was determined to be beneficial areas covered by the Wadi Jizzi river basin. The gross cultivated area within this coastal plain is about 3,830 hectares and its net cultivation area is 2,640 hectares. The water resources for irrigation in these areas are wells with a depth of 20 meters to 30 meters on average, and numerous small scale pumps are installed for lifting groundwater as irrigation water.

Since the estimation of present annual amounts used for irrigation by means of pumping operation during a year was found out to be difficult due to numerous pump installation as explained above, the amounts were estimated based upon the cropping areas and consumptive use of each crop.

### a) Present Cropping Area and Cropping Pattern

According to the field survey and collected data on the existing land use, the existing cropping acreage is summarized as shown below and its cropping pattern is shown in Figure A-5-1 and A-5-2.

## Existing Cropping Acreage

<u>Vegetable</u>			
Oion	<b>‡</b>	96 ha	(3.4%)
Garlic	:	25	(0.9)
Tomato	:	16	(0.6)
Potato	:	15	(0.5)
0kra	:	8	(0.3)
Others	:	23	(8.0)
Fruit Crop			
Dates	:	1,820	(64.2)
Lime	:	309	(10.9)
Banana	:	128	(4.5)
Mango		83 -	(2.9)
Feed Crop	1		. •
Alfalfa	:	211	(7.4)
Sorghum	:	101	(3.6)
<u>Total</u>	:	2,835	(100 %)

### b) Irrigation Water Requirement

## 1) Potential Evapo-transpiration

The reference crop potential evapo-transpiration (ETPc), which is generally recognized as a fairly reliable index for calculating consumptive use, can be determined by a number of methods, such as the evaporation measurement with evaporation pan and the application of empirical formula based on the climatological data. In the Project, the monthly evapo-transpiration was estimated on the monthly basis, by applying modified Penman method based on the climatological data observed at Sohar meteorological station.

- 2) Crop Water Supply Requirements
- (a) Crop Evapo-transpiration

Crop evapo-transpiration (ETc) (consumptive use of crops) of each crop has been estimated by multiplying the estimated ETPc values by crop coefficients which express the relationship between the reference crop potential and the actual evapo-transpiration during the vegetative stage of the crops. The crop evapo-transpiration (ETc) of each crop was estimated based upon the above procedures and is shown in Tables A-5-2(1) to A-5-2(4).

### (b) Water Supply Requirements

For the estimation of water supply requirement for each crop, the following equation has been applied:

$$V = \frac{10}{EP} \left[ \frac{A (ETc - RE)}{1 - LR} \right]$$

where; A : area (ha)

ETc : crop evapo-transpiration (mm)

RE : effective rainfall (mm)

EP : project irrigation efficiency

LR : leaching requirement

In the estimation, effective rainfall and leaching requirements have been accounted based upon the following bases.

## Effective Rainfall

The effective rainfall for crops was estimated on the monthly basis by applying the criteria provided in FAO Irrigation and Drainage Paper No. 24, Table 34 "Average Monthly Effective

Rainfall as related to Average Monthly EPTc and Mean Monthly Rainfall": (see Table A-5-3)

### Leaching Requirements

With regard to Water Quality, the groundwater in the Wadi Jizzi plain, which is the water sources of irrigation for the Project, has been found to need some countermeasure for salinity control such as leaching. The leaching requirements are the minimum amount of irrigation water that must be supplied and drained through the root zone to keep the soil salinity at a specific level. For sandy loam soils with good drainage and in the areas where rainfall is small, the leaching requirements can be obtained from the following equation.

Leaching requirement (LR) for drip and high frequency sprinkler:

$$LR = \frac{EC(W)}{2MAX \cdot EC(E)} \cdot \frac{1}{LE}$$

where, EC(W) : electric conductivity of irrigation
water, 0.56 mmhos/cm (average of five
water samples)

Max. EC(E): maximum tolerable electircal conductivity of the soil saturation extract for crops, derived from FAO Irrigation and Drainage Paper No. 24, Table 36.

LE : leaching efficienty, 0.8 (sandy loan)

Water supply requirement for each crops were calculated employing the above mentioned procedures as shown in Tables A-5-2(1) to A-5-2(4), and the results are summarized with respect

to crops and land use as given in Tables A-5-4 and A-5-5, respectively. Consequently the annual consumption of irrigation water for the existing cultivated areas of 2,835 ha is estimated at 11.0 MCM, taking present status of irrigation water application into consideration.

Table A-5-1 Penman's Reference Crop Evapotraspiration

Penman's Reference Crop Evapotranspiration

Place: SOHAR Latitude: 24°20' Longitude: 56°40' Altitude: 15m

). Cu Lu Lu Lu Lu Lu Lu Lu Lu Lu Lu Lu Lu Lu	; ;	, c		1 0 2	) (	West	į	[14]	φφ	Ç	4	N.	ç Ç	Note /Forestion
Thean	200	28.	,	22.0	25.9	30.1	32.1	32.6	31.6	29.6	25.9	21.7	9 6 5	oiven data
RHnean	6%	68.5		67.4	59.9	55.2	61.8	70.5	74.0	9.69	63.7	6.99	70.5	1 00
69	mbar	20.9		26.4	33.4	42.7	47.9	49.2	46.5	41.5	33.4	26.0	22.8	
e D	mbar	14.3		17.8	20.0	23.6	29.6	34.7	34.4	28.9	21.3	17.4	16.1	ea x RHmaen/100
(ea-ed)	mbar	9-9		8.6	13.4	19.1	18.3	14.5	12.1	12.6	12.1	9.6	6.7	(1)
บร	Km/day	56.4	61.0	67.2	70.3	73.3	75.2	85.3	84.9	71.0	58.6	9.67	46.3	given data
f (n)		0.42		0.45	97.0	0.47	0.47	0.50	0.50	0.46	0.43	07.0	0.40	0.27(1+U2/100)
1 - W		0.34		0.29	0.25	0.22	0.20	0.19	0.20	0.22	0.25	0.29	0.32	
(1-W)f(u)(ea-ed)		0.94		1.12	1.54	1.97	1.72	1.38	1.21	1.28	1.30	1.00	0.86	(2)
Ra		10.1		13.9	15.4	16.4	16.6	16.5	15.8	14.5	12.6	10.6	9.6	
C.	hr/day	7.2		8.1	8.7	10.3	9.0	8.7	8.5	9-3	9.1	4.8	7.6	given data
×	hr/day	10.7		12.0	12.7	13.3	13.6	13.5	13.0	12.3	11.6	10.9	10.6	
Rs	mm/day	5.92		8.17	9.12	10.45	10.19	9.44	9.11	9.11	8.09	6.73	5.84	(0.25_0.50n/N)Ra
Rns	mm/day	7.7		1.9	6.8	7.8	7.6	7.1	8.	8.9	6.1	5.1	7-7	0.75Rs
fImean	ı	14.2		15.0	15.9	16.7	17.2	17.4	17.1	16.6	15.9	14.9	14.5	
f(ed)	1	0.17		0.15	0.14	0.13	0.10	0.08	0.08	0.10	0.14	0.16	0.16	0.34-0.044ved
f(n/N)	1	0.70		0.70	0.71	0.79	0.76	0.69	0.69	0.78	0.81	0.79	0.75	N/E6.9+1.0
Rnl	mm/day	1.7		1.6	7.6	1.7	1.3	1.0	6.0	1.3	1.8	1.9	1.7	f(Tmean)f(ed)f(n/N)
Rn	mm/day	2.7		4.5	5.2	6.1	6.3	6.1	5.9	5.5	4.3	3.2	2.7	Rns-Rn1
Z.	t	99.0		0.71	0.75	0.78	0.80	0.81	08.0	0.78	0.75	0.71	0.68	1-(1-W)
W.Rn	тт/day	9		3.2	3.9	4.8	5.0	4.9	4.7	4.3	3.2	2.3		(3)
(1) + (3)	-	2.7		4.3	5.4	8.9	6.7	6.3	5.9	9.6	4.5	9.3	2.7	(2)+(3)
U		1.06		1.10	1.12	1.13	1.13	1.13	1.13	1.12	1.10	1.07	1.06	
ETO	nm/day	2.9		4.7	6.0	7.7	7.6	7.1	6.7	6.3	5.0	3.5	2.9	
- qo -	mm/month	89.9		145.7	180.0	238.7	228.0	220.1	207.7	189.0	155.0	105.0	89.9	

Table A-5-2 (1) Unit Irrigation Water Requirement

Onion:	MEC(E) =	8, LR =	0.044				
Month	Rain	ETo	KC	ET(C)	R(E)	_A	V
Sep.	0.4	189.0	1.00	189.0		0.10	395
Oct.	4.1	155.0	1.00	155.0	2.8	0.40	1,274
Nov.	3.2	105.0	0.70	73.5	2.3	1.00	1,489
Dec.	9.0	89.9	0.75	67.4	6.4	1.00	1,276
Jan.	9.3	89.9	0.65	58.4	6.3	0.66	720
	26.0				17.8	* .	5,154
:							
Garlic:	MEC(E)	= 15, LR	= 0.023			•	
Mont h	Rain	ЕТо	<u>KC</u>	ET(C)	<u>R(E)</u>	A	v
Sep.	0.4	189.0	0.35	66.1	. ~	0.50	677
Oct.	4.1	155.0	0.50	77.5	3.0	1.00	1,525
Nov.	3.2	105.0	0.70	73.5	2.3	1.00	1,457
Dec.	9.0	89.9	0.75	67.4	6.4	1.00	1,249
Jan.	9.3	89.9	0.80	71.9	6.8	1.00	1,332
Feb.	44.9	100.8	0.80	80.6	32.6	1.00	983
Mar.	15.0	145.7	0.80	116.6	12.1	1.00	2,138
Apr.	13.1	180.0	0.70	126.0	10.9	1.00	2,356
May	2.1	238.8	0.50	119.3	1.5	0.50	1,206
	$1\underline{01.1}$				75.8		12,923
Tomato:	MEC(E)	= 13, LR	= 0.027				
Month	Rain	ETo	KC	ET(C)	R(E)	A	V
Sep.	0.4	189.0	1.00	189.0	-	0.10	388
Oct.	4.1	155.0	1.00	155.0	2.8	0.10	313
Nov.	3.2	105.0	0.45	47.2	1.4	1.00	942
Dec.	9.0	89.9	0.55	49.4	5.7	1.00	899
Jan.	9.3	89.9	0.85	76.4	6.9	1.00	1,428
Feb.	44.9	100.8	1.00	100.8	33.5	1.00	1,383
Mar.	15.0	145.7	0.85	123.8	12.4	1.00	2,290
Apr.	13.1	180.0	0.70	126.0	10.9	1.00	3,366
May	2.1	238.7	0.65	155.2	0.9	0.50	1,585
	101.1		5		74.6		11,594
Potato:	MEC(E)	= 10, LR	= 0.035				
Month	Rain	ЕТо	KC	ET(C)	R(E)	A	ν.
Nov.	3.2	105.0	0.45	47.2	1.4	1.00	949
Dec.	9.0	89.9	1.05	94.4	6.5	1.00	1,821
Jan.	9.3	89.9	1.35	121.4	7.6	1.00	2,358
Feb.							
100.	44.9	100.8	1.25	126.0	35.6	0.50	937

Table A-5-2 (2) Unit Irrigation Water Requirement

EC(E) =	10, LR =	0.035									
Rain 44.9 15.0 13.1 2.1 0.0 75.1	ETO 100.8 145.7 180.0 238.7 228.0	KC 0.40 0.70 0.85 0.80 0.60	ET(C) 40.3 102.0 153.0 191.0 136.8	R(E) 30.1 11.5 11.1 0.5 - 53.3	A 0.50 1.00 1.00 1.00 1.00	V 105 1,875 2,941 3,947 2,835 11,704					
Others: MEC(E) = 10, LR = $0.035$											
Rain 3.2 9.0 9.3 44.9 15.0 81.4	ETO 105.0 89.9 89.9 100.8 145.7	KC 0.40 0.70 0.85 0.80 0.60	ET(C) 42.0 62.9 76.4 80.6 87.4	R(E) 1.6 6.2 6.9 32.6 11.3 58.7	A 1.00 1.00 1.00 1.00 0.50	837 1,175 1,440 995 788 5,236					
MEC(E) =	32, LR -	0.011									
Rain 9.3 44.9 15.0 13.1 2.1 0.0 0.3 0.6 0.4 4.1 3.2 9.0 102.0	ETO 89.9 100.8 145.7 180.0 238.7 228.0 220.1 207.7 189.0 155.0 105.0 89.9	KC 0.70 0.70 0.70 0.70 0.70 0.70 0.70 0.70 0.70 0.70 0.70	ET(C) 62.9 70.6 102.0 126.0 167.1 159.6 154.1 145.4 132.3 108.5 73.5 62.9	R(E) 6.5 32.0 11.5 10.9 1.0 2.7 2.3 6.2 73.2	A 1.00 1.00 1.00 1.00 1.00 1.00 1.00 1.0	V 1,142 779 1,829 2,328 3,359 3,228 3,116 2,940 2,675 2,139 1,439 1,439 1,146 26,120					
EC(E) =	8, LR = 0	-044									
Rain 9.3 44.9 15.0 13.1 2.1 0.0 0.3 0.6 0.4 4.1 3.2 9.0	89.9 100.8 145.7 180.0 238.7 228.0 220.1 207.7 189.0 155.0 89.9	KC 0.50 0.50 0.50 0.50 0.50 0.50 0.50 0.50 0.50 0.50	ET(C) 44.9 50.4 72.8 90.0 119.3 114.0 110.1 103.8 94.5 77.5 52.5 44.9	R(E) 6.0 30.3 11.1 9.8 1.5 - - 3.0 1.5 5.8	A 1.00 1.00 1.00 1.00 1.00 1.00 1.00 1.0	V 815 420 1,291 1,677 2,465 2,385 2,302 2,173 1,977 1,559 1,068 820					
	Rain 44.9 15.0 13.1 2.1 0.0 75.1  MEC(E)  Rain 3.2 9.0 9.3 44.9 15.0 81.4  MEC(E) =  Rain 9.3 44.9 15.0 13.1 2.1 0.0 0.3 0.6 0.4 4.1 3.2 9.0 102.0  EC(E) =  Rain 9.3 44.9 15.0 13.1 2.1 0.0 0.3 0.6 0.4 4.1 3.2 9.0 102.0	Rain 44.9 100.8 15.0 145.7 13.1 180.0 2.1 238.7 0.0 228.0 75.1 MEC(E) = 10, LR  Rain 5.0 105.0 9.0 89.9 9.3 89.9 44.9 100.8 15.0 145.7 81.4 MEC(E) = 32, LR - Rain 69.3 89.9 44.9 100.8 15.0 145.7 13.1 180.0 2.1 238.7 0.0 228.0 0.3 220.1 0.6 207.7 0.4 189.0 4.1 155.0 3.2 105.0 9.0 89.9 102.0 EC(E) = 8, LR = 0  Rain 6.0 ETO 89.9 100.8 15.0 145.7 13.1 180.0 2.1 238.7 0.0 228.0 0.3 220.1 0.6 207.7 0.4 189.0 4.1 155.0 3.2 105.0 9.0 89.9 102.0 105.0 9.0 89.9 102.0 105.0 9.0 89.9 105.0 9.0 89.9 105.0 9.0 89.9 105.0 9.0 89.9 105.0 9.0 89.9 105.0 9.0 89.9 105.0 9.0 89.9 105.0 9.0 89.9 105.0 9.0 89.9 105.0 9.0 89.9 105.0 9.0 89.9	Hain   ETO   KC   Rain   ETO   Rain   ETO   Rain   ETO   Rain   Rain   ETO   Rain   Rain	Rain   ETO   KC   ET(C)   44.9   100.8   0.40   40.3   15.0   145.7   0.70   102.0   13.1   180.0   0.85   153.0   2.1   238.7   0.80   191.0   0.0   228.0   0.60   136.8   75.1	Rain         ETO         KC         ET(C)         R(E)           44.9         100.8         0.40         40.3         30.1           15.0         145.7         0.70         102.0         11.5           13.1         180.0         0.85         153.0         11.1           2.1         238.7         0.80         191.0         0.5           0.0         228.0         0.60         136.8         -           75.1         53.3     MEC(E) = 10, LR = 0.035   MEC(E) = 32, LR - 0.60  MEC(E) = 32, LR - 0.60  MEC(E) = 32, LR - 0.011   MEC	$\begin{array}{c ccccccccccccccccccccccccccccccccccc$					

Table A-5-2 (3) Unit Irrigation Water Requirement

Banana	: MEC(E)	= 8, LR =	0.044				
Month	$\frac{\text{Rain}}{9.3}$	ETo 89.9	KC 0.90	ET(C) 80.9	$\frac{R(E)}{6.9}$	$\frac{A}{1.00}$	V 1,548
Jan.	44.9	100.8	0.90	90.7	33.0	1.00	1,207
Feb.	15.0	145.7	0.90	131.1	12.5	1.00	2,481
Mar.	13.1	180.0	0.90	162.0	11.5	1,00	3,149
Apr.	2.1	238.7	0.90	214.8	0.8	1.00	4,477
May Jun.	0.0	228.0	0.90	205.2	<b></b> ,	1.00	4,293
Jul.	0.3	220.1	0.90	198.1	_	1.00	4,144
Aug	0.6	207.7	0.90	186.9	_	1.00	3,911
Sep.	0.4	189.0	0.90	170.1	<b>-</b>	1.00	3,559
Oct.	4.1	155.0	0.90	139.5	3.0	1.00	2,856
Nov.	3.2	105.0	0.90	94.5	1.8	1.00	1,939
Dec.	9.0	89.9	0.90	80.9	6.7	1.00	1,553
Dec.	102.0	03.3	0.50	00.7	76.2	1.00	35,117
Mango:	MEC(E) =	8, LR =	0.044				
Month	Rain	ЕТо	KC	ET(C)	R (E)	A	V
Jan.	9.3	89.9	0.70	62.9	6.5	$\frac{1.00}{1.00}$	1,181
Feb.	44.9	100.8	0.70	70.6	32.0	1.00	806
Mar.	15.0	145.7	0.70	102.0	11.5	1.00	1,893
Apr.	13.1	180.0	0.70	126.0	10.9	1.00	2,408
May	2.1	238.7	0.70	167.1	1.0	1.00	3,475
Jun	0.0	228.0	0.70	159.6	_	1.00	3,339
Jul.	0.3	220.1	0.70	154.1	_	1.00	3,223
Aug.	0.6	207.7	0.70	145.4	_	1.00	3,042
Sep.	0.4	189.0	0.70	132.3	_	1.00	2,767
Oct.	4.1	155.0	0.70	108.5	2.7	1.00	2,213
Nov.	3.2	105.0	0.70	73.5	2.3	1.00	1,489
Dec.	9.0	89.9	0.70	62.9	6.2	1.00	1,186
	102.0				73.2		27,021
Alfarf	a: MEC(E)	= 16, LR	= 0.022				
36 h	Dain	Diff.	W.C	ET(C)	R (E)	٨	V
Month Jan.	Rain 9.3	ETo 89.9	KC 0.85	76.4	6.9	$\frac{A}{1.00}$	$\frac{\sqrt{1,421}}{1,421}$
Feb.	44.9		0.85	85.7	32.8	1.00	1,081
Mar.	15.0	100.8 145.7	0.85	123.8	12.4	1.00	2,278
Apr.	13.1	180.0	0.85	153.0	11.1	1.00	2,902
May	2.1	238.7	0.85	202.9	0.3	1.00	4,143
Jun.	0.0	228.0	0.85	193.8	-	1.00	3,963
Jul.	0.3	220.1	0.85	187.1		1.00	3,826
Aug.	0.6	207.7	0.85	176.5		1.00	3,610
Sep.	0.4	189.0	0.85	160.6		1.00	3,285
Oct.	4.1	155.0	0.85	131.7	3.2	1.00	2,628
Nov.	3.2	105.0	0.85	89.3	2.0	1.00	1,785
Dec.	9.0	89.9	0.85	76.4	6.7	1.00	1,425
	102.0	0,00	0.05	, , , ,	75.5		32,348

Table A-5-2 (4) Unit Irrigation Water Requirement

Sorghum: MEC(E) = 18, LR = 0.019

			120	ET(C)	R(E)	Α	V
Month	Rain	ETO	KC		$\frac{107}{2.2}$	0.50	531
Oct.	$\frac{-4.1}{}$	155.0	0.35	54.3			
Nov.	3.2	105.0	0.60	63.0	1.9	1.00	1,246
		89.9	1.00	89.9	6.6	1.00	1,699
Dec.	9.0	89.9				1.00	1,694
Jan.	9.3	89.9	1.00	89.9	6.8		•
Feb.	44.9	100.8	0.75	75.6	32.4	1.00	880
				87.4	11.3	0.50	776
Mar.	15.0	145.7	0.60	07.4		0.50	6 025
	85.5				61.2		6,825

Table A-5-3 Average Monthly Effective Rainfall as Related to Average Monthly Effective and Mean Monthly Rainfall

c, 125. S Average Monthly Effective Rainfall as Related to Average Monthly 112. 쫎 'n 75.0 62.5 ETcrop and Mean Monthly Rainfall 50.0 37.5 25.0 12.5  $\infty$  $\infty$ σı 검 검 Rainfall (mm) Monthly Mean Average Monthly ETCrop in mm

= Effective rainfall as calculated × Correction factor Source: FAO Irrigation and Drainage Paper No. 24 (1977 Edition), Table 34 (IIII Correction factor = 1.04 (for effective storage of 130Remarks: Effective rainfall

verage Monthly Water Supply Requirement per Sectar by Crops

	Average	312ha	1,509	1,016	1,792	1,963	2,802	2,680	2,587	2,441	2,222	1,949.	1,611	1,514	24,086
Feed Crop	Sorghum	101ha	1,694	880	776							531	1,246	1,699	6,825
	Alfalfa	211na	1,421	1,081	2,278	2,902	4,143	3,963	3,826	3,610	3,285	2,628	1,785	1,425	32,348
	Average	2,340ha	1,122	756	1,796	2,290	3,306	3,179	3,069	2,895	2,634	2,104	1,419	1,127	25,697
do	Mango	83ha	1,181	808	1,893	2,408	3,475	3,339	3,223	3,042	2,767	2,213	1,489	1,186	27,021
Fruit Crop	Banana	128ha	1,548	1,207	2,481	3,149	4,477	4,293	4,144	3,911	3,559	2,856	1,939	1,553	35,117
	Line	309ha	815	420	1,291	1,677	2,465	2,385	2,302	2,173	1,977	1,559	1,068	820	18.951
	Dates	1,820ha	1,142	779	1,829	2,328	3,359	3,228	3,116	2,940	2,675	2,139	1,439	1,146	26,120
:	Average	183ha	1,059	462	673	657	476	124			334	904	1,246	1,216	7,151
	Others	23ha	1,440	995	788								837	1,175	5,236
	Okra	8ha		105	1,875	2,941	3,947	2,835							11,704
Vegetable	Potato	15ha	2,358	937									676	1,821	6,066
	Tomato	16ha	1,428	1,383	2,290	2,366	1,583				388	313	942	668	11,594
	Garlic	25ha	1,332	983	2,138	2,356	1,205		1		677	1,525	1,457	1,249	12,923
	Onion	96ha	720							:	395	1,274	1,489	1,276	5,154
	Month		Jan.	Feb.	Mar.	Apr.	May	Jun.	Jul.	Aug.	Sep.	Oct.	Nov.	Dec.	Total

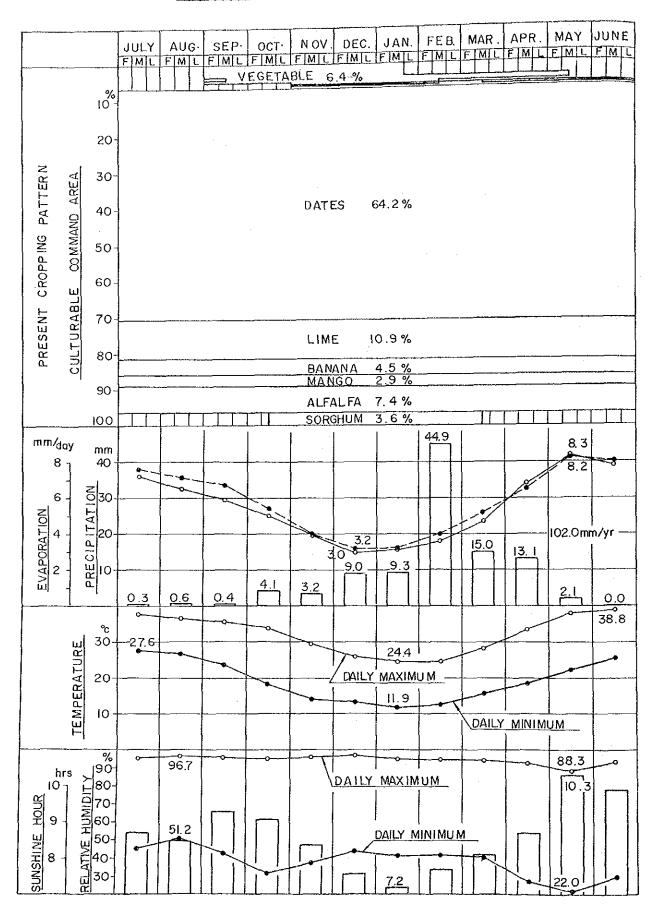
Table A-5-5 Unit Irrigation Water Supply Requirement by Land Use

Average Monthly Water Supply Requirement per Hectare by Land Use

(Unit: cu.m/ha)

				100											
	Average	495hz	1,343	811	1,370	1,480	1,942	1,735	1,631	1,539	1,524	1,563	1,476	1,403	17,825
	Sorghum	101ba	1,694	880	776				٠			531	1,246	1,699	6,825
	Alfalfa	211ha	1,421	1,081	2,278	2,902	4,143	3,963	3,826	3,610	3,285	2,628	1,785	1,425	32,340
8	Others	23ha	1,440	966	788								837	1,175	5,236
Upland Crop	Okra	8ha		105	1,875	2,941	3,947	2,835							11,704
	Potato	15ha	2,358	. 937			·						676	1,821	6,066
	Tomato	16ha	1,428	1,383	2,290	2,366	1,585				388	313	942	668	11,594
	Garlic	25ha	1,332	983	2,138	2,356	1,206				677	1,525	1,457	1,249	12,923
	Onion	96ha	720								395	1,274	1,489	1,276	5,154
	Average	520ha	1,054	675	1,680	2,156	3,121	3,007	2,902	2,740	2,493	1,983	1,350	1,058	24,219
Orchard	Mango	83ha	1,181	808	1,893	2,408	3,475	3,339	3,223	3,042	2,767	2,213	1,489	1,186	27,021
Ö	Banana	128ha	1,548	1,207	2,481	3,149	4,477	4,293	4,144	3,911	3,559	2,856	1,939	1,553	35,117
	Lime	309ha	815	420	1,291	1,677	2,465	2,385	2,302	2,173	1,977	1,559	1,068	820	18,951 35,117
	Dates	1,820ha	1,142	677	1,829	2,328	3,359	3,228	3,116	2,940	2,675	2,139	1,439	1,146	26,120
	Month		Jan.	Feb.	Mar.	Apr.	May	Jun.	Jul.	Aug.	Sep,	Oct.	Nov.	Dec.	Total

Figure A-5-1 Present Cropping Pattern



Present Cropping Calender

June		
Мау		
A pr.		
Mar.		
Feb.		
Jan.		
Dec.		
Nov.		
0ct.		
Sep.		
Aug.		
July		
Crop	Vegetable Onion Gattic Tomato Potato Okra Others	Lime Banana Mango Alfalfa Sorghum

# ANNEX B GROUNDWATER

# ANNEX B

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в-26	- Do - (12-months Moving Average: Case 1)	5 9
B-27	Simulated Groundwater Tables (Case 2)	6 0
B-28	Deviation of Groundwater Level (Case 2)	6 1
B-29	- Do - (12-months Moving Average: Case 2)	6 2
B-30	Simulated Groundwater Tables (Case 3)	6 3
B-31	Deviation of Groundwater Level (Case 3)	6 4
B-32	- Do - (12-months Moving Average: Case 3)	6 5

#### Annex B Groundwater

#### Groundwater

#### B-1. Hydrogeological Unit

The Project Area consists of the following three hydrogeological units; the impervious formations, terrace deposits, and alluvial deposits.

The impervious formations consist of the Hawasina group, basic volcanic rocks, and tertiary sedimentary formations forming the main central ranges and their flanks in the upstream of the Wadi Jizzi.

The Hawasina group consists of silicified limestone, mudstone, and chart with the well stratified beds of some ten centimeters in thickness. The basic volcanic rocks are mainly composed of Oman Ophiolite forming the main central ranges in the mountainous area in the mid-basin of the Wadi Jizzi.

The tertiary sedimentary formations consist of mudstone and limestone, etc., forming the basemant of terrace deposits and low hills in the west edges of the gravel plain. The formations were found by the exploratory well drillings at the gravel plain in the depth of 40 to 50 meter below ground surface in weakly consolidated condition.

The Terrace deposits have a large exposure in the middle stream of the Wadi Jizzi and the west edge of the gravel plain but their distribution is limited to an upper stream of the Wadi. The deposits are divided into four kinds of sediments based on height of their platform, three of which are distributed in the Wadi Jizzi basin and the lower one is distributed in the mouth of the catchment of the Wadi Bani Umar, forming the alluvial fan.

The deposits are composed of partially cemented sand and gravel of fluvial origin with various sizes of grains of the basic volcanic and sedimentary rocks.

Although the deposits seem aquifure, occasionally their uncemented thin layers of sand and granule among the deposits function as aquifers. Therefore, they act as part of aquifer in terms of hydrogeology.

The alluvial deposits are exposed in a limited area along the wadi course in the catchment; however, they have a large exposure in the coastal plain. The deposits consist of sand and gravel with partially cemented beds of alluvial origin. Thickness of the deposits range from several meters at the river beds in the catchment to 10 m in the mouth of gravel plain and more than 80 m in coastal plain where the deposits grow excellent as unconfined aquifers.

#### B-2 Aquifer Characteristics

The main aquifers in the Project Area are restricted to the terrace deposits and alluvial deposits. The aquifer characteristics in the coastal plain, especially of alluvial aquifers, have been obtained by the aquifer tests since 1973. Specific capacity and transmissivity of the alluvial aquifers in the east edge of the gravel plain range from 30 to 60 cu.m/hr/m and from 4,000 to 50,000 sq.m/day, respectively. Storativity which was obtained by aquifer tests at production well No. 1 of Sohar Expansion Farm was calculated at 0.05 in an average.

Permeability coefficient ranges from  $2.5 \times 10^{-3}$  to  $8.0 \times 10^{-1}$  m/min with an average of  $2.0 \times 10^{-2}$  m/min.

#### B-3 Hydrogeological Structure

The groundwater basin comprising the terrace deposits and alluvial deposits is enclosed by the impervious formations at the north and west edges with depth less than 40 m and its thickness to the east reaches more than 100 m at the coast. The basin ends near Qabail and Majis where the impervious formations crop out near the sea. Location of the south end of the basin developing in the downstream of the Wadi Jizzi is found at the north of Wadi Kadaq where trenches of groundwater table extend. The entire groundwater basin extends along the coastal plain over an area of about 10 km long and 10 to 20 km wide. The groundwater being recharged from the Wadi Jizzi has been considered to flow down along two ridges of groundwater table that extend from Fasiqah and Qabail toward inland.

#### B-4 Movement of Groundwater

Groundwater tables observed in the Wadi Jizzi area are shown in Table B-1. The transition of groundwater tables obtainable from 1974 to 1984 data is given in Figure B-2. In eleven years during when groundwater level observations are available, the years 1977 and 1978 are considered to be high-water years. On the contrary the years in and after 1983 are recognized to be low-water years. Groundwater movement in both the high-water period and low-water period are briefly summarized as follows:

## 4-1 Groundwater Movement in Low-water Period

Groundwater contour lines drawn from the data observed in

April, 1985 are shown in Figure B-3. As is seen in the said figure, the groundwater table during the low-water period is characterized by three ridges extending from Qabail, Fasiqah and Sohar to the Wadi Jizzi and Wadi Khadaq and by two trenches. Among those ridges, two ridges which extend from Qabail and Fasiqah toward inland ranges along the middle reaches of the Wadi Jizzi, and that extending from Sohar toward inland ranges the Wadi Khadaq. Therefore these ridges are considered to form channels of groundwaters being recharged from both the Wadis.

Groundwaters flowing along such channels are usually provided in the direction of trenches. Due to a small amount of groundwater provided toward trenches, the shape of such trenches become clear especially during the low-water period. The groundwater level is normally kept higher than 1.0 m above mean sea level around Fasiqah, while it is commonly lower than 0.5 m above msl around Amq and Sohar where groundwater trenches exist. The average hydraulic gradient of groundwater in the coastal plain is calculated at 1:930.

#### 4-2 Groundwater Movement in High-water Period

Figure B-4 presents groundwater contour lines observed in 1978. The contour lines run almost parallel with the coast line and the groundwater tables are maintained higher than 1.0 m above msl at all places along the coast line. The average hydraulic gradient of groundwater is measured at 1:480.

#### 4-3 Mean Groundwater Table

The standardized groundwater tables obtainable by means of arithmetical mean are shown in Figure B-5. The contour lines generally run parallel with the coast line with a little uneven

portion. The contour of 1.0 m above msl is distributed at a distance of about 1.0 km from the coast line. Average hydraulic gradient of groundwater is also calculated at 1:590.

To express annual changes of groundwater tables, standard deviations are calculated as shown in Figure B-6. As a general tendency, deviations of less than 0.5 m are observed in the area along the coast whereas values exceeding 2.0 m are generally found inland. However even along the coast, a larger value of 1.5 m is collected from Amq where a trench of groundwater table exists.

#### $_{\mbox{\footnotesize{B-5}}}$ Groundwater Flow in View of Iso-Ec Contour Line

Distribution of EC at the surface layer of groundwater along the Wadi Jizzi is presented in Figure B-7. The Iso-Ec contour lines also run parallel with the coast line indicating conductivities ranging from 5,000 to 3,000 micro mho/cm. 500 mmho/cm contour line is distributed at a distance about 3.5 km from the coast line, and a stable value of conductivity of 470 mmho/cm is observed elsewhere inside of 500 mmho/cm line. The groundwater indicating conductivity higher than 500 mmho/cm is considered to be affected by sea water intrusion.

According to the EC loggings collected at OA-2, the vertical section of groundwater along the coast line can be divided into three layers. The upper and middle layers are formed by groundwater runoff, while the lower layer is characterized by sea water intrusion. Variations of depth and conductivity in cach layer are shown in Table B-2 as well as in Figure B-8. As is seen in the said table and figure, no significant difference has been observed in the last eleven years. It is therefore concluded that continuous intrusion of sea water has not occurred yet in the Project area.

### B-6 Chemical Quality of Groundwater

The result of chemical analysis conducted in 1982 and 1985 are presented in Table B-3 and are plotted on a key diagram of hydrochemistry as shown in Figure B-9 and B-10. The majority of groundwaters are categorized in type I field indicating fresh water. The groundwater sampled at the coast is classified in type III field, which characterizes hydrochemical contamination caused by sea water intrusion. Waters collected from Wadis and groundwater channels are categorized in type IV field.

#### B-7 Change of Groundwater Storage

A change of groundwater storage can be defined as the change of groundwater table multiplied by storativity. Storativity of 0.05 was adopted for the study. An average annual change of groundwater table in the recent eleven years from 1974 to 1984 is calculated at 43 mm in defect, corresponding to a volume decrease of 1.02 MCM per annum. Annual changes of groundwater storage so computed are shown in Table B-4 as well as in Figure B-11. The maximum increase of groundwater storage is found at 6.36 MCM per annum between two years 1975 and 1976, while the maximum decrease at 11.54 MCM/yr between 1981 and 1982.

#### B-8 Groundwater Runoff to the Sea

Groundwater runoff to the sea is obtainable by use of Darcy's formula. Darcy's formula is given in the following equation:

Q = K. A. I

where Q : Flow volume

K: Permeability coefficient =  $2.0 \times 10^{-2}$  m/min.

A : Flow area

 $A = L \cdot H$ 

L: Length of coast line = 10 km

H: Depth of upper layer by EC logging = 17.5 m

Thus,  $A = 1.75 \times 10^5 \text{ sq.m}$ 

I : Hydraulic Gradient (Refer to in Table B-5)

The hydraulic gradient of groundwater table in the coast area is calculated at  $1.70 \times 10^{-3}$  (1:590) as an average in eleven years from 1974 up to 1984. Groundwater to the sea is thus computed at 3.13 MCM per annum

Figure B-12 presents annual change of groundwater runoffs to the sea. From this, the maximum runoff is seen at 4.26 MCM in 1977 and the minimum runoff at 1.15 MCM in 1983.

Groundwater Level

Standard	Deviation(TD)			0.35	1.58	0.98	0.59	0.82	1.37	0.52	0.24	0.72	0.29	0.78	0.25
Average	(m)		,	6.23	23.30	10.99	9.77	8, 99	8,56	1,69	1.72	9,55	6.14	2.22	4.29
	1985	.1	7.26	6.60	25.60	11.78	9,99	9,54	10, 41	2.14	2.02	10.55	6.37	2.52	4.58
	1984	1	4.72	6.47	24.47	12.36	10.70	9.32	9.91	2.28	1.86	10.65	9.00	2.05	4.21
	1983	ı	2.81	6.45	(24.67)	(11.82)	10, 18	9,63	9.07	2.14	1.89	8.43	(6.37)	4.06	4.77
	1982		1	6.26	25.60	11.54	(10, 17)	9.82	90.6	1.69	(1.87)	10.49	(6.32)	2.26	(4.46)
(m)	1981		11.88	6.25	22.05	11.15	(9.85)	9.05	(8.82)	(1.79)	(1.74)	(9.71)	6.56	2.44	(4.33)
(Depth)	1980	ı	6.53	6.03	22.85	9.45	(9.47)	8.56	(8.00)	(1.46)	1.58	(6.03)	6.29	1.70	(4.16)
Mean Groundwater Level (Depth	1979	1	6.62	6. 12	23.31	10.12	(9.55)	8.59	(8.11)	(1.52)	(1.62)	(9.34)	5.87	1.88	(4.19)
oundwater	1978	22.30	4.48	5. 79	21.65	10.16	8.57	8.23	7.29	1.33	1.47	9.03	5.82	2.06	4.01
Mean Gr	1977	17.42	2.11	6.07	21.38	(10.11)	9, 49	8.51	7.28	1, 19	1.54	8.84	5.74	1.78	3.96
	1976	11.42	6.40	6.05	(22.33)	10.37	9.57	8.85	7.70	1.28	1.52	60.6	6.35	1.74	4.11
	1975	21.32	10.27	6.15	(24.35)	(10.95)	9.88	8.69	8.29	1.71	1.82	9.80	6.16	2.11	4.37
	1974	1	ì	67.9	24.38	11.10	(8.85)	(6.07)	(8.78)	(1.78)	(1.73)	(6.69)	5.85	2.03	(4.32)
		0A-1	AE-49	OA-2	EA-1	WS I -26	AE-61	AE-62	AE-91	AE-93	A E —99	AE-101	AE-104	AE-142	AE-162

Table B-2 Change of EC during 1974-1985

 $1,150 \sim 6,400$ 6,400~50,000 6, 400 1, 150 50,000 Apr. 85  $24 \sim 36$  $36 \sim 42$ 42~57 (soqu)  $\sim$ 24 Depth  $\sim$ 25 Change of EC during  $1974 \sim 1985$  at 0A-2 $900 \sim 5,500$ 5,500~32,000 四3/四万 50,000+ 900 5, 500 Mar. 82  $40 \sim 61$  $\sim$ 22  $22 \sim 34$  $34 \sim 40$ (spam) Depth 61 3,500~17,000  $880 \sim 3,300$ MJ/WZ 18,000 880 3,500 Feb. 74  $61 \sim 147$  $32 \sim 40$  $30 \sim 61$  $27 \sim 32$  $\sim$ 27 (spdm) Depth Transitional Layer Transitional Layer 3rd Layer ist Layer 2nd Layer

Chemical Analysis Data

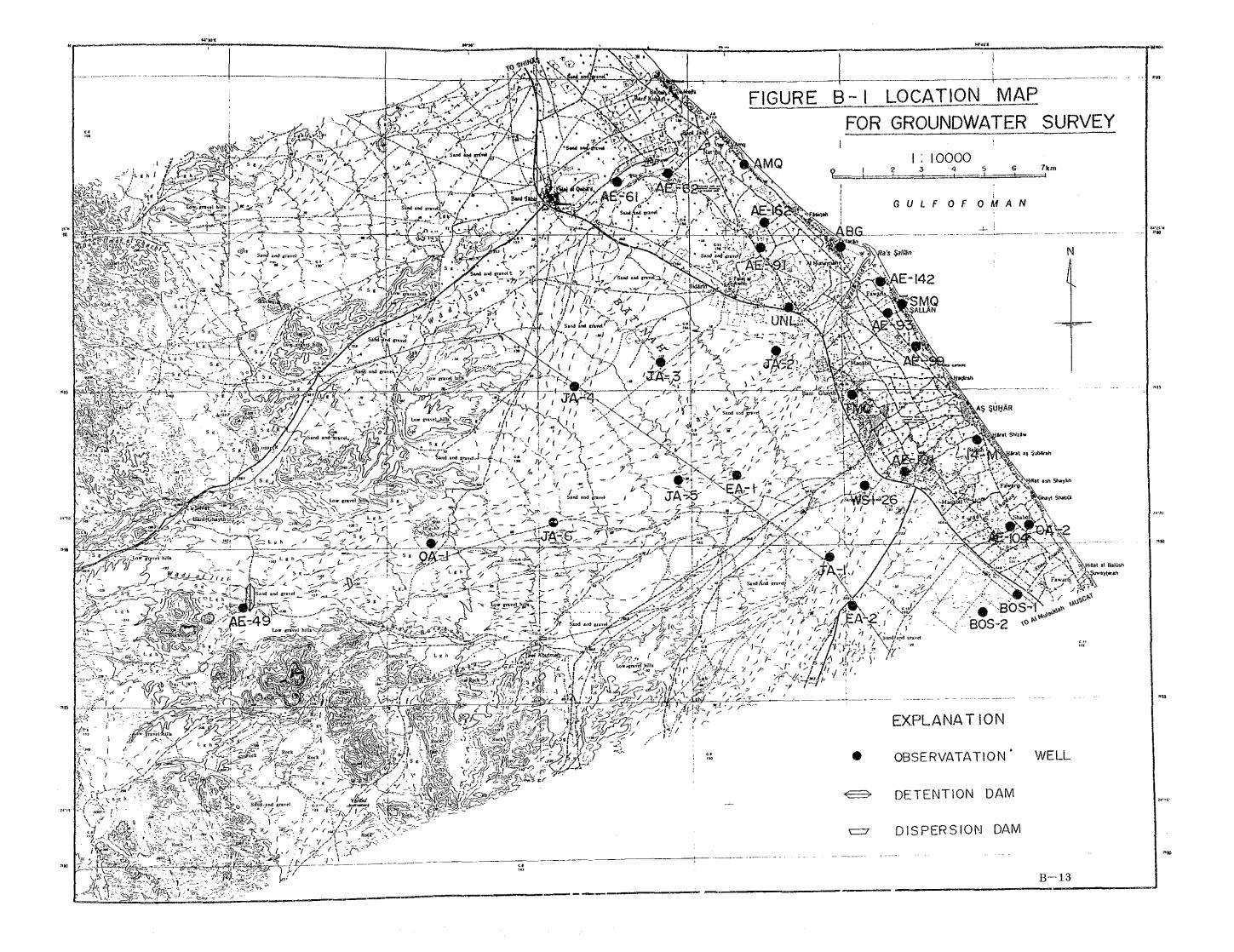
		Total	9.31	3.31	4.76	4.62	3.92			Total	9.46	3.66	4.22	5.51	4.11	3.56	9.86	4.15	6.51	11.46	9.91	8.64	14.74	24.33	5.65
		S0.	1.66	0.16	0.66	0.42	0.42			- <b>3</b> 05	1.66	0.16	0.42	0.66	0.66	0.68	1.16	1	0.66	3.91	1.91	0.39	0.99	7.33	1
	Anions, me/L	_าว	4.25	1.10	1.80	1.30	1.25		Anions, me/L	ر اد	4.40	1.20	1.40	2.05	1,15	09.0	6.20	1.75	2.35	3.35	5.20	2.45	11.40	13.90	2 05
	Anions	HCO =	2.80	1.45	2.30	2.30	1.85	-	Anion	100F	3.40	2.30	2.40	2.80	2.30	2.30	2.50	2.40	3.50	4.20	2.80	5.20	2.40	3.10	03.0
q	-	£00	9.0	9.0	1	9.0	0.4				1	1	1	ı	ı	ı	1	-	l	1		ı	. ]		1
Analysis Data		Totai	9.66	3.74	4.83	5.02	4, 48			Total	9.69	3.84	4.40	5.87	4.19	3.82	9.88	4.22	6.54	11.67	9.93	8.80	15.43	24.08	5 70
		‡ <b>6</b> %	4. 10	2.50	2.80	3, 45	2.90		Cations , me/L	‡ \$	3, 10	2.14	2.43	3.05	2.15	0.97	4.57	1.57	2.48	2.90	5.65	3.67	3.60	13.35	00 0
Chemical	Cations, me/L	‡ ¢3	1.10		0,40	0.20	0.20			‡g	1.70	0.26	0.27	0.55	0.55	1.48	1.18	0.58	1.07	2.60	1.05	0.53	1.05	2.60	1 05
	Catio	+_=	0.11	90.0	0.07	0.07	0.07		Catio	+*	0.08	0.04	0.05	0.05	0.05	0.10	0.08	0.08	90.0	90.0	0.08	0.08	0.09	0.08	000
		+ pu	4.35	1.74	1.55	1.30	1.30			+ 52	4. 78	1.35	1.65	2.22	1.44	1.27	4.05	1.99	2.93	6.11	3.15	4.52	10.67	8.04	2 27
	EC 254	25°C	1,006	367	518	489	413		EC	Umbos/cm 25°C	953	379	440	562	391	367	1,002	416	623	1,063	996	831	1, 467	2,078	677
	ž	Ξ.	7.0	7.4	7.6	7.7	7.45			<del></del>	8.1	6.7	8.0	8.0	8. 75	8.05	7.95	7.90	8.15	8.10	8.05	8,35	8.05	8.10	31.0
2)	Mell	₽	JA-1	JA-2	JA-3	JA-4	JA-5	5)	- III	Q.	JA-1	JA-2	JA-3	JA-4	JA-5	JA-6	0A-2	EA-1	4SI-26	AE-49	AE-62	AE-91	AE-93	AE-101	6-34
(1982	Sample	£	-	2	co.	7	3	(1985)	Sample	S.		2	8	4	5	9	7	8	6	10	=	12	13	14	1

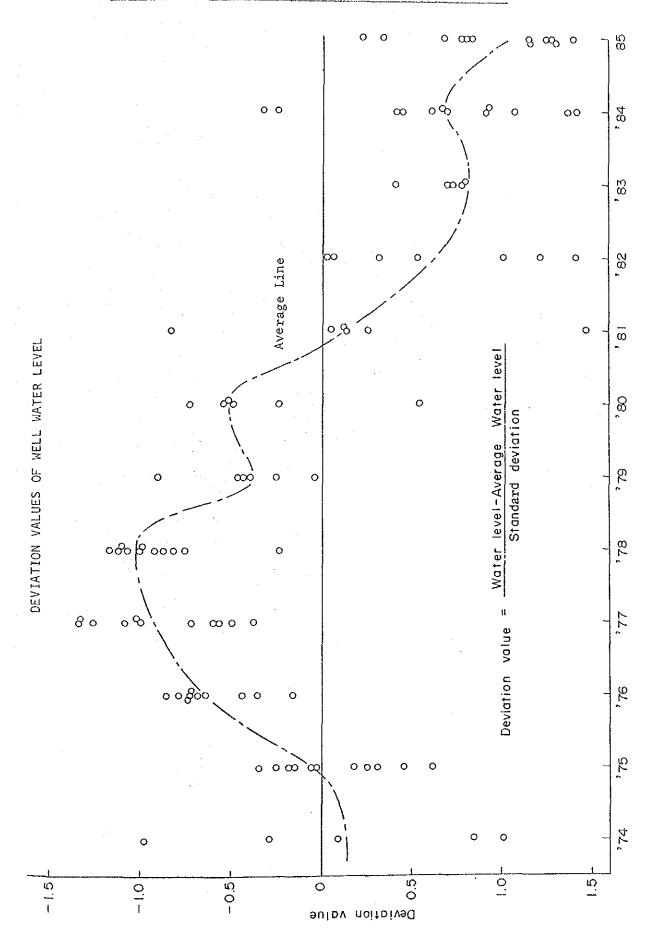
Water Level Fluctuation and Change of Storage

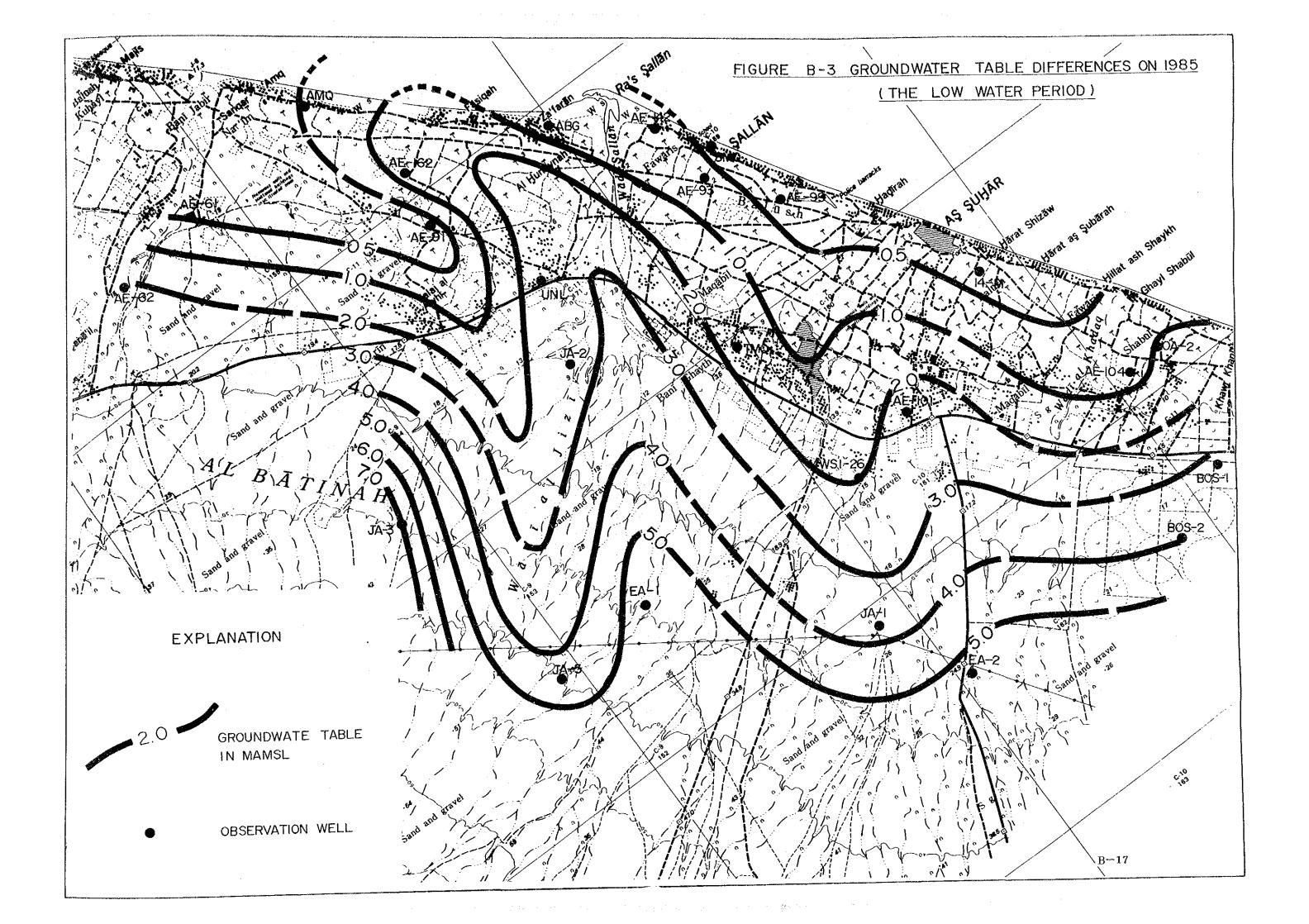
Average	1974~1984													-0.04	-1.02	
1983	~1984	-0.02	+0.20	-0.54	-0.52	+0.31	-0.84	-0.14	+0.03	-2.22	+0.37	+2.01	+0.56	-0.07	-1.65	
1982	$\sim$ 1983	-0.09	+0.93	-0.28	-0.01	+0.19	-0.01	-0.45	-0.02	+2.06	-0.05	-1.80	-0.31	+0.01	+0.24	
1981	~1982	-0.01	-3.55	-0.39	-0.32	-0.80	-0.24	+0.10	+0.13	-0.73	+0.24	+0.18	-0.13	-0.49	-11.64	
1930	~1981	-0.22	+0.80	-1.70	-0.38	-0.46	-0.82	-0.33	-0.16	-0.68	-0.27	-0.74	-0.17	-0.43	-10.13	
1979	~1980	+0.19	.+0.46	+1.17	+0.08	+0.03	+0.11	+0.06	+0.04	+0.31	-0.42	+0.18	+0.03	+0.18	+4.24	
1978	~1979	-0.33	-1.66	-0.46	-0.98	-0.36	-0.82	-0.19	-0.15	-0.31	-0.05	+0.18	-0.18	-0.44	-10.36	
1977	~1978	+0.28	-0.27	-0.05	+0.92	+0.28	-0.01	-0.14	+0.07	-0.19	-0.08	-0.28	-0.05	+0.04	+0.94	
1976	~1977	-0.05	+0.95	+0.27	+0.08	+0.34	+0.42	+0.09	-0.02	+0.25	+0.61	-0.04	-0.15	+0.23	+5.42	
1975	~1976	+0.10	+2.02	+0.58	+0.31	-0.16	+0,59	+0.43	+0.30	+0.71	-0.19	+0.37	+0.26	+0,43	+10.13	
1974	~1975	+0.34	+0.03	+0.15	-0.05	+0.38	+0.49	+0.07	+0.09	-0.11	-0.31	-0.08	-0.05	+0.11	+1.88	(+2.51)
		0A-2	EA-1	WS1-26	A E -61	A E -62	AE-91	AE-93	AE-99	AE-101	AE-104	AE-142	A E - 162	Average (m)	Change of Storage	(MOM/yr)
			Mater Level Fluctuation (M/yr)											Aver	Chan	

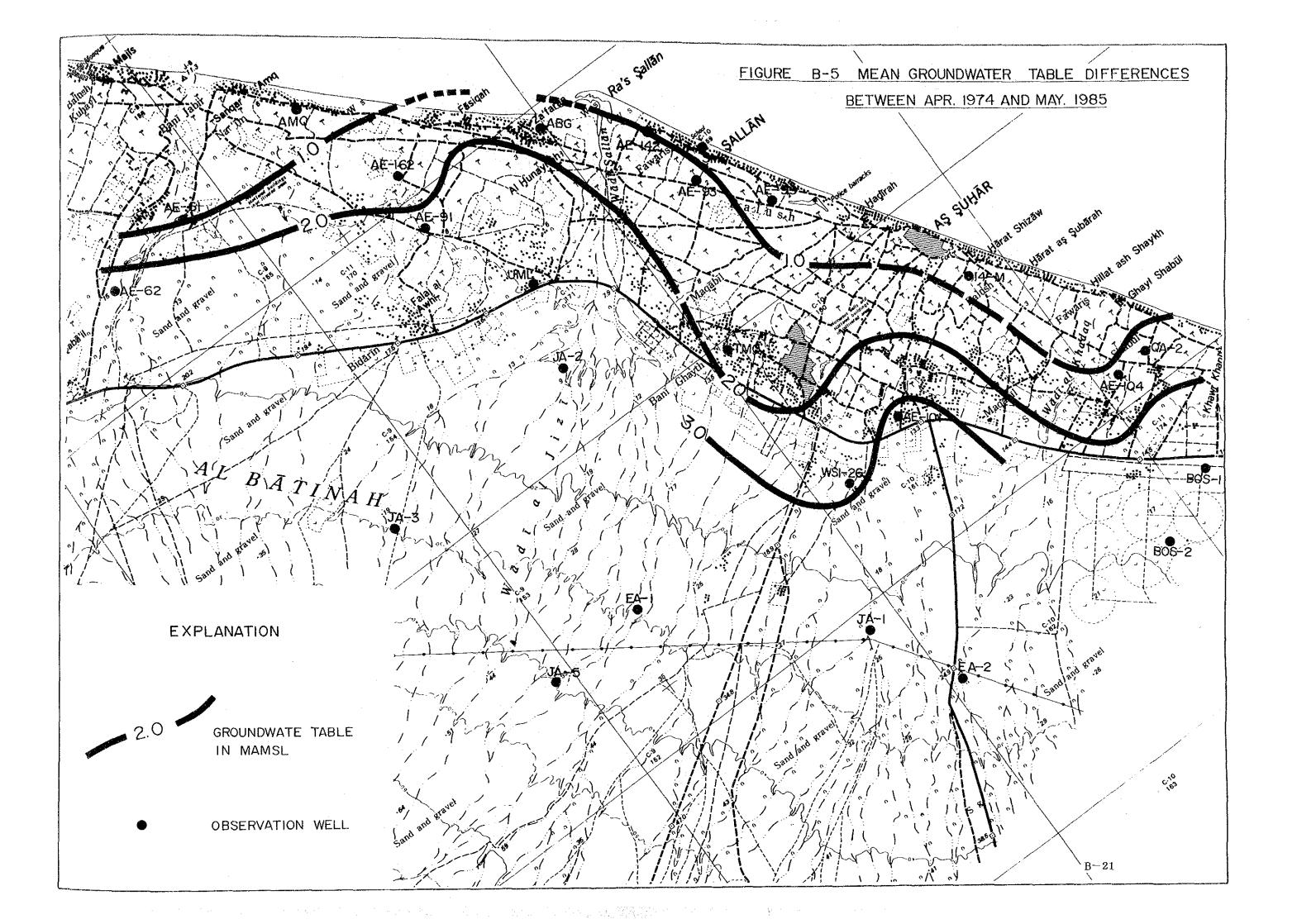
Hydraulic Gradient and Groundwater Run-off

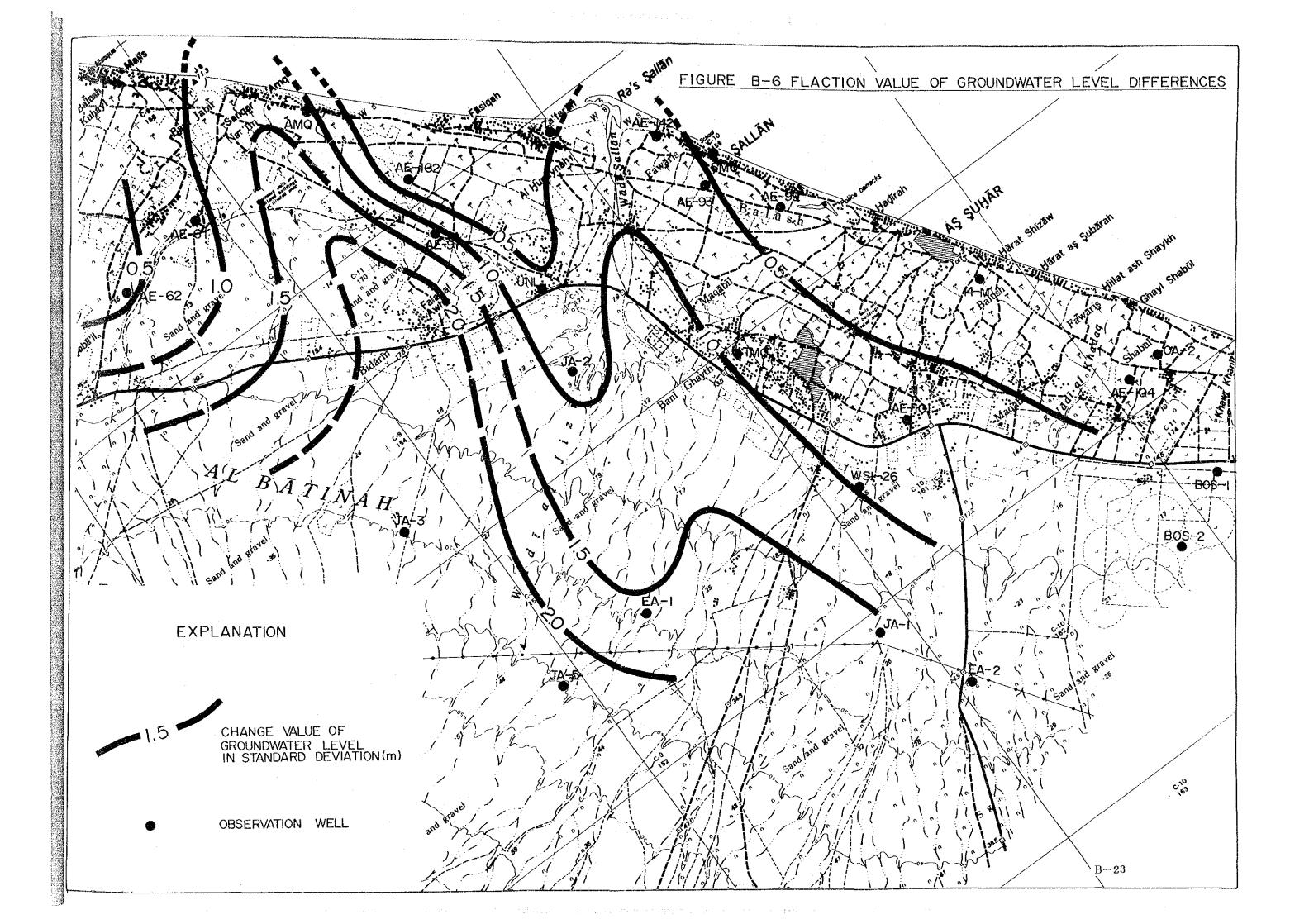
Average 1974~1984	2.17	0.44	1.11	1.21	0.81	3.32	1.07	1.04	1.51	1.70	3, 13
1984	1.77	0.27	0.39	0.58	0.64	2.13	1.19	1.18	1.56	1,43	2.63
1983	1.80	0.14	0.83	0.72	0.61	4.35	0.82	-0.83	1.07	0.59	1.09
1982	2.07	90.0	0.83	1.17	0.63	2.29	0.87	0.99	1.34	1.45	2.67
1981	2.09	0.39	0.95	1.07	0.76	3.07	0.63	0.79	1.45	1,46	2.69
1980	2.40	0.58	1.37	1.40	0.92	3.75	06.0	1, 53	1.60	2.05	3.77
1979	2.27	0.57	1.32	1.34	0.88	3.44	1.32	1.35	1.57	1.96	3,61
1978	2.74	0.72	1.74	1.53	1.03	3.75	1.37	1,17	1.73	2.10	3.86
1977	2.34	09.0	1,74	1.67	96.0	3.94	1.45	1.45	1.77	2.19	4.03
1976	2.37	0.46	1.53	1.58	0.98	3.69	0.84	1.49	1.64	2.08	3.83
1975	2.23	0.53	1.23	1.15	0.68	2.98	1.03	1. 12	1.42	1.69	3.11
1974	1.74	0.49	0.97	1.08	0.77	3.09	1.34	1.20	1.46	1.68	3.09
	0A-2	AE-62	AE-91	AE-93	AE-99	AE-101	AE-104	AE-142	AE-162	Average	Groundwater Run-off (MCM/yr)
	Hydraulic Gradient (10×-3)										Groundw.

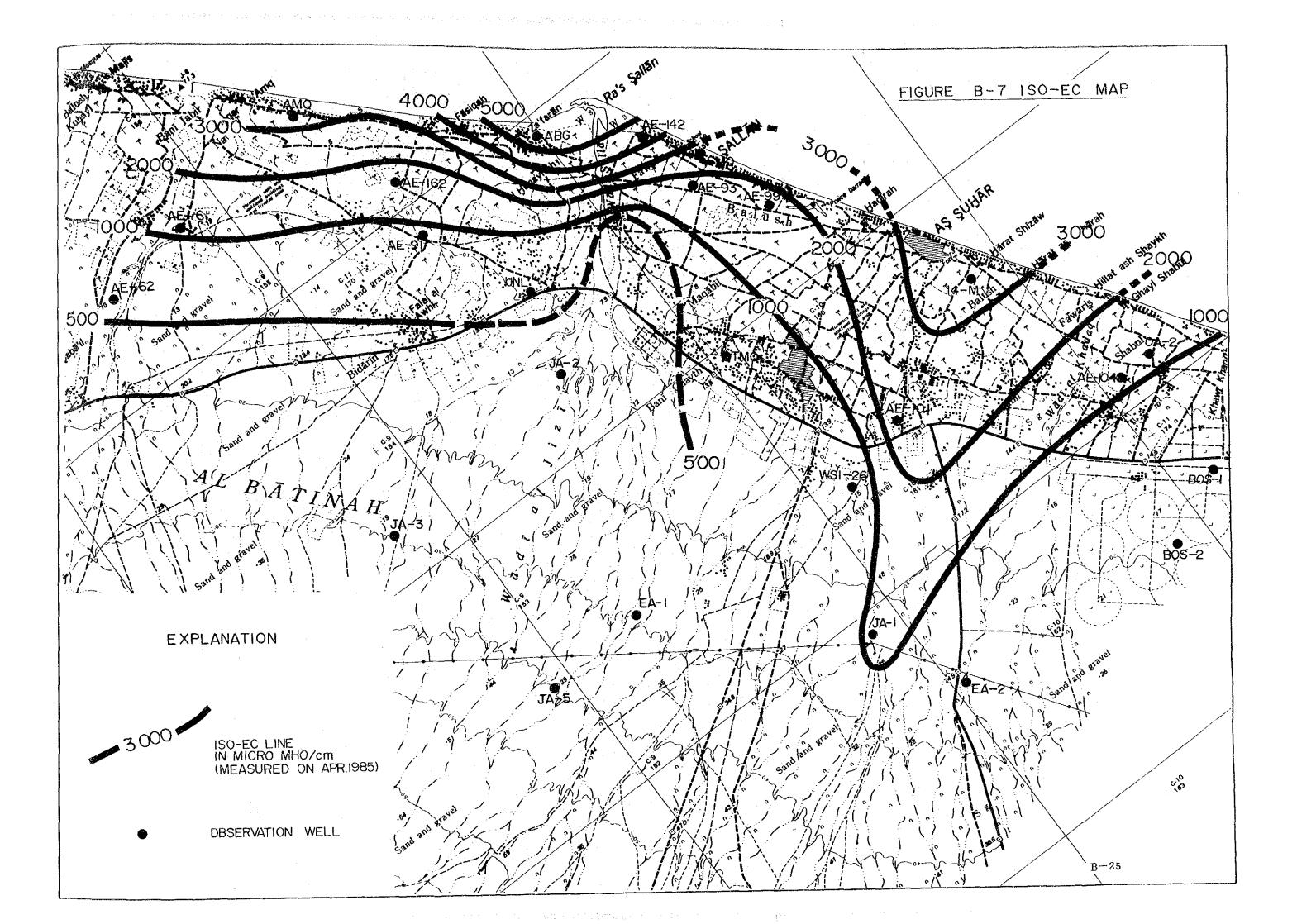


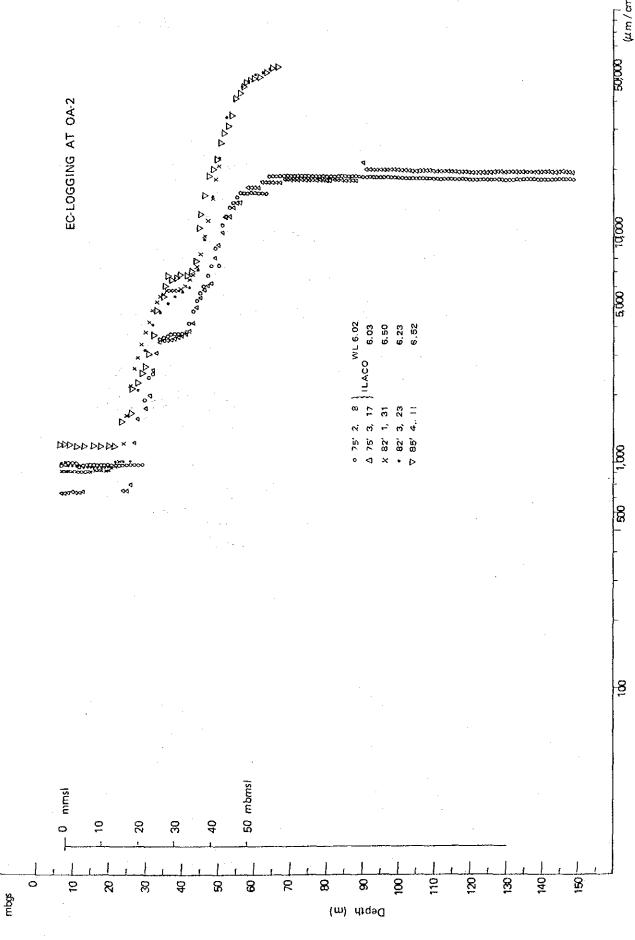




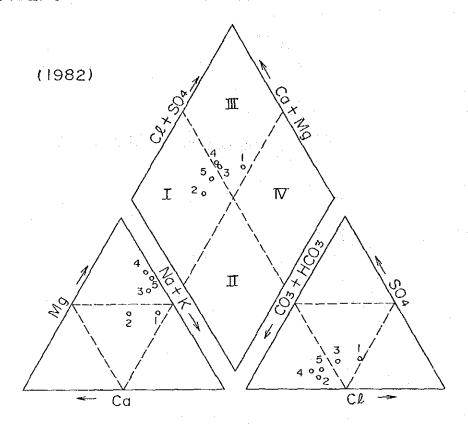


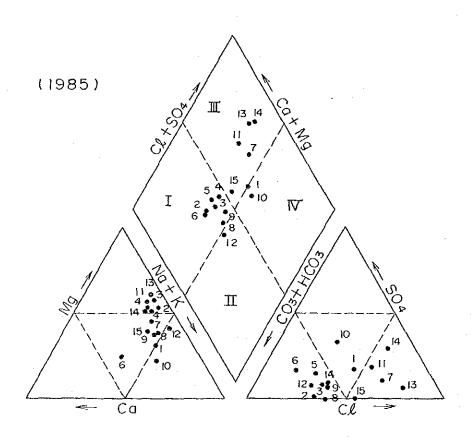


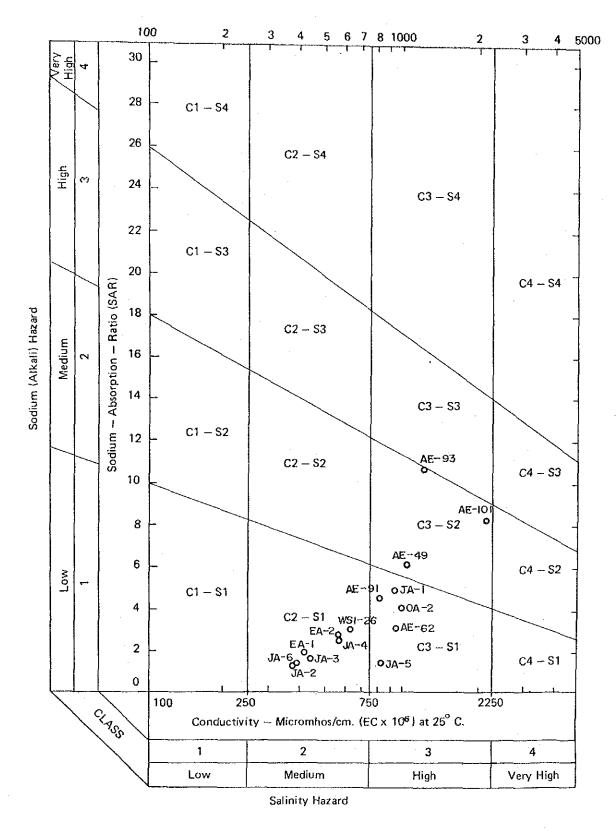




FIGERU B-9 WATER ANALYSIS DIAGRM







Source: Agricultural Handbook 60, U.S. Dept. of Agriculture

Class

#### Salinity or Conductivity

#### Sodium-Absorption Ratio

Low

Low-Salinity Water (C1) can be used for irrigation with most crops on most soils with little likelihood that soil salinity will develop. Some leaching is required, but this occurs under normal irrigation practices except in soils of extremely low permeability.

Low-Sodium Water (S1) can be used for irrigation on almost all soils with little danger of the development of harmful levels of exchangeable sodium. However, sodium-sensitive crops such as stone-fruit trees and avocados may accumulate injurious concentrations of sodium.

Medium

Medium-Salinity Water (C2) can be used if a moderate amount of leaching occurs. Plants with moderate salt tolerance can be grown in most cases without special practices for salinity control.

Medium-Sodium Water (S2) will present an appreciable sodium hazard in fine-textured soils having high cation-exchange capacity, especially under low-leaching conditions, unless gypsum is present in the soil. This water may be used on coarse-textured or organic soils with good permeability.

High 3 High-Salinity Water (C3) cannot be used on soils with restricted drainage. Even with adequate drainage, special management for salinity control may be required and plants with good salt tolerance should be selected.

High-Sodium (S3) may produce harmful levels of exchangeable sodium in most soils and will require special soil management-good drainage, high leaching, and organic matter additions. Gypsiferous soils may not develop harmful levels of exchangeable sodium from such waters. Chemical amendments may be required for replacement of exchangeable sodium, except that amendments may not be feasible with waters of very high salinity.

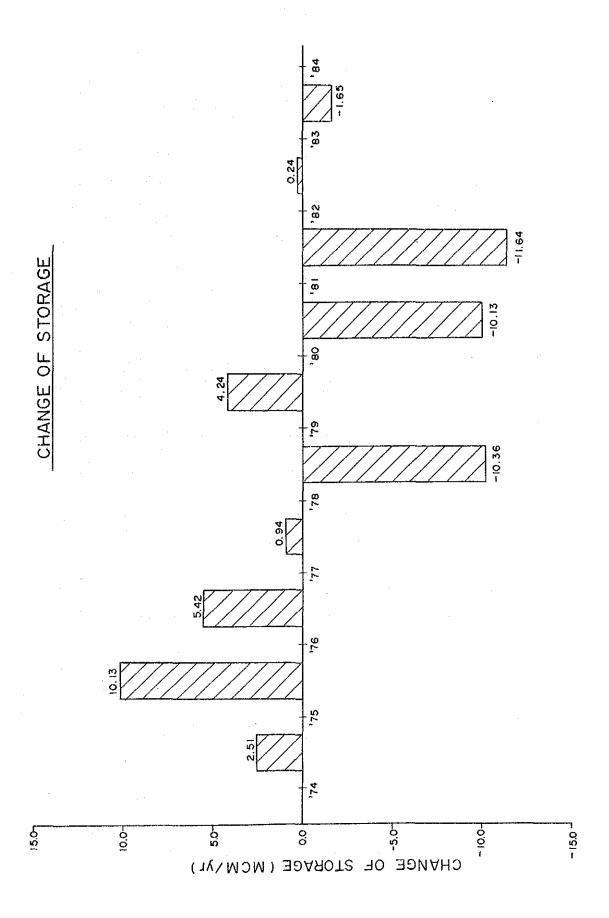
Very High 4

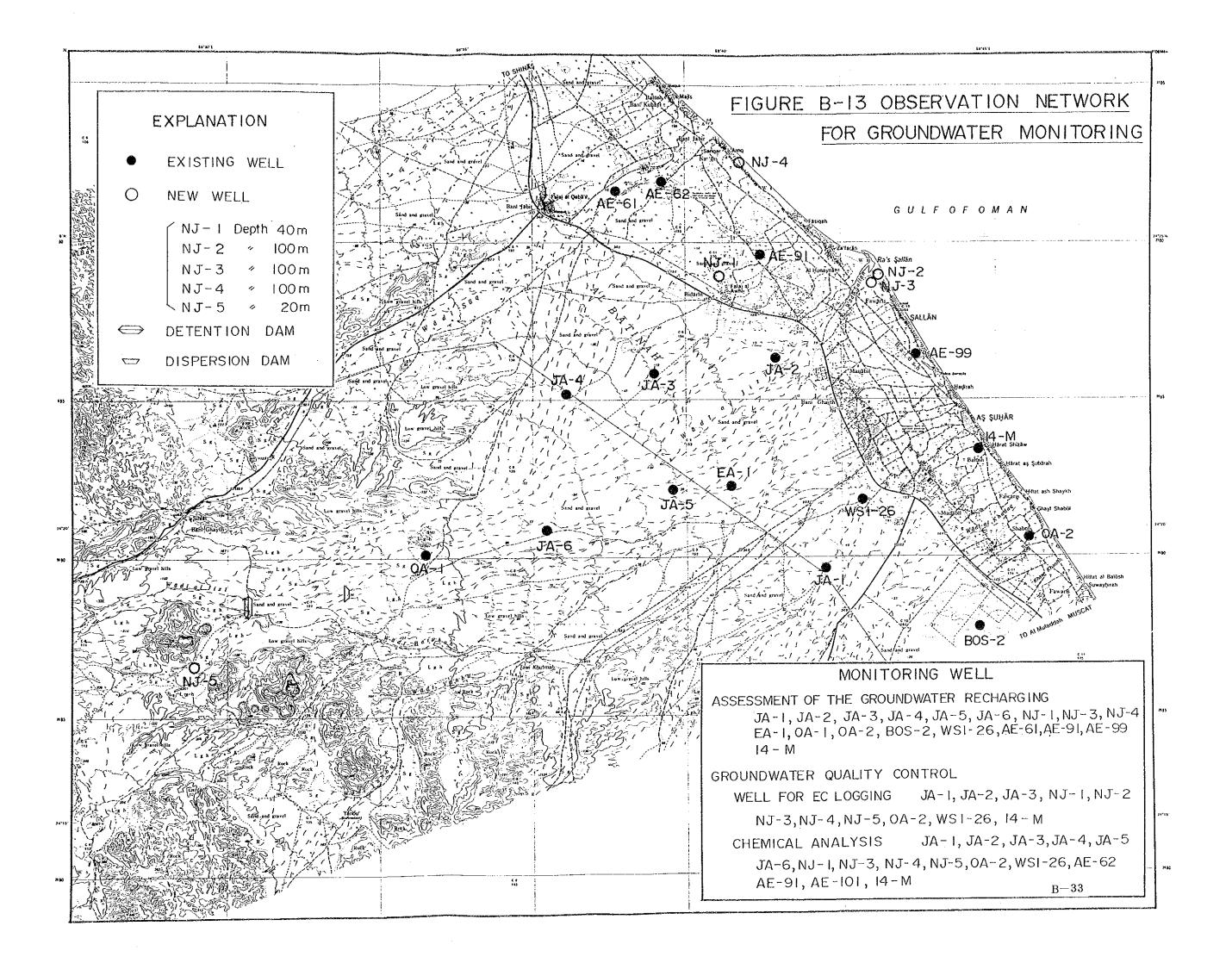
High Very High Salinity Water (C4) is not suitable for irrigation under ordinary conditions, but may be used occasionally under very special circumstances. The soils must be permeable, drainage must be adequate, irrigation water must be applied in excess to provide considerable leaching, and very salt-tolerant crops should be selected.

Very High Sodium Water (S4) is generally unsatisfactory for irrigation purposes except at low and perhaps medium salinity, where the solution of calcium from the soil or use of gypsum or other amendments may make the use of these waters feasible

FIGURE 8-10 DIAGRAM FOR CLASSIFICATION OF IRRIGATION WATER

B-29





B-9 Computer Simulation of Groundwater Basin

# 9-1 Simulation Model\*/

A general purpose digital computer is used in developing and testing a two-dimensional diffusion model of a groundwater basin, aiming at (1) the development of the model and (2) the use of the model for analyzing and predicting dynamic behavior of the basin under imposed conditions.

Currently as in the past, the extractions from the coastal plain have exceeded its natural replenishment. As a result, water levels continue to recede. In addition, the groundwater quality near the sea-coast is impaired by the intrusion of sea water. These difficulties can be surmounted by constructing a detention dam for additional recharge of greater quantities of fresh water, and by controlling extractions. These and other remedial measures are evaluated and applied through groundwater basin management. Optimum results will be achieved when the dynamics of the basin are better understood.

In order to study the dynamic behavior of the groundwater basin, the details of its geometry, physical characteristics, and its associated replenishment and extraction data must be obtained. Unfortunately, the cost of determining such physical characteristics (transmission and storage properties) in the requied detail by field measurements is prohibitively large. On the other hand, the geometric details can, in many instances, be readily determined from a variety of records that accrue from geophysical studies, water quality analyses, well logs, etc. If such records are available for a given basin, and if replenishment and extraction data can be obtained in detail, then the computer can be used to find a representative set of transmission and

<sup>\*/</sup> Ground-water Management for the Nation's Future - Computer

Simulation of Ground-water Basins; by H.N. Tyson and E.M. Weber

Journal of the Hydraulics Division, Proceedings of ASCE, July, 1964

storage properties or coefficients. This is accomplished by a cut-and-try adjustment of the coefficients until the dynamic responses of the model satisfactorily matches the historical water level elevations measured in the field.

## 9-2 Physical Situation of Ground-water Basin

The groundwater basin comprising the terrace deposits and alluvial deposits coincides with depth of the impervious formations beneath the gravel plain.

The groundwater basin is enclosed by the impervious formations at the north and west edges with depth of less than 80 m and it thickens to the east up to more than 100 m at the coast. The basin ends near Qabail and Majis where the impervious formations crop out near the sea. Location of the south end of the basin is estimated at the south of Wadi Ahin where the impervious formations are croped out near the sea. An entire area of the groundwater basin mainly developing in the downstream of Wadi Jizzi extends about 20 km in length along the coast with 8 km width. Furthermore, depth of it is estimated 50 to 60 m at the west edge of the basin and it deepens to the sea up to more than 100 m. Depth of the basin, especially at the west edge is verified by the exploratory drilling at JA-5 and 6, the production well for Mining Co., TS-6, 7, 8 and 9, and geo-electric survey at lines ES-1 and ES-V4.

The groundwater basin in the west edge of the gravel plain is composed aquifers of the terrace deposits with depth ranging 40 m in maximum to less than 20 m at an outlet of catchment and it extends to the wadi beds in the catchment decreasing thickness of aquifer.

For the initial computer analysis, the groundwater complex is replaced by a simplified model. This model consists of a single equivalent aquifer whose local properties are composites of the corresponding properties of the several aquifers that make up the actual

structure. The thickness of the single aquifer is allowed to vary with position, and yet it is considered to be small compared to its lateral dimensions. The bounding edges of the model are irregular in shape. The model is divided into small polygonal zones which enclose observation wells. The dynamic response of the portion of the model included within each zone is represented by a single water level elevation. The size of the zones is dependent on the variations in replenishment, extraction, transmission, storage, and water level data. For purposes of testing the model against historical water level data, provision is made for the extraction or injection of time-varying flow rates from each of the zones.

## 9-3 Mathematical Model

The equation of continuity of an unconfined aquifer, in which there is no vertical variation of properties, is given by;

$$\nabla \overline{V} + S \frac{\partial h}{\partial t} + Q = 0$$
 ....(1)

in which 
$$h = \delta + z \qquad (2a)$$
 and 
$$\delta = \delta_0 + \overline{\delta} \qquad (2b)$$

 $\boldsymbol{\delta}_0$  and  $\overline{\boldsymbol{\delta}}$  are, respectively, the mean and perturbation components of  $\boldsymbol{\delta}.$ 

Darcy's Law provides the equation of motion;

$$\overline{v} = -\rho g \frac{k}{\mu} \nabla h$$
 (3)

In Equations 1, 2a, 2b and 3, z is the reference elevation, h is the head,  $\delta$  refers to the local thickness of saturated portion of the aquifer,  $\vec{V}$  denotes the velocity, S refers to the storage coefficient, Q describes the volumetric flow rate per unit area,  $\rho$  is the density, g refers to the acceleration of gravity, k is the permeability,  $\mu$  is the absolute viscosity, and t denotes the time.

Equations 1 and 3 are combined and linearized to yield a single equation that, subject to appropriate boundary conditions, describes the dynamics of flow in the aquifer. The thickness of the aquifer is assumed to be small compared to its lateral dimensions. This equation is;

$$\nabla T \nabla h - S \frac{\partial h}{\partial t} - Q = 0 \qquad .... \qquad (4)$$

The quantities T and S are, respectively, the local transmissibility and storage coefficient of the aquifer. The source flow rate, Q, is in most cases time dependent. This flow rate is the algebraic sum of several component extraction and replenishment flows. The replenishment flows are precipitation, imported water, stream percolation, artificial recharge and subsurface inflow across boundaries. The extraction flows consist mainly of the water pumped from the aquifer for consumptive use and subsurface outflow across boundaries.

Herein, the Equation 4 is replaced by an equivalent system of difference - differential equations, the simultaneous solution of which gives the wanted function, h, at a finite number of node points lying within the boundaries of the aquifer.

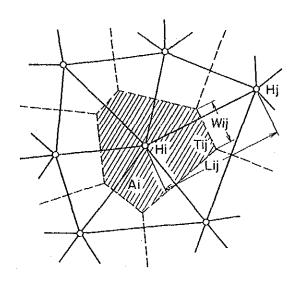
The attendant system of difference - differential equations is;

$$\sum_{j} (h_{j} - h_{i}) \cdot Y_{j,i} = A_{i} \cdot S_{i} \cdot \frac{dh_{i}}{dt} + A_{i} \cdot Q_{i}$$
and
$$Y_{j,i} = \frac{W_{j,i} \cdot T_{j,i}}{L_{j,i}}$$
.... (6)

in which  $A_i$  denotes the area associated with node i in sq $\cdot$ km,  $Y_j$ , is the conductance of pass between nodes j and i, in MCM per month per

meter,  $S_i$  is the storage coefficient of polygonal zone associated with node i,  $Q_i$  is the volumetric flow rate per unit area at MCM per month per sq\*km,  $T_{j,i}$  denotes the transmissibility at midpoint between nodes j and i, in MCM per month per meter,  $L_{j,i}$  refers to the distance between nodes j and i, and  $W_{j,i}$  describes the length of perpendicular bisector associated with nodes j and i, in km.

A typical node point, its associated polygonal zone, and contiguous node points are shown as follows:



The terms in the first equation of the Equation 6 are interpreted physically with the aid of the above figures as follows:

The left side represents the sum of the flows within the aquifer to the polygonal zone,  $A_{\underline{i}}$ , and the first term on the right side represents the rate of water storage within  $A_{\underline{i}}$ . The remaining term represents the extraction or replenishment flow from  $A_{\underline{i}}$ .

# 9-4 Digital Computer Solution

The Equation 6 is solved on the general purpose digital computer by an implicit numerical integration technique:

$$\sum_{i} (H_{i}^{t+1} - H_{i}^{t+1}) \cdot Y_{j,i} = \frac{A_{i} \cdot S_{i}}{\Delta t} (H_{i}^{t+1} - H_{i}^{t}) + A_{i} \cdot Q_{i}^{t+1} \dots (7)$$

in which the superscripts t denotes points along the time coordinate. This method of integration has the advantage that the magnitude of the time step,  $\Delta t$ , does not depend on a stability criterion.

The computation procedure is as follows:

- (1) Initial water levels,  $H_{H0i}$ , are impressed at the terminals labeled  $H_i^t$  (i=1,2,...,N; N=number of nodes)
- (2) Then, for a given set of coefficients L, W, T, A, S and Q, the values  $H_{\mathbf{i}}^{t+1}$  are implicitly determined at the end of a step in time,  $\Delta t$ .
- (3) Once determined, these value become the initial water elevations for the next succeeding step in time.
- (4) Consecutive numbers are given to all nodes and line segments connecting such nodes, as illustrated in Figure B-14 for the present condition of the groundwater basin as well as in Figure B-15 for the proposed condition (Three observation wells are added, as proposed, for the proposed condition).
- (5) All the source flow rates at each node (A·Q) are calculated first on the basis of findings obtained through hydrological analyses.
- (6) Then, all the node-to-node subsurface flows (0) and all the storage flows (ST) are calculated.
- (7) Next, all the flows (subsurface, storage and extraction) are balanced at each node by setting their sum equal to a residual term (RES).

(8) Water level elevation at each node is adjusted by the magnitude of the residual, attenuated by a relaxation coefficient. The relaxation coefficient, RELAX, is defined as;

$$RELAX_{i} = \frac{1}{\sum_{j} Y_{i,j} + \frac{A_{i} \cdot S_{i}}{\Delta t}}$$

- (9) After all the values of h have been adjusted in this manner, a sum is formed of the nodal residuals. This sum is compared with a threshold value (ERROR), which is defined as the maximum acceptable sum of nodal flow residuals at any time step.
- (10) If the sum of the residuals is less than or equal to the threshold value, the calculation of the values of h is complete for that time step. Otherwise the calculation is repeated as many times as is required to reduce the sum of residuals to a value less than or equal to the threshold value.

The digital computer programs for the groundwater basin model are written, provisionally, using the FORTRAN system. The FORTRAN programs developed and used for the project are attached at the end of the Appendix B-9, and a part of its corresponding flow chart is shown in Figure B-16.

### 9-5 Modeling of the Groundwater Basin

Based on the conception previously mentioned, the groundwater basin is divided into a number of polygonal zones as shown in Figure B-14 for the present condition of the basin before construction of the proposed detension dam, as well as in Figure B-15 for the proposed condition after construction of the dam.

#### 9-6 Input Data

Among the basic parameters to furnish the model, Falaj Use and base flow runoff in and from the catchment of the proposed dam, direct surface runoffs from the dam catchment as well as from residual area below the dam, and direct recharge in the area are referred to in the hydrological investigations.

Rainfal	1-Runoff	Relation

Areal	Rain	(mm)	Runoff	at Dam	(mm)	Runoff	at Rivermouth	(mm)
	8.0			0.0			*****	
•	10.0			0.13			· •	
٠	11.5			( - )			0.0	
• •	20.0			0.83			0.43	
	30.0			1.98			0.90	
	40.0			3.45			1.30	
* :	50.0	٠		5.23	,		1.70	
	60.0		•	7.29			2.08	
	70.0	•		9.74	7		2.47	
٠	80.0			12.55			2.82	

These values are obtainable from the areal rainfall in the areas above and below the dam, and hence the model requires spot rainfall records at Daqiq, Kitnah, Hayl (Wadi Jizzi), Hayl (Wadi Hayl), Farfar and Sohar to be inputted.

Groundwater recharges in the polygonal zones from the surface flow are estimated in proportion to the length of pass of percolation along the main course of the Wadi Jizzi.

. . .

Zone	No.	Length	Rate of			
Present	Proposed	of pass (km)	Percolation			
1	1	3.80	0.33			
2	2	2.05	0.18			
4	4	1.20	0.11			
5	5	1.10	0.10			
10	10	2,.35	0.21			
13	15	0.85	0.07			
	Total	11.35	1.00			

Seasonal variation of irrigation water requirement to be extracted from the groundwater storage is obtained from the study made for evaluation of irrigation water consumption.

Seasonal Variation of Irrigation Consumption

(Unit: MCM/month/1000ha)

Month	Consumption
January	0.220
February	0.140
March	0.283
April	0.336
May	0.468
June	0.438
July	0.419
August	0.396
September	0.371
October	0.327
November	0.258
December	0.226

The FAO method of estimating effective rainfalls is also used to examine the balance of rain waters. Table B-6 presents the monthly average crop evapo-transpiration so prepared to calculate effective rainfalls through the criteria provided in FAO Irrigation and Drainage Paper No.24.

The initial water level in each observation well, permeability coefficients, storativity, and topographic parameters such as thickness of the permeable layors are also referred to in the findings of hydro-geological investigations, as shown in Figures B-17 to B-19.

#### 9-7 Groundwater Simulation

Four cases of the groundwater simulation study are conducted as under;

Case No.	Description of Simulation Condition
0	Present condition with water extraction of 11.0
	MCM/yr for irrigation and 0.2 MCM/yr for
	domestic use.
1	Proposed condition with Project, with extraction
	same as in the present conditon.
2	In addition to the Case 1, additional water
	extraction of 1.06 MCM/yr for domestic uses
	(in total 1.26 MCM/yr) is considered.
3	In addition to the Case 2, additional water
	extraction of 1.39 MCM/yr for irrigation purpose
•	is considered.

All computations are progressed for the entire period of available data of 11 years from 1974 up to 1984 employing a monthly time interval. The data used, simulated result of the groundwater balance and well water levels, and the simulated water levels in terms of the deviation value, with their 12-months moving averages, are presented in Figures B-21 to B-23 and Tables B-7 to B-10, Figures B-24 to B-26 and Tables B-11 to B-14, Figures B-27 to B-29 and Tables B-15 to B-18, Figures B-30 to B-32 and Tables B-19 to B-22, respectively for the Cases 0, 1, 2, and 3.

Figure B-23 is compaired with Figure B-2 for the purpose of model verification, and the followings may bring a conclusion.

- As compaired with observed movement of groundwater levels, the simulated behavior of groundwater levels almost follows its actual pattern, with exception in and after 1982.
- ii) The above differences may be caused by the differences in the amount of water extraction between the actual field performance and the estimated water consumption of a uniform rate throughout a number of years.
- iii) The model parameters and input data may be modified and updated after several years of study and experience to simulate more accurate behavior of the groundwater movement.
- iv) Consequently, the model simulation will be useful to predict the groundwater behavior under the inposed condition, subsequently making the model of a useful tool for the management of groundwater resources.

Figure B-20 summarizes the simulated changes of groundwater storage for the respective computation cases, resulting in a conclusion that the Project would provide the required water demand associated with the Case 3 without producing any undesired result on the groundwater resources, and proving that an expected value of permeability of  $2.0 \times 10^{-2} \frac{1}{}$  m/min. and storativity of  $0.05\frac{2}{}$  are, as a whole, reasonable.

 $\underline{1}/$ ,  $\underline{2}/$ : The values of permeability and storativity were computed by using the data of pumping up tests.

Table B-6 Average Crop Evapo-Transpiration (For calculation of effective Rainfall)

Average	2835	64.0	68.8	97.4	115.3	152.4	143.6	138,3	130,5	126.8	106.1	72.2	63.8
Sorghum	101	6.68	75.6	87.4	1	ı	I	ı	1	1	54.3	63.0	6.68
Alfalfa	211	76.4	85.7	123.8	153.0	202.9	193.8	187.1	176.5	160.6	131.7	89.3	76.4
Others	23	76.4	9.08	87.4	ı	I	ì	ı	1	i .	I	42.0	62.9
Okra	8	ı	40.3	102.0	153,0	191.0	136.8	1	1	ì	ì	1	1
Potato	15	121.4	126.0	ì	1	1	ı	ı	ŀ	ì	ı	47.2	7.76
Tomato	16	76.4	100.8	123.8	126.0	155.2	ı	ı	2	189.0	155.0	47.2	7.67
Garlic	25	71.9	80.6	116.6	126.0	119.3	t	1	1	T.99	77.5	73.5	67.4
Onion	96	58.4	1 -	ı	I	1	ı	1	ı	1.89.0	155.0	73.5	67.4
Mango	83	67.9	70.6	102.0	126.0	167.1	159.6	154.1	145.4	132.3	108.5	73.5	62.9
Banana	128	6.08	7.06	131.1	162.0	214.8	205.2	198:1	186.9	170.1	139.5	94.5	80:9
Lime	309	6.44	50.4	72.8	0.06	119.3	114.0	110.1			77.5	52.5	6.44
Dates	1820	62.9	70.6	102.0	126.0	167.1	159.6	154.1	145.4	132.3	108.5	73.5	62.9
Crop	Area(ha)	Jan.	Heb.	Mar.	Apr.	May	Jun.	Jul.	Aug.	Sep.	Oct.	Nov.	Dec.

Note: Unit = mm/month

Figure B-14 Modeling of Groundwater Basin
(Present Condition)

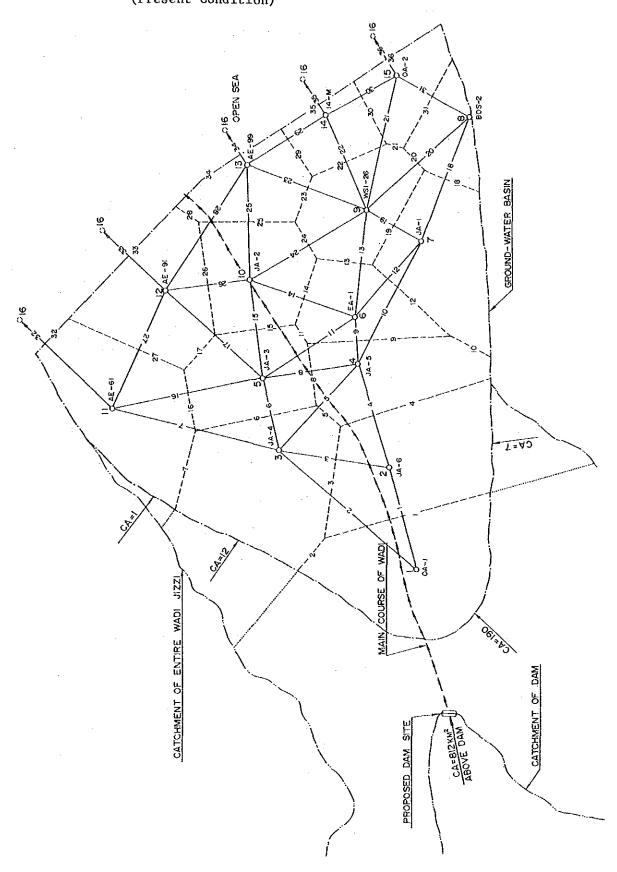


Figure B-15 Modeling of Groundwater Basin (Proposed Condition)

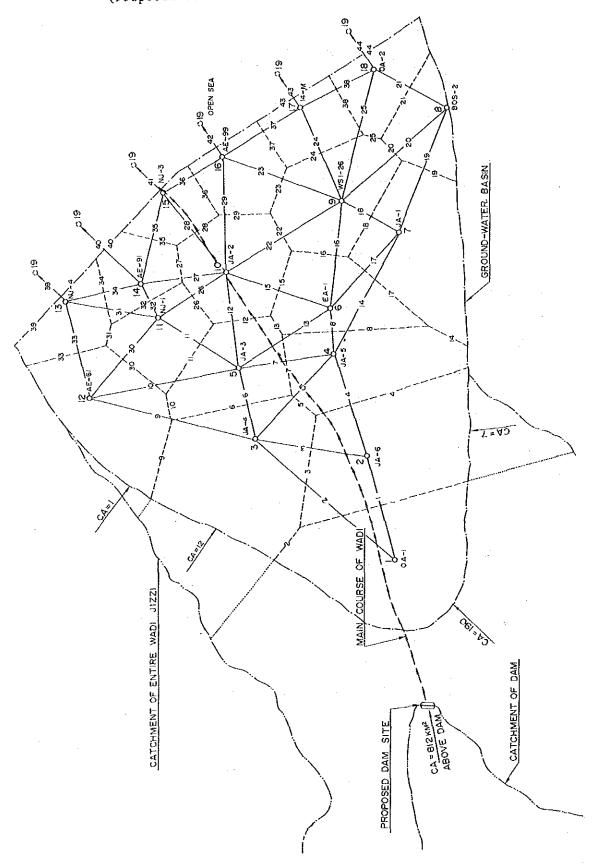
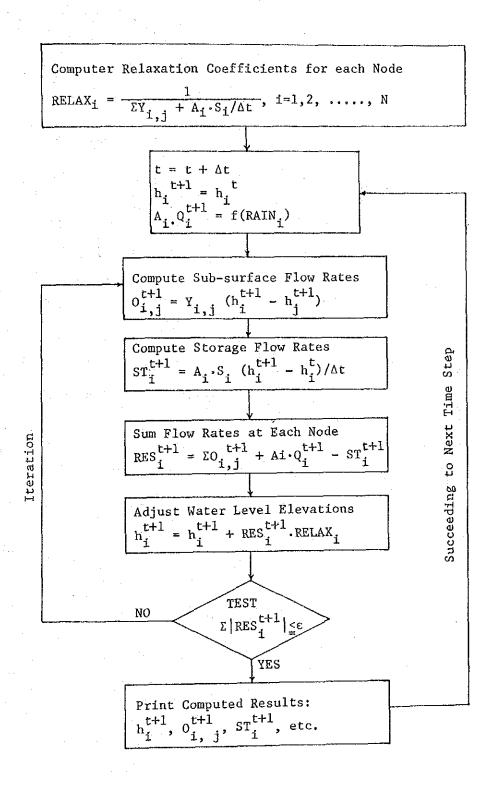
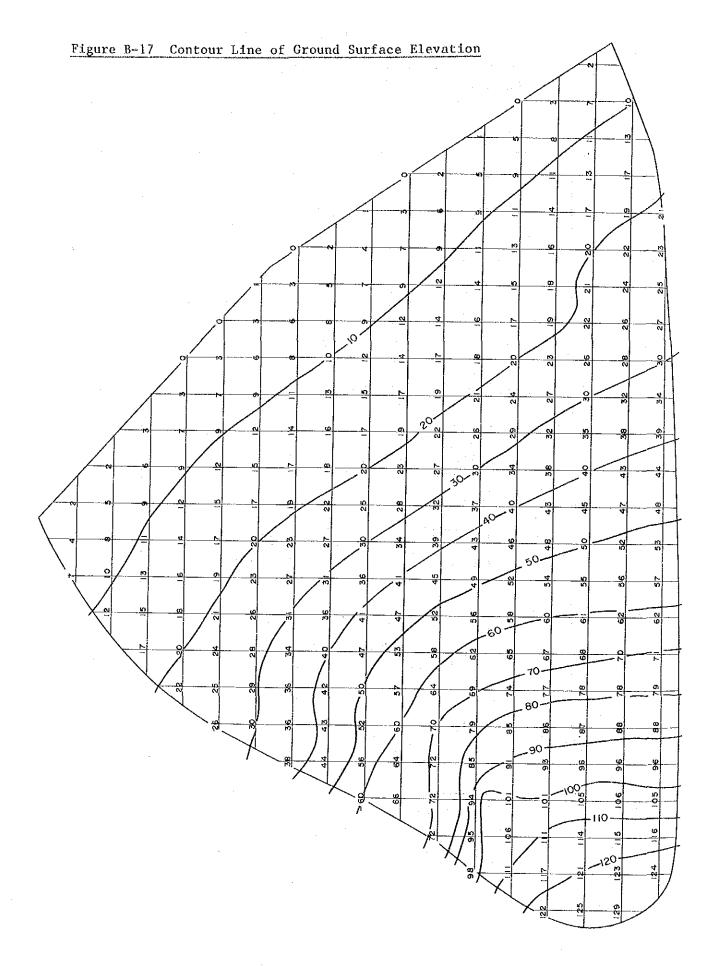
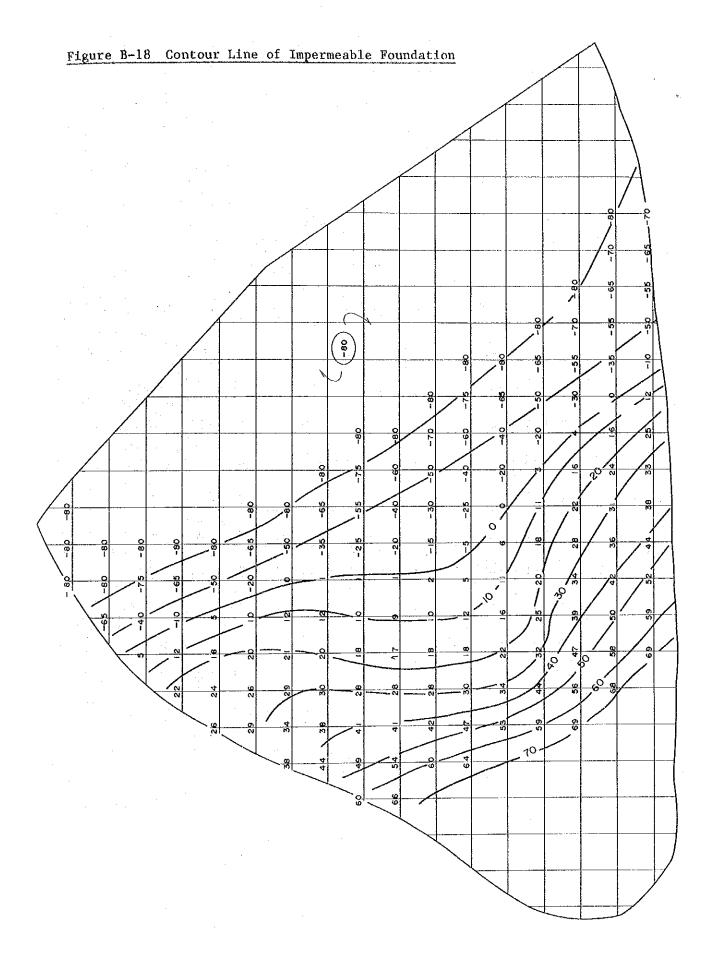


Figure B-16 Simplified Flow Chart for Digital Computer Solution of Ground-Water Problem

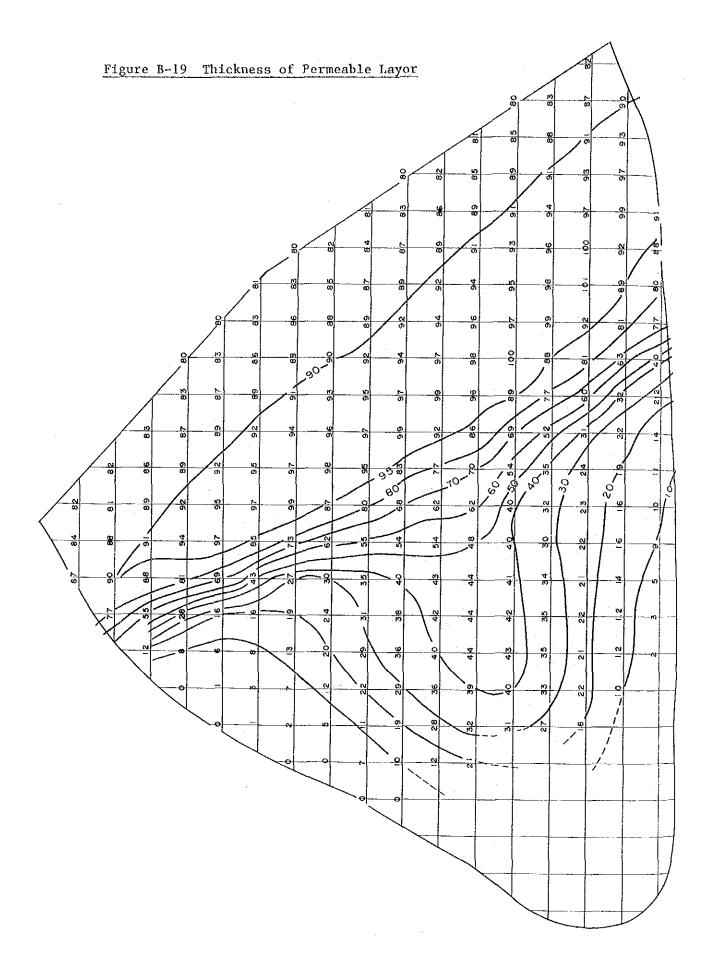




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Figure B-20 Simulated Change of Groundwater Storage

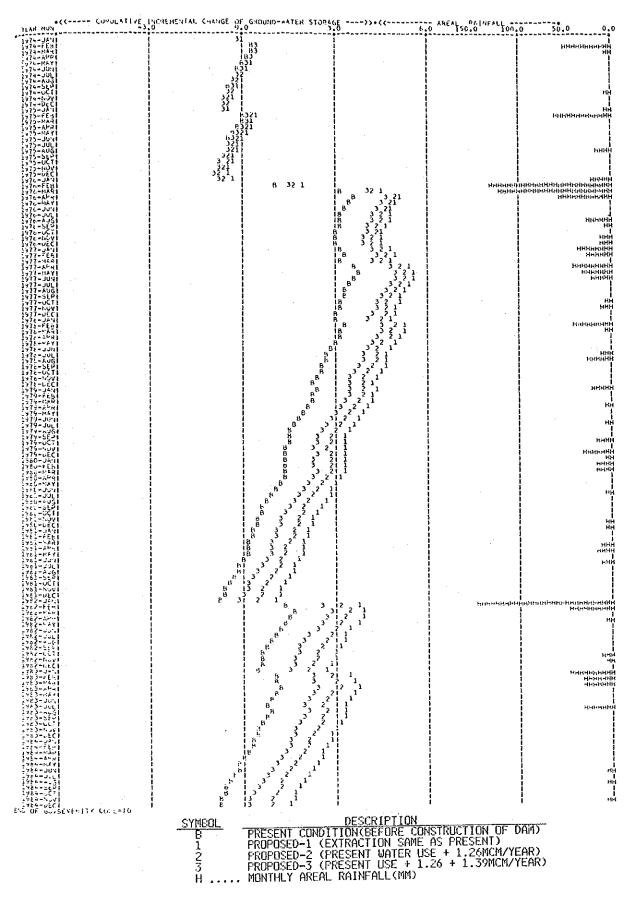


Figure B-21 Simulated Groundwater Tables (Case 0)

WELL SO, (1) (2) (3) (4) (5) (6) (7) (8) (9) (10) (11) (12) (13) (14) (15) (16) SYMBOL 1 5 9 C SYMBOL 3 7 A E SYMBOL 4 8 8 F WELL JA-5 BOS-2 AE-91 SEA WELL JA-6 EA-1 JA-2 14-M WELL JA-4 JA-1 AE-61 OA-2

Figure B-22 Deviation of Groundwater Level (Case 0)

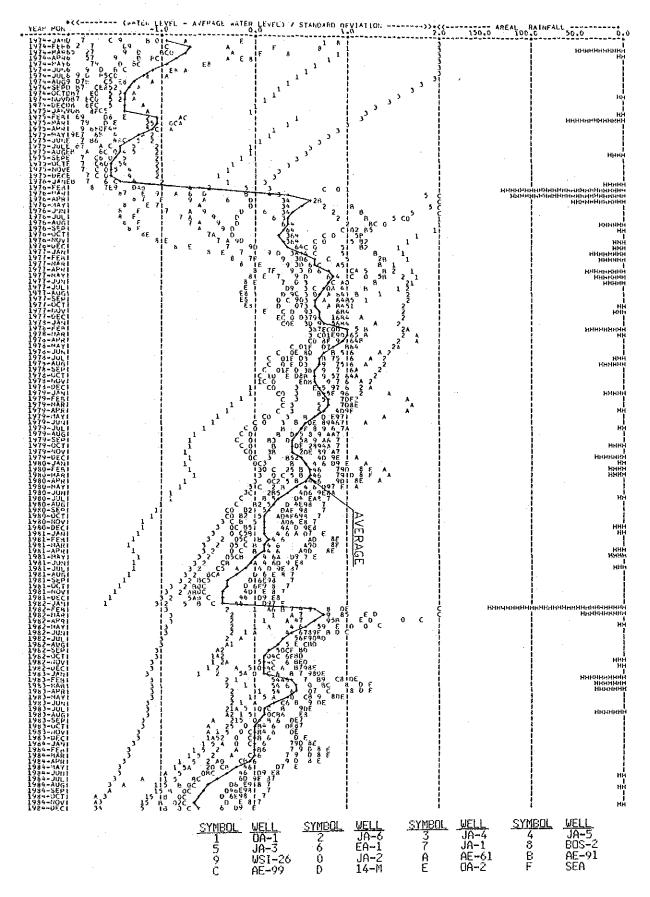


Figure B-23 Deviation of Groundwater Level (Case 0)
(12-months Moving Average)

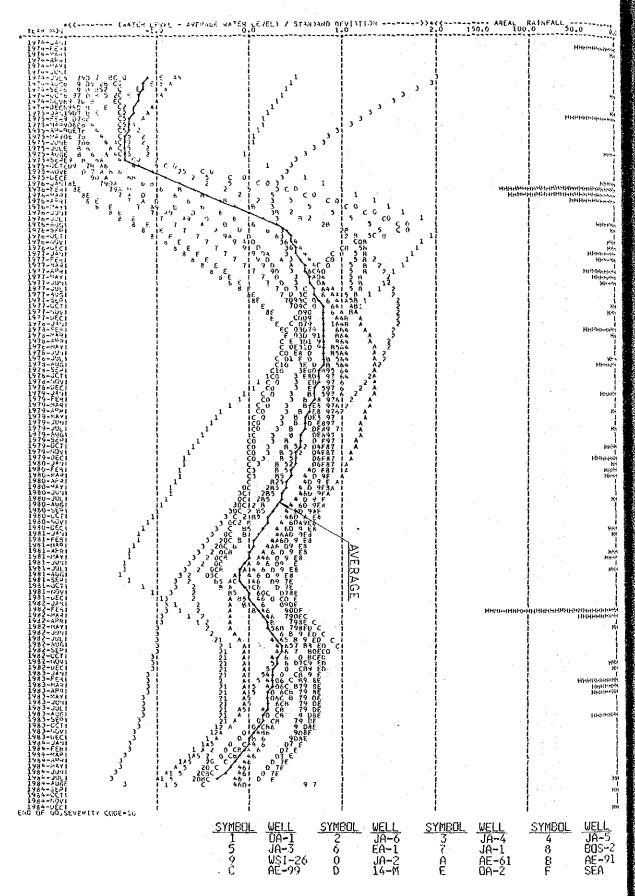


Figure B-24 Simulated Groundwater Tables (Case 1)

# 돌아내이다. ( 1) ( 3) ( 3) ( 4) ( 2) ( 6) ( 1) ( 8) ( 3) (10) (11) (12) (12) (12) (12) (12) (13) (13) (13) WELL JA-5 BOS-2 AE-61 AE-99 SYMBOL 2 6 0 D H WELL JA-4 JA-1 NJ-1 NJ-3 SEA WELL DA-1 JA-3 WSI-26 NJ-4 14-M <u>SYMBOL</u> 4 8 B F SYMBOL 1 5 9 C G

Figure B-25 Deviation of Groundwater Level (Case 1)

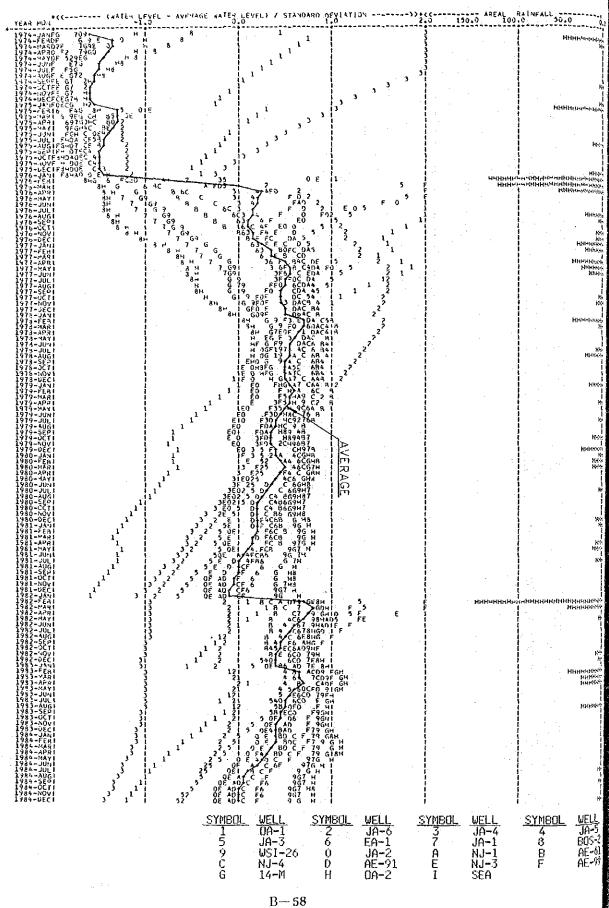


Figure B-26 Deviation of Groundwater Level (Case 1)

